

TINY RHIZOMORPHIC ROOTING SYSTEMS FROM THE EARLY PERMIAN ABO FORMATION OF NEW MEXICO, USA

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Premise of research. Extant *Isoetes* species, which all develop small pseudoherbaceous habits, are the only living remnants of the once diverse clade of rhizomorphic lycopsids, which included trees that grew to towering heights of 50 m. Although the rhizomorphic lycopsids evolved a range of diverse body plans in the Paleozoic, it is thought that the evolution of the small pseudoherbaceous habit, with small rooting systems similar to modern *Isoetes* species, was a late event in the clade's history, occurring in the Mesozoic. Here we describe small fossilized rooting systems from the early Permian Abo Formation of New Mexico, increasing our knowledge of the diversity of small Paleozoic rhizomorphic rooting systems.

Methodology. Ten fossilized rooting systems are described from a slab collected from the early Permian Abo Formation of New Mexico.

Pivotal results. Here we report 10 rooting systems interpreted as those of rhizomorphic lycopsids due to the presence on each of a large number of isotomously branched rootlets radiating from a central rhizomorph and two associated microphyll leaf impressions. Because of the fossils' tiny size (the largest rhizomorph is only 1.5 cm in diameter, and the largest rooting system, including rootlets, is only 6 cm in diameter), we interpret these fossils either as juvenile plants or as adult morphologies with a small rooting system. Given the paucity of mature rhizomorphic lycopsids in the Abo Formation and the limited fossil record of juvenile rhizomorphic lycopsids from any geological period, we predict that these are most likely to be adult plants with small rooting systems.

Conclusions. The small size of the specimens described here increases our knowledge of the diversity of small rhizomorphic rooting systems in the Paleozoic before the radiation of the modern *Isoetes* growth habit during the Mesozoic.

Keywords: *Isoetes*, lycopsids, Permian, roots, rhizomorph.

Introduction

The rhizomorphic lycopsids (DiMichele and Bateman 1996) are a hugely diverse group of plants with an evolutionary history spanning the mid-Devonian (ca. 387 Ma) to the present day (Phillips and DiMichele 1992; Pigg 1992, 2001; Skog and Hill 1992; Stewart and Rothwell 1993; Grauvogel-Stamm and Lugardon 2001; Xu and Wang 2016). Although extant *Isoetes* are small pseudoherbaceous plants only centimeters in height (Paolillo 1963), their Paleozoic relatives could grow to colossal heights of 50 m (Thomas and Watson 1976; Stewart and Rothwell 1993) and were anchored to the substrate by giant

rooting axes extending 15 m from the trunks of large trees and covered in rootlets (Williamson 1887; Stewart and Rothwell 1993; Hetherington et al. 2016). The group reached its ecological dominance during the Carboniferous Period (ca. 360–300 Ma) when many species were key components of wetland forests during the Early and Middle Pennsylvanian (Phillips and DiMichele 1992; DiMichele 2014).

After the Carboniferous, rhizomorphic lycopsids still played key roles in terrestrial ecosystems (Pigg 1992, 2001; Skog and Hill 1992; Stewart and Rothwell 1993; Grauvogel-Stamm and Lugardon 2001), but they did not achieve the same stature or ecological importance as they had earlier. In a general sense, therefore, the evolution of the rhizomorphic lycopsids is marked by a decrease in size from the giants of the Paleozoic that were tens of meters high, to the small arborescent forms in the Mesozoic that were only a couple of meters high (e.g., *Pleuromeia*), and finally to the forms that resembled modern *Isoetes* (e.g., *Isoetites*), which have been only a few centimeters high (Ash and

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Pigg 1991) from the Mesozoic to the present. This decrease in stature is seen most strikingly with the decrease in stem elongation and the transition of the rooting system from large axial branching rhizomorph axes (Williamson 1887; Frankenberg and Eggert 1969) to both unbranched axial axes and cormose forms (such as extant *Isoetes* species; Paolillo 1963; Rothwell and Erwin 1985; Pigg 1992, 2001; Grauvogel-Stamm and Lugardon 2001).

We now know that the evolution of the rhizomorphic lycopsids was more complex than previously assumed and did not involve a direct reduction series from giant arborescent trees to small pseudoherbaceous forms (reviewed by Pigg 1992). Instead, the fossil record preserves evidence of the coexistence of lycopsids with diverse growth habits and of both unbranched axial and cormose rhizomorphs having earlier origins than previously thought (Phillips and Leisman 1966; Pigg and Rothwell 1979; Jennings et al. 1983; Pigg and Taylor 1985; Pigg 1992, 2001; Xu and Wang 2016). However, despite the diversity of growth habits known from the Paleozoic, it appears that the origin of the small pseudoherbaceous habit typical of modern *Isoetes* was a comparatively late innovation. Plants with a habit similar to extant *Isoetes* with cormose rhizomorphs of only a few centimeters in diameter and reduced stem axial elongation are known only from the Triassic (Pigg 1992, 2001; Retallack 1997; Grauvogel-Stamm and Lugardon 2001). This late origin of the pseudoherbaceous growth habit of *Isoetes* is important as it hints that plants with modern growth habits, and therefore the ecologies that accompany them, evolved only in the Mesozoic despite a long Paleozoic history of the clade. The Mesozoic rhizomorphic lycopsids may therefore be thought of as “weedy survivors” of the Permian-Triassic extinctions (Retallack 1997). Here we report the discovery of small fossilized rhizomorphic rooting systems from the lower Permian of New Mexico. These rooting systems increase our knowledge of the diversity of rhizomorphic lycopsids with small rooting systems from the Paleozoic, before the evolution of the modern growth habit of *Isoetes* during the Mesozoic.

Material and Methods

Geological Locality

In central New Mexico, siliciclastic red beds of the lower Permian Abo Formation (Wolfcampian to early Leonardian) record fluvial deposition during the last stages of the ancestral Rocky Mountain orogeny (Lucas et al. 2014). In the Cerros de Amado–Joyita Hills area of Socorro County, the Abo Formation contains a significant record of continental plant fossils (DiMichele et al. 2013), vertebrate fossils (Berman et al. 2015), and prolific assemblages of red bed (*Scoyenia* ichnofacies) trace fossils representing both vertebrates and invertebrates (Hunt and Lucas 2015).

The slab documented here was collected in the Cerros de Amado–Joyita Hills area. It is housed in the New Mexico Museum of Natural History and Science (NMMNH) in Albuquerque, New Mexico, and catalogued as NMMNH P-25887. It was collected at NMMNH locality 3250, which is a sandstone bed stratigraphically high in the upper member (Cañon de Espinoso Member) of the Abo Formation, ca. 8 m below the overlying Arroyo de Alamillo Formation of the Yeso Group (fig. 1).

Given its stratigraphic position, NMMNH locality 3250 is almost certainly of early Leonardian age (Lucas et al. 2015). Other fossils found at, and near, NMMNH locality 3250 are impressions of walcian conifers and tetrapod footprints (fig. 1).

DiMichele et al. (2013) reviewed the Abo Formation paleoflora based on collections made from ca. 200 localities. This paleoflora is of low diversity and dominated by fossils of walcian conifers and of the peltasperms *Supaia*.

Results

The slab containing the fossils is roughly rectangular (ca. width 28 cm, height 21 cm) and is ca. 1 cm deep (fig. 2). The most prominent features on the surface of the slab are 10 roughly circular, and in some cases overlapping, structures (fig. 2B, labeled 1–10). Because of their apparent in situ position on the bedding plane, we interpret the structures as rooting systems. Each rooting system consists of a similar overall organization (figs. 2, 3): roughly circular and ca. 4–6 cm in diameter (based on measurement of the specimens that are completely preserved; fig. 2B, 1–5, 8, 10) and protruding slightly above the surface of the bedding plane. At the center of each rooting system is a roughly circular region of ca. 1–1.5 cm in diameter that either protrudes farther from (fig. 2B, 1, 3–5, 7, 8; fig. 3) or is indented into (fig. 2B, 2, 10; fig. 3) the rock surface. The central region, which we will refer to as the rhizomorph (fig. 3), is not fully preserved (fig. 2B, 6, 9). Where the central rhizomorph forms a depression in the surface, there are no further defining features (fig. 2B, 2, 10; fig. 3A, 3B). By contrast, in the examples where the central rhizomorph protrudes above the rock surface, further detail is preserved (fig. 2B, 1, 3, 4, 7, 8; figs. 3C, 3D, 4). The rhizomorph (fig. 3C, 3D) appears as a raised tier and is the highest part of the rooting system above the bedding surface (fig. 4B, 4E, 4I). The details of the central rhizomorph are not well preserved. However, in each case the region is marked by an angular indent (fig. 3C, 3D) that is either triangular (fig. 4C, 4F, 4G) or a four-point-star shape (fig. 4J) at the center.

Radiating in all directions from each central rhizomorph is a large number of frequently branched linear structures, preserved as indentations in the rock surface; these structures are darker in color than the surrounding matrix. These linear structures are blunt tipped. They are relatively uniform in diameter between branch points and decrease in diameter with each dichotomy, which is isotomous. The central rhizomorph lacks these branched structures. Because of their point of initiation at the periphery of the central rhizomorph and their highly branched nature, we interpret these linear structures as rootlets (fig. 3). The lateral extent of the rootlets marks the periphery of the circular rooting systems (fig. 3). Rootlets frequently branch as they radiate from the central rhizomorph, and one rootlet developed at least three orders of branching (fig. 2B, black arrow; fig. 5A, 5B). Rootlet networks may overlap where they occur at high densities on one plant or where they emanate from plants preserved close to each other (fig. 2B, 3–5, 7–9). The rootlets radiate in all directions from the rhizomorph, forming a concentric circle (fig. 3). Unfortunately, the connection of the majority of rootlets cannot be traced back to the rhizomorph, making it impossible to determine the structure of the meristem from which they developed or their rhizotaxy.

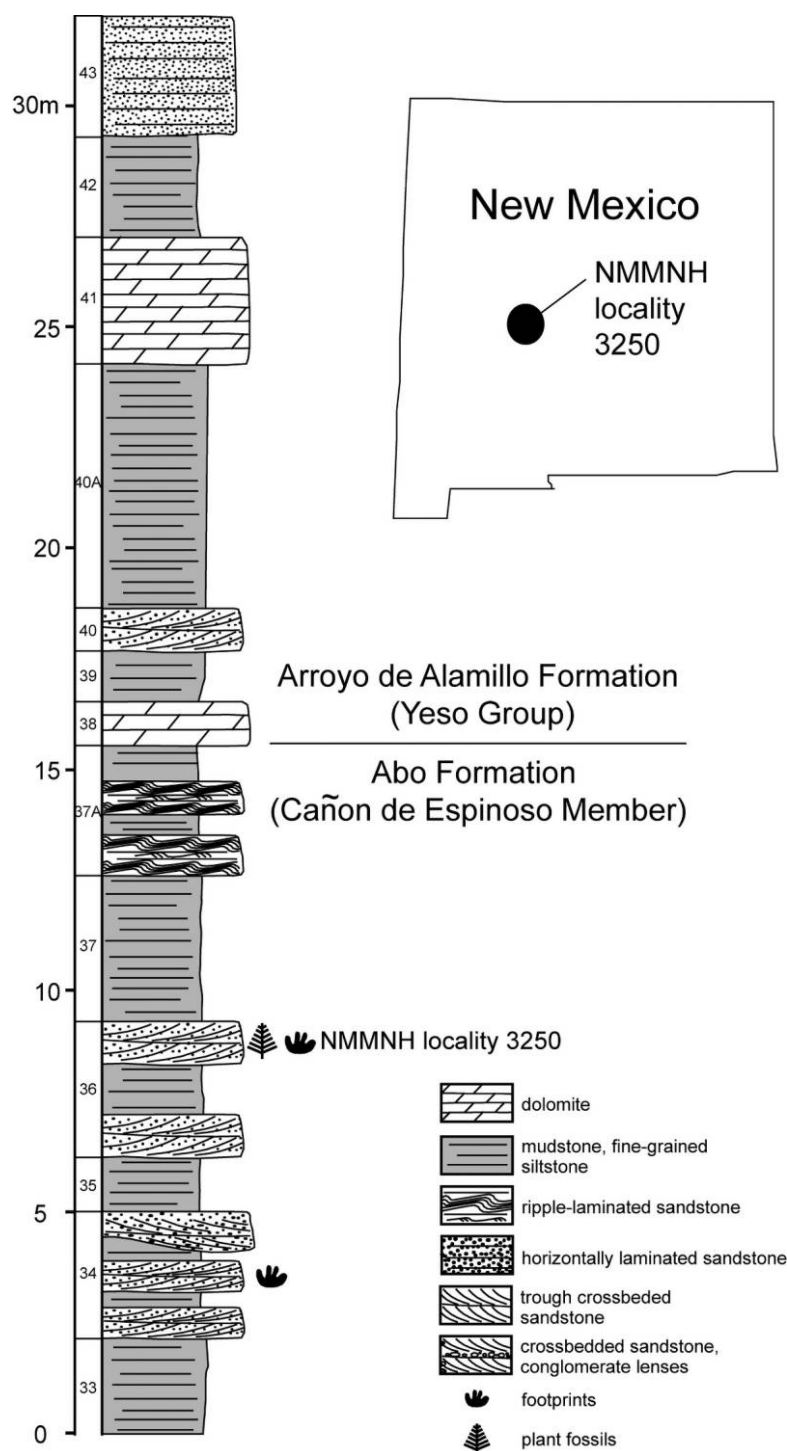


Fig. 1 Stratigraphy of the Abo Formation (New Mexico) showing the location of NMMNH locality 3250 where the slab NMMNH P-25887 was collected.

The lower surface of the slab displays no circular rooting structures. The regions that directly correspond to the rooting structures on figure 2A (labeled 1–10 on fig. 2B) do not show evidence of the continuation of the rhizomorph or of the surrounding network of rootlets. This suggests that the rooting sys-

tems described in figure 2A do not penetrate the depth of the slab.

In summary, the rooting systems preserved on the slab are roughly circular overall and consist of two parts: first, a roughly circular central rhizomorph that varies in its preservation be-

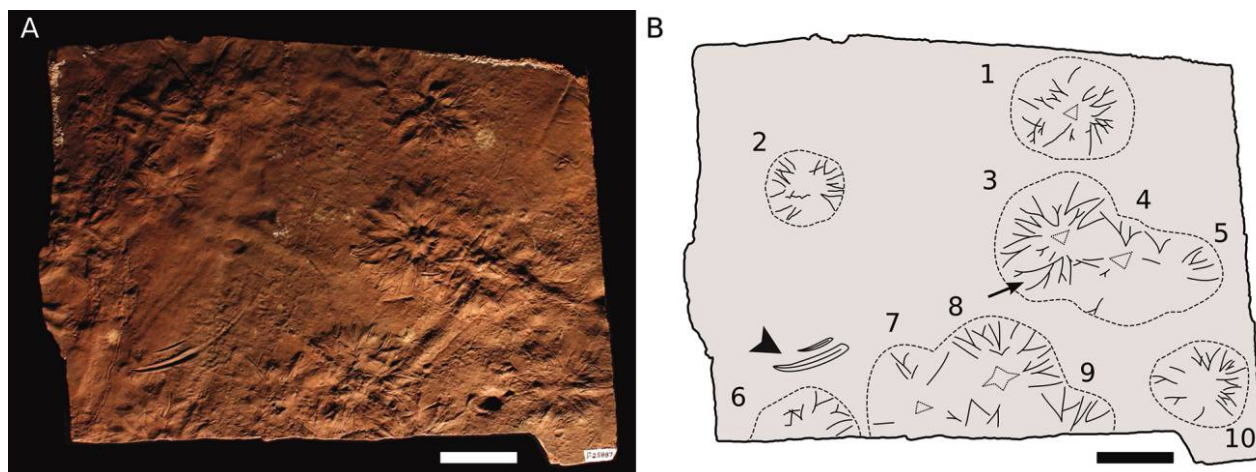


Fig. 2 Permian rhizomorphic rooting systems. *A*, Surface of the slab NMMNH P-25887, showing evidence of 10 roughly circular rooting systems and two associated leaf impressions. *B*, Schematic of the major features of the bedding surface, showing impressions of 10 roughly circular rooting systems (1–10). Note that rooting systems 6 and 9 are only partially preserved. The impressions of two leaves shown in figure 5C, 5D are highlighted with an arrowhead. The branched rootlet shown in figure 5A, 5B is highlighted with an arrow. Lighting: *A*, diffuse lighting from multiple directions with predominant lighting from bottom right of the image. Scale bars = 4 cm.

tween the specimens and, second, radiating from this central rhizomorph, a large number of isotomously branched rootlets (fig. 3).

In addition to the rooting systems preserved on the slab (fig. 2A), there are impressions of two tapered linear structures with slightly curved shapes and pointed tips, which we interpret as leaves (highlighted with an arrowhead in figs. 2B, 5C, 5D). Both leaf impressions have a similar structure, although they differ in size (the larger leaf is ca. 3.5 cm long, the smaller leaf ca. 1.5 cm; fig. 5C, 5D). Both are marked at the center by a central linear impression running the length of the leaf (fig. 5C, 5D). This central keel is consistent with the leaf having a single central vascular trace, a characteristic feature of the microphylls of rhizomorphic lycopsids. Apart from these detached microphylls, there is no other evidence of the anatomy of the above-ground shoot systems.

Discussion

The diagnostic feature of the rooting systems described here is a large number of isotomously branched rootlets radiating in all directions from a central rhizomorph (fig. 3). Roots that branch strictly isotomously are a characteristic feature of both extinct and extant lycopsids (Hetherington and Dolan 2017). Thus, the highly isotomously branched rootlets support the hypothesis that the fossils are the rooting systems of lycopsids. The prediction of lycopsid affinity is further supported by the associated leaf impressions (fig. 5C, 5D), both of which possess a single central keel consistent with a single leaf trace running along the center of the leaves. This trace is a diagnostic feature of the microphylls of lycopsids (Kenrick and Crane 1997); leaves with this morphology are not typical of any other group of plants from the Euramerican Pennsylvanian or Permian age. Therefore, by association, these leaves support the lycopsid affinity of the rooting systems. A large number of rootlets, developed in high density from a central rhizomorph, is a charac-

teristic feature of the rhizomorphic lycopsids (DiMichele and Bateman 1996). We therefore interpret these fossilized rooting systems as having affinities with that group of plants. Given that the central rhizomorphs show no evidence of branching (figs. 2–4), we interpret the central rhizomorphs as either unbranched and axial or cormose in structure. Furthermore, because the central rhizomorphs do not penetrate the slab, this suggests that these may have been shallowly buried rooting systems. Rhizomorphic lycopsids that developed axial or cormose rhizomorphs are known to both pre- and postdate the early Permian age of these fossils (Phillips and Leisman 1966; Pigg and Rothwell 1979; Jennings et al. 1983; Pigg and Taylor 1985; Pigg 1992, 2001; Skog and Hill 1992; Stewart and Rothwell 1993; Retallack 1997; Grauvogel-Stamm and Lugardon 2001; Xu and Wang 2016). However, these rooting systems are of interest due to their particularly small size.

The most striking feature about the rooting systems preserved on the slab, when compared to the rooting systems of many other rhizomorphic lycopsids, is their small size. The maximum diameter of the largest rhizomorph is only 1.5 cm, and the maximum diameter of the largest rooting system, including rootlets, is only 6 cm. These rooting systems are, therefore, much smaller than most of the rhizomorphic rooting systems known from the Carboniferous (Jennings et al. 1983; Pigg and Rothwell 1983; Pigg 1992, 2001). The closest similarity in size is the small rhizomorph described for *Cormophyton mazonensis*, with a width of 2.6 cm (Pigg and Taylor 1985; Pigg 1992). In their size, the Abo Formation specimens are similar to rooting systems of the pseudoherbaceous forms known from the Triassic onward (Pigg 1992, 2001; Retallack 1997; Grauvogel-Stamm and Lugardon 2001).

We propose two explanations for the small size of these plants: either they are juvenile plants of much larger rhizomorphic lycopsids or they are adult rhizomorphic lycopsids with small rooting habits. If they are juvenile plants, they represent some of the few examples discovered that preserve the early ontogeny

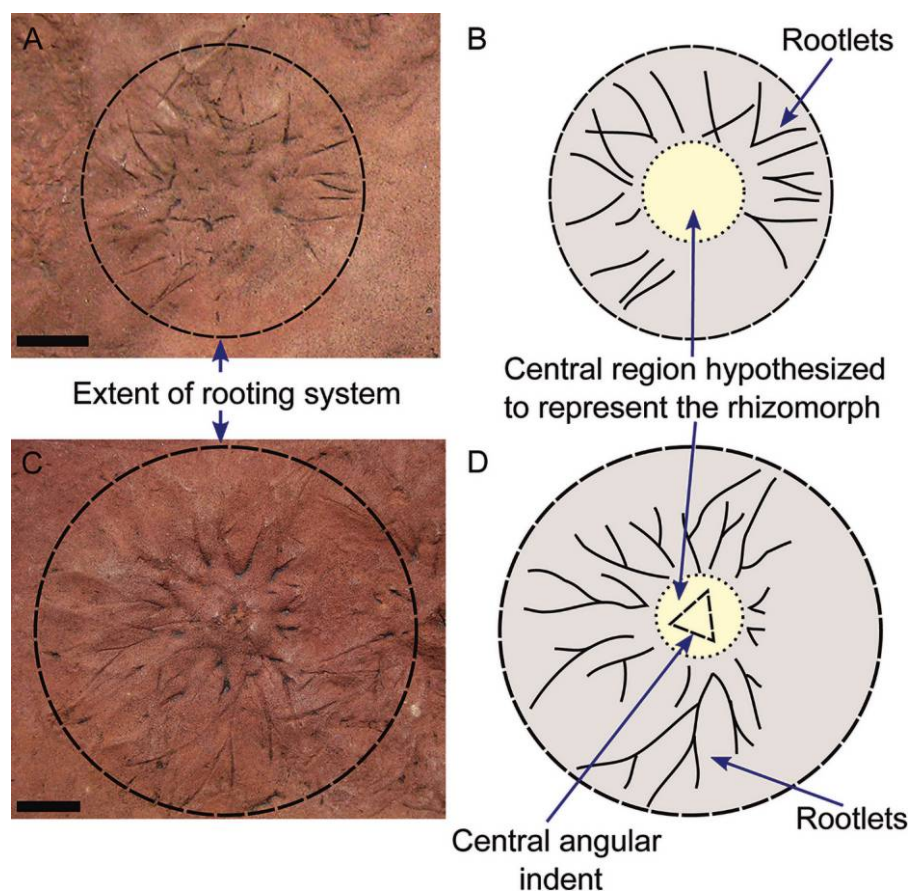


Fig. 3 Interpretive diagram of rhizomorphic rooting systems. A, Image of specimen 2 from figure 2B with the extent of the rooting system highlighted. B, Interpretive diagram of A, highlighting the lateral structures interpreted as rootlets and the central region hypothesized to represent the rhizomorph. C, Image of specimen 3 from figure 2B with the extent of the rooting system highlighted. D, Interpretive diagram of C, highlighting the lateral structures interpreted as rootlets and the central region hypothesized to represent the rhizomorph, including the central angular indent. Lighting: A, C, vertical lighting from above. Scale bars = 1 cm.

of rhizomorphic lycopsids (Karrfalt 1984; Boyce and DiMichele 2016), connecting developing spores (Phillips 1979) with larger mature specimens (MacGregor and Walton 1948; DiMichele and DeMaris 1987; Stewart and Rothwell 1993). The lack of arborescent lycopsid remains (in which periderm and wood would more likely be preserved than in immature specimens) in the Abo Formation (DiMichele et al. 2017), combined with the elusive nature of juvenile rhizomorphic lycopsids in the fossil record (Boyce and DiMichele 2016), makes the juvenile-arborescent hypothesis the less likely of the two options. Alternatively, if the fossils are adult plants, they increase our knowledge of Paleozoic rhizomorphic lycopsids with small rooting systems. Unfortunately, the rhizomorph morphology is not well preserved, making it impossible to determine whether these plants had a cormose or axial rhizomorph. Furthermore, although the rhizomorphs do not extend through the depth of the slab, implying that they were shallowly buried, we have no evidence of the nature of the aboveground shoot system. Based on these lines of evidence, we predict that these plants were shallowly buried adult morphologies with small rhizomorphic rooting systems.

Despite the poor preservation of the fossils, it is possible to speculate on the structure of their rhizomorphs. In support of

the rhizomorphs being unbranched and axial is the observation that the rootlets form a concentric circle around the roughly circular region interpreted as the rhizomorph (figs. 2–4). If the plants developed unbranched axial rhizomorphs, they may have had a growth habit similar to *Paurodendron fraipontii* (Phillips and Leisman 1966; Rothwell and Erwin 1985). Alternatively, it is possible that these rooting systems are small cormose forms more typical of the Mesozoic (Pigg 1992, 2001; Grauvogel-Stamm and Lugardon 2001). Very tentative evidence for the cormose nature of these fossils comes from the triangular and star-shaped marks at the center of the rhizomorphs (figs. 3, 4C, 4F, 4G, 4J). The base of a three-lobed rhizomorph such as *Isoetes nuttallii* (Karrfalt and Eggert 1978) or of a four-lobed rhizomorph such as *Pleuromeia sternbergii* (Grauvogel-Stamm 1993) would leave a triangular or star-shaped impression, respectively. However, this interpretation seems unlikely as we would expect to find evidence of the meristematic furrows within the triangular or star-shaped indentation (Karrfalt and Eggert 1978; Grauvogel-Stamm 1993). Until further specimens are discovered, it is not possible to reach conclusions about the nature of the rhizomorphs or the complete interpretation of the habit of these plants.

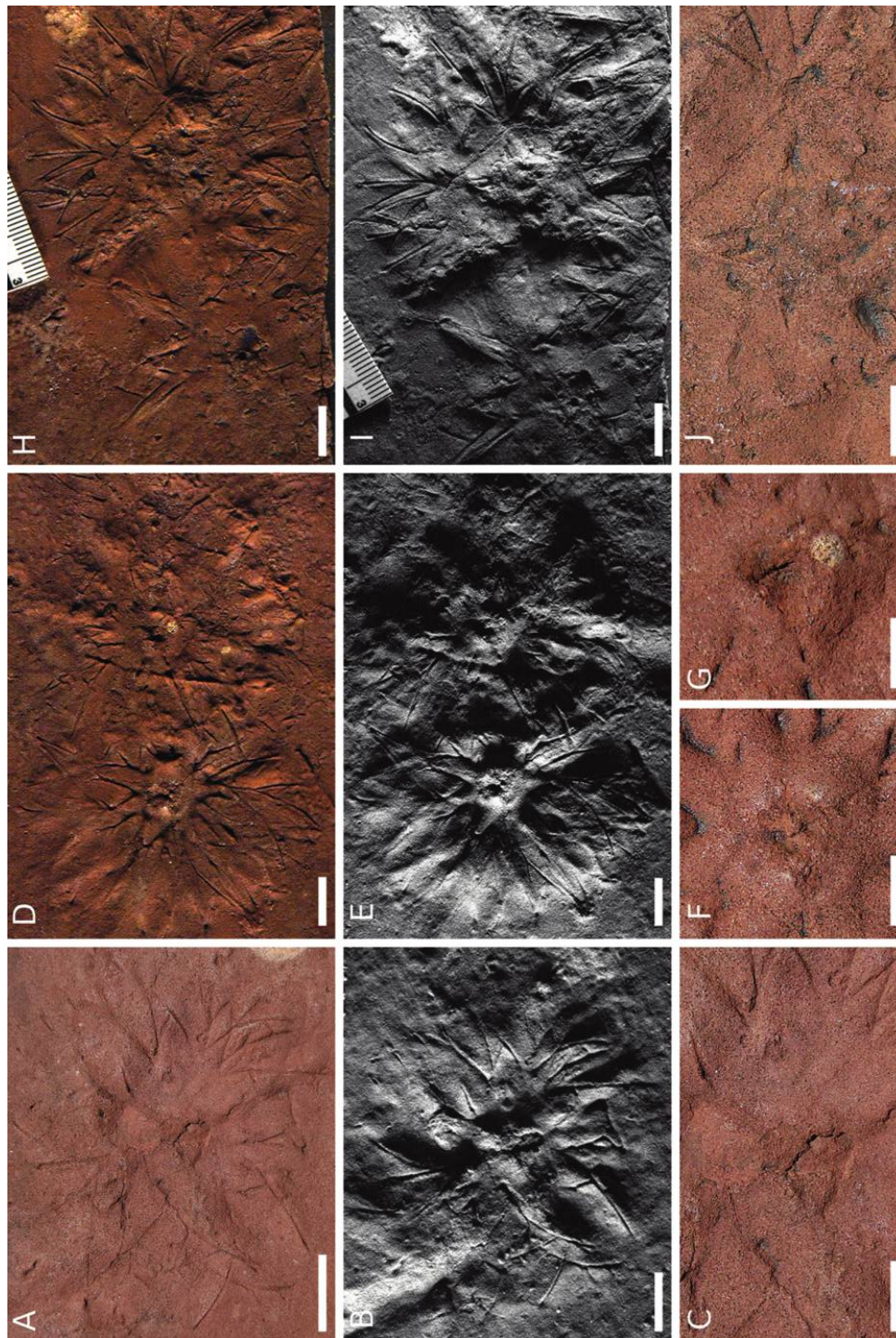


Fig. 4 Details of well-preserved rhizomorphic rooting systems. A–C, Images of specimen 1 from figure 2B. D–G, Images of specimens 3 and 4 from figure 2B. H–J, Images of specimens 7 and 8 from figure 2B. A, D, H, Images illustrating the overall shape of the rooting systems including the central rhizomorph and surrounding network of rootlets. B, E, I, Low-angle-light images highlighting the topography of the rooting systems—in each case the rooting system appears as a raised mound on the surface of the slab. C, F, G, J, Higher-magnification images of the center of the rhizomorphs of specimens 1, 3, 4, and 8, respectively. Lighting: A, C, F, G, J, vertical lighting from above; B, E, I, low-angled light from left (B, E) and right (I); D, H, diffuse lighting from multiple directions. Scale bars = 1 cm (A, B, D, E, H, I), 5 mm (C, F, G, J).

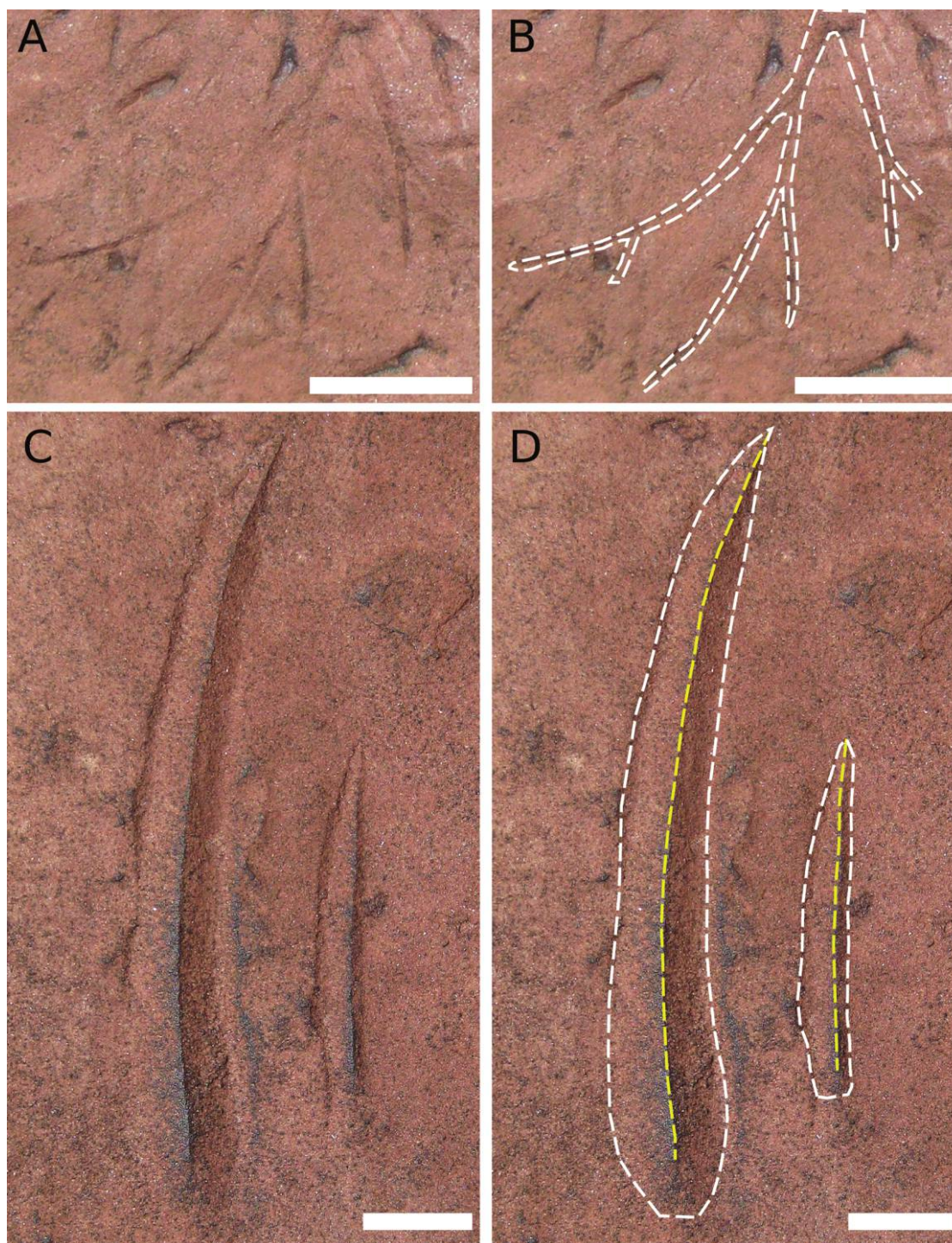


Fig. 5 A, Highly branched rootlet highlighted with an arrow in figure 2B. B, Same image as A with dashed white lines highlighting the outline of the branched rootlet. C, Impression of two microphylls highlighted with an arrowhead in figure 2B. D, Same image as C with dashed white lines highlighting the outline of the microphyll and dashed yellow lines highlighting the position of the central keel. Lighting: vertical lighting from above. Scale bars = 1 cm.

Despite the limited preservation of these small rhizomorphic rooting systems, their discovery is intriguing as it suggests that a cryptic diversity of rhizomorphic lycopsids with small rooting systems was present during deposition of the Abo Formation.

This cryptic diversity may reflect, in part, the low preservation potential and high likelihood of early erosion of so much of the terrestrial landscape, permitting a great deal of evolutionary history to escape capture in the geological record (Looy et al.

2014). The Abo Formation itself has been intensely sampled for plant, animal, and ichnofossils for several decades, and just recently, the first occurrence of the calamitalean foliage *Annularia* was discovered there (DiMichele and Lucas 2017). This type of foliage is very common and widely distributed throughout the Euramerican tropics during the Pennsylvanian and Permian, and yet it remained unknown in the Abo Formation until this discovery, demonstrating the complexities of rarity in nature versus rarity in the fossil record. The findings reported here as well as the previous discovery of *Annularia* (DiMichele and Lucas 2017) highlight the importance of continued fieldwork to document the entirety of a fossil flora in the Abo Formation.

Conclusions

Here we report the discovery of a slab containing 10 rooting systems preserved in early Permian strata. Based on the diagnostic anatomy of the rooting systems (a large number of isotomously branched rootlets radiating from a structurally complex central region; fig. 3) and associated leaf impressions (with a single leaf trace per leaf; fig. 5), we interpret the rooting systems as belonging to members of the rhizomorphic lycopsids. The central rhizomorphs show no evidence of branching, leading us to conclude that the rooting systems were from plants that produced either cormose or unbranched axial rhizomorphs.

Although both cormose and axial rhizomorphs pre- and post-date the occurrence of these fossils, it is the size of the fossils that makes them interesting. The small size of the specimens indicates that these plants were either juveniles of larger tree forms or mature plants with small rooting systems. Given the absence of mature, arborescent rhizomorphic-lycopsid remains in red beds of the Abo Formation and the elusive nature of juvenile rhizomorphic lycopsids in the fossil record, we suggest that these remains most likely represent mature plants of a very small size. The small size of these specimens increases our knowledge of small rhizomorphic rooting systems in the Paleozoic before the origin of the small Mesozoic forms (Pigg 1992, 2001), as well as broadening the list of taxa present in the Permian flora of the Abo Formation.

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