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3 **The temporal and spatial scale of rocky coast geomorphology: A Commentary**
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26 **Abstract**
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28 Rocky shores are complex landforms which are the result of marine erosion and subaerial
29 weathering. They are time-integrated features where their present day form is the result of
30 instantaneous erosion, often on the millimetre to sub-metre scale, occurring for centuries to
31 millennia. As a result, research on rocky coasts focuses on a range of temporal and spatial scales
32 from granular-scale swelling of a rock surface and instantaneous wave impact to modelling
33 millennial-scale sea level drivers. The challenge for rocky coast researchers is to either upscale or
34 downscale their results to human-timescales of greatest interest to managers. The research
35 presented in Earth Surface Processes and Landforms over the past 3 years highlights the range of
36 spatial and temporal approaches to the study of coastal cliffs and shore platforms. We identify a key
37 temporal and spatial gap in current research. Seasonal – annual timeframes over hundreds of metres
38 to kilometre scale studies appear to be lacking and are likely critical in understanding the future
39 evolution of rocky coasts, especially their response to climate change.
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45 **Introduction**
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47 Rocky coasts are erosional landscapes. Their form at a particular instant in time is a function of how
48 they resist erosive processes as they retreat landwards. Erosion is driven by marine processes acting
49 on a surface formed in materials that have been weakened through subaerial weathering. The
50 inherent strength of the lithology in which the coast is formed provides a critical boundary condition
51 on the potential landscape that can be formed. That is, rock mechanical strength properties will
52 determine the size of the steep cliff edifice that can form, as well as the presence and
53 geomorphologic character of fronting shore platforms.
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3 The timescale on which these processes occur is highly variable as also is the spatial location on
4 which they act within the coastal zone. For example, wave energy on a microtidal shore platform
5 tends to be concentrated at its seaward edge (Ogawa et al., 2014) with subaerial weathering of
6 greater importance where wetting and drying cycles are most frequent and where the mediating
7 effects of biological activity may be greatest (Gowell et al., 2015; Moses et al., 2015). These
8 erosional processes occur over the instantaneous scale (minutes – hours) with waves removing
9 material already weakened by subaerial weathering or directly plucking blocks from the cliffs and
10 platform surface. Erosion over the instantaneous scale for millennia leads to the current form
11 observed on the coast.
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15 The fundamental difficulty in understanding rock coast evolution directly relates to the issue of the
16 scale of the driving processes and the age of landform systems. This is because numerous feedbacks
17 exist between erosive and weathering processes making it difficult to reconcile field measurements
18 with current geomorphic form. For this reason, geomorphologists have resorted to a range of
19 techniques from time-space substitution to theoretical models to explain landform evolution.
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22 The research reported in Earth Surface Processes and Landforms (ESPL) over the past 3 years on
23 shore platforms and sea cliffs highlights the range of scales over which these landforms systems
24 operate. In this special issue of ESPL we explore the evolution of rock coast systems at a range of
25 spatial and temporal scales (Figure 1).
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28 29 30 **Instantaneous Scale**

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32 At the shortest temporal scale, wave breaking occurs and is the primary energy input that both
33 removes material loosened by subaerial weathering as well as directly eroding unweathered rock. In
34 the work of Ogawa et al. (2015), they show that wave energy transfer across a semi-horizontal
35 platform is a function of the relative submergence of the platform edge. While wave energy on the
36 edge was a function of offshore conditions, at the base of the landward cliff energy was independent
37 of the deep water wave climate and more a function of the water depth on the platform surface.
38 The energy expenditure of waves breaking at the rear of the platform is not however solely confined
39 to the platform itself. Microseismic energy is transferred into the landward cliff with cliff top
40 vibrations being correlated to significant wave height and water maximum depth at the cliff base
41 (Young et al., 2016). While wind-generated waves of gravity to infragravity frequency are the
42 primary energy input, the instantaneous impact of longer period tsunami waves also can drive
43 landscape formation in some circumstances. Such is observed in the erosional scars on the platforms
44 and cliffs of the Maltese Islands (Mottershead et al., 2015).
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49 While waves are the dominant erosive agent, the resistance of most rock surfaces is weakened by
50 subaerial weathering and erosion to become more sensitive to wave breaking and water movement.
51 While rock plucking and the resulting movement of the erosion products can be observed in situ on
52 platforms by direct surveying with, for example, unmanned aerial vehicles (Pérez-Alberti and
53 Trenhaile, 2015), granular-scale changes are more difficult to visualise. The very small-scale wearing
54 of rock surfaces can be accurately quantified with micrometers and microerosion meters . Such
55 methodologies have become more common in recent decades, yet upscaling annual measurements
56 to timeframes of significance for landform shape is difficult. Only a few studies have attempted to
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3 monitor microerosion for longer than a few years, and in this special issue Moses et al. (2015)
4 present one of the first decade-length records of downwearing from a tropical carbonate shore. At
5 the small granular level scale, weakening of the rock surface can be influenced by many factors,
6 particularly the amount of wetting and drying. Wetting and drying cycles in turn can be affected by a
7 wide range of factors, ranging from waves to cover by macroalgae. For example, laboratory
8 experiments by Gowell et al. (2015) simulating the influence macroalgae on mudstone platforms
9 found a reduction in the rate of mineral breakdown by as much as 79%, attributed to thermal
10 insulation and moisture retention at the rock surface. It is due to many local environmental factors
11 and feedback loops that microerosion metre-derived downwearing cannot be simply extrapolated
12 indefinitely into the future. This is because as a rock surface lowers it will support different biotic
13 communities and experience greater tidal inundation which will all affect wetting and drying cycles
14 and other subaerial processes. It therefore remains a challenge for the discipline to extrapolate
15 short-term field observations and laboratory experiments to the millennial scale over which many
16 rock landforms are created.
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23 **Event Scale**

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25 Infrequent weather events can also cause major landform change at the sub-metre scale. This was
26 observed on the coast of Mesnil-Val, in NW France, where a sudden frost event in 2009 caused
27 temperatures to drop to -9.5° , causing widespread frost shattering of the shore platform (Dewez et
28 al., 2015). In this study the authors found that six episodes of frost led to platform lowering of $0.8 \pm$
29 0.5 mm in a single winter. Such extreme events highlight how low frequency events can lead to high
30 magnitude of change in the system at some locations.
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34 Cliff failure is another example of the complexities of the mismatch between observational
35 timescales and those of landform evolution. Cliff collapse leads to a sudden landward shift in the cliff
36 top of many metres. The frequency of such events is not necessarily high, which means a single
37 event cannot be used to characterise the evolutionary trajectory of a cliffed shore. New statistical
38 techniques and utilisation of catalogues of observed cliff failure now provide great promise in
39 reconciling these different timescales. In this Special Issue Letortu et al. (2015) developed an
40 inventory of 331 individual landslides between 2002 and 2009 on the Upper Normandy coast of
41 France. A thorough statistical analysis of the boundary conditions for each of these events allowed
42 “a starting point for the prediction and prevention of the hazard of coastal chalk cliff rock falls” (p
43 1371). Baldassarri and Sapoval (2015) take the analysis further using Power Law statistics. Their
44 theoretical approach yields important results as it highlights that many different physical
45 phenomena can lead to the same type of statistical results – an important consideration for
46 predicting rocky coast hazards where a range of boundary conditions can determine final landscape
47 form.
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53 **Geomorphic/Geological Scale**

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55 The classic approach to understanding long term rocky coast evolution has been to use space-for-
56 time substitution. Long sections ($>$ kilometre-scale) are studied to quantify the variation in
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3 geomorphology to provide inferences on what drives land formation. Such large-scale studies can
4 vary boundary conditions through investigations of sites of different geology, tide or wave climate.
5 By confining a study to a specific boundary condition, while varying others, it is possible to both
6 define the natural variation within a system as well as how changes in boundary conditions can
7 determine the evolutionary pathway of a system. The spatial scale of these studies tends to be
8 determined by the specific research question. For example Bini et al. (2014) investigated a hundred-
9 metre stretch of the Argentinian coast to specifically quantify the relationship of erosive notches to
10 sea level. Kennedy et al. (2014) on the other hand focussed on long term landform evolution of
11 granitic rocky coasts, with a field site almost 100 km long in order to vary wave exposure while
12 keeping the boundary conditions of rock type and tidal range constant. These two papers also
13 highlight the methodological advances within the discipline. Bini et al. (2014) used 3D point clouds
14 derived from field photographs. Kennedy et al. (2014) on the other hand utilised terrestrial and
15 bathymetric aerial LiDAR combined with high resolution multibeam sounding. Such technological
16 advances are providing a revolution in the study of rocky coasts.
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21 Numerical modelling is a key technique for reconciling the different spatial and temporal scales of
22 rock coast evolution. In this Special Issue there is a focus on sea stack evolution (Limber and Murray,
23 2015) and cliff retreat (Carpenter et al., 2015) over centuries to millennial timeframes. Trenhaile
24 (2014) further extends the timeframe of investigation by modelling rocky coasts over millions of
25 years, back to the Pliocene. Modelling techniques have the advantage of being able to test
26 landscape response through experimentally modelling each boundary condition systematically. They
27 do however require to be grounded in field data, which is often difficult as there are still large
28 swathes of the Earth's coastline that have yet to be systematically described.
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34 **Future Directions**

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36 A key temporal and spatial gap in research is evident in the literature and also appears in this special
37 issue; studies over seasonal – annual timeframes for hundreds of metres to kilometre scale appear
38 to be lacking. Some studies do include these spatial and temporal scales within their wider
39 frameworks, yet research specifically focussed on how rocky coast systems behave at this scale
40 appears to be lacking. A reason for this is likely that current methodological approaches are not
41 suited for measuring change at this scale and that landform response is difficult to quantify (Figure
42 2). However, for models to be applicable, and for weathering and wave impacts to be upscaled in
43 order to link the instantaneous and the long-term, a continuum of research over the entire spatial
44 and temporal scale of land formation is required.
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50 **References**

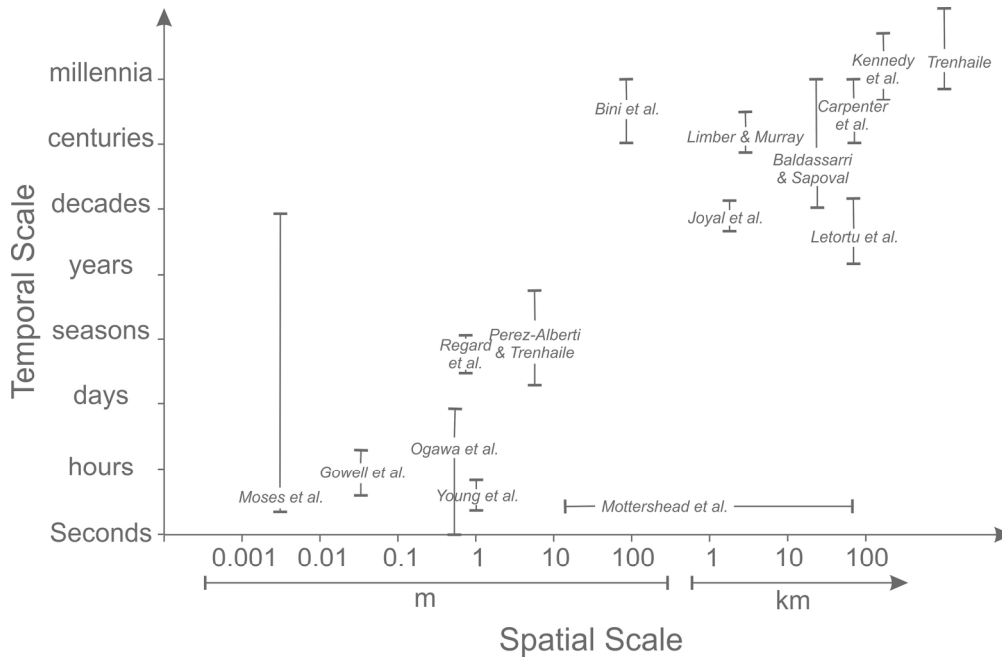
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47 **Figure Captions**

48 Figure 1: The spatial and temporal scale of articles published on rocky coasts in *Earth Surface*
49 *Processes and Landforms* from 2014 – 2017.

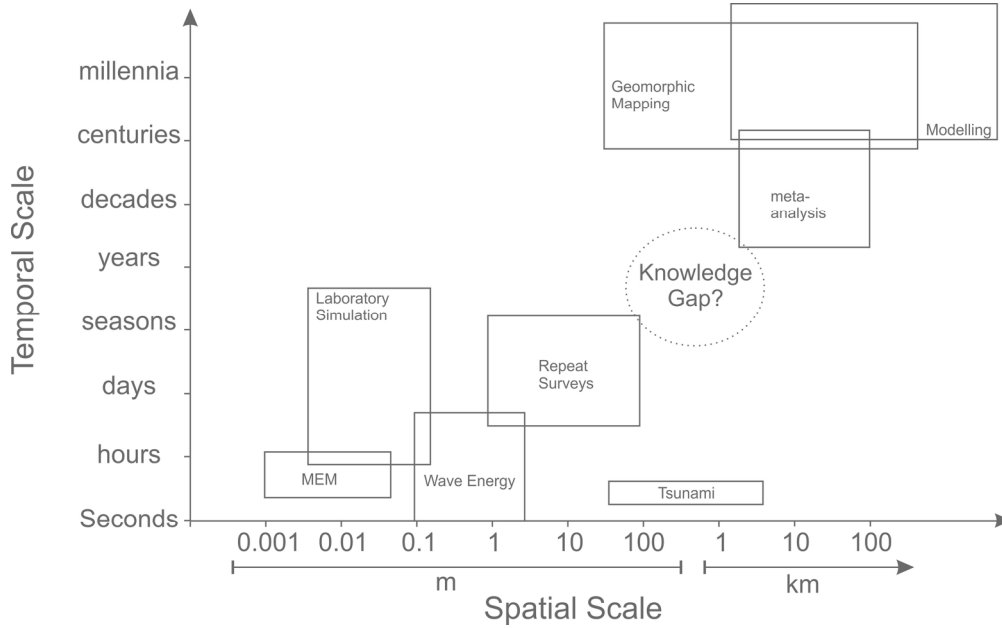
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52 Figure 2: The temporal and spatial scale of the primary methodological approaches used to research
53 cliffs and shore platforms. A gap in knowledge appears at approximately the kilometre scale over
54 annual timescales.
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The spatial and temporal scale of articles published on rocky coasts in *Earth Surface Processes and Landforms* from 2014 – 2017

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The temporal and spatial scale of the primary methodological approaches used to research cliffs and shore platforms. A gap in knowledge appears at approximately the kilometre scale over annual timescales.

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