



Can homogeneous nucleation resolve the inner core nucleation paradox?

Alfred J. Wilson^{a,*}, Dario Alfè^{b,c,d}, Andrew M. Walker^e, Christopher J. Davies^a

^a School of Earth and Environment, University of Leeds, Woodhouse, Leeds, LS2 9JT, UK

^b Department of Earth Sciences, University College London, 5 Gower Place, London, WC1E 6BS, UK

^c London Centre for Nanotechnology, Thomas Young Centre, University College London, 17–19 Gordon Street, London, WC1H 0AH, UK

^d Dipartimento di Fisica “Ettore Pancini”, Università di Napoli “Federico II”, Monte S. Angelo, Napoli, 80126, Italy

^e Department of Earth Sciences, University of Oxford, S Parks Rd, Oxford, OX1 3AN, UK



ARTICLE INFO

Article history:

Received 9 February 2023

Received in revised form 18 April 2023

Accepted 20 April 2023

Available online 16 May 2023

Editor: J. Badro

Keywords:

inner core

nucleation

molecular dynamics

ABSTRACT

The formation of Earth's solid inner core is thought to mark a profound change in the evolution of the deep Earth and the power that is available to generate the geomagnetic field. Previous studies generally find that the inner core nucleated around 0.5–1 billion years ago, but neglect the fact that homogeneous liquids must be cooled far below their melting point in order for solids to form spontaneously. The classical theory of nucleation predicts that the core must be undercooled by several hundred K, which is incompatible with estimates of the core's present-day temperature. This “inner core nucleation paradox” therefore asserts that the present inner core should not have formed, leaving a significant gap in our understanding of deep Earth evolution. In this paper we explore the nucleation process in as yet untested iron-rich systems which may comprise the Earth's early core. We find that 1 mol.% Si and S increase the supercooling required to freeze the inner core compared to pure iron by 400 K and 1000 K respectively. 10 mol.% O reduces the required inner core nucleation supercooling to 730 K and 3 mol.% C to only 612 K, which is close to resolving the paradox but still requires that the inner core formed recently.

Crown Copyright © 2023 Published by Elsevier B.V. This is an open access article under the CC BY license (<http://creativecommons.org/licenses/by/4.0/>).

1. Introduction

The Earth's magnetic field is produced by the geodynamo in the liquid outer core. The majority of the convective power which drives the present dynamo is from inner core growth (Labrosse, 2015; Nimmo, 2015a; Davies, 2015), where light elements partitioned to the liquid create a positive buoyancy anomaly at the innermost outer core. This field shields the Earth's surface from potentially harmful space weather events and solar radiation. Palaeomagnetic records suggest that the geodynamo could have been extant for at least the last 3.4 Gyrs (Tarduno et al., 2010). Prior to inner core growth, the geodynamo must have been powered by other means such as secular cooling, radiogenic heating or precipitation of light elements (O'Rourke and Stevenson, 2016; Hirose et al., 2017; Badro et al., 2018; Wilson et al., 2022). Because it presents such a fundamental change in regimes, the nucleation of the inner core is perhaps the most important event in the thermal history of the core and might present an observable signature in the palaeomagnetic record (Biggin et al., 2015; Bono et al., 2019; Zhou et al., 2022; Davies et al., 2022). Despite this the age of

the inner core is unknown and controversy over the thermal conductivity of the core has led to a wide range of inner core age estimates (e.g. Nimmo, 2015b; Labrosse, 2015; Driscoll and Davies, 2023) from 1 Ga to 500 Ma.

Adding to the controversial timing of inner core formation, a more recent problem has come to light. Theory and atomic scale simulations predict that there is a substantial barrier to the formation of new solid in liquid iron under core conditions (Huguet et al., 2018; Davies et al., 2019; Wilson et al., 2021) that means substantial supercooling is expected to be needed before inner core formation. Such large supercooling has consistently be found to be required for nucleation in metallic systems both in extremely large simulations and experiment (e.g. Shibuta et al., 2017; Kelton et al., 2006).

Classical nucleation theory (CNT, e.g. Christian, 2002) describes the thermodynamics of nucleation and states that for a liquid to freeze it must be supercooled. This is because whilst the liquid will be thermodynamically unfavourable compared to the solid for a system below its melting temperature, the interface between the first solid and the remaining liquid comes with an energetic penalty. Only when a critical nucleus size is exceeded will the energetic preference for the solid phase outweigh the energetic penalty due to the interface. Nuclei which grow larger than this

* Corresponding author.

E-mail address: a.j.wilson1@leeds.ac.uk (A.J. Wilson).

will become increasingly likely to continue to grow, leading to the system freezing. Huguët et al. (2018) used CNT to describe the supercooling needed to freeze liquid iron in the core. Applying existing calculations of interfacial energy and enthalpy of fusion of iron (Zhang et al., 2015) to define the balance of energies, the authors found that a supercooling on the order of 1000 K was needed for spontaneous nucleation. Huguët et al. (2018) estimated the maximum permissible supercooling to be 220 K by finding the largest feasible present-day separation of isentrope and melting curve at the centre of the Earth whilst preserving an intersection at the inner core boundary (ICB). Even satisfying this supercooling would imply that the inner core froze rapidly and recently, raising additional problems for the thermal evolution of the deep Earth. It is this mismatch between the predicted and allowed supercooling of the core which is the inner core nucleation paradox.

CNT is the most commonly applied framework for defining nucleation, but does not match observations completely. CNT predicts that the most stable solid phase should form from the supercooled liquid, but theoretical and experimental studies of metallic systems at ambient pressure have shown that local ordering in the liquid produces quasi-crystals first (Schenk et al., 2002; Kelton et al., 2003, 2006; Shibuta et al., 2017). These observations are predicted by the Frank hypothesis (Frank, 1952) which suggests that these structures nucleate more easily than the pure crystal predicted by CNT and act as a catalyst to nucleation in general. Wilson et al. (2021) and Sun et al. (2022) examine the nucleation of pure Fe in the core and both conclude that an intermediary phase precedes the formation of the thermodynamically stable hexagonally close packed (hcp) phase. This shows that the Frank hypothesis applies to high pressure and temperature systems as well. In our previous study (Wilson et al., 2021) we test the application of CNT at large supercooling at to supercooling relevant to inner core nucleation. We found that whilst some of the assumptions of CNT are not valid and can be discarded, pseudo-classical nucleation theory (so called for the inclusion of the Frank hypothesis) is sufficient for describing nucleation under core conditions. Because of this, we continue with this approach here.

Following the discovery of the inner core nucleation paradox, several different approaches using atomic scale simulation have been used to examine the problem in hopes of a resolution. Davies et al. (2019) directly observed homogeneous nucleation in molecular dynamic simulations of Fe and FeO systems at extreme supercooling and extrapolated results to Earth-like conditions, confirming the existence of the paradox. Others have probed the relevant conditions with molecular dynamic simulations of pure Fe to characterise the size distribution of sub-critical (those which re-melt) nucleation events (Wilson et al., 2021). Both approaches, find that these simple systems reproduce the original prediction of Huguët et al. (2018) with a 675–807 K supercooling requirement for spontaneous homogeneous nucleation of the inner core. A metadynamic approach has suggested that metastable iron phases may lower the nucleation barrier in a two-step nucleation process. Sun et al. (2022) show that a nucleus of body-centred cubic iron with low interfacial energy could produce a smaller free energy barrier than hexagonally close packed, although this metastable phase has not been reported in molecular dynamic studies of the same systems (Davies et al., 2019; Wilson et al., 2021).

All prior studies have focused almost exclusively on pure iron systems with the exception of Davies et al. (2019) who found $\text{Fe}_{0.9}\text{O}_{0.1}$ had little effect on the required supercooling due to almost equal but opposite effects of reducing the nucleation barrier and melting point. The composition of the core is to be expected far more complex, with an overall density deficit of $\sim 10\%$ attributed to dissolved light elements (Anderson, 2002) and the density contrast between outer and inner core requiring compositional variation within the core (Davies et al., 2015). Silicon, sulphur, car-

bon and oxygen are all candidate light elements in the core (Hirose et al., 2021) due to their solubility in liquid iron at high temperature and their abundance in the mantle, although their effects on nucleation are unlikely to be similar. For example, because silicon and sulphur do not strongly partition to the solid (Alfè et al., 2002b, 2000), they are expected to depress the melting curve of Fe less than carbon and oxygen, which do (Li et al., 2019; Alfè et al., 2007). The exact proportion of light elements in the core is disputed, but geophysical and geochemical constraints suggest likely abundances of Si, S, C and O to be in the range of 0–10, 0–3, 0–5 and 0–13, respectively (Badro et al., 2015; Davies et al., 2015; Dreibus and Palme, 1996; Bajgain et al., 2021).

This study will examine whether iron-rich binary alloys with compositions thought to be relevant to Earth's core are capable of spontaneous homogeneous nucleation at supercooling, which would resolve the inner core nucleation paradox. We only consider homogeneous nucleation because there are no obvious solid surfaces on which iron can first nucleate at the centre of the core. We will first describe the methods used to simulate supercooled liquids and characterise nucleation within them following our previous work (Wilson et al., 2021). We will then present predictions of critical nucleus sizes for $\text{Fe}_x\text{O}_{1-x}$, $\text{Fe}_x\text{C}_{1-x}$, $\text{Fe}_x\text{Si}_{1-x}$ and $\text{Fe}_x\text{S}_{1-x}$ at $x = 1$ and 3 mol.%. Finally, we will compare the rate at which the critical events are achieved to a revised estimate of the geophysically viable supercooling in the core.

2. Methods

This study aims to define the supercooling required to freeze iron-rich systems under the conditions of Earth's core. CNT is applied to describe the thermodynamic process of supercooled liquids freezing. In order to characterise nucleation, we require models which accurately describe molecular dynamic behaviour whilst having the computational efficiency to perform simulations of many thousands of atoms running for long periods of time and thus generate useful statistics. Embedded atom models (EAMs) are used for these simulations and first principles calculations define the parameters of these models. To frame these calculations at relevant pressure and temperatures, equations of state and melting curves must also be calculated for these models. We use the methodology of Wilson et al. (2021) expanded to binary systems to identify nuclei, calculate nucleation rates and predict waiting times for systems to freeze.

2.1. Classical nucleation theory

The rate per unit volume at which nuclei spontaneously form (I) in a supercooled liquid is

$$I(r) = I_0 \exp\left(\frac{-\Delta G(r)}{k_B T}\right), \quad (1)$$

where r is the radius of the nucleus, I_0 is a prefactor scaling the kinetics of the system, ΔG is the free energy associated with forming the nucleus, k_B is the Boltzmann constant and T is temperature. As discussed earlier, ΔG is comprised of a favourable term associated with converting liquid to solid and an unfavourable term associated with forming an interface between the states. For a sphere

$$\Delta G(r) = \frac{4}{3}\pi r^3 g^{sl} + 4\pi r^2 \gamma, \quad (2)$$

where γ is the interfacial energy and g^{sl} is the difference between the free energy of the solid and the liquid ($g^{sl} = g^s - g^l$). g^{sl} can be approximated through the enthalpy of fusion, h_f and an accommodation for second order non-linearity in the temperature dependence, h_c ,

$$g^{sl} = h_f \frac{\delta T}{T} (1 - h_c \delta T). \quad (3)$$

g^{sl} varies with temperature and supercooling ($\delta T = T - T_m$) whilst γ is constant to a first approximation. This means that under CNT, the scaling of g^{sl} with T is what drives the exponential relation of nucleation rate with temperature in Eq. (1). Liquids must be supercooled to freeze because no part of ΔG is favourable otherwise. Furthermore, Eq. (1) shows that at the melting temperature, the nucleation rate of all nuclei is infinitesimally small.

The nucleation barrier (Eq. (2)) is dominated by γ at small r because of high surface area to volume ratio. All nuclei must grow from a single atom through all smaller nuclei sizes before a system can be completely frozen. The value of ΔG increases with r to a peak at which point the probability of continued growth is equal to that of remelting. This is the critical size, r_c , beyond which, the continued growth of a nucleus becomes exponentially more likely and so will usually result in the system freezing. r_c is found by evaluating the peak of the barrier, where the gradient of ΔG with respect to radius is zero

$$\frac{\delta \Delta G}{\delta r} = 0 \quad (4)$$

gives

$$r_c = \frac{-2\gamma}{g^{sl}}. \quad (5)$$

Combining Eq. (1)–(3) with Eq. (5) then gives the rate at which the critical event occurs, the inverse of which is the average waiting time between critical events

$$\tau_w = \tau_0 \exp\left(\frac{\Delta G(r_c)}{k_B T}\right), \quad (6)$$

where

$$\tau_0 = \frac{z}{NS}, \quad (7)$$

and

$$z = \left(\frac{4\pi r_c^3 g^{sl}}{k_B T}\right)^{-1/2}. \quad (8)$$

Here, S , N and z are the rate of nuclei growth, number of available nucleation sites and Zeldovich factor, respectively. With this formulation, once we know the thermodynamic properties of the system we can evaluate the value of δT compatible with the available incubation time for the inner core. We use classical molecular dynamic (CMD) simulations to observe sub-critical ($r < r_c$) nuclei and record I , the distributions of which inform the thermodynamic quantities in Eq. (2) (see section 2.4 for details).

2.2. Molecular dynamics

Simulations must contain tens of thousands of atoms and be observed for several nanoseconds in order to provide useful statistics (many orders of magnitude larger and longer than possible with first principles calculations) because larger nuclei are significantly more rare than smaller ones (Eq. (1)). Embedded atom models define these large scale CMD simulations and are fit to first principles molecular dynamics simulations for high accuracy. First principles calculations provide trajectories, energies, and pressures, which are fit using embedded atom models. We follow the work of Davies et al. (2019) and Wilson et al. (2021) using existing EAM parameters for pure iron (Alfè et al., 2002a) and fitting for the additional components (Eqs. (9)–(15)).

EAMs define the total energy of a system (E) through the sum of energies contributed by each atom (i) from the pairwise interaction with other atoms (j)

$$E = \sum_{i=1}^{N^{Fe}} E_i^{Fe} + \sum_{i=1}^{N^X} E_i^X + \sum_{i=1}^{N^{FeX}} E_i^{FeX}. \quad (9)$$

For the binary systems considered here this consists of iron-iron, iron-solute and solute-solute interactions. Each of these energies includes a repulsive term (Q), which depends on the separation of the pair (r_{ij}), and an embedded term (F) which depends on the electron density between the pair (ρ_{ij})

$$E_i^{Fe} = Q_i^{Fe} + F^{Fe}(\rho_i^{Fe}) = \sum_{j=1, j \neq i}^{N^{Fe}} \epsilon^{Fe} \left(a^{Fe}/r_{ij}\right)^{n^{Fe}} - \epsilon^{Fe} C^{Fe} \sqrt{\rho_i^{Fe}}, \quad (10)$$

$$E_i^X = Q_i^X + F^X(\rho_i^X) = \sum_{j=1, j \neq i}^{N^X} \epsilon^X \left(a^X/r_{ij}\right)^{n^X} - \epsilon^X C^X \sqrt{\rho_i^X}, \quad (11)$$

$$E_i^{FeX} = Q_i^{FeX} = \frac{1}{2} \sum_{i=1}^{N^{Fe}} \sum_{j=1, i \neq j}^{N^X} \epsilon^{FeX} \left(a^{FeX}/r_{ij}\right)^{n^{FeX}}, \quad (12)$$

where ϵ , a , n and C are free parameters specific to each interaction. The electron densities are also defined in terms of a radial separation

$$\rho_{ij}^{Fe} = \sum_{j=1, j \neq i}^{N^{Fe}} \left(a^{Fe}/r_{ij}\right)^{m^{Fe}} + \rho_i^{FeX}, \quad (13)$$

$$\rho_{ij}^X = \sum_{j=1, j \neq i}^{N^X} \left(a^X/r_{ij}\right)^{m^X} + \rho_i^{FeX}, \quad (14)$$

$$\rho_{ij}^{FeX} = \sum_{j=1, j \neq i}^{N^X} \left(a^{FeX}/r_{ij}\right)^{m^{FeX}}, \quad (15)$$

and include an additional parameter m^{Fe} , m^X and m^{FeX} for each class of interaction. Details of first principles calculations, fitting of EAMs and subsequent validation can be found in the supplementary information.

Davies et al. (2019) present a model for the FeO system which we adopt here, negating the need for fitting this system. EAMs were further validated through the mean square root of the fluctuations in energy differences between configurations evaluated through EAM potentials and independent DFT simulations (not used for the fitting procedure). For Fe_{0.97}Si_{0.03}, Fe_{0.97}S_{0.03} and Fe_{0.97}C_{0.03} at 6000 K these are 0.245, 0.325 and 0.360 eV per cell respectively, which is less than $k_B T$ (0.517 eV).

CMD simulations were conducted using the LAMMPS package (Plimpton, 1995) within the canonical ensemble. Systems contain 21296 atoms with periodic boundary conditions and trajectories calculated via the Verlet algorithm. Uniquely random initial configurations of iron atoms interspersed with each impurity are evolved for 2 ps at 10000 K to remove any pre-existing structure in the liquid. The system is then cooled to the desired temperature over 1 ps and then evolved for 1 ns, recording atom positions every 100 steps. Volumes are adjusted for each temperature and composition according to Birch–Murnaghan 3rd order equations of state (Birch, 1947) also calculated using the EAMs.

Table 1

Thermodynamic parameters fit to $r_c(T)$ for each composition tested where all evaluations for this study were carried out at 360 GPa. τ_0 varies with temperature but is given here as the value at the temperature which coincides with the supercooling required for inner core nucleation. ^aMelting temperature is adjusted from the value at 323 GPa to agree with the pure Fe result of (Alfè et al., 2002a). ^aWilson et al. (2021), ^bDavies et al. (2019).

	T_m K	h_f (J m ⁻³ × 10 ⁸)	h_c	γ (J m ⁻²)	τ_0 (s m ⁻³)	ICN δT K
Fe _{0.99} Si _{0.01}	6439	9.8	3.1×10^{-3}	1.2	2.69×10^{22}	1230(346)
Fe _{0.97} Si _{0.03}	6362	3.7	9.6×10^{-3}	1.4	6.43×10^{20}	2040(233)
Fe _{0.99} S _{0.01}	6441	47.8	1×10^{-6}	1.1	1.64×10^{21}	1837(1125)
Fe _{0.97} S _{0.03}	6347	31.9	1×10^{-4}	1.1	4.29×10^{20}	2131(786)
Fe _{0.99} C _{0.01}	6444	57.0	1×10^{-3}	1.005	2.93×10^{23}	711(55)
Fe _{0.97} C _{0.03}	6348	130.0	1×10^{-6}	1.005	4.63×10^{23}	612(139)
Fe ^a	6522	7.119×10^{10}	6.069×10^{-5}	1.02	5.742×10^{44}	807
Fe _{0.90} O _{0.10} ^b	5987 [*]	9.8×10^9	7.05×10^{-5}	1.02	1.26×10^{45}	730

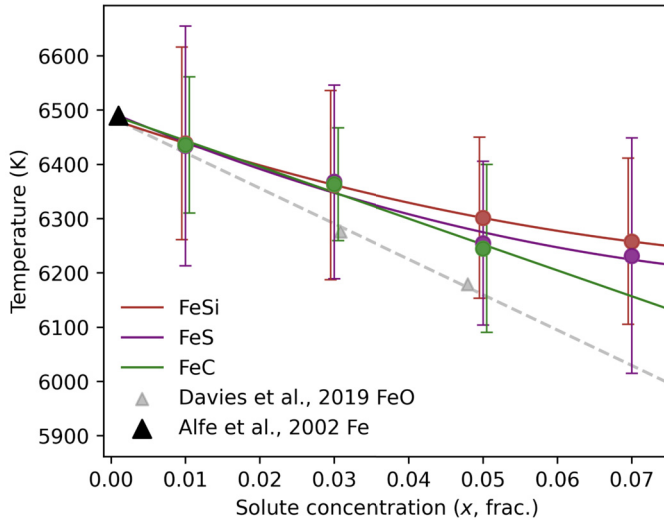


Fig. 1. Melting temperatures of FeSi, FeS, and FeC (maroon, purple and green circles respectively) systems calculated via coexistence simulation at 360 GPa. All systems take a pure Fe melting temperature from Alfè et al. (2002a) (black triangle) which also provides the same Fe EAM used here. Davies et al. (2019) is shown for comparison (grey triangles and line) and is adjusted to match the pure Fe 360 GPa value. Lines are polynomial fits of melting points.

2.3. Melting temperatures and equations of state

Melting temperatures are necessary to evaluate Eq. (3) and frame the supercooling of a system more generally. We calculate self-consistent melting temperatures for the systems studied here through coexistence simulations. These involve simulating conditions close to a point on the melting curve with regions of both solid and liquid within a single system. When the system is allowed to evolve under the microcanonical ensemble some portion of either phase will convert to the other in order to establish an equilibrium whilst maintaining constant energy and the temperature of the system will adopt a point on the melting curve. Melting point depression for Si, C and O is presented in Table 1 and Fig. 1. The value of O is taken from Davies et al. (2019) and other systems are polynomials fit to our coexistence simulation results.

2.4. Nuclei identification

The conditions of interest for nucleation the Earth's inner core involve a maximum volume of 17.62×10^{18} m³ and a maximum waiting time on the order of 1 Gyrs. These are not volumes or waiting times which are accessible to molecular dynamic simulation and so the critical event cannot be observed. Instead we record sub-critical nuclei, the frequency of which is directly sam-

pling I . Solid-like arrangements of atoms are identified within the supercooled liquid as in our previous study (Wilson et al., 2021) which in turn uses a previously developed method for categorising local bonding environments via local order parameters (Van Duineveldt and Frenkel, 1992; Rein ten Wolde et al., 1996; Persson et al., 2011). Nuclei are comprised of solid like atoms within bonding distance of one-another (as defined by the full width of the first peak of the radial distribution function). Nuclei sharing at least half of their atoms with another in an adjacent timestep are considered to be a time evolution of the same nuclei. This information allows the frequency and growth rate of each nuclei size to be recorded.

Following Wilson et al. (2021), we use I recorded from CMD simulations to predict r_c . By considering a single T , Eq. (1) can be expressed as

$$-\ln(I_T(r)) \propto \Delta G_T(r). \quad (16)$$

Using this approach we can interpret nucleation rates as the portion of Eq. (2) where $r < r_c$ (because the critical event will never occur within practical durations). The absolute magnitude of ΔG remains poorly constrained, meaning that g^{sl} , γ and I_0 cannot be calculated yet. Instead, we fit this distribution via

$$\Delta G_T(r) = 4/3\pi r^3 A + 4\pi r^2 B, \quad (17)$$

where A and B are variables at each T . Once fit, this distribution then predicts $r_c = -2B/A$ in the same way as Eq. (5). Once r_c is known at all temperatures, the temperature dependence of r_c is described by combining Eq. (5) and (3) to give

$$r_c(T) = \frac{-2\gamma}{h_f \frac{\delta T}{T_m} (1 - h_c \delta T)}. \quad (18)$$

Predictions of $r_c(T)$ are fit with h_f , h_c and γ being free parameters. This leaves only τ_0 remaining to populate Eq. (6), where N and S are recorded directly from simulations and z is calculated from h_f and h_c (Eq. (7)).

3. Results

CMD simulations were used to calculate melting curves, equations of state and to characterise nucleation rates. Coexistence simulations of liquid and solid were performed at 13 volumes (corresponding to 200–400 GPa), 10 temperatures (5000–7000 K) and 4 solute concentrations (1–7 mol.%) for each system (Si, S and C). O bearing systems were analysed at 10 mol.% for direct comparison to Davies et al. (2019), who used the same EAM, in terms of critical radius. Calculation of melting temperature (Fig. 1) is necessary due to the lack of relevant published melting curves (in terms of

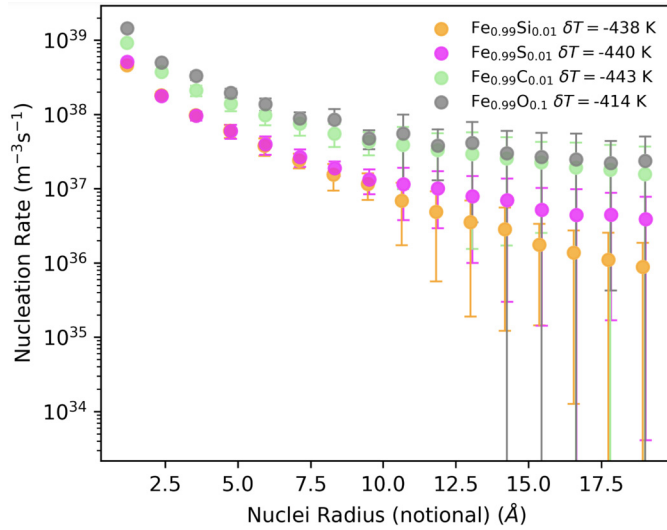


Fig. 2. Nucleation rates of sub-critical nuclei for 1 mol.% Si, S, O and C systems (orange, magenta, grey, light green) at similar supercooling. Size is shown as notional radius (Eq. (19)). Si and S bearing systems nucleate slower than those containing O and C for all nuclei sizes.

precise composition) and the requirement to frame sub-critical nuclei information in terms of supercooling. For the FeO system we adapt the result of Davies et al. (2019) at 323 GPa, extrapolating all points such that the pure Fe result agrees with the 6490 K value of Alfè et al. (2002b) at 360 GPa (from which the Fe EAM originates). Results in this study are evaluated at 360 GPa because the centre of the Earth will have experienced the longest incubation time for nucleating the inner core.

Fig. 1 shows melting temperatures calculated from coexistence simulations. At low solute concentration (1 mol.%) all systems see a melting point depression of ~ 50 K. At 7 mol.% concentration, FeSi and FeS systems have a similarly small $\frac{dT_m}{dx}$ whilst FeO and FeC remain approximately linear over this compositional range. This is because Si and S are partitioned equally between solid and liquid (Alfè et al., 2007), meaning the effect on free energies of solid and liquid is similar and balanced. The opposite is true of O, which is strongly partitioned to the liquid (Alfè et al., 2002b) implying greater melting point depression. The intermediate result of C suggests that it is favourably partitioned to the liquid but not as strongly as O (Li et al., 2019). Melting point depression for O bearing systems found by Davies et al. (2019) is greater than the effect we find for systems with C. We do not use results from concentrations above 3 mol.% in this study (more details below).

To characterise nucleation we perform CMD calculations of supercooled systems at 5–10 temperatures and 1–3 solute concentrations in order to record the properties of sub-critical nuclei. For sufficient statistics to be gathered, we run these calculations with between 80 and 200 random initial configurations per temperature and composition, resulting in ~ 6000 calculations totalling more than 50 million cpu hours.

Fig. 2 shows nucleation rates in systems containing Si, S, O and C. Whilst the functional form of these results is not important, a first order observation is that systems with C nucleate faster than others at the same concentration and those containing 10 mol.% O nucleate fastest. Those with C produce nuclei $\sim 20\%$ faster than those containing Si and S for similar supercooling, suggesting that the nucleation barrier is lower (ΔG is smaller) whilst being comparable to that of systems with a higher concentration of O. It is helpful to express these results in terms of a notional radius

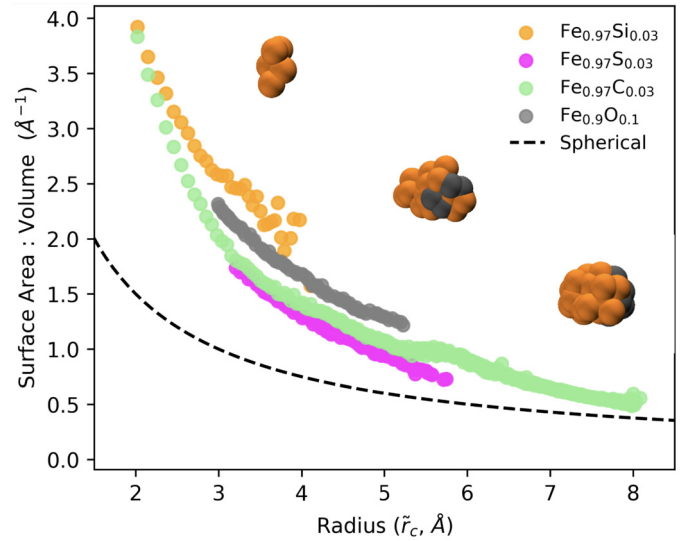


Fig. 3. Surface area to volume ratio for sub-critical nuclei at ~ 400 K supercooling ($r_c > 20$ Å). Systems containing 3 mol.% Si, S, O and C (orange, pink, grey and green circles) are shown as well as the spherical case (black dashed line, $3/r$) for comparison. Also shown are example nuclei from the C bearing system for reference. Surface area to volume ratios are similar for all systems and approach spherical before the critical size.

$$\tilde{r} = \left(\frac{N^{nuc} v_{par}}{\frac{4}{3}\pi} \right)^{1/3}, \quad (19)$$

where $v_{par} = V/N^{atoms}$, V is the volume of the system, N^{atoms} is the number of atoms in the system and N^{nuc} is the number of atoms in the nuclei. Whilst \tilde{r} is framed in terms of spherical nuclei, shapes can vary from this significantly as we discuss below.

Simulations containing ≥ 5 mol.% Si, S and C often produced liquid regions enriched in the solute. This means that the system then contains at least two liquids, one Fe rich and another solute rich. The formulations of CNT and free energy differences applied here are not appropriate to describe these conditions. Additionally, our EAMs were not trained or validated with this phenomena. Because of this we choose not to include these concentrations here and focus on the cases where the liquid compositions remain consistently homogeneous. Despite not including these results, we observe the same phenomena of liquid phase separation in first principles calculations of smaller systems at the same conditions, suggesting that liquid-liquid immiscibility may occur in these systems.

All simulations see non-spherical nuclei at small sizes (Fig. 3). CNT typically assumes a spherical geometry (e.g. Christian, 2002) despite this formulation being intended for vapour-liquid systems. Spheres are energetically preferred due to a minimisation of surface area compared to other geometries. Many solids exhibit preferred growth directions in crystal lattices where the energetic benefit of forming a non-spherical crystal can outweigh the minimised surface area of a spherical nucleus. The anisotropic nature of these situations can largely be ignored (Christian, 2002) provided that a consideration of geometry is still applied. In a previous study, we retained the spherical treatment of CNT equations, however, the distributions produced here via Eq. (16) from CMD simulations produce greater $\frac{\delta \Delta G}{\delta r}$ at small r compared to spherical geometry which is assumed by standard CNT.

Whilst small non-spherical nuclei were apparent with the pure Fe system, we find the departure from sphericity to be more pronounced in impure systems. Despite this, as nuclei grow, they incorporate a greater number of defects, randomising the preferred growth direction and becoming spherical before reaching the critical size. This is true in both pure and light element bearing

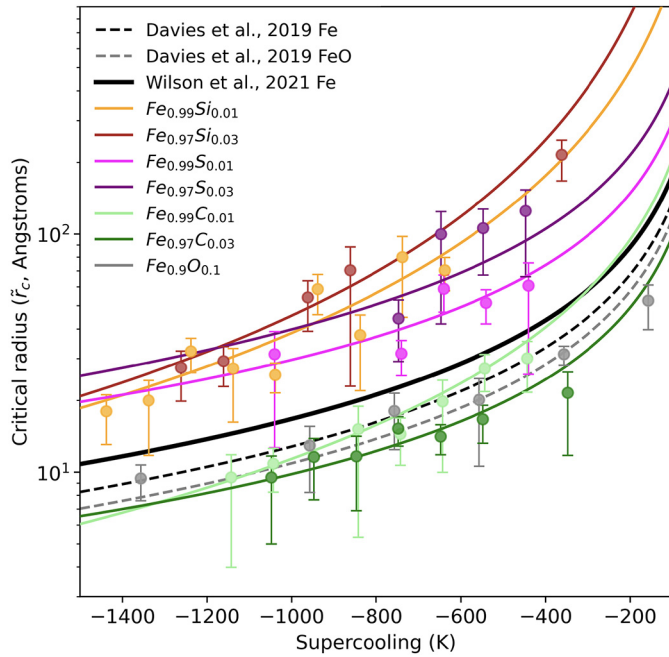


Fig. 4. Critical radii sizes predicted at different δT from distributions of nuclei sizes in CMD simulations. Coloured points are the results of this study and the fits are shown as solid lines. The pure Fe results of Wilson et al. (2021) (thick black line) are shown for comparison as well as the results of Davies et al. (2019) (dashed lines). Larger critical radii than the pure Fe system are found in Si and S bearing systems, which represents a reduced nucleation rate and an increased barrier to nucleation. O and C bearing systems present reduced critical radii relative to the pure system. Davies et al. (2019) and Wilson et al. (2021) are shown for comparison where the latter uses the same methods applied here.

systems, where the pure Fe systems which freeze are best described as defect rich hcp structure (Wilson et al., 2021). These defects are randomly distributed and disrupt the structure of the nuclei but are able to relax from the structure to form the energetically favoured phase given time. When nuclei first form, they contain few defects and so are most likely to have a single preferred growth in the direction of the basal plane as platelets, as is generally the case with hcp metals and alloys (Bergman et al., 2003). This is the mechanism which promotes dendritic growth in hcp structured materials and leads to small nuclei becoming elongate here.

For completeness we include a description of non-spherical geometries. The surface area to volume ratio of these geometries follows a power law decay, the same as a sphere, only with a greater initial gradient. We therefore apply Eq. (16) with $\Delta G = VA + \omega VB$ in place of Eq. (2), where V is volume of the nuclei, A and B are proxies for the free energy contributions and ω is the surface area to volume ratio $\omega = \frac{s}{V} = \alpha V^{-\beta/3}$, using α and β as fitting parameters. All results here include this modification. Despite this accommodation, the geometry of larger nuclei becomes increasingly spherical towards the critical size and the temperature dependence of r_c remains appropriately described by Eq. (18).

Through Eq. (16), nucleation rates recorded from MD simulations allow the prediction of critical radius (Fig. 4) following the methods of Wilson et al. (2021). For Si and S systems, the critical nuclei predicted are 10–100% larger than in the pure case at the same supercooling. The O bearing system matches the extrapolated result of Davies et al. (2019) where nuclei are $\sim 10\%$ smaller than the pure Fe case for 10 mol.% O. This further confirms the result of Wilson et al. (2021), where the methods here agree with direct observation of nucleation in molecular dynamics simulation. C is more efficient at producing smaller nuclei than O. 1 mol.% C gives a similar result to 10 mol.% O for moderate supercooling, but is

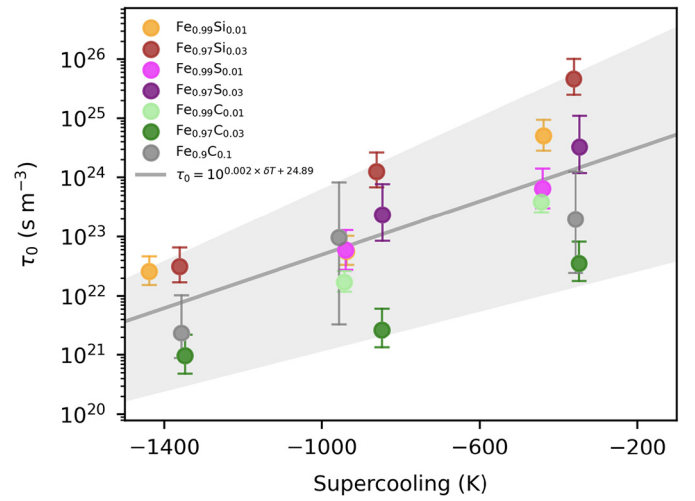


Fig. 5. Prefactor to critical event waiting times calculated from values in Table 1 and the growth rate (S) and number of nucleation sites (N) recorded from CMD simulations. Temperature dependence is a linear fit (grey line) of all data and the shaded region captures the uncertainty of this fit.

less efficient at small supercooling. Extrapolation of these results to 200 K supercooling (where Huguet et al. (2018) proposed the paradox would be resolved) suggests that the small C concentrations would not reduce the barrier to nucleation compared to Fe. 3 mol.% C provides smaller critical nuclei over the 10 mol.% O case at all T studied, being $\sim 20\%$ smaller than the pure Fe case. Because the results here for O confirm those of our previous study (Davies et al., 2019), we do not perform calculations to complete further analysis and include our previous results subsequent comparisons.

Fits to the temperature dependence of r_c using Eq. (18) give the thermodynamic quantities for each system, shown in Table 1. Fig. 5 shows τ_0 where these quantities are used to evaluate z . The remaining components of z are the rate at which nuclei grow (S) and the number of nucleation sites present at any given time (N , unrelated to terms in EAM models). It should be noted that τ_0 is a poorly known quantity (Christian, 2002), but the exact value is unimportant due to the scale of the exponential term in τ_w . The variability across compositions and the temperatures relevant to inner core incubation is within several orders of magnitude, and far less than the variance between values applied by previous studies (Christian, 2002; Huguet et al., 2018; Davies et al., 2019). Holding τ_0 constant (as a mean of all systems and temperatures) does not greatly impact the waiting time results presented here and has been the approach of most applications of CNT previously. For completeness, and to illustrate that the variance in Fig. 5 is unimportant, we choose to include a temperature dependence in our calculation of Eq. (6). This is a linear fit to the exponent of all τ_0 ($\tau_0(T) = 10^{0.02 \times \delta T + 24.89}$).

The difference in free energy between solid and liquid defines the energetic benefit to freezing the liquid. A more negative g^{sl} is seen for C bearing systems compared to the pure Fe system and those containing Si and S (Fig. 6). 1 mol.% C produces a similar g^{sl} to 10 mol.% O (Davies et al., 2019) albeit with greater non-linearity. S and Si see a smaller free energy difference at all temperatures when compared to other systems (Fig. 6), agreeing with previous finding that partitioning is approximately evenly between solid and liquid iron (Alfè et al., 2007).

4. Discussion

Our results shown that nucleation rates in Fe rich liquids containing C and O are faster than those containing Si or S (Fig. 2).

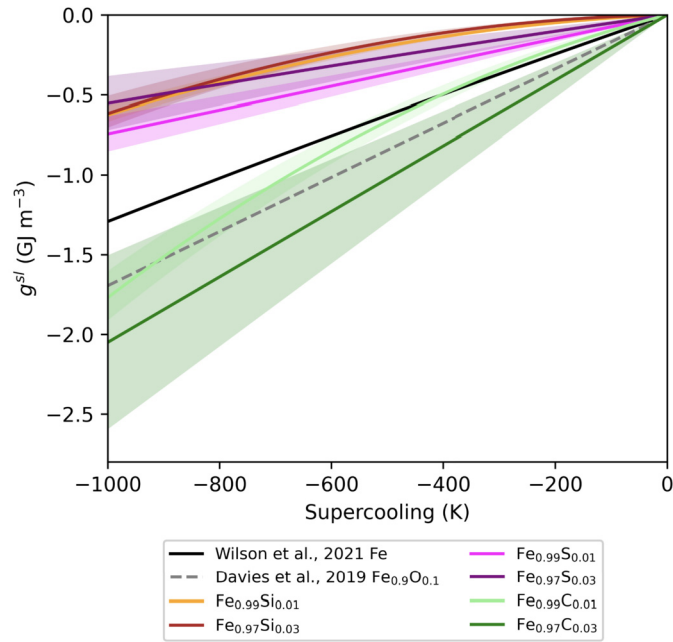


Fig. 6. Comparison of $g^{sl}(T)$ from fits to $r_c(T)$ for each system studied. Dashed black line is the pure Fe case from Wilson et al. (2021) using the same methods applied here. Large differences in g^{sl} between the different cases imply that the structure of nucleating material and the composition of the liquid are largely responsible for differing nucleation behaviour.

Compared to the pure Fe system, critical nuclei sizes are larger in systems containing Si and S and smaller in those containing C and O (Fig. 4). These findings suggest that systems containing C and O should freeze at higher temperatures (lower supercooling) than a pure Fe system. We find that the methods employed here, which allow interpolation of the relevant supercooling for the core, validate the results from Davies et al. (2019) for $\text{Fe}_{0.9}\text{O}_{0.1}$ where extrapolation was required.

To assess whether the systems studied here might resolve the nucleation paradox, we must compare the time taken to nucleate (τ_w) with the available time to nucleate in the core, the incubation time ($\tau_i = v_{ic}t_{ic}$, where v_{ic} is the volume of the inner core which spontaneously freezes and t_{ic} is the supercooled time). Huguët et al. (2018) estimated a maximum allowable supercooling of the core as $\tau_i \sim 200$ K by calculating the separation between the isentrope and melting curve at the centre of the Earth whilst preserving the intersection of melting curve and temperature profile at the present inner core boundary. To do this, the authors defined a melting curve with Lindemann's law and the result of Anzellini et al. (2013), and isentrope from Labrosse (2003), both populated with material properties of the core. By varying the parameters of these functions within their uncertainty, the authors found that the maximum supercooling at the centre of the core is ~ 200 K. This would translate to a τ_i of $2.4 \times 10^{35} \text{ s m}^3$ if the core is cooling at $\sim 200 \text{ K Gyr}^{-1}$.

We consider two cases, the extreme case τ_i^E and a more moderate version τ_i^M . In the extreme case, the inner core grew instantaneously to its current size, whereas for the moderate case, the core grew to half of its current radius instantaneously and then grew slowly. For the extreme case we take a similar approach to Huguët et al. (2018) (blue lines, Fig. 7) but explore alternative melting curves and temperature profiles. If the thermal conductivity of the core is both high and depth dependent then the centre of the core can become thermally stratified (Gomi et al., 2013). In this case the temperature profile of the core would not be isentropic and could be isothermal at an extreme. We vary the parameters of the adiabat and the effect of melting point de-

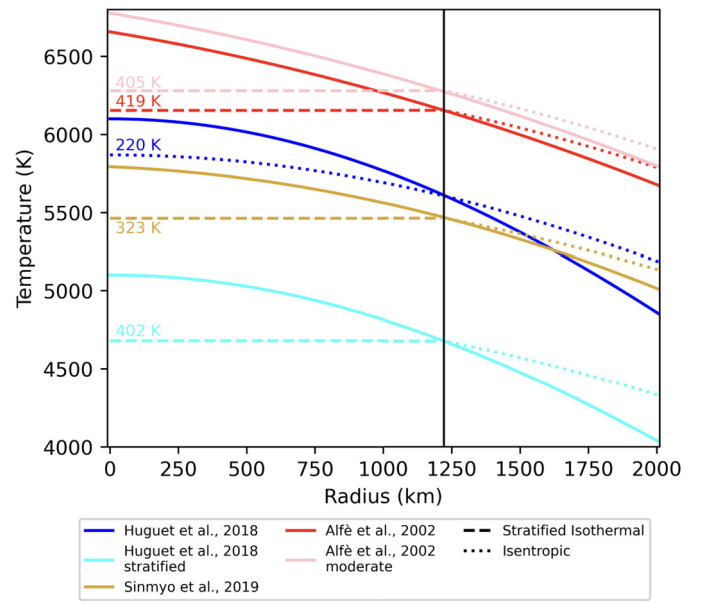


Fig. 7. Melting curves (solid lines) and core temperature profiles (dotted and dashed lines) with radius. Dotted (dashed) lines represent isentropic (isothermal) regions of the core. Temperature profiles are obtained by varying the material properties of the core used by Huguët et al. (2018) within uncertainty and also considering thermal stratification of the innermost core. Melting curves from Huguët et al. (2018) (blue, cyan), Alfè et al. (2002a) (red, pink) and Sinmyo et al. (2019) (brown) are also applied (with varying degrees of melting point depression, causing the apparent separation) to find the largest plausible separation of temperature and melting point, and therefore supercooling, at 360 GPa whilst preserving an intersection at the ICB (solid black line). Also shown are the numerical δT for each case.

pression on the melting curves of Alfè et al. (2002c) and Sinmyo et al. (2019). These melting curves are chosen to explore different predicted gradients at the ICB. Fig. 7 presents some of these combinations, including a case showing the maximum permissible supercooling of 419 K (red lines, Fig. 7). This means that the extreme case $\tau_i^E = 5.04 \times 10^{35} \text{ s m}^3$, more than double that used by previous studies. For a moderate case, which might offer a resolution to the paradox whilst being plausible to incorporate into thermal histories of the core, we begin by taking the extreme case without exploring the uncertainties or melting point depression. With an isothermal inner core the maximum permissible supercooling is 405 K (pink lines, Fig. 7). Most importantly, we consider that the incubation volume is half the radius of the present inner core, implying that the remaining half (87.5% of volume) of the inner core froze slowly as the core cooled. For this moderate case, $\tau_i^M = 3.1 \times 10^{34} \text{ s m}^3$, and represents the time available to not just resolve the paradox, but do so with the inner core age being compatible with thermal history models of the core.

The duration before a supercooled system will produce a critical event and freeze is presented here as waiting time (Fig. 8). As predicted by a lower nucleation rate, larger critical nuclei and less favourable thermodynamic properties, Si and S bearing systems require significantly greater supercooling than the pure iron system. In reality these systems would simply freeze via alternate mechanisms; the fluctuations of composition we observe in high solute concentration systems would produce Fe rich regions spontaneously. This would result in an elevated nucleation rate due to a more pure Fe system, meaning that freezing would occur at supercooling closer to that described by the pure Fe case. In the case of the $\text{Fe}_{0.9}\text{O}_{0.1}$ systems, our results of r_c confirm those of Davies et al. (2019), where 730 K of supercooling is needed to nucleate the inner core.

Compared to all other cases considered so far, the FeC system shows a far more efficient reduction of the nucleation barrier,

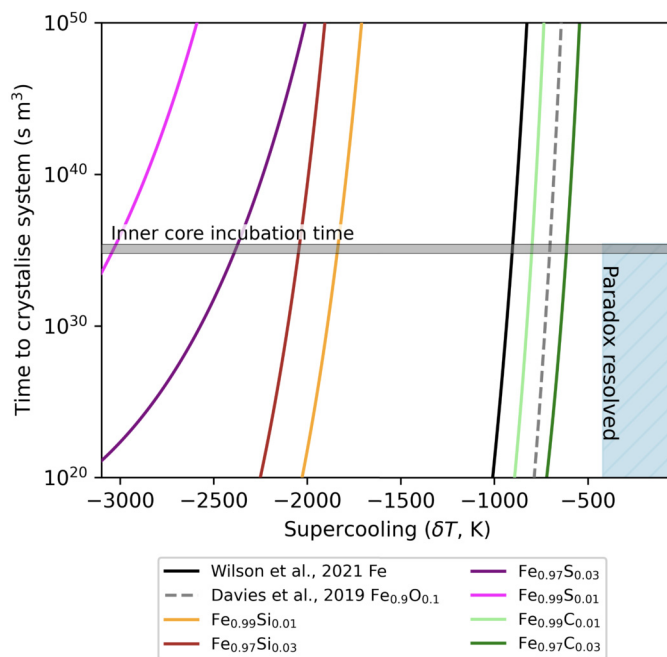


Fig. 8. Average waiting time to observe freezing time against δT for different iron-rich liquids. Solid coloured lines are results of this study for systems containing Si (orange, maroon), S (pink, purple) and C (greens) where dark colours are higher solute concentration. Black (grey) line is the Fe ($\text{Fe}_{0.9}\text{O}_{0.1}$) systems from previous studies (Wilson et al., 2021; Davies et al., 2019). Grey shaded region represents plausible incubation times available to nucleate the inner core assuming a cooling rate of 200 K Gyr^{-1} and is bounded by the extreme and moderate cases (upper and lower bounds respectively).

partly due to a smaller depression of the melting curve compared to O. 1 mol.% C requires 711 K (± 55 K) of supercooling and is comparable to the effect of 10 mol.% O (730 K found by Davies et al., 2019, and confirmed by our results). 3 mol.% C requires cooling to 612 (± 139) K below melting in order to nucleate the inner core for the extreme case, close to the 419 K of permissible supercooling of core to avoid a nucleation paradox. For the moderate case, where incubation time accounts for the core being several hundred million years old, this system requires 615 (± 148) K of supercooling. We find a corresponding moderate value for the maximum supercooling permissible in the core to be 405 K.

5. Conclusion

This study examines the effect of Si, S, O and C on the nucleation of the inner core, using CNT as an already proven theoretical framework for these conditions (Wilson et al., 2021). Both oxygen and carbon can make a reduction to the supercooling required to produce the first solids in the core. The best conceivable solution from the binary systems tested here is within ~ 50 K of resolving the most optimistic rendition of the paradox. Other higher concentration or ternary systems may surpass these results and present possible resolutions to the paradox but testing of these scenarios is beyond the capability of the methodology applied here. Combinations of Si or S and O or C might have a neutralising effect whilst O and C together could offer a low concentration resolution to the paradox, although complex, non-linear interactions are likely. It should also be noted that a minimum viable resolution to the paradox still presents significant challenges for the thermal history of the core as it implies a very young inner core which is incompatible with a high thermal conductivity core and consistent geodynamo output. We find that for a more reasonable case, where half of the present inner core radius grew slowly, the paradox is

~ 20 K more difficult to resolve than our extreme case where the entire inner core grew instantaneously.

There remains the possibility that CNT and the processes explored here are less favourable than others which might revolve the paradox. Ternary systems will involve complex chemical interactions where unfavourable elements studied here could form beneficial quasi-crystals or chemical heterogeneities. Such interactions may not fit within the confines of CNT and would likely require explicit chemical potentials to define the free-energy balances at play.

The presence of compositionally distinct regions in our simulations means a breakdown of the thermodynamic theory we apply here. If explored appropriately, these may provide alternate resolutions to the paradox; for example, through local enrichment in elements which reduce the nucleation barrier.

CRediT authorship contribution statement

Alfred Wilson: Conceptualization, Methodology, Software, Validation, Formal Analysis, Writing, Visualisation. **Dario Alfè:** Conceptualization, Methodology, Software, Writing, Funding acquisition. **Andrew Walker:** Conceptualization, Supervision, Writing, Funding acquisition. **Christopher Davies:** Conceptualization, Supervision, Writing, Project administration, Funding acquisition.

Declaration of competing interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

Data availability

Data will be made available on request.

Acknowledgements

We thank the reviewers for their thoughtful assessment of the manuscript and helpful comments. We acknowledge a Natural Environment Research Council grant, reference NE/T000228/1, which supports this project. Calculations were performed on the UK National supercomputing service ARCHER2 (via allocation through the Mineral Physics Consortium).

Appendix A. Supplementary material

Supplementary material related to this article can be found online at <https://doi.org/10.1016/j.epsl.2023.118176>.

References

- Alfè, D., Gillan, M., Price, G., 2000. Constraints on the composition of the Earth's core from ab initio calculations. *Nature* 405, 172–175.
- Alfè, D., Gillan, M., Price, G., 2002a. Complementary approaches to the ab initio calculation of melting properties. *J. Chem. Phys.* 116, 6170–6177.
- Alfè, D., Gillan, M., Price, G., 2007. Temperature and composition of the Earth's core. *Contemp. Phys.* 48, 63–80.
- Alfè, D., Gillan, M., Price, G.D., 2002b. Composition and temperature of the Earth's core constrained by combining ab initio calculations and seismic data. *Earth Planet. Sci. Lett.* 195, 91–98.
- Alfè, D., Price, G., Gillan, M., 2002c. Iron under Earth's core conditions: liquid-state thermodynamics and high-pressure melting curve from ab initio calculations. *Phys. Rev. B* 65, 165118.
- Anderson, D.L., 2002. The case for irreversible chemical stratification of the mantle. *Int. Geol. Rev.* 44, 97–116.
- Anzellini, S., Dewaele, A., Mezouar, M., Loubeyre, P., Morard, G., 2013. Melting of iron at Earth's inner core boundary based on fast x-ray diffraction. *Science* 340, 464–466.

- Badro, J., Aubert, J., Hirose, K., Nomura, R., Blanchard, I., Borensztajn, S., Siebert, J., 2018. Magnesium partitioning between Earth's mantle and core and its potential to drive an early exsolution geodynamo. *Geophys. Res. Lett.* 45, 13–240.
- Badro, J., Brodholt, J.P., Piet, H., Siebert, J., Ryerson, F.J., 2015. Core formation and core composition from coupled geochemical and geophysical constraints. *Proc. Natl. Acad. Sci.* 112, 12310–12314.
- Bajgain, S.K., Mookherjee, M., Dasgupta, R., 2021. Earth's core could be the largest terrestrial carbon reservoir. *Commun. Earth Environ.* 2, 165.
- Bergman, M.I., Agrawal, S., Carter, M., Macleod-Silberstein, M., 2003. Transverse solidification textures in hexagonal close-packed alloys. *J. Cryst. Growth* 255, 204–211.
- Biggin, A.J., Piispa, E., Pesonen, L.J., Holme, R., Paterson, G., Veikkolainen, T., Tauxe, L., 2015. Palaeomagnetic field intensity variations suggest mesoproterozoic inner-core nucleation. *Nature* 526, 245–248.
- Birch, F., 1947. Finite elastic strain of cubic crystals. *Phys. Rev.* 71, 809.
- Bono, R.K., Tarduno, J.A., Nimmo, F., Cottrell, R.D., 2019. Young inner core inferred from Ediacaran ultra-low geomagnetic field intensity. *Nat. Geosci.* 12, 143–147.
- Christian, J.W., 2002. *The Theory of Transformations in Metals and Alloys*. Newnes.
- Davies, C., 2015. Cooling history of Earth's core with high thermal conductivity. *Phys. Earth Planet. Inter.* 247, 65–79.
- Davies, C., Pozzo, M., Alfè, D., 2019. Assessing the inner core nucleation paradox with atomic-scale simulations. *Earth Planet. Sci. Lett.* 507, 1–9.
- Davies, C., Pozzo, M., Gubbins, D., Alfè, D., 2015. Constraints from material properties on the dynamics and evolution of Earth's core. *Nat. Geosci.* 8, 678–685.
- Davies, C.J., Bono, R.K., Meduri, D.G., Aubert, J., Greenwood, S., Biggin, A.J., 2022. Dynamo constraints on the long-term evolution of Earth's magnetic field strength. *Geophys. J. Int.* 228, 316–336.
- Dreibus, G., Palme, H., 1996. Cosmochemical constraints on the sulfur content in the Earth's core. *Geochim. Cosmochim. Acta* 60, 1125–1130.
- Driscoll, P., Davies, C., 2023. The “new core paradox:” challenges and potential solutions. *J. Geophys. Res., Solid Earth* e2022JB025355.
- Frank, F.C., 1952. Supercooling of liquids. *Proc. R. Soc. Lond. Ser. A, Math. Phys. Sci.* 215, 43–46.
- Gomi, H., Ohta, K., Hirose, K., Labrosse, S., Caracas, R., Verstraete, M.J., Hernlund, J.W., 2013. The high conductivity of iron and thermal evolution of the Earth's core. *Phys. Earth Planet. Inter.* 224, 88–103.
- Hirose, K., Morard, G., Sinmyo, R., Umemoto, K., Hernlund, J., Helffrich, G., Labrosse, S., 2017. Crystallization of silicon dioxide and compositional evolution of the Earth's core. *Nature* 543, 99–102.
- Hirose, K., Wood, B., Vočadlo, L., 2021. Light elements in the Earth's core. *Nat. Rev. Earth Environ.* 2, 645–658.
- Huguet, L., Van Orman, J.A., Hauck II, S.A., Willard, M.A., 2018. Earth's inner core nucleation paradox. *Earth Planet. Sci. Lett.* 487, 9–20.
- Kelton, K., Gangopadhyay, A., Kim, T., Lee, G., 2006. A case for local icosahedral order in undercooled metallic liquids and the influence on the nucleation barrier. *J. Non-Cryst. Solids* 352, 5318–5324.
- Kelton, K., Lee, G., Gangopadhyay, A.K., Hyers, R., Rathz, T., Rogers, J., Robinson, M., Robinson, D., 2003. First x-ray scattering studies on electrostatically levitated metallic liquids: demonstrated influence of local icosahedral order on the nucleation barrier. *Phys. Rev. Lett.* 90, 195504.
- Labrosse, S., 2003. Thermal and magnetic evolution of the Earth's core. *Phys. Earth Planet. Inter.* 140, 127–143.
- Labrosse, S., 2015. Thermal evolution of the core with a high thermal conductivity. *Phys. Earth Planet. Inter.* 247, 36–55.
- Li, Y., Vočadlo, L., Alfè, D., Brodholt, J., 2019. Carbon partitioning between the Earth's inner and outer core. *J. Geophys. Res., Solid Earth* 124, 12812–12824.
- Nimmo, F., 2015a. Energetics of the core. In: Schubert, G. (Ed.), *Treatise on Geophysics*, 2nd edn., vol. 8. Elsevier, Amsterdam, pp. 27–55.
- Nimmo, F., 2015b. Thermal and compositional evolution of the core. In: *Treatise on Geophysics*, vol. 9, pp. 201–219.
- O'Rourke, J.G., Stevenson, D.J., 2016. Powering Earth's dynamo with magnesium precipitation from the core. *Nature* 529, 387–389.
- Persson, J., Desgranges, C., Delhommelle, J., 2011. Polymorph selection during the crystallization of iron under the conditions of Earth's inner core. *Chem. Phys. Lett.* 511, 57–61.
- Plimpton, S., 1995. Fast parallel algorithms for short-range molecular dynamics. *J. Comput. Phys.* 117, 1–19.
- Schenk, T., Holland-Moritz, D., Simonet, V., Bellissent, R., Herlach, D., 2002. Icosahedral short-range order in deeply undercooled metallic melts. *Phys. Rev. Lett.* 89, 075507.
- Shibuta, Y., Sakane, S., Miyoshi, E., Okita, S., Takaki, T., Ohno, M., 2017. Heterogeneity in homogeneous nucleation from billion-atom molecular dynamics simulation of solidification of pure metal. *Nat. Commun.* 8, 10.
- Sinmyo, R., Hirose, K., Ohishi, Y., 2019. Melting curve of iron to 290 GPa determined in a resistance-heated diamond-anvil cell. *Earth Planet. Sci. Lett.* 510, 45–52.
- Sun, Y., Zhang, F., Mendelev, M.I., Wentzcovitch, R.M., Ho, K.M., 2022. Two-step nucleation of the Earth's inner core. *Proc. Natl. Acad. Sci.* 119, e2113059119.
- Tarduno, J.A., Cottrell, R.D., Watkeys, M.K., Hofmann, A., Doubrovine, P.V., Mamajek, E.E., Liu, D., Sibeck, D.G., Neukirch, L.P., Usui, Y., 2010. Geodynamo, solar wind, and magnetopause 3.4 to 3.45 billion years ago. *Science* 327, 1238–1240.
- Van Duineveldt, J., Frenkel, D., 1992. Computer simulation study of free energy barriers in crystal nucleation. *J. Chem. Phys.* 96, 4655–4668.
- Wilson, A.J., Pozzo, M., Alfè, D., Walker, A.M., Greenwood, S., Pommier, A., Davies, C.J., 2022. Powering Earth's ancient dynamo with silicon precipitation. *Geophys. Res. Lett.* 49, e2022GL100692.
- Wilson, A.J., Walker, A.M., Alfè, D., Davies, C.J., 2021. Probing the nucleation of iron in Earth's core using molecular dynamics simulations of supercooled liquids. *Phys. Rev. B* 103, 214113.
- Rein ten Wolde, P., Ruiz-Montero, M.J., Frenkel, D., 1996. Numerical calculation of the rate of crystal nucleation in a Lennard-Jones system at moderate undercooling. *J. Chem. Phys.* 104, 9932–9947.
- Zhang, W.J., Liu, Z.Y., Liu, Z.L., Cai, L.C., 2015. Melting curves and entropy of melting of iron under Earth's core conditions. *Phys. Earth Planet. Inter.* 244, 69–77.
- Zhou, T., Tarduno, J.A., Nimmo, F., Cottrell, R.D., Bono, R.K., Ibanez-Mejia, M., Huang, W., Hamilton, M., Kodama, K., Smirnov, A.V., et al., 2022. Early Cambrian renewal of the geodynamo and the origin of inner core structure. *Nat. Commun.* 13, 1–7.