

The formation of lithium-rich
melts during early crustal
evolution — a case study from the
Archean Zimbabwe Craton



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A thesis submitted for the degree of
Doctor of Philosophy

Trinity 2025

Acknowledgements

This project was completed because of the immense support I received during my time at Oxford, and because of that I have many acknowledgements to give! I am very grateful for support I received from NERC and the Oxford–Radcliffe scholarship for funding this research. The project also greatly benefited from support, either by funding or in-kind, by the Geological Society of London, the Society of Economic Geologists, Premier African Minerals and Prospect Mineral Resources, to which I am very grateful. I would also like to thank the Research Council of Zimbabwe for enabling this research, and Midlands State University for supporting me during my fieldwork in Zimbabwe.

I would also like to sincerely thank Richard Palin, Nick Gardiner, and Laurence Robb for being fantastic supervisors as we all jumped into the pegmatite world together. Thank you for giving me the freedom to explore my interests and supporting the direction the project went.

A huge thank you to Tania Martins and Bob Linnen for being open to discussing every crazy question I had about pegmatites, and for making me feel part of your community since we met in Maine. A special thanks also goes to Catriona Breasley for the fantastic discussions we have had about all things pegmatite.

I am indebted to Brian Mappingire, Antony Mamuse, Tony Martin, Brayden St. Pierre, and Liberty Ndlovu for their support in the field during my field seasons. I am also very grateful for the support and assistance Bruce Cumming, George Roach, Jabulani Chirasha, Tinashe Muzondo, Tendai Gumi, Rutendo Musinga, Dereck Muchingami, and the rest of the team at Premier African Minerals have given me during my stays on-site and beyond. A very big thank you also goes to James Winch, Adam Moodley, and especially Roger Tyler for their help around Arcadia and sample logistics.

I want to extend my sincerest gratitude to Jane Barling for diligently teaching me everything there is to know about digestions in the clean lab. The geochemistry could not have been done without you. Owen Green, Matthew Beverly-Smith, and Emily Donald, thank you for teaching me all things rock prep, and keeping me sane during dark days in the basement. Sebastian (Batzi) Fischer and Martin Mangler

are sincerely thanked for their time (and patience) whilst teaching me how to use the laser during my analytical runs at the StAGE lab in St Andrews. I am also very grateful to Cyril Chelle-Michou for inviting me to ETH in Zurich to do additional geochronology, and Lorenzo Tavazzani for assisting me during my time there. My thanks also to Brendan Dyck and Kyle Larson for being fantastic and running my Rb–Sr geochronology in the Fipke lab at UBC Okanagan.

To my officemates – Hannah, Ally, Lewis, Harri, and Kate – thank you for making the past four years as fun (and chaotic) as they were, and thanks for being such fantastic support. My appreciation also goes out to Nikki, Bei Bei, Ian, Toby, Rebecca, Vicky, and everyone else in the department who have made my time in Oxford so special. Thanks also to the Assynt team every year, this trip was a highlight of my time here. Claire Nichols is especially thanked for being a sounding board during the car rides in and around Assynt, Marc St Onge for his infinite geological wisdom, and the undergraduates for keeping the trips fun and a pleasure to teach. To my family, thank you for supporting me through thick and thin, and always being there for me. Thank you also for entertaining my interesting passion for rocks, which I (understandably) still think hasn't rubbed off on any of you!

And most importantly, thank you to Charlotte for being the best. Your endless support and affection have constantly encouraged me to push myself beyond what I imagined possible, and I am eternally grateful for you always being there for me.

Thank you all!

Statement of Authorship

I confirm the work submitted here is my own and has not been submitted in whole or in part for consideration for any other degree or qualification in this, or any other university.

Whilst this thesis is my own work, as with any science, parts were completed in collaboration with others as outlined below.

The work in **Chapter 3** has been published in the journal *Geology* under the following title: “*The formation of lithium-rich pegmatites through multi-stage melting*”. **Koopmans, L.**, Martins, T., Linnen, R., Gardiner, N.J., Breasley, C.M., Palin, R.M., Groat, L.A., Silva, D. Robb, L.J. doi:10.1130/G51633.1

- *Conceptualisation and project design*: L. Koopmans, T. Martins, R. Linnen. *Modelling*: L. Koopmans. *Data analysis*: L. Koopmans. *Manuscript drafting*: L. Koopmans. *Editing*: R.M. Palin, N.J. Gardiner, C.M. Breasley, L.A. Groat, D. Silva.

The work in **Chapter 4** has been published in the journal *Mineralium Deposita* under the following title: “*Structural controls on lithium mineralization in shear-zone hosted granitic pegmatites of the Zulu pegmatite field, Zimbabwe – implications for exploration*”. **Koopmans, L.**, Gardiner, N.J., St. Pierre, B., Palin, R.M., Musinga, R., Robb, L.J. doi:10.1007/s00126-025-01371-x

- *Conceptualisation and project design*: L. Koopmans. *Sampling and data collection*: L. Koopmans, R. Musinga, B. St. Pierre. *Data analysis*: L. Koopmans. *Manuscript drafting*: L. Koopmans. *Editing*: L. Koopmans, N.J. Gardiner, R.M. Palin, B. St. Pierre.

The work in **Chapters 5.2.1, 5.2.2, and 5** has been invited for submission in *Earth Science Reviews*, with the following author list: **Koopmans, L.**, Palin, R.M., Mappingire, B., Gardiner, N.J., Mamuse, A., Robb, L.J., Martin, A.R.

- *Conceptualisation and project design*: L. Koopmans, N.J. Gardiner, R.M. Palin. *Sampling and fieldwork*: L. Koopmans, B. Mapingire, A. Mamuse, A.R. Martin. *Analytical campaign*: L. Koopmans, N.J. Gardiner. *Data analysis*: L. Koopmans. *Project advisor*: R.M. Palin, N.J. Gardiner, L.J. Robb.

Abstract

The exponential increase in demand for electric vehicles as part of the renewable energy transition is driving a surge in production of the critical metal lithium. Several forecasts have predicted a deficit in supply owing to declining rates of new economic discoveries at a time where over 50 new mines are needed to come online by 2030 to meet demand. The primary source of lithium is hard rock (pegmatite) deposits, which account for 70% of current reserves. Given their critical importance in securing sufficient lithium supply, a better understanding of how these deposits form is of vital importance to aid new discoveries. This thesis aims to contribute towards a mineral system model for lithium-rich pegmatites, by integrating new petrological modelling to detailed field-based, geochemical, and geochronological studies of the Zimbabwe craton – a relatively understudied Archean craton which supplied 9% of global lithium production in 2024. In this thesis, I propose that the Zimbabwe craton grew through episodic felsic magmatism between 3.8 and 2.5 Ga, and that major siliciclastic supracrustal successions formed through erosion as a result of continental emergence which was driven by a major orogenic episode c. 2.68 Ga. Through petrologic modelling, I show that whilst conventional pegmatite formation models are feasible, they do not sufficiently account for the necessary enrichment required to generate economic hard rock lithium deposits. Importantly, the absence of potential metasedimentary sources in source regions for lithium-rich pegmatites in Zimbabwe suggests an alternative process may have occurred. I show that the formation of intermediate, K-rich, granitoids through melt-driven crustal recycling can sufficiently enrich the source region to form the geochemical signatures found in lithium-rich pegmatites. Finally, transcurrent shear zones are proposed as fundamental components of the lithium mineral system, and future exploration should focus on delineating these structures to identify new targets.

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Chapter 1

Introduction

1.1 Lithium as a critical metal

Lithium (Li) is the lightest metal and lowest density solid element in the periodic table (Rumble, 2025). It forms the basis of various pharmaceuticals (Kesler and Simon, 2015), is used as a treatment for psychiatric disorders (Rybakowski, 2020), doped in specialty ceramics to improve their thermal properties (Kesler and Simon, 2015), and is mixed into lubricating greases to improve the useable temperature range and water resistance (Delgado et al., 2006). Most importantly, lithium's very high electrochemical potential (-3.04 V vs. standard hydrogen electrode, Xie and Lu, 2020), and charge:weight ratio, make it prized in high energy density batteries. This has driven a surge in demand because of two main use-cases: firstly, efforts to increase the proportion of renewable energy in national grids are coupled with the need to balance the poorly aligned peak supply from renewable sources (e.g., wind, solar) to the consumer demand. Whilst this is currently largely done by combined cycle gas generators (CCGT) in the UK, batteries are well positioned to replace CCGTs in national grids to help achieve net-zero goals (Chowdhury et al., 2020). Secondly, the electrification of the global vehicle fleet requires a significant increase in the production of large but lightweight batteries (Benson et al., 2025). The exponential increase in electric vehicle (EV) production over the past five years has therefore caused an explosion in demand for lithium, with production attempting to follow suit (Figure 1.1). Supply is currently c. 1 Mt lithium carbonate equivalent (LCE, Figure 1.1), and demand is forecast to double in the next four years, before reaching 3 Mt by 2032 (Benchmark Mineral Intelligence, 2025). In order to meet this demand, Benchmark Mineral Intelligence (2025) predict 52 additional mines (assuming new mines are equivalent to the current average mine size) need to come online by 2030.

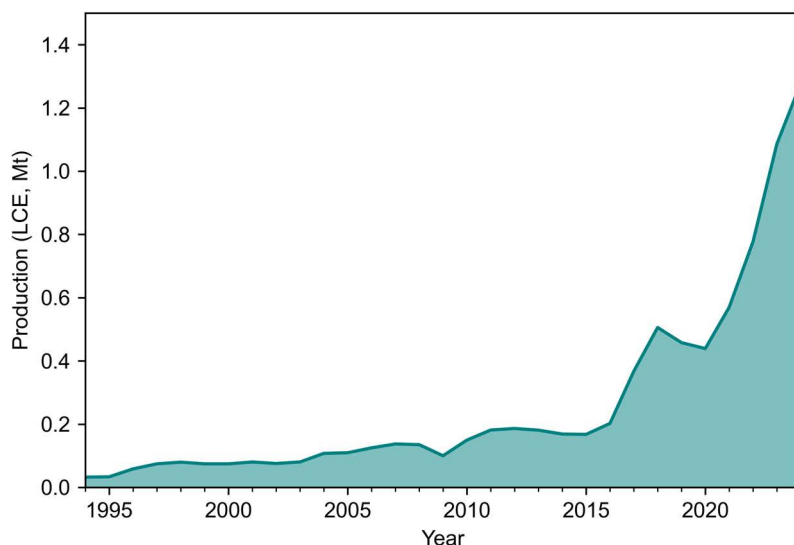


Figure 1.1: Lithium production (all sources) through time. Data reported by the United States Geological Survey (2025).

Because no metal has better physicochemical properties, lithium is unlikely to be substituted and will remain a key component of battery technology (Benson et al., 2025). Lithium is therefore one of the most critical metals in national and organisational critical minerals reports (Gardiner et al., 2024a; McNulty and Jowitt, 2021; European Commission, 2020; Nassar and Fortier, 2021), and significant research has been stimulated to better understand the ore deposit geology (Benson et al., 2025).

Lithium is sourced primarily from hard rock (pegmatite) deposits (Černý and Ercit, 2005; Müller et al., 2025; Benson et al., 2025) and closed-basin brine (“salar”) deposits (Munk et al., 2016), however unconventional deposit types are also being explored (e.g., oilfield brines, clay, Putzolu et al., 2025, Benson et al., 2025). Hard rock deposits account for 40% of the currently known lithium resources, and 70% of current reserves (Figure 1.2). Given hard rock deposits are uniquely positioned to rapidly come online (Yao, 2022), a better understanding of how these bodies form, in order to find new deposits, is critical (Gardiner et al., 2024a).

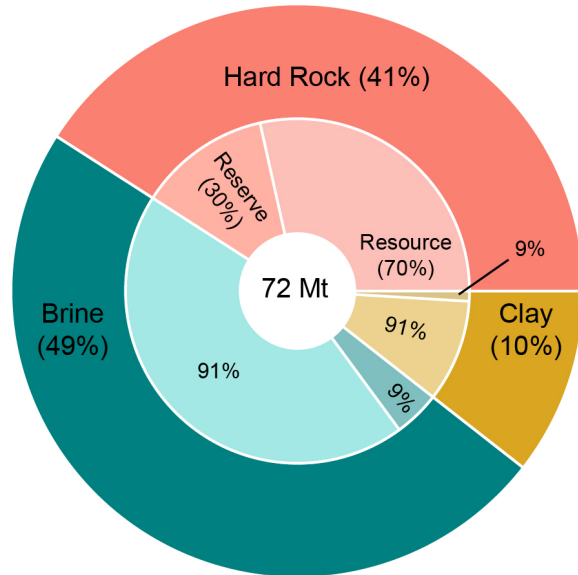


Figure 1.2: An overview of reported lithium resources and reserves (as of September 2024), classified by deposit type. Darker and lighter portions of the inner circle are reserves and resources, respectively.

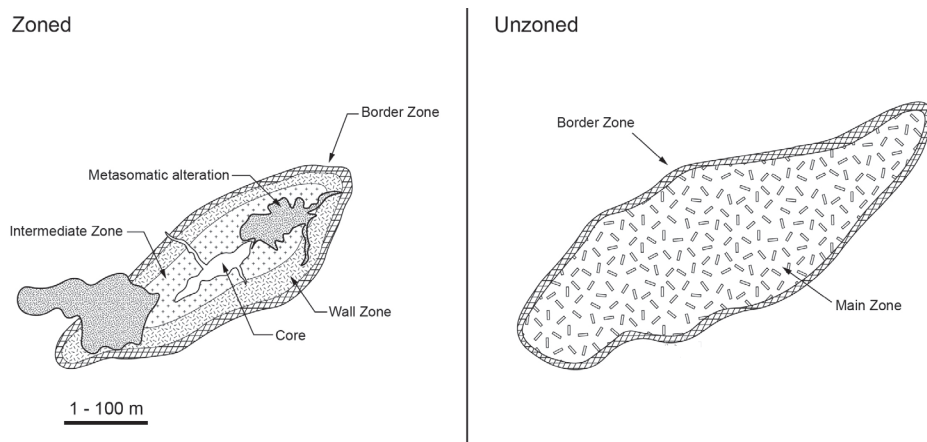


Figure 1.3: Schematic diagrams for typical zoned (left) and unzoned (right) pegmatites, outlining the standard zonation scheme. Adapted from Černý (1991a).

1.2 Hard rock lithium deposits

1.2.1 What is a pegmatite?

Granitic pegmatites (hereafter pegmatites) are small-scale, highly evolved igneous rocks which exhibit textures indicative of significant disequilibrium during crystallisation (London, 2008, 2014; Simmons and Webber, 2008). Disequilibrium results in the many distinctive features of pegmatites; large variability in crystal sizes, graphic intergrowths, and abrupt changes in mineralogy across pegmatite intrusions (London, 2014).

Experimental work has demonstrated that the pressure–temperature (P–T) conditions during crystallisation have a significant influence on which, and in what order, minerals crystallise (e.g., London, 1984; Sirbescu et al., 2017). Estimates for crystallisation conditions using pre-existing mineral geothermobarometers, however, did not always match regional P–T estimates, which led to the recognition that pegmatites can begin crystallising below the haplogranite solidus (e.g., London, 2005; Simmons and Webber, 2008). This “undercooling” is driven by the significant proportions of fluxing components, such as H₂O, Li, B, F, and P in the melt, which inhibit polymerisation of the silicate melt and greatly diminish the number of nucleation sites available in the system (London, 2014; Linnen, 1998; Sirbescu et al., 2017; Nabelek et al., 2010). The resultant melt is highly sensitive to perturbations, and changes in melt conditions and/or degassing of fluxing components drives rapid crystallisation and various textures observed in pegmatites (Simmons and Webber, 2008; Nabelek et al., 2010; Sirbescu et al., 2017; London, 2014).

A common feature of pegmatites is a strong mineralogical zonation (although unzoned varieties also exist, Figure 1.3). The zoning scheme of a simple pegmatite typically includes; a border zone (considered to be a chilled margin, cf. London 2014), a mineralogically equivalent but coarser wall zone, an intermediate zone (defined by the appearance of exotic, incompatible element-rich phases), and a core zone that is predominantly composed of late-stage quartz (\pm accessory phases). This magmatic zonation is thought to be the result of a process called *constitutional zone refinement*, where rapid crystallisation from the margin of a pegmatite body expels incompatible elements into a boundary layer between the crystallising front and the residual melt, which can become periodically saturated in exotic phases (e.g., tourmaline, phosphates, etc.) thereby generating a mineralogic zoning (London, 2014, 2018). A pegmatite is considered to be complexly zoned when metasomatic processes such as albitisation (e.g., Ballouard et al., 2020; Kaeter et al., 2018; Černý, 1991a), and

greisenisation (e.g., Černý, 1989; Fuchsloch et al., 2018; Goodenough et al., 2025) lead to a patchy and irregular replacement of earlier (magmatic) phases (Černý, 1991a, , Figure 1.3). Although zoned pegmatites can be economic (e.g., Tanco, Canada; Stilling et al., 2006), the majority of economic deposits of lithium are poorly- or un-zoned (and referred to as simple albite-spodumene pegmatites, Černý, 1991b).

1.2.2 Pegmatite classification schemes

Pegmatites come in many shapes and sizes, and are commonly found in swarms of 10s to 100s of individual bodies (Černý, 1991b). Several attempts have been made to group pegmatites into meaningful classes. This has been attempted on the basis of formation depth (e.g., Ginsburg and Rodionov, 1960; Ginsburg, 1984; Ye. Zagorsky et al., 2003), mineralogy (e.g., Wise, 1999; Wise et al., 2013), and geochemistry (e.g., Niggli, 1920; Landes, 1933). Those that have caught most traction incorporate all components (Černý and Ercit, 2005; Wise et al., 2022b). Pegmatites are therefore typically divided into two major groups based on geochemical affinity and mineralogy (thought to result from variations in source and depth, Muller et al., 2025).

- The LCT (Černý and Ercit, 2005) or Group 1 (Wise et al., 2022b) pegmatites are commonly enriched in lithium, caesium, and tantalum, however can also have economic quantities of rubidium, beryllium, and tin (McCaffrey and Jowitt, 2023). Group 1/LCT pegmatites are classically thought to crystallise from melts sourced (either indirectly, and staged through a parental granite, or directly) by partial melting of a metasedimentary source, and therefore have S-type (peraluminous) geochemical affinities (Chappell and White, 1992; Černý et al., 2012).
- The NYF (Černý and Ercit, 2005) or Group 2 (Wise et al., 2022b) pegmatites are commonly enriched in niobium, yttrium, and fluorine, and may additionally contain economic quantities of REEs (McCaffrey and Jowitt, 2023). Group2/NYF pegmatites are considered to have a metagneous (i.e. amphibolite) source and occur in either anorogenic (A-type) or orogenic (I-type) settings (Černý et al., 2012).

An uncommon third group (Group 3) is distinguished by Wise et al. (2022a), which is moderately enriched in aluminium, beryllium, and boron. Other classification schemes (e.g., Černý and Ercit, 2005) often include an “Abyssal” or “Barren” group defined by an absence in exotic minerals that may enable a classification into a

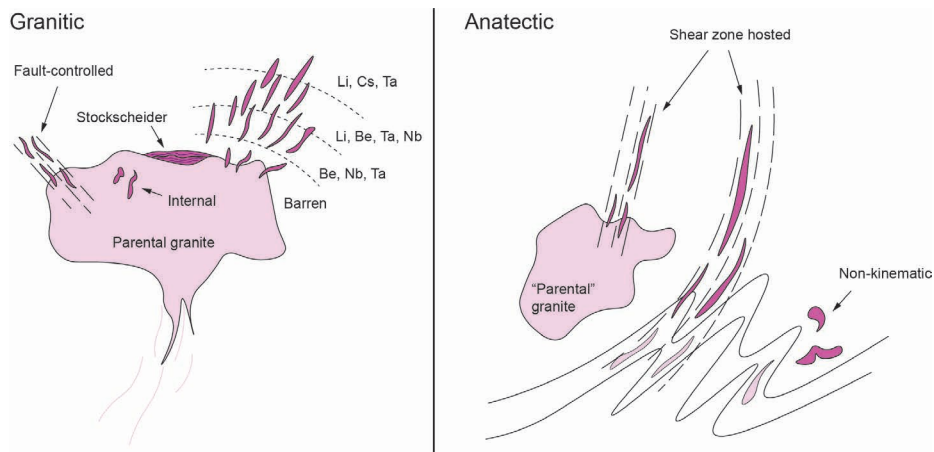


Figure 1.4: Schematic diagram outlining the two main models proposed for pegmatite formation. In the granitic model (left), fault controlled, internal, stockscheider, and regionally zoned (after Černý, 1991b) varieties are shown. In the anatectic model (right), both meta-igneous and metasedimentary sources are shown. Not to scale.

different group, however the distinction as to when a pegmatite becomes sufficiently enriched so as to become an LCT or NYF is then unclear. Wise et al. (2022b) suggests that all pegmatites can be classified based on the mineralogy of their border zone (Figure 1.3), and should therefore always be classed by their mineralogical and geochemical affinity.

Hard rock lithium deposits (including albite-spodumene pegmatites) fall under the LCT/Group 1 pegmatite class. These pegmatites typically contain quartz, alkali-feldspar, plagioclase (albite), and muscovite. Economic lithium deposits will also contain the lithium ore minerals spodumene ($\text{LiAlSi}_2\text{O}_6$, with an ideal Li_2O content of 8.0 wt. %), and/or petalite ($\text{LiAlSi}_4\text{O}_{10}$, with an ideal Li_2O content of 4.5 wt. %) (Bowell et al., 2020). Whether spodumene or petalite is the dominant lithium aluminosilicate is thought to be dependent on the P–T conditions of crystallisation (London, 1984). In either case, the melt must become saturated in spodumene or petalite, which occurs at c. 5000–10000 ppm Li (Maneta et al., 2015). Whether an economic lithium pegmatite is zoned or unzoned (Figure 1.3) is therefore dependent on the initial Li concentration of the melt which controls the amount of crystallisation required to achieve saturation.

1.2.3 Formation models

Two main formation models exist for LCT pegmatites; a granite fractionation model (cf. London, 2005; Černý et al., 2012; Cameron et al., 1949; Černý, 1991b; Černý and

Ercit, 2005; Jahns, 1953; London, 2018), and an anatectic model (cf. Stewart, 1978; Shaw et al., 2016; Müller et al., 2017; Simmons et al., 1995).

Under the granite fractionation model (Figure 1.4), pegmatites are considered to be the result of extensive (>90%) fractionation of a parental granitic melt, resulting in a volatile- (flux) and incompatible-rich residual melt (London, 2005). The residual melt may rise to the top of the granitic magma chamber, forming enriched cupolas and stockscheider intrusions (e.g., Müller et al., 2025; Roda-Robles et al., 2018, Figure 1.4). Under a suitable stress regime, low viscosity pegmatitic melts (London, 2014) can be ejected from the parental granite and generate a regional zonation pattern around a parental granite (Brisbin, 1986; Černý, 1991b). Similarly, shear zones may focus melt away from the source granite in a spatially restricted direction (Müller et al., 2025). In exceptional scenarios, fractionation can thus be traced from the primitive cumulate at the base of large granitic intrusions through to genetically related pegmatites using geochemical tracers in accessory phases (e.g., Liu et al., 2023).

By contrast, the anatectic model proposes that pegmatites can form directly from low-degree partial melting of a fertile (in the case of LCT, metasedimentary) source, forming a similarly volatile- and incompatible-rich melt (Simmons et al., 1995; Shaw et al., 2016; Müller et al., 2017, Figure 1.4). The low degree of melt promotes high degrees of enrichment, and the composition of the melt is controlled by restitic assemblage (e.g., Kunz et al., 2022; Knoll et al., 2023). The melt can then either migrate passively, or be concentrated along a shear zone prior to emplacement (Bhatt et al., 2019; Müller et al., 2025; Demartis et al., 2011). This model is mostly invoked in areas with no apparent parental granite (Shaw et al., 2016; Müller et al., 2017), or in areas where crystallisation ages of pegmatites are younger (10–50 Myr) than neighbouring granites (Brou et al., 2022; Goodenough et al., 2014; Zhang et al., 2016; Liu et al., 2025).

Although strong cases have been made for both models for pegmatites in general, the application of these models to economically extractable pegmatite deposits is uncertain, as the degree of enrichment either model can invoke is not well constrained. Given the extreme levels of enrichment required to generate spodumene (5000–10000 ppm Li, Maneta et al., 2015) from incipient crystallisation (in the case of albite-spodumene pegmatites), a petrogenetic model which can suitably explain the formation of economic hard rock lithium deposits must be able to account for this (Gardiner et al., 2024b).

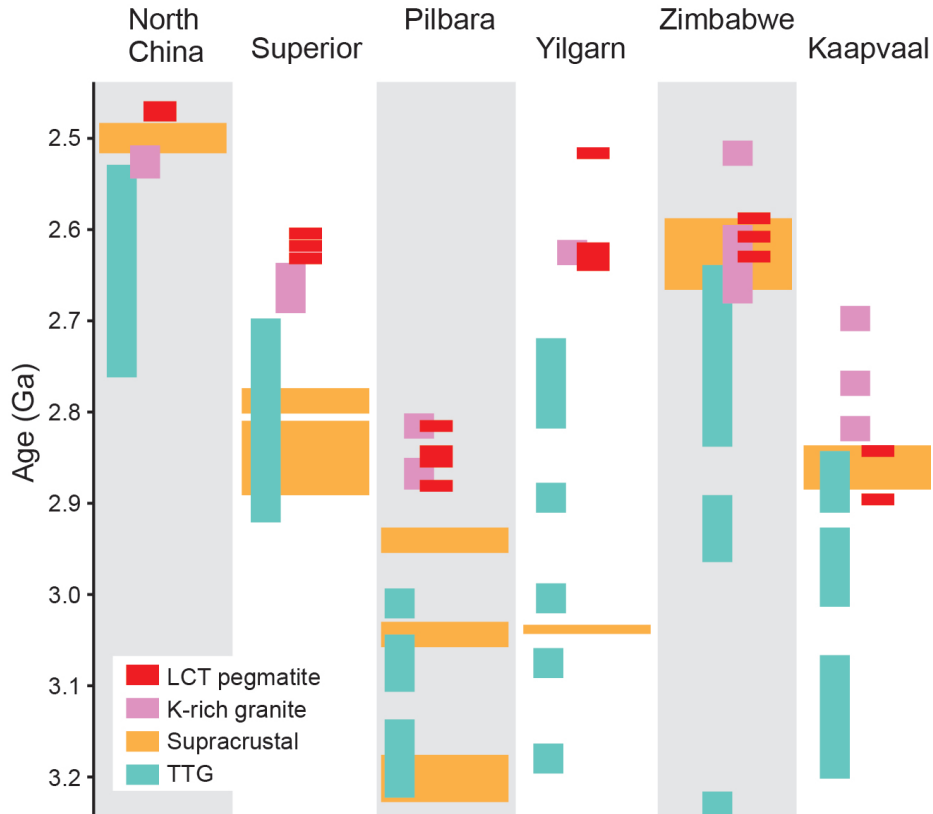


Figure 1.5: Time–space plot for felsic intrusive and siliciclastic supracrustal events in the North China, Superior, Pilbara, Yilgarn, Zimbabwe, and Kaapvaal cratons. Modified from Gardiner et al. (2024b). TTG: tonalite–trondhjemite–granodiorite.

1.3 The Archean paradox for lithium enrichment

Archean terranes are the current largest single source of battery grade lithium, and are therefore arguably the most prospective lithium provinces (Bowell et al., 2020). Understanding Archean crustal formation and evolution is therefore a necessary prerequisite to understanding how and why these terranes are so well endowed in lithium.

Archean cratons fundamentally develop through a protracted (or multiple discrete) period(s) of juvenile continental crust (termed tonalite–trondhjemite–granodiorite, or TTG) formation during their early history to form cratonic nuclei (Cawood et al., 2022; Brown et al., 2020, Figure 1.5). These cratonic nuclei subsequently assembled, thickened, and stabilised during the late Archean to form stable cratonic blocks, which is commonly linked to horizontal plate motions, a defining feature of plate tectonics (Palin et al., 2020; Cawood et al., 2018). This cratonic assembly coincides with the first appearance of potassium-rich granites in the late Meso- to Neoproterozoic, which is possibly driven by a shift in geodynamics (Gardiner et al., 2024b; Nebel et al., 2018,

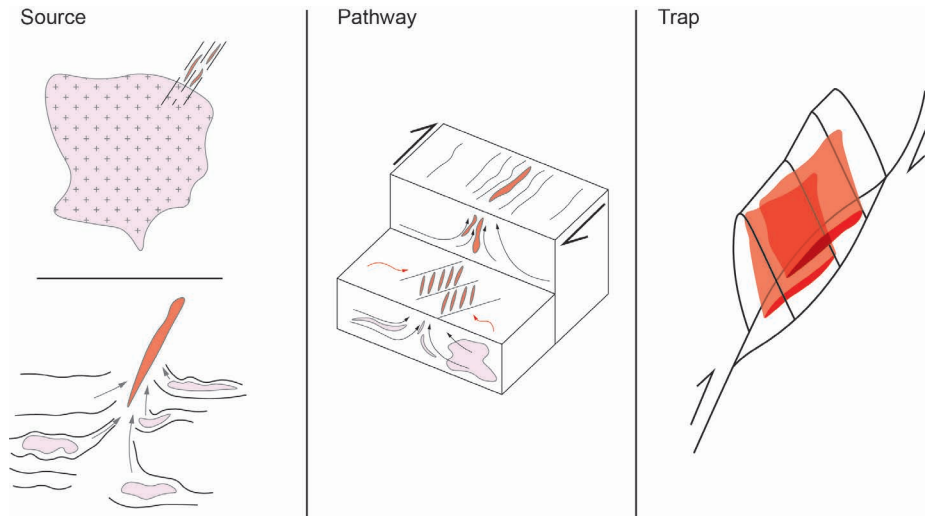


Figure 1.6: A simplified diagram denoting the three main stages (source, pathway, and trap) of the mineral system model (in sequential order from left to right) studied in this thesis.

Figure 1.5). In many cases, this is also when LCT/Group 1 pegmatites first occur (Figure 1.5). Assuming LCT/Group 1 pegmatites must incorporate metasedimentary material into their source, they may therefore signify that crustal recycling has matured in the cratonic block (Kendall-Langley et al., 2020; Gardiner et al., 2024b; Laurent et al., 2014; Dittrich et al., 2019).

However, in some cratons, such as the Zimbabwe or Kaapvaal, pegmatites form contemporaneously with (or precede) the formation of major sedimentary cover successions (Figure 1.5). This raises questions as to how sedimentary material could have been incorporated into the melt source.

This is the Archean lithium paradox — do LCT/Group 1 pegmatites record recycling of continental crust through a cycle of erosion and deposition, followed by burial and melting, or are there other mechanisms to generate LCT/Group 1 signatures?

1.4 Towards a mineral system model for lithium

To better explore for lithium-rich pegmatites, a mineral system model (Wyborn et al., 1994; Hagemann et al., 2016) must be developed which accounts for the paradox described above (Figure 1.6). The mineral system model approach for use in exploration for other commodities arose due to the observation that many mineral deposits have distinct signatures on various geologic scales, and are commonly clustered into provinces that are particularly enriched in select elements (e.g., Carlson, 1991). In

its simplest sense, a mineral system model for lithium must appropriately delineate a source which can sufficiently enrich a melt in lithium to form albite–spodumene pegmatites; a pathway that enables extraction and migration of the melt from the source region to final emplacement; and a trap that controls the spatial distribution and orientation of individual pegmatite intrusions (Figure 1.6). Subsequent mineralisation and preservation of the ore-forming minerals of interest during later processes are pivotal for the ore deposit to persist through time.

1.5 Research motivation

In this thesis, I use the Zimbabwe craton as a case study to test our understanding of lithium pegmatite petrogenesis, and thus contribute towards a robust mineral system model for economic hard rock lithium deposits. Given Zimbabwe is particularly well endowed in lithium (supplying 9% of global production in 2024; US Geological Survey, 2025), and it contains a craton in which siliciclastic-dominant supracrustal belts are unlikely to have been a source component in the melting environment (Figure 1.5), the Zimbabwe craton is an excellent natural laboratory to challenge our current petrogenetic models for lithium-rich pegmatites.

In order to do this, a more holistic understanding of the broader Archean crustal evolution in Zimbabwe is required, which better integrates the intrusive (TTG) history to the tectonic models developed primarily from studies on the supracrustal successions (e.g., Wilson et al., 1995; Jelsma et al., 2021). New tools have recently been developed that allow us to do this more efficiently, and in this thesis I will integrate detailed geochronological and geochemical analyses and state-of-the-art thermodynamic models (e.g., MAGEMin, Riel et al., 2022), with traditional field-based structural studies to develop an understanding of the relative timing of lithium-rich pegmatite formation in the broader cratonic evolutionary context.

More broadly, this research will therefore provide insight into whether sedimentary material is a necessary component in the source of LCT/Group 1 pegmatites, and demonstrate how a more robust understanding of the regional geologic (and structural) history can enable more efficient exploration targeting for deposits with small surface footprints.

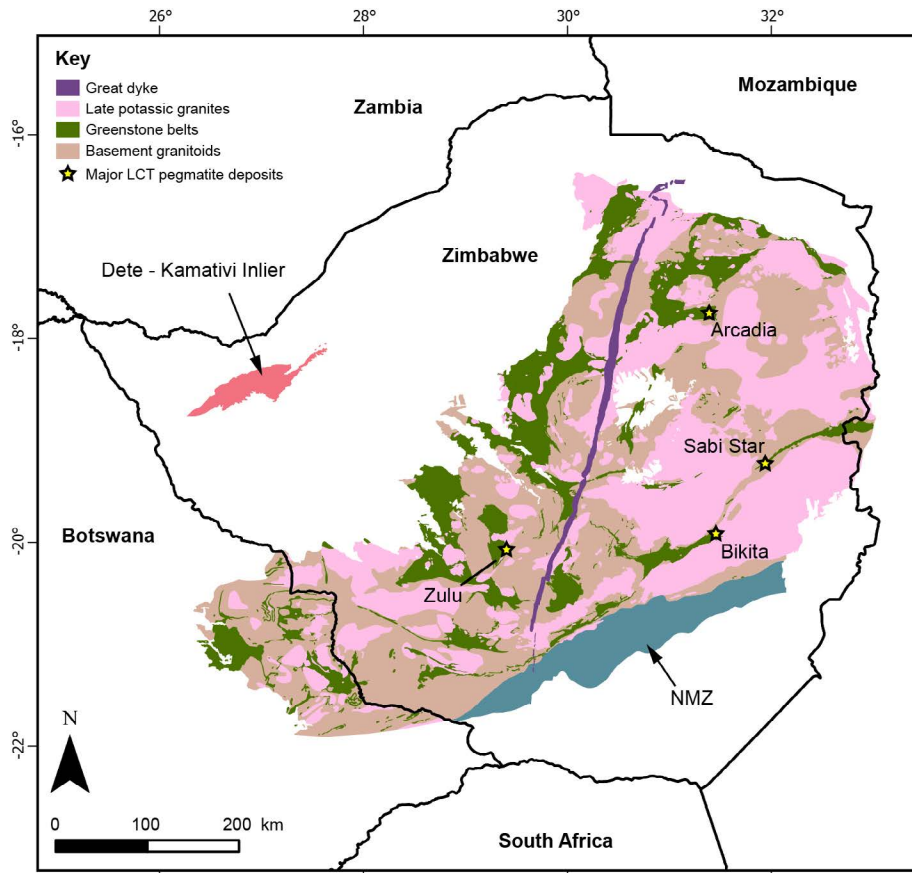


Figure 1.7: Simplified geological map of the Zimbabwe craton (modified from Ncube, 1994). Both the central craton, northern marginal zone (NMZ) and the Dete–Kamativi inlier are highlighted. Cover sequences and bounding orogens are white.

1.6 The Zimbabwe craton

The Zimbabwe craton is an Archean cratonic block with a present-day surface area of c. 300 000 km² (Figure 1.7, Jelsma et al., 2021; Ncube, 1994). Extending from north-east Botswana, through Zimbabwe, into western Mozambique, the craton is composed of Eo- to Neoproterozoic TTGs, Neoproterozoic potassic granites, and a series of greenstone belts that are scattered across the craton (Hofmann et al., 2022; Macgregor, 1951; Moor bath et al., 1976; Taylor et al., 1991; Wilson et al., 1995; Blenkinsop et al., 1997). The Zimbabwe craton is notably one of the most fertile terranes in Africa for hard rock lithium deposits (Goodenough et al., 2025).

The craton is bound by five orogenic belts; the Limpopo belt (c. 2000 Ma) to the south (Kamber et al., 2015), the Magondi belt (c. 1800-2000 Ma) to the west (Master et al., 2010), the Pfunzi (2600 Ma, Vinyu et al. 2001) and Zambezi belts (c. 500-850

Ma) to the north (Hanson et al., 1994; Goscombe et al., 2020; Kuribara et al., 2019), and the Mozambique belt to the east (Chauue et al., 2017). Three major periods of TTG magmatism have been identified during cratonic growth; a 3.6–3.2 Ga magmatic episode (Horstwood et al., 1999; Dodson et al., 2001; Taylor et al., 1991; Wilson et al., 1995; Hofmann et al., 2022), a 3.0–2.85 Ga episode (Wilson et al., 1995; Jelsma et al., 2021), and major craton expansion at 2.7–2.6 Ga (Wilson et al., 1995; Chagondah et al., 2023; Jelsma et al., 1996; Blenkinsop et al., 1997). Cratonisation culminated in the emplacement of voluminous potassic granites (2.68–2.54 Ga, Jelsma et al. 2021), and finally the emplacement of the Great Dyke (2.575 Ga, Oberthur et al., 2002).

Each intrusive event has a corresponding extrusive (predominantly mafic volcanic) succession preserved in the greenstone belts throughout the craton. These were initially subdivided into the Magnesian (Sebakwian), Lower Series (Bulawayan), and Upper Series (Shamvaian) supergroups (in younging order, Macgregor, 1947, Stageman, 1978), Later revision by Wilson (1979) further subdivided the Bulawayan supergroup into the Lower and Upper Greenstones, proposing a regional unconformity between the two. Further work, most notably in the Belingwe greenstone belt (Bickle et al., 1993; Martin et al., 1993) introduced a Belingwean succession, which encompassed the majority of the Lower Greenstone succession of Wilson (1979) (Wilson et al., 1995). More recent work has proposed a two final stratigraphic divisions; in the upper Bulawayan, a distinction was introduced between the Western and Eastern succession, largely driven by difficulties in correlating the stratigraphy between greenstone belts along the western margin and central parts of the craton (e.g., Wilson et al., 1995; Jelsma et al., 2021), and a separation of the Shamvaian (the major supracrustal siliciclastic cover sequence, Figure 1.5) into an upper and lower Shamvian (Wilson et al., 1995). Given inconsistent use of the stratigraphic nomenclature of Zimbabwe, an outline of the stratigraphy as used in this study is outlined in Section 1.6.1.

Over the past half century, the tectonic evolution of the Zimbabwe craton has been the focus of considerable debate, with two contrasting models having been proposed. A vertical accretion and growth model is favoured, in which the stratigraphy of the greenstone belts can be correlated across the craton (e.g., Bickle et al., 1993; Blenkinsop et al., 1993; Hunter, 1998; Macgregor, 1951; Wilson et al., 1995), whereas an allochthonous, horizontal accretion model was spearheaded by Kusky and Kidd (1992, and supported by other authors (e.g., Dirks et al., 2002; Hofmann et al., 2001; Jelsma and Dirks, 2002; Kusky, 1998; Hofmann and Kusky, 2004).

Further uncertainty in the Archean evolution of the craton is the nature of the Northern Marginal Zone (NMZ), which is exposed along the south-eastern margin of the craton and has classically been considered part of the Limpopo belt (Fedó et al., 1995; Kamber and Biino, 1995; Rollinson and Blenkinsop, 1995). However, reinterpretation of earlier work (Frei et al., 1999; Berger et al., 1995; Blenkinsop, 2004; Rollinson and Whitehouse, 2011) by Rollinson (2022) proposed that the NMZ is a deep-crustal equivalent of the Zimbabwe craton. Similarly, recent geochronology has revealed that the Dete-Kamativi inlier (Figure 1.7), previously considered to have exclusively Magondi-age protoliths (Master et al., 2010) also contains Archean protoliths, potentially expanding the footprint of the craton (Glynn et al., 2020).

Whilst recent reviews of geology of the Zimbabwe craton have primarily focused on reconciling the supracrustal stratigraphy (e.g., Jelsma et al., 2021), attempts at fully integrating the intrusive history of the craton remain limited, largely due to the limited number of coupled geochronological and geochemical analyses on the intrusive suites. Given the ongoing debates concerning the genesis of TTG magmas during the Archean, which form major components of all cratons worldwide, alongside their implications for the geodynamic characteristics of the early Earth (cf. Palin et al., 2020), improving our understanding of TTG formation in the Zimbabwe craton may shed light on evolution and stabilisation of the Zimbabwe craton as a whole (e.g., Bédard, 2018; Moyen and Martin, 2012; Wyman, 2013).

Central to the debate around the stabilisation of Archean cratons is whether potassic granites, are preceded by continental emergence and sedimentation, or else form due to crustal thickening, which promotes sedimentation and the formation of supracrustal cover sequences (Chowdhury et al., 2021; Reimink and Smye, 2024; Roberts, 2024). In the Zimbabwe craton potassic granites cover c. 50% of the presently exposed area (Blenkinsop and Treloar, 2001); thus, investigation of the genesis of these evolved melts will shed light on the nature of tectonics in and around the craton, and has implications for lithium enrichment. Given lithium-rich pegmatites are thought to require metasedimentary rocks in their source region (in both granite-fractionation and anatexis models) to generate their peraluminous signatures (Černý, 1991a; Černý et al., 2012; Martin and De Vito, 2005), the absence of voluminous sedimentary successions that represent potential source material (Jelsma et al., 2021, Figure 1.5) provides an interesting conundrum that needs to be resolved (Section 1.3).

1.6.1 Adopted stratigraphic nomenclature

This study adopts the following lithostratigraphic classifications (adapted from Blenkinsop et al., 1997 and Jelsma et al., 2021):

- Sebakwian group: The oldest greenstone belt group, comprised of tightly folded remnants within the Tokwe and Rhodesdale terranes.
- Belingwean group: Named after the Belingwe greenstone belt which contains the type section, this group encompasses stratigraphy deposited or erupted between c. 3.1–2.8 Ga. Here, this group is expanded to include the Buhwa stratigraphy, which otherwise is excluded from traditional stratigraphic models.
- Bulawayan group: Largely consistent with that defined by Wilson et al. (1995) to comprise of the Koodoovale–Manjeri–Reliance–Zeederberg–Cheshire succession (as seen in the Belingwe greenstone belt, Martin et al., 1993), and equivalent formations in other greenstone belts that were deposited c. 2.75–2.70 Ga (Hofmann et al., 2024; Prendergast and Wingate, 2013).
- Western Succession: Broadly consistent with that defined by Wilson et al. (1995) and Jelsma et al. (2021), and comprises of the successions along the western margin of the Sebakwe protocraton which do not directly correlate to the Bulawayan group (discussed in detail below).
- Shamvaian group: The final supracrustal succession in the stratigraphy of the Zimbabwe craton, largely comprised of siliciclastic sediments deposited unconformably above earlier groups within syntectonic basins (discussed in detail in Chapter 5).

1.7 Thesis structure

This thesis will develop the building blocks for a mineral system model by working towards unraveling the mechanisms which can sufficiently enrich lithium in the source, and how these melts can subsequently migrate and be emplaced into sufficiently large deposits to form economic bodies. This will be done by using the Zimbabwe craton as a case study.

I will first outline the methods that were used to do this work (Chapter 2), followed by a modelling-based approach to testing various source models for lithium-rich pegmatites (Chapter 3). In Chapter 4, I will present a focused case study on the Zulu

pegmatite field (Figure 1.7), constraining the timing of and controls of emplacement relative to regional deformation events. In Chapter 5, I will present the results of the geochronological and geochemical campaigns completed on samples from the Zimbabwe craton. The complete work is synthesised into a revised tectonic model for the Zimbabwe craton, integrating the new data collected with previous work. Finally, the main findings of the thesis are summarised in Chapter 6.

Chapter 2

Methods

2.1 Dataset compilation

A dataset of published geochemical and geochronological data was compiled on relevant samples collected from the Zimbabwe craton. Samples were grouped into basement (i.e., granitoids) and greenstone (i.e., volcanic and sedimentary cover successions), and where available all possible context (including ages) for the sample retained. Geochemical data were compiled into a standardized reporting style. In some instances, the same samples were measured multiple times in different studies – in these cases, the most precise analyses were incorporated into the compilation. I also expanded the geochronological compilation presented in Jelsma et al. (2021), with analyses published since 2021, and included this into the new compilation

In total, the present compilation (including samples analysed in this study) comprises of 851 greenstone samples, and 300 basement samples. A total of 382 ages have been compiled, including 141 U–Pb magmatic crystallization ages.

2.2 Sample collection and preparation

Samples of basement units were collected during field campaigns to Zimbabwe in 2022, 2023, and 2024 (Appendix A). Sample collection focused on regions with limited published U-Pb geochronology and/or geochemistry and aimed to capture the full suite of magmatic events within the craton based on literature and observed field relations. Individual samples were 2–3 kg, and sampled localities logged using a Garmin GPSMap 62s and photographed using a Nikon D7200 DSLR camera.

Initial sample preparation was conducted at the Department of Earth Sciences, University of Oxford, UK. Samples were catalogued and imaged before being split using a hydraulic press. An off cut of each sample was crushed using a steel jaw

crusher to produce fragments <5 mm in diameter. Thirty grams of material was separated using a riffle splitter and placed into agate mortars with 10-mm agate balls for pulverisation. Samples were milled to <200 μm using a planetary ball mill, running for 20 minutes at 200 rpm and switching rotation direction every five minutes.

For *ex situ* geochronology, 500 g aliquots of crushed (not pulverised) material were dry-sieved to +100/-212 and +212/-500 μm and washed to remove fine residue. The sieved material was then panned under ethanol, and the heavy fraction picked under a binocular microscope. The mineral separates were then placed neatly on double sided electrical tape and mounted in Metprep EPO-FLO resin to produce standard 25 mm mineral mounts. Once cured, the mineral mounts were ground down to expose the cores of the mineral separates using grit paper, and polished in stages using 9 μm , 6 μm , 3 μm , and finally 1 μm polycrystalline diamond suspension paste on cloth lapping plates.

2.3 Geochemistry

2.3.1 X-ray fluorescence

X-ray fluorescence (XRF) was performed at Franklin and Marshall College, USA. One gram of each sample was dehydrated at 950 °C in a muffle furnace to calculate initial loss on ignition. A 400 mg aliquot was subsequently hand-mixed with 3600 mg of lithium tetraborate flux. The mixtures are then transferred to a Spex Mixer Mill for 10 minutes to homogenise. The mixtures were heated over a Meeker burner in a ~95% Pt – 5% Au capped crucible with the addition of a 2% solution of lithium iodide (to reduce viscosity) until fully liquefied. The samples were then quenched on the crucible lid to form the flux beads for subsequent analyses. X-ray fluorescence was performed on a Malvern PANalytical Zetium XRF to calculate the major oxide composition of each sample – i.e. SiO_2 , Al_2O_3 , $\text{Fe}_2\text{O}_3(\text{T})$, MnO , K_2O , Na_2O , TiO_2 , P_2O_5 , CaO , MgO .

2.3.2 ICP-MS

Trace element concentrations of samples were obtained via inductively-coupled-plasma mass spectrometry (ICP-MS) at the Department of Earth Sciences, University of Oxford, UK. Fifty milligram powdered aliquots of each sample underwent a HF:HNO₃ Parr Bomb digestion method (adapted from Okina et al., 2016). Aliquots were measured on a Sartorius five-digit mass balance (precision \pm 0.1 mg) three times, and

average masses were taken forward. Twelve sample aliquots, three blanks, and three standards (USGS standards GSP-2 or G-2) at a time were subsequently placed into 3mL octagonal Teflon vials (hex nuts) and 400 μL 29M HF + 600 μL 16M HNO_3 added to each vial prior to capping. Three hex nuts were then placed into each PTFE liner, to which 5 mL 8M HNO_3 was added to equalize pressure within the hex nuts during digestion. The PTFE liners were capped and placed into the steel jacket of Parr digestion vessels, screwed tight, and heated to 196 °C for 18 hours.

Once cooled, hex nuts were extracted from the digestion vessels, uncapped and heated on a hot plate to dryness. 500 μL 16M HNO_3 was then added to the dried residue and heated to dryness three times to dissolve any fluorides produced during the initial digestion phase. Dry downs were performed in ISO Class 5 laminar flow hoods within ISO Class 7 clean laboratories to minimize contamination to the samples. Following the drying steps, the residue was dissolved by adding 1 mL 8M HNO_3 to the hex nuts and capped, placed back into PTFE liners (to which 5 mL 8M HNO_3 was added), and sealed into Parr digestion vessels. The digestion vessels were left in an oven for 12 hours at 180 °C. Upon extraction from the digestion vessels, the sample solutions were visually assessed for clarity. If solutions were not clear, the digestion was re-run. Otherwise, the solution was dried down and re-diluted in 2% (w/w) HNO_3 for ICP-MS analyses.

An ESI PrepFAST autosampler and autodiluter was used to manage throughput, additional dilution, and the addition of internal analytical standards (rhenium and rhodium) to correct for instrumental drift. Samples, standards, and blanks were analysed for a suite of 43 trace and rare earth elements using a PerkinElmer NeXION 350D quadrupole ICP-MS. Precision across the run was calculated using the relative standard deviation (RSD), which was consistently <5% for most elements, and all below 10%. USGS standards GSP-2 and G-2 were used to verify complete digestion of refractory phases (e.g., zircons), as these minerals typically control the whole-rock budget of key trace elements (e.g., REEs, Zr, Hf). The detection limit was determined as ten times the standard deviation of ten back-to-back blank measurements and is typically <0.1 ppm. Analytical accuracy was assessed using the river reference standard SLRS-6, with external analytical uncertainties of <5%.

2.4 Geochronology

2.4.1 BSE Imaging

For *ex situ* analyses, samples were imaged to aid spot-selection using an FEI Quanta 650 FEG scanning electron microscope (SEM) equipped with a lens-mounted concentric backscatter detector (CBS) and a single energy dispersive (ED) Oxford Instruments X-Max50 spectrometer. The detectors are controlled by Oxford Instruments AZtec software, and spectral resolution is calibrated periodically to the position and intensity of a Mn $K\alpha$ peak (5.895 keV) achieving <124 eV at 20 kcps (kilo-counts per second). A specimen working distance of 10 – 12 mm was used. Mosaic-stitched backscatter electron (BSE) images were collected at an accelerating voltage of 20 kV, spot 4.0 with the final lens aperture of 30 μm , yielding a spot size of ca. 6 μm and X-ray output rates up to 35 kcps. BSE images were stitched using the AZtec software which controls the instrument.

2.4.2 Zircon geochronology

Zircons from basement units in the Zimbabwe Craton were analysed across seven analytical sessions *ex situ* from resin mounts by laser ablation inductively-coupled-plasma mass spectrometry (LA-ICP-MS). Analyses were carried out at the St Andrews Geochronology (StAGE) Laboratory using a New Wave UP-213 Nd:YAG 213 nm ablation system. The ablated material was sent to an Agilent 8800 ‘triple quadrupole’ QQQ-ICPMS for simultaneous U–Pb isotopic and trace element analysis. The aerosol produced during ablation was carried to the mass spectrometer via an Ar-He (880 ml min^{-1} Ar, 775 ml min^{-1} He) mixture with N_2 gas (10 ml min^{-1}) added to the aerosol stream to enhance sensitivity. Individual zircons were ablated using a 30 or 40 μm static circular spot, with the laser repetition rate set to 10 Hz, 45% power, yielding an approximate fluence of 3.5–4.5 J cm^{-2} measured at the sample.

Natural zircon reference materials were measured approximately every 10 unknowns during each analytical run and the drift between reference material brackets was corrected using a step–forwards function.

Dwell times for U, Th, and Pb isotopes on the Agilent were 10 ms for ^{202}Hg and ^{232}Th , 30 ms for ^{204}Pb , 40 ms for ^{206}Pb , 70 ms for ^{207}Pb , and 20 ms for ^{208}Pb and ^{238}U . Measurements were collected using 2 s of background measurement, 30 s of ablation time, with a 30 s washout between samples. Isotopic data were reduced with the Iolite 4 software package (Paton et al., 2011) using the U-Pb data reduction scheme (DRS; Paton et al., 2010) with a smooth cubic spline to model down-hole fractionation.

Natural zircon 91500 (Horstwood et al., 2016; Wiedenbeck et al., 1995) was used as a primary standard to calibrate downhole fractionation of U and Pb. Reference zircons Maniitsoq (Marsh et al., 2019), OG1 (Stern et al., 2009) and Plešovice (Sláma et al., 2008) were used as secondary standards to validate U–Pb data. Final U–Pb ages were calculated using IsoplotR (Vermeesch, 2018), and analyses were considered if these defined a statistical uniform population. The ages for unknowns are reported with 2 SE propagated uncertainties calculated by *iolite* and are uncorrected for common Pb.

Dwell times for trace elements were 10 ms for ^{85}Rb , ^{153}Eu , and ^{157}Gd , 20ms for ^7Li , ^{118}Sn , ^{139}La , ^{140}Ce , ^{141}Pr , ^{146}Nd , and ^{147}Sm , and 50 ms for ^{31}P , ^{43}Ca , ^{47}Ti , ^{49}Ti , ^{89}Y , ^{91}Zr , ^{93}Nb , ^{159}Tb , ^{163}Dy , ^{165}Ho , ^{166}Er , ^{169}Tm , ^{172}Yb , ^{175}Lu , ^{178}Hf and ^{181}Ta . Measurements were collected using 2s of background measurement, 30s of ablation time, with a 30s washout between samples. Isotopic data were reduced with the *Iolite 4* software package (Paton et al., 2011) using their trace element DRS, normalized to natural zircon 91500.

2.4.3 Titanite geochronology

In-situ analyses on syn-kinematic titanite crystals were conducted on 30 μm polished thin sections of samples from within the Irisvale–Lancaster shear zone. Titanite crystals were analysed for simultaneous U–Pb geochronology and trace element geochemistry by LA-ICP-MS at the Institute for Geochemistry and Petrology, ETH Zurich, Switzerland, using an ASI RESOLUTION S-155 193-nm ArF excimer laser ablation system connected to a Thermo Element XR sector-field ICP-MS. The titanites were ablated using a 29 μm static circular spot, with laser fluence of $\sim 2.5 \text{ J cm}^{-2}$, at a repetition rate of 4 Hz. Individual analyses began with 28 s of gas blank followed by 30 s of ablation. The ablated material was mixed in the S-155 Laurin Technic ablation cell with a carrier gas composed of helium (ca. 0.5 L min^{-1}) and a make-up gas composed of argon (ca. $0.9\text{--}0.99 \text{ L min}^{-1}$) and nitrogen (2 mL min^{-1}), homogenized, and introduced to the plasma torch.

The following masses were measured: ^{25}Mg , ^{27}Al , ^{29}Si , ^{31}P , ^{43}Ca , ^{45}Sc , ^{49}Ti , ^{51}V , ^{55}Mn , ^{57}Fe , ^{65}Cu , ^{66}Zn , ^{88}Sr , ^{89}Y , ^{90}Zr , ^{93}Nb , ^{115}In , ^{118}Sn , ^{138}Ba , ^{139}La , ^{140}Ce , ^{141}Pr , ^{146}Nd , ^{147}Sm , ^{153}Eu , ^{157}Gd , ^{159}Tb , ^{163}Dy , ^{165}Ho , ^{166}Er , ^{169}Tm , ^{172}Yb , ^{175}Lu , ^{178}Hf , ^{181}Ta , ^{182}W , ^{202}Hg , ^{204}Pb , ^{206}Pb , ^{207}Pb , ^{208}Pb , ^{232}Th , ^{235}U , and ^{238}U . Dwell times for ^{206}Pb and ^{207}Pb was 50 ms, 20 ms for ^{238}U , and for all other masses was 11 ms.

The U- and Pb-isotopic ratios were derived using MKED-1 ($1517.32 \pm 0.1 \text{ Ma}$; Spandler et al., 2016) as primary reference material and the UcomPbine data reduction scheme available in *Iolite 4* (Chew et al., 2014; Paton et al., 2011; Petrus

and Kamber, 2012). Final titanite U–Pb ages were determined using unanchored regressions in IsoplotR (Vermeesch, 2018) via linear regression on Tera-Wasserburg concordia plots, with the lower intercepts interpreted as crystallization ages. Analytical accuracy for U–Pb was assessed by measuring BLS (1049 ± 1.3 Ma; Aleinikoff et al., 2007), 94-35 (51.5 ± 0.7 Ma; Dana et al., 2023), Otter West (147.9 ± 1.2 Ma; Butler et al., 2002), Ecstall West (91.5 ± 1.0 Ma; Butler et al., 2002), Khan River (516.3 ± 1.3 Ma; Mazoz et al., 2022), and Bear Lake (1067.8 ± 0.7 Ma; Mazoz et al., 2022) as secondary standards. Trace elements were calibrated against NIST SRM610 (Jochum et al., 2011) as primary reference material, and validated by measuring GSD-1G (Guillong et al., 2005) as the secondary reference material.

2.4.4 Apatite geochronology

Apatite grains from samples GW06 and NG01 were analysed from mineral mounts as the samples had no identifiable and/or measurable zircon crystals. Apatite crystals were analysed for U–Pb geochronology and trace element geochemistry by LA-ICP-MS using the same analytical equipment described in section 2.4.3. Apatite crystals were ablated using a $43 \mu\text{m}$ static circular spot, with a laser fluence of $\sim 3 \text{ J cm}^{-2}$ and a repetition rate of 4 Hz. Analyses followed the same protocol (28 s of gas blank, 30 s of ablation) as titanite, with a make-up gas composed of argon (ca. 0.25 L min^{-1}) and nitrogen (2 mL min^{-1}).

The following masses were measured: ^{24}Mg , ^{27}Al , ^{29}Si , ^{31}P , ^{43}Ca , ^{49}Ti , ^{51}V , ^{53}Cr , ^{55}Mn , ^{57}Fe , ^{85}Rb , ^{88}Sr , ^{89}Y , ^{90}Zr , ^{93}Nb , ^{137}Ba , ^{139}La , ^{140}Ce , ^{141}Pr , ^{146}Nd , ^{147}Sm , ^{151}Eu , ^{157}Gd , ^{159}Tb , ^{163}Dy , ^{165}Ho , ^{166}Er , ^{169}Tm , ^{172}Yb , ^{175}Lu , ^{178}Hf , ^{181}Ta , ^{202}Hg , ^{204}Pb , ^{206}Pb , ^{207}Pb , ^{208}Pb , ^{232}Th , ^{235}U , and ^{238}U . Dwell times for ^{206}Pb and ^{207}Pb was 50 ms, 20 ms for ^{238}U , and for all other masses was 11 ms.

The U- and Pb-isotopic ratios were calibrated using the Durango apatite (31.44 ± 0.18 Ma; McDowell et al., 2005) as primary reference material using the UcomPbine data reduction scheme in Iolite 4 (Chew et al., 2014; Paton et al., 2011; Petrus and Kamber, 2012). Final ages were determined using unanchored regressions in IsoplotR (Vermeesch, 2018), interpreting the lower intercept of a linear regression in Terra-Wasserburg space as the crystallization age. Analytical accuracy for U–Pb was assessed using Kovdor (377.5 ± 3.5 Ma; Chew et al., 2011), MAD ($473.46 \text{ Ma} \pm 0.0$ Ma; (Thomson et al., 2012), and Ap42 (CITE) as secondary standards. Trace elements were calibrated against G-NIST610 (Jochum et al., 2011) as primary reference material, and validated against GSD-1G (Guillong et al., 2005).

2.4.5 Mica geochronology

Lithium-rich muscovite from within the Zulu pegmatite field (Chapter 4), and biotite from the surrounding host rocks were analysed in-situ on 30 μm polished thin sections. Mica crystals were simultaneously analysed for Rb–Sr geochronology and trace elements by LA-ICP-MS at the Fipke Laboratory for Trace Element Research (FiLTER) at the University of British Columbia Okanagan, Canada following the methodology described in Larson et al. (2023, 2024). Analyses were conducted using an ESL Excimer 193 nm ArF laser equipped with a TwoVol3 sample chamber connected to an Agilent 8900 triple quadrupole ICP-MS. Individual crystals were ablated using a 50 μm diameter circular static spot, a laser fluence of 3.5 J cm^{-2} , and a repetition rate of 8 Hz. The ablated material was carried by a helium carrier gas (0.3 L min^{-1}), and mixed with a make-up gas composed of argon (0.9 L min^{-1}) and nitrogen (25 mL min^{-1}). Analyses consisted of 25 s of ablation, followed by 15 s of washout by a gas blank.

The following masses were measured: ^7Li , ^{24}Mg , ^{47}Ti , ^{49}Ti , ^{56}Fe , ^{57}Fe , ^{85}Rb , ^{133}Cs , ^{139}La , and ^{172}Yb , as well as mass shifted ^{86}Sr , ^{87}Sr , ^{88}Sr (86-102, 87-103, and 88-104) post-reaction with ^{16}O (see Hoggmalm et al. 2017). The dwell time was 120 ms for ^{86}Sr , and ^{87}Sr , 50 ms for ^{85}Rb , ^{87}Rb , and ^{88}Sr , 10 ms for ^7Li , ^{47}Ti , ^{49}Ti , ^{56}Fe , ^{57}Fe , ^{133}Cs , ^{139}La , and ^{172}Yb , and 5 ms for ^{24}Mg .

Primary Rb–Sr isotope normalization was carried out using Mica Mg nanopowder (Hoggmalm et al., 2017), with $^{87}\text{Sr}/^{86}\text{Sr}$ homogeneity verified against the NIST610 glass reference material (Woodhead and Hergt, 2001). Data reduction was done in Iolite 4 (Paton et al., 2011), using an in-house developed data reduction scheme (Larson, 2024) based in part by that described in Redaa et al. (2021). The method was verified through analyses of multiple unknowns for both brown mica and white mica runs, which yielded results within uncertainty of expected ages.

For trioctahedral mica analyses, Mica Fe nanopowder yielded isochron ages of 307 ± 5 Ma (ZU01, MSWD = 1.8, n = 12/14) and 306 ± 6 Ma (ZU05, MSWD = 1.3, n = 10/10) in agreement with the expected age (ca.305 Ma; Rösel and Zack, 2022). The sample measured in the same run as the standard is indicated in parentheses). Mica 1B yielded isochron ages of 996 ± 17 (ZU01, MSWD = 0.7, n = 14/14) and 989 ± 20 Ma (ZU05, MSWD = 0.29, n = 10/10), within error of the reported age (990 ± 5 Ma, Camacho et al., 2020). Mica 1O yielded isochron ages of 1003 ± 15 Ma (ZU01, MSWD = 1.9, n = 13/14) and 984 ± 15 Ma (ZU05, MSWD = 0.98, n = 10/10), within error of the reported age (986 ± 5 Ma; Camacho et al., 2020). QT yielded an isochron age of 1716 ± 20 Ma (ZU01, MSWD = 0.38, n = 14/14), which

overlaps with the corresponding 1720 Ma Ar/Ar age (internal unpublished reference material), and TQ yielded an isochron age of 1737 ± 23 Ma (ZU01, MSWD = 1.4, n = 12/14), which overlaps with the 1720 Ma Ar/Ar age (internal unpublished reference material).

For dioctahedral mica analyses, Mica Fe nanopowder yielded an isochron age of 306 ± 2 Ma (MSWD = 1.9, n = 25/25), in agreement with the expected age. MA1 yielded an isochron age of 349 ± 2 Ma (MSWD = 1, n = 25/25), in agreement with internally accepted ages (340 – 360 Ma). Tanco high-Li white mica yielded isochron ages of 2690 ± 55 Ma (MSWD = 0.35, n = 25/25), within error of the accepted age (2620 ± 20 Ma; Camacho et al. 2012).

Final ages were determined using isochrons in IsoplotR (Vermeesch, 2018).

2.5 Modelling

2.5.1 Petrologic modelling

For Chapter 3, petrological calculations were performed using the modelling program Theriak-Domino (De Capitani and Petrakakis, 2010) in the compositional system MnO — Na₂O — CaO — K₂O — FeO — MgO — Al₂O₃ — SiO₂ — H₂O — TiO₂ — O₂ (MnNCKFMASHTO), which is most suitable for calculating phase equilibria in metasedimentary rock types. MnO and TiO₂ were excluded from modelling of the second stage of melting due to the absence of MnO and TiO₂ as components in the haplogranite melt model (Holland and Powell, 2011; White et al., 2014a). For metasediments, modelling considered the internally consistent thermodynamic dataset ds62 of Holland and Powell (2011), considering the activity-composition (a-x) relations for feldspars (Holland et al., 2022), sapphirine (Wheller and Powell, 2014), magnetite-spinel, ilmenite-hematite (White et al., 2000), garnet, epidote, margarite, muscovite-paragonite, biotite, orthopyroxene, cordierite, staurolite, chlorite, chloritoid, and melt (Holland and Powell, 2011; White et al., 2014c,a). Pure phases considered were kyanite, andalusite, sillimanite, quartz, rutile, titanite, and aqueous fluid (H₂O).

In chapter 3, the modelled melts considered the internally consistent thermodynamic dataset ds633 of Holland et al. (Holland et al., 2018), with the a-x relations for feldspars (Holland et al., 2022), cordierite (White et al., 2014b; Holland et al., 2018), ilmenite-hematite (White et al., 2000), epidote (Holland and Powell, 2011), muscovite-paragonite, biotite (White et al., 2014b), clinopyroxene (Green et al., 2016), garnet, olivine, orthopyroxene, clinopyroxene, spinel, and melt considered (Holland et al.,

2018). Pure phases considered were kyanite, andalusite, sillimanite, quartz, rutile, titanite, and aqueous fluid (H₂O).

In chapter 5, petrological calculations were performed using a Gibbs free energy minimisation approach in the modelling program MAgEMin (Riel et al., 2022) in the Na₂O – CaO – K₂O – FeO – MgO – Al₂O₃ – SiO₂ – H₂O – TiO₂ – O₂ – (Cr₂O₃) (NCKFMASHTO(Cr)) compositional system. Chromium oxide (Cr₂O₃) was only included when modelling (ultra)-mafic rock types to accurately predict the stability of spinel phases. Modelling considered the internally consistent thermodynamic dataset ds636 of Green et al. (2025). Activity-composition (a-x) relations for epidote (Holland and Powell, 2011), clinoamphibole (Green et al., 2016), muscovite (White et al., 2014a), ilmenite/hematite, orthopyroxene, clinopyroxene, garnet (Weller et al., 2024), ternary feldspars (Holland et al., 2022), olivine (Holland et al., 2018), spinel-group minerals (Tomlinson and Holland, 2021) and biotite, cordierite, aqueous fluid, and silicate melt (Green et al., 2025). Kyanite, andalusite, sillimanite, nepheline, quartz, cristobalite, tridymite, coesite, rutile, sphene, and H₂O were treated as pure phases.

For modelling partial melting of hydrated basalts and granitoids, the water content was fixed using the H₂O-saturation tool in MAgEMin in order to minimally saturate the system at the point of incipient melting (cf. Palin et al., 2016a). Modelling phase equilibria in primitive mantle was conducted under anhydrous conditions.

2.5.2 Trace element modelling

Trace element modelling was performed by using phase proportions calculated using the model setups described above. The trace element concentration of the melt during partial melting was calculated using a batch melt model:

$$C_{liq} = \frac{C_{bulk}}{D_{bulk} + F(1 - D_{bulk})} \quad (2.1)$$

where C_{liq} is the concentration of the trace element in the melt, C_{bulk} is the bulk concentration of the trace element, D_{bulk} is the bulk partition coefficient of the residuum assemblage, and F is the melt weight fraction in equilibrium with the residuum. The bulk partition coefficient was calculated through the following equation:

$$D_{bulk} = \sum_{k=0}^n K_n X_n \quad (2.2)$$

where K_n is the mineral/melt partition coefficient of the element of interest for the given phase, and X_n is the weight fraction of the given phase in the system. For modelling of lithium in the melting environment, we used partition coefficients compiled by Simons et al. (2017) and Chakraborty and Upadhyay (2020) (Table 2.1).

2.5.3 Fractional crystallisation modelling

Fractional crystallization was modelled using the Raleigh fractionation equation (Shaw, 1970):

$$C_{liq} = C_0 F^{D_{bulk}-1} \quad (2.3)$$

where C_{liq} is the concentration of the trace element in the residual melt, C_0 is the initial concentration of the trace element in the melt, F is the melt fraction remaining, and D_{bulk} is the bulk partition coefficient of the crystallizing phases. The bulk partition coefficient was calculated using Equation 2.2, using the corresponding stable assemblage near the solidus modelled in Theriak.

Mineral-melt pairing	Partition coefficient	Reference
plagioclase/melt	0.02	Jolliff et al., 1992
alkali feldspar/melt	0.01	Jolliff et al., 1992
cordierite/melt	0.44	Evensen and London, 2003
biotite/melt	1.67	Icenhower and London, 1995
muscovite/melt	0.82	Icenhower and London, 1995

Table 2.1: Mineral–melt partition coefficients for Li used in modelling

Trace element partitioning between melt and residuum for all other elements (Rb, Ba, Th, U, Nb, Ta, La, Ce, Pb, Pr, Sr, Nd, Zr, Hf, Sm, Eu, Gd, Tb, Dy, Y, Ho, Er, Tm, Yb, Lu, V, Sc) was modelled using the trace element modelling tool in MAGEMin. The trace element calculations in this tool are performed using the same calculations described above, with the addition of predicting the stability of zircon using the zircon saturation model of Crisp and Berry (2022). Partition coefficients for other elements were taken from Laurent (2012).

Chapter 3

The formation of lithium-rich pegmatites through multi-stage melting

Foreword

This chapter has been published in the journal *Geology* under the following title: “*The formation of lithium-rich pegmatites through multi-stage melting*”. **Koopmans, L.**, Martins, T., Linnen, R., Gardiner, N.J., Breasley, C.M., Palin, R.M., Groat, L.A., Silva, D. Robb, L.J. **doi:10.1130/G51633.1**. Author contributions are described in the statement of authorship.

Abstract

Lithium-cesium-tantalum type pegmatites – the primary source of lithium – crystallize from highly evolved, volatile- felsic melts that incorporated crustal material in their source. Pegmatites are classically thought to form either from extreme fractionation of a parental granite body or via low-degree partial melting of a metamorphic rock (anatectic-origin). However, the processes that lead to the formation of economic lithium pegmatite deposits remain enigmatic, since precipitation of lithium ore minerals requires melt lithium concentrations in excess of 5000 ppm – approximately 500 times upper crustal abundances. Here, we use petrological modeling to quantify lithium enrichment in an anatectic-origin scenario and show that it is primarily driven by the relative stability of residual biotite and muscovite at medium to high pressures (~ 8 kbar), and biotite and cordierite at low pressures (~ 3 kbar). We show anatexis of an average lithium-enriched metasedimentary source cannot sufficiently elevate the lithium content of the ensuing melt to form economic deposits; however, if this first-generation melt – now crystallized as granitic crust – is re-melted, the second-generation melt will be sufficiently concentrated in lithium to crystallize lithium ore minerals. We propose a petrogenetic model for anatectic-origin lithium pegmatites, in which a region experiences at least two stages of partial melting, ultimately generating lithium-rich melts without invoking extensive fractional crystallization. This mechanism can account for both the occurrence of unzoned lithium pegmatites and explains why economic pegmatites in many terranes are younger than their inferred source granites.

3.1 Introduction

Lithium–cesium–tantalum (LCT) pegmatites are formed from highly fractionated melt (Černý, 1991b) and dominate global lithium (Li) production (U.S. Geological Survey, 2023). LCT pegmatites first appeared in the Neoproterozoic (c. 2.8 Ga), are dominantly hosted in greenschist- to amphibolite-grade supracrustal rocks (Dittrich et al., 2019), and retain geochemical signatures which imply their genesis involved reworking of existing (typically metasedimentary) crust (Černý, 1991b). However, fundamental questions remain regarding how these evolved rocks become highly concentrated in Li, since the primary extractable Li ore minerals – the Li-aluminosilicates spodumene and petalite – only crystallize when melt Li contents exceed 5000 ppm (London, 1984; Maneta et al., 2015), representing an enrichment of ca. 500 times

above ordinary crustal abundances (Taylor and McLennan, 2003). Furthermore, to be economically viable, an economic Li-pegmatite deposit must contain a sufficiently large mineralized zone (Bradley and Mccauley, 2017), the extent of which is controlled by the initial Li concentration in the emplaced pegmatitic melt (London, 2014).

In general, concentrating incompatible elements such as Li within a small-volume melt fraction occurs through either (i) high-degree crystal fractionation of a larger body of magma, or (ii) siphoning low-degree (e.g., 7–10%) partial melts away from a source rock undergoing anatexis. These contrasting processes underpin two proposed LCT pegmatite petrogenetic models: (1) extreme fractionation of a granitic magma that carries a typical (crustal) Li concentration, generating a minor volatile- and metal-rich melt fraction which migrates from the source region and ultimately crystallizes as a pegmatite (Cameron et al., 1949; Jahns, 1953; Černý, 1991b; Černý and Ercit, 2005), or (2) low-degree partial melting of Li-rich host rocks during prograde metamorphism (‘anatectic origin’), generating a volatile-rich melt that does not require significant additional fractionation (Stewart, 1978; Simmons et al., 1995; Shaw et al., 2016; Müller et al., 2017). Recent research has focused on the anatectic model as a viable mechanism to form economic pegmatite deposits (Müller et al., 2017; Kunz et al., 2022; Knoll et al., 2023). Here, we use petrological modelling to quantify the extent of Li enrichment during crustal melting of metasedimentary protoliths, and then assess how subsequent melting of their crystallized magmatic products might serve to further concentrate Li in newly formed melt fractions. Our results imply that multi-step anatexis in long-lived orogenic systems can act as a highly efficient mechanism to produce economic grade LCT pegmatites.

3.2 Modeling Li enrichment during crustal anatexis

3.2.1 Lithium in sedimentary rocks

The concentration of Li in a given melt fraction is principally controlled by the original Li concentration in the protolith undergoing partial melting. Reported geochemical estimations of an ‘average’ metapelite proposed a Li concentration of ~ 55 ppm (Shaw, 1956); however, the upper limits of Li concentrations in different sedimentary lithologies are unconstrained. To address this knowledge gap, we filtered the Sedimentary Geochemistry and Paleoenvironments Project (SGP) database (Farrell et al., 2021) to

produce a compilation ($n = 11,634$) of the measured Li contents of global sediments, attached in Supplementary file 2.

The filtered dataset shows an average Li concentration of 50.06 ppm in siliciclastic sediments, 18.83 ppm in carbonate sediments, and 18.04 ppm in chemical sediments (e.g. cherts). Nonetheless, all sediment types recorded outliers, with Li concentrations exceeding 1500 ppm in some cases (Figure 3.1). These data likely represent geological anomalies; for example, Li may become enriched in metasedimentary aureoles surrounding LCT pegmatites, Li is highly mobile under hydrothermal conditions and therefore diffuses rapidly through host rocks (Morgan and London, 1987; Linnen et al., 2012; Roza Llera et al., 2019), or else Li may be enriched in evaporitic sediments, such that these anomalously high values represent primary concentrations (Roda Robles et al., 1999).

There is reason to suggest that Li can be enriched in sedimentary rocks. Surface weathering processes preferentially enrich Li in the washed sediment, inferring surface interaction in enriched metasedimentary lithologies (Teng et al., 2010). Given the potential for enrichment above a baseline value, we take the statistical upper bound of outliers as our theoretical maximum for Li concentration (125.5 ppm for siliciclastic rocks) in metasediments that have not been affected by extraneous metasomatic or hydrothermal processes.

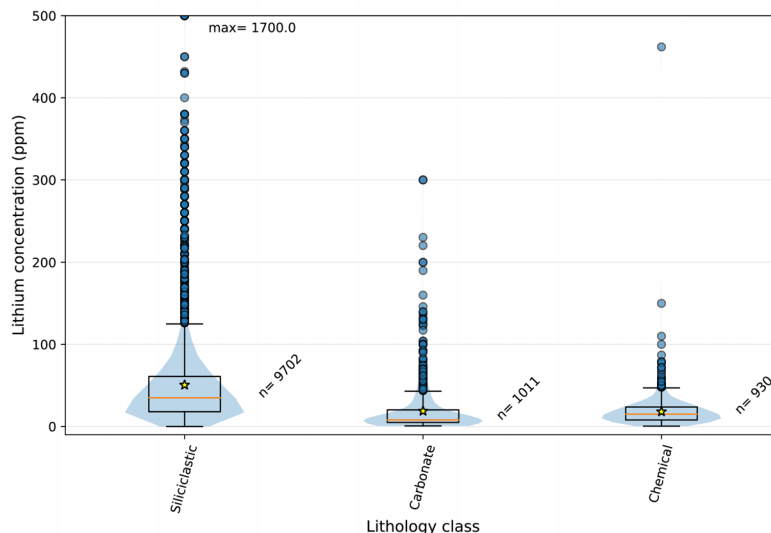


Figure 3.1: Violin plots of Li concentrations in the filtered SGP dataset binned by lithological class. Number of data points in each class is given as n.

3.2.2 Lithium in the melting environment

Lithium in metasedimentary rocks is primarily hosted in biotite, muscovite, cordierite, and staurolite (Simons et al., 2017; Kunz et al., 2022; Knoll et al., 2023). During anatexis, the Li content of a melt is controlled both by the breakdown of these minerals, and by the partitioning of Li between restitic minerals and melt as they re-equilibrate. We examined Li mineral–melt partitioning during crustal anatexis via a batch melting geochemical model, which assumes that generated melt remains in equilibrium with the source rock until extraction (Rosenberg and Handy, 2005); an assumption validated in experimental studies (e.g., Acosta-Vigil et al., 2012). The onset of melting, as well as the stability and thus presence of restitic minerals in equilibrium with the melt, depends on the source composition and thermobarometric conditions of melting. As such, we performed Gibbs free energy minimization calculations using the petrological modeling program Theriak-Domino (De Capitani and Petrakakis, 2010) to determine the stable mineral–silicate melt assemblages at different pressure (P, i.e. depth) and temperature (T) conditions within the continental crust for two putative metasedimentary source rocks: the average metapelite of Ague (1991) and a greywacke (Composition I of Pettijohn (1963)). Although Li contents vary between metasedimentary rocks, for ease of comparison we assume both compositions have a starting Li composition of 125.5 ppm, the maximum (excluding outliers) of 11,634 siliciclastic sedimentary rocks worldwide (Figure 3.1).

We calculated P–T phase diagrams for both metasedimentary protoliths and examined the crystallized products of extracted melts. Initial anatexis was examined along isobaric heating paths at 3 kbar (~ 9 km depth, i.e., intrusion-related heating) and 8 kbar (~ 24 km, i.e., a standard orogenic geotherm intersecting the solidus) under both dehydration melting (minimal saturation) and flux melting scenarios to quantify how Li behaves during partial melting. Our approach allowed the calculation of changes in melt Li content during crustal anatexis and then the behavior of that melt fraction after extraction (set at 7 vol. % melt, the threshold at which melt can escape its source, after Rosenberg and Handy, 2005).

Melt Li concentration was then modeled for two end-member scenarios: (1) the melt extracted undergoes fractional crystallization in a closed system; and (2) the resulting extracted melt crystallizes to form a granitic body which then undergoes a second (later) stage of anatexis. Bulk Li partition coefficients for fractional crystallization were calculated using stable assemblages present at the solidus and published mineral/melt partition coefficients (Table 2.1). All models were run through TDMelts (Koopmans et al., 2023).

For non-flux melting models, the water content in the bulk composition for each protolith was fixed in order to provide minimal free (aqueous) fluid at the point of first melting, defined here as having less than 1 mol. % H_2O at the solidus (cf. Palin et al., 2016b). We note that the absolute amount of H_2O in the bulk composition to allow minimal saturation was determined independently for runs at 3 kbar and 8 kbar, given the different parageneses in the subsolidus domain. The modelled bulk compositions, given in terms of atomic proportions, were as follows:

Protolith	Pressure (kbar)	Bulk composition
Greywacke	3	SI(73.45)AL(18.28)CA(2.34)MG(3.09)FE(2.71)K(3.64)NA(8.02)TI(0.2)H(4.00)O(190.77)
Granite	3	SI(71.34)AL(14.85)CA(0.23)MG(0.47)FE(0.92)K(7.15)NA(6.19)TI(0)H(1)O(173.815)
Greywacke	8	SI(73.45)AL(18.28)CA(2.34)MG(3.09)FE(2.71)K(3.64)NA(8.02)TI(0.2)H(5.14)O(191.34)
Granite	8	SI(63.69)AL(15.51)CA(0.42)MG(0.49)FE(0.86)K(5.98)NA(7.5)TI(0)H(2)O(160.23)
Pelite	3	SI(59.92)AL(19.85)CA(1.54)MG(3.98)FE(5.51)K(4.61)NA(2.98)TI(0.57)MN(0.08)H(5.06)O(168.81)
Granite	3	SI(72.13)AL(14.44)CA(0.22)MG(0.45)FE(0.9)K(6.96)NA(6.02)TI(0)H(1.00)O(174.55)
Pelite	8	SI(59.92)AL(19.85)CA(1.54)MG(3.98)FE(5.51)K(4.61)NA(2.98)TI(0.57)MN(0.08)H(9.20)O(170.88)
Granite	8	SI(59.16)AL(14.59)CA(0.49)MG(0.07)FE(0.15)K(6.3)NA(5.83)TI(0)H(2)O(147.985)

Table 3.1: Theriak-Domino inputs for all model runs (compositions in atomic proportions).

Modebox plots were produced by running an isobaric pressure–temperature (P–T) path in Theriak. The outputs keep track of both the melt composition and modal abundances of the residuum phases to use as inputs for the trace element modelling. Flux melting was modelled by introducing excess H_2O into the system until 7 mol. % partial melt was generated at the P–T point of interest on the solidus. Then, the composition of melt predicted at that point was ‘extracted’ from the model, used as a starting bulk composition in a new calculation with minimal fluid saturation, and allowed to cool isobarically until fully crystallized.

The starting lithium concentration for each model is shown in Table 3.2.

Protolith	Pressure (kbar)	Lithium concentration (ppm)
Metasedimentary rocks	3, 8	125.5
Granite – Greywacke source	3	538
Granite – Greywacke source	8	528
Granite – Pelite source	3	411
Granite – Pelite source	8	255

Table 3.2: Lithium concentration in different protoliths under varying pressures.

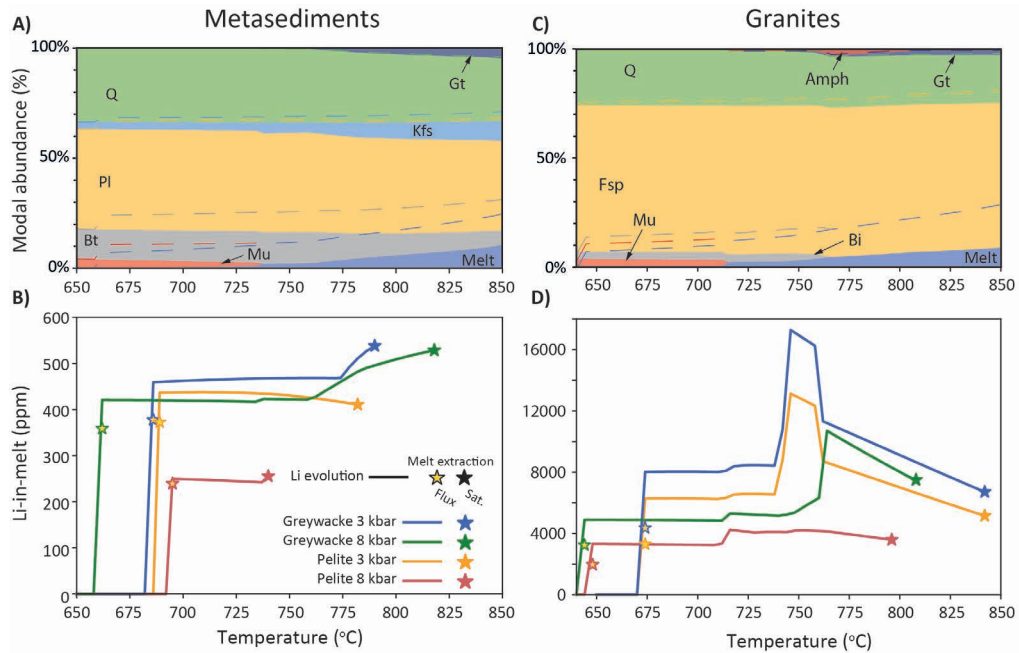


Figure 3.2: Representative results of petrologic modeling of a metasedimentary rock (a, b, here the greywacke at 8 kbar) and granite (c, d, sourced from greywacke at 8 kbar) composition undergoing isobaric heating. Modebox diagrams (a, c) show the equilibrium assemblages across the modelled temperature range. The dashed lines highlight the stable assemblage under flux melting conditions. The graph in (b) traces the concentration of lithium in the melt component of the modelled metasedimentary scenarios. (d) as (b), though of the melt components of the corresponding metasedimentary rocks of (b). Stars denote melt extraction points for each condition, where Flux corresponds to flux melting and Sat. corresponds to dehydration melting. Mineral abbreviations: Amph: amphibole; Bt: biotite; Cd: cordierite; Fsp: feldspar; Gt: garnet; Kfs: Alkali feldspar; Mu: muscovite; Pl: plagioclase; Opx: orthopyroxene; Q: quartz.

3.3 Results

During melting, both metasedimentary compositions (Figure 3.2) stabilize cordierite and biotite at low pressures, whereas muscovite and biotite are stable at higher pressures. Figure 3.2 (b) shows the predicted partitioning of Li between metasediment-derived melt and residuum for both starting compositions. At fluid-present conditions, the solidus occurs between 660 and 690 °C and the release of Li into the melt is largely controlled by the initial breakdown of muscovite and later breakdown of biotite.

Metapelites are typically enriched in Al compared to greywackes and stabilize higher proportions of cordierite and mica at equivalent P–T conditions. At equivalent

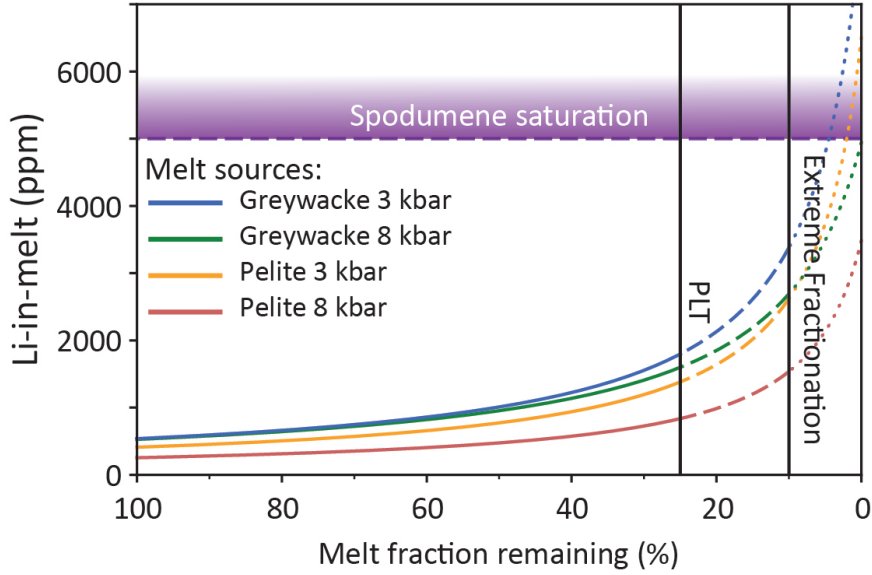


Figure 3.3: The behavior of lithium (Li) in the extracted melt during Rayleigh fractionation. The initial Li content of each melt corresponds to the concentration of Li at extraction in each of the modelled scenarios. The partition coefficients are taken from the model near the solidus of each individual melt. The particle-locking threshold (PLT) of Vigneresse et al. (1996) and the extreme fractionation boundary of Zhao et al. (2022) are highlighted, as well as the minimum threshold for spodumene saturation (Maneta and Baker, 2014).

pressures, greywacke-derived melts contain higher Li contents (by a factor of 1.3–2.1) than metapelite-derived melts (Figure 3.2b); this difference is amplified at higher pressure due to the extended stability of mica compared to cordierite.

Micas are the main hosts of Li in the subsolidus, and our models confirm that Li enrichment during partial melting of metasedimentary rocks is primarily controlled by biotite and muscovite stability, with cordierite having a minor influence, and feldspar playing an insignificant role (Figure 3.2), agreeing with previous work (Kunz et al., 2022; Knoll et al., 2023). Li is released into the melt upon mica breakdown, but importantly, the persistence of micas during melting sequesters Li and serves to inhibit melt Li enrichment.

At a melt proportion of ~ 7 vol.%, we calculate maximum Li concentrations of 411 ppm at 788 °C and 538 ppm at 785 °C for metapelites and metagreywackes, respectively. These values represent a potential maximum Li enrichment factor of ~ 4.3 at extraction compared to the source. Additionally, flux melting (i.e., ingress of external H_2O) also serves to depress melt Li enrichment under all conditions, largely due to the higher proportion of micas present at melt extraction conditions that have

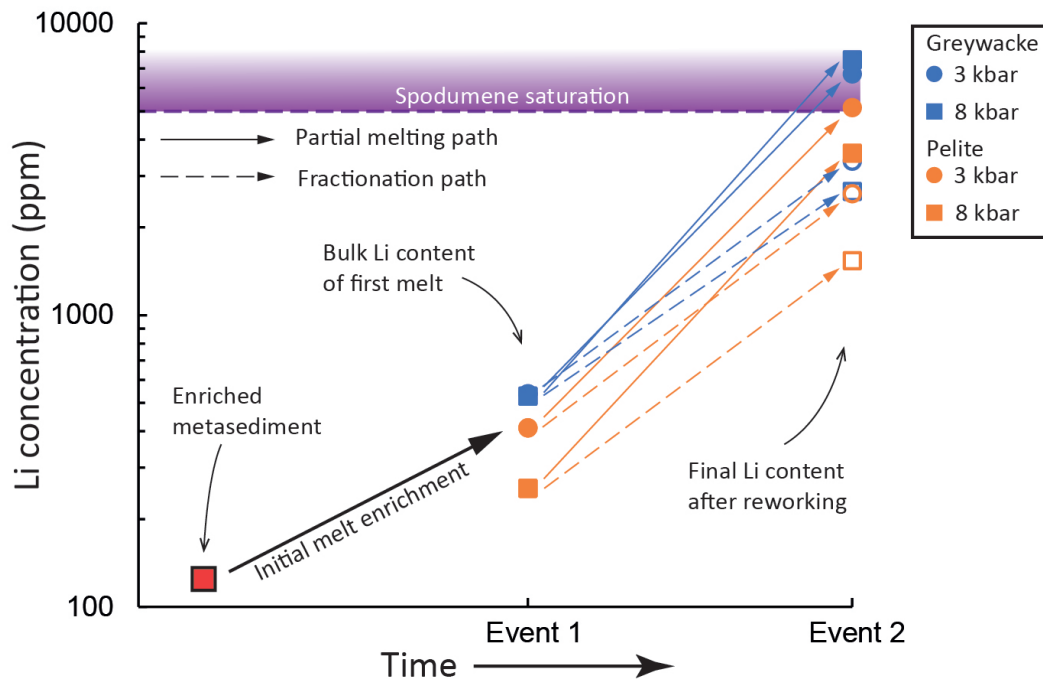


Figure 3.4: Evolution of Li concentrations during melting and fractionation of melt components generated during crustal anatexis – with the spodumene saturation boundary highlighted (Maneta et al., 2015). After initial melt extraction, the granite either: 1) undergoes fractional crystallization (hollow symbols, the compositions of fractional crystallization was taken as the composition at % crystallization); or (2) is assumed to crystallize and then remelted and extracted at the 7 vol. % melt (filled shapes).

not broken down during melting reactions, although we note melts are extractable at a significantly lower temperature (Figure 3.2b).

Fractional crystallization (our scenario 1) during cooling may further enrich the melt phase in incompatible elements including Li. Modeled Rayleigh fractionation trends for melts extracted at the 7 vol. % threshold are shown in Figure 3.3. In all cases, the concentration of Li reaches 800–1800 ppm at 75% fractionation and 1500–3400 ppm at extreme fractionation conditions (Zhao et al., 2022), still below the minimum saturation point for spodumene. Spodumene saturation is only achieved at low pressure (3 kbar) during the final stages of crystallization (Figure 3.3), when the melt fraction is reduced below 5 vol. %, at which point the melt is not physically extractable (Vigneresse et al., 1996).

Alternatively (scenario 2), the melt produced during anatexis of a metasedimentary protolith may crystallize as a granitic body in the crust and experience a separate, later stage of melting. We modeled the remelting of the granitic compositions pro-

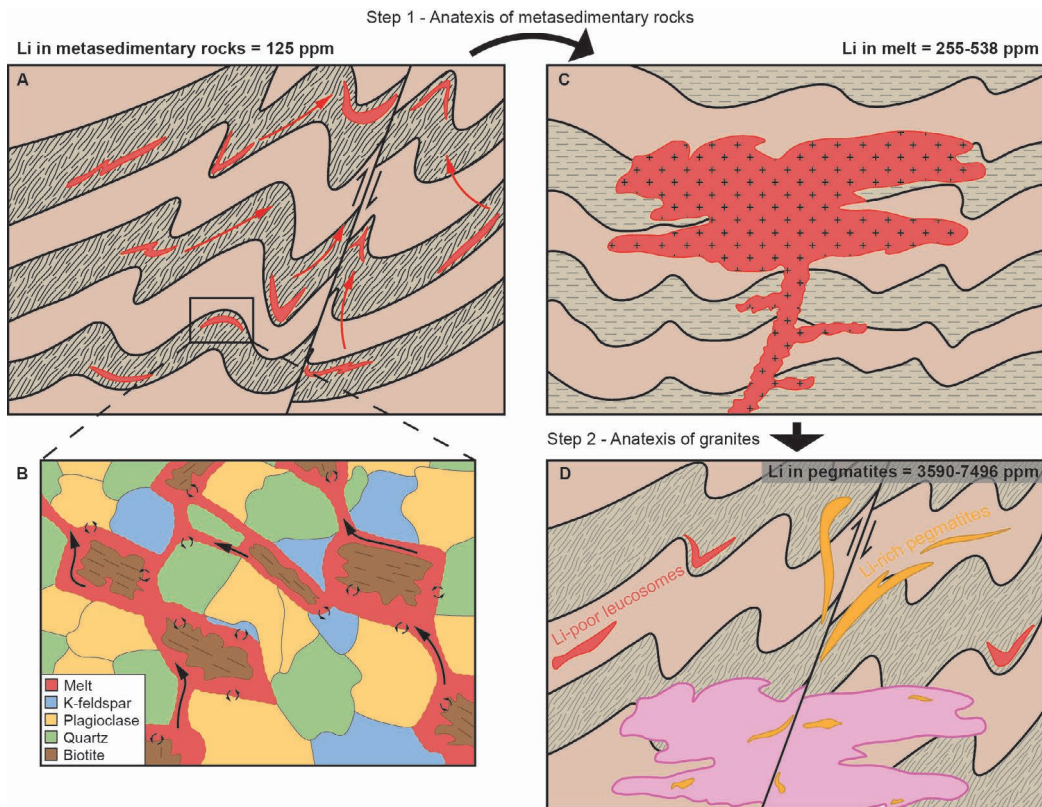


Figure 3.5: Schematic petrogenetic model for generating lithium (Li)-rich pegmatites in the via crustal anatexis. (A) Partial melting of a metasedimentary rock generates melt lenses, which periodically lose melt. (B) grain-scale diagram visualizing melt films connecting and driving melt out of the system in (A). The breakdown of hydrous phases and mineral-melt diffusion controls the Li budget of the resultant melt. (C) The melt accumulates and crystallizes as a granite structurally above the anatexis zone. (D) A subsequent melting event, either associated with the primary metamorphic event, or during a later orogenic cycle melts the granite, generating Li-rich pegmatites (orange) in their vicinity.

duced through primary anatexis (Figure 3.2c, d) to calculate potential second stage melt Li concentrations. In this scenario, melt first appears between 640 °C and 675 °C, and coexists with a smaller proportion of micas (<8 wt. %) relative to metasedimentary rocks, and no stable cordierite in any of the modeled scenarios (Figure 3.2c). Accordingly, the secondary melting stage generates significantly higher melt Li concentrations of 3590–7496 ppm at a 7 vol. % extraction point (Figure 3.2d), equal to a potential Li enrichment factor of ~ 14.2 (Figure 3.2d, 3.4). These results importantly indicate that a second stage of melting substantially enriches Li; in two stages driving Li concentration from 125.5 to 7496 ppm, with the resultant secondary melts having economic potential without invoking fractional crystallization.

3.4 Multi-step distillation model for Li pegmatites

Our results show that partial melting of typical metasedimentary sources followed by closed-system fractional crystallization of those melt fractions cannot concentrate Li to the saturation limit required (>5000 ppm) to generate a large zone of Li-aluminosilicates within a pegmatite. Thus, production of a melt capable of saturating Li ore minerals in economic quantities within a single stage melt scenario would require a rare, extremely enriched protolith (e.g., 1160 ppm in the metasedimentary rock). While such concentrations have been documented in reworked Li-rich volcanic sediments (e.g., Kadir et al., 2023), they are typically not documented near Li-rich pegmatite provinces (e.g., Roda Robles et al., 1999) and therefore their role in Li pegmatite formation is equivocal. By contrast, additional processing (i.e., remelting) of a package of continental crust can increase the Li concentration in a second-generation in order to allow spodumene/petalite saturation very early in the fractionation history, obviating the need for extensive fractional crystallization.

We therefore propose the following petrogenetic model for anatectic-origin Li pegmatites with economic potential (Figure 3.5):

- (1) Partial melting of a metasedimentary succession during prograde metamorphism, which produces a granitic melt that is modestly enriched in Li;
- (2) This melt crystallizes as a granitic intrusion structurally above the migmatite zone;
- (3) A subsequent melting event, either an extension of the initial metamorphic event or later during an unrelated orogenic cycle, reheats and remelts the granite forming a highly enriched melt that ultimately crystallizes as a Li-rich pegmatite.

Our two-stage anatectic model indicates that the size of the pegmatite generated is approximately 200 times smaller than the metasedimentary protolith; for example, a deposit the size of Tanco (0.021 km^3 , Stilling et al., 2006) likely formed from an initial metasedimentary package of c. 4.5 km^3 .

This mechanism also accounts for two particular features of pegmatites. Firstly, pegmatites can be unzoned with economic minerals homogeneously distributed throughout, requiring Li-bearing minerals to crystallize from the margin inwards (e.g., Kings Mt pegmatites, Swanson (2012); Mt Cattlin pegmatites, Sweetapple et al., 2019). This suggests the emplaced melt has sufficiently high Li concentrations prior to crystallization.

Secondly, where pegmatite fields have a spatial relationship with granites, they are often significantly younger (e.g., Stilling et al., 2006), which is at odds with arguments for the rapid crystallization of pegmatites from the granitic melt when standard magmatic fractionation is invoked (Simmons and Webber, 2008). Alternatively, many orogenic events are long-lived and experience multiple cycles of melting (e.g., Mulcahy et al., 2014). Together, this suggests the possibility of granites remelting as a mechanism for pegmatite genesis (e.g., Issia Zone, Brou et al., 2022).

A multi-step anatectic model for the formation of Li pegmatites therefore satisfies these geochemical and geochronological constraints, providing an efficient mechanism to elevate Li concentrations and generate economic Li deposits.

Chapter 4

Structural controls on lithium mineralization in shear-zone hosted granitic pegmatites of the Zulu pegmatite field, Zimbabwe – implications for exploration

Foreword

This chapter has been published in the journal *Mineralium Deposita* under the following title: “*Structural controls on lithium mineralization in shear-zone hosted granitic pegmatites of the Zulu pegmatite field, Zimbabwe – implications for exploration*”. **Koopmans, L.**, Gardiner, N.J., St. Pierre, B., Palin, R.M., Musinga, R., Robb, L.J. doi:10.1007/s00126-025-01371-x. Author contributions are described in the statement of authorship.

Abstract

Granitic pegmatites are a significant source of critical metals including tin, tantalum, and most notably lithium. To meet future demand, a comprehensive exploration model is required to assist in the discovery of new hard rock deposits. Whereas recent work has largely focused on understanding the source and mineralization processes of pegmatites, the structural controls on the distribution and size of individual deposits remains poorly understood and understudied. In this contribution, we present a structural study on the Zulu pegmatite field in Zimbabwe, which provides a good example of the influence of shear zones, host rock rheology, and lithological competency contrasts on the orientation, size, and distribution of pegmatite bodies within a pegmatite field. At Zulu, we observe both structural and petrographic evidence for two types of pegmatite emplacement within an active shear zone during D_2 strike-slip dominated deformation. An early generation (Type 1) was emplaced syn-kinematic to D_2 within dilational jogs subparallel to the shear fabric, and continued ductile shearing also drove significant recrystallization which affected the primary magmatic phases and therefore influenced the preserved mineralogy. A later generation (Type 2) was emplaced syn-to-late-kinematic to D_2 along tension gashes and subordinate fracture sets oblique to the shear fabric, which served to truncate the cooling history and preserve a primarily magmatic mineralogy within this pegmatite group. By comparing Zulu to other large pegmatite deposits, we conclude that geologic structures are critical to source-to-sink connectivity in lithium pegmatite systems, and affect the mineralization potential of individual deposits by driving recrystallization.

Assessing the structural history and relative timing of emplacement within a pegmatite field, in conjunction with detailed (micro)textural observations from within pegmatite bodies, is essential to understanding pegmatite emplacement geometries. A more systematic approach in constraining these relationships will therefore aid in generating new exploration targets in both greenfield and brownfield settings.

4.1 Introduction

Granitic pegmatites are the primary source of several key metals critical to green technologies including tin, tantalum, and niobium, but, most importantly, lithium (Bowell et al., 2020; McCaffrey and Jowitt, 2023). To meet projected lithium demand, pegmatite deposits are forecast to account for more than 80% of future supply (Yao, 2022). Granitic pegmatites result from small-volume granitic melts through either

extreme fractionation (Jahns and Wayne Burnham, 1969; Černý, 1991b; London, 2005) or low-degree partial melting of a fertile source (anatectic pegmatites) in the mid-crust (Simmons et al., 1995; Müller et al., 2017, , Chapter 3). These melts migrate upwards to be emplaced at shallower depths in the upper crust (Plunder et al., 2022), typically as bodies up to 1 km in length and in swarms of 10–100 individual intrusions (Černý, 1991b).

A comprehensive exploration model for economic-grade lithium pegmatites must encompass a mineral system approach (Wyborn et al., 1994; Hagemann et al., 2016), which at its simplest defines a source, a migration pathway and a trap, as well as the preservation mechanisms that enable a deposit to persist. Much recent work has focused on the source of pegmatites (Černý, 1991b; Simmons et al., 1995; Shaw et al., 2016; Müller et al., 2017; London, 2018, , Chapter 3) and preservation of their magmatic mineralogy, including the primary lithium ore minerals spodumene and petalite (Maneta et al., 2015; Ballouard et al., 2020; Wilde et al., 2021; Shaw et al., 2022; Pfister et al., 2023). However, the control that pathways such as faults and folds play on the distribution and size of economic pegmatites is relatively poorly understood (Hall and Kisters, 2012; Silva et al., 2023; Gardiner et al., 2024b).

Globally, pegmatites have a clear spatial relationship with regional structures such as faults and shear zones on scales up to 10s of km (Kontak et al., 2005; Dill et al., 2012; Deveaud et al., 2013). Most importantly for exploration, the majority of large lithium pegmatite deposits are spatially associated with significant shear zones (Partington et al., 1995; Sweetapple, 2000; Selway, 2005; Morissette et al., 2022). Further, detailed studies of pegmatite districts have shown how such structures may affect the distribution and remobilization of mineralized pegmatites within an individual district (Keyser et al., 2023; Silva et al., 2023), highlighting how regional stresses, and the resultant structures, affect the orientation, shape, and (re)crystallization of pegmatites. A key question relates to whether structures such as brittle faults and ductile shear zones play a role in pegmatitic melt migration and emplacement by providing a fundamental tapping mechanism, or do pegmatitic melts simply migrate passively through pre-existing crustal weaknesses (cf. Vanderhaeghe, 1999; Brown et al., 2011; Hall and Kisters, 2012)?

Here, we address this question through a comprehensive structural study of the Archean Zulu Pegmatite Field (Main pegmatite: 24.75 Mt at 0.43% Li₂O, Premier African Minerals Limited, 2024), situated within the Fort Rixon–Shangani greenstone belt in the south of the Zimbabwe Craton. The Zulu Pegmatite Field offers an excellent demonstration of how a combination of the structural history, host rocks,

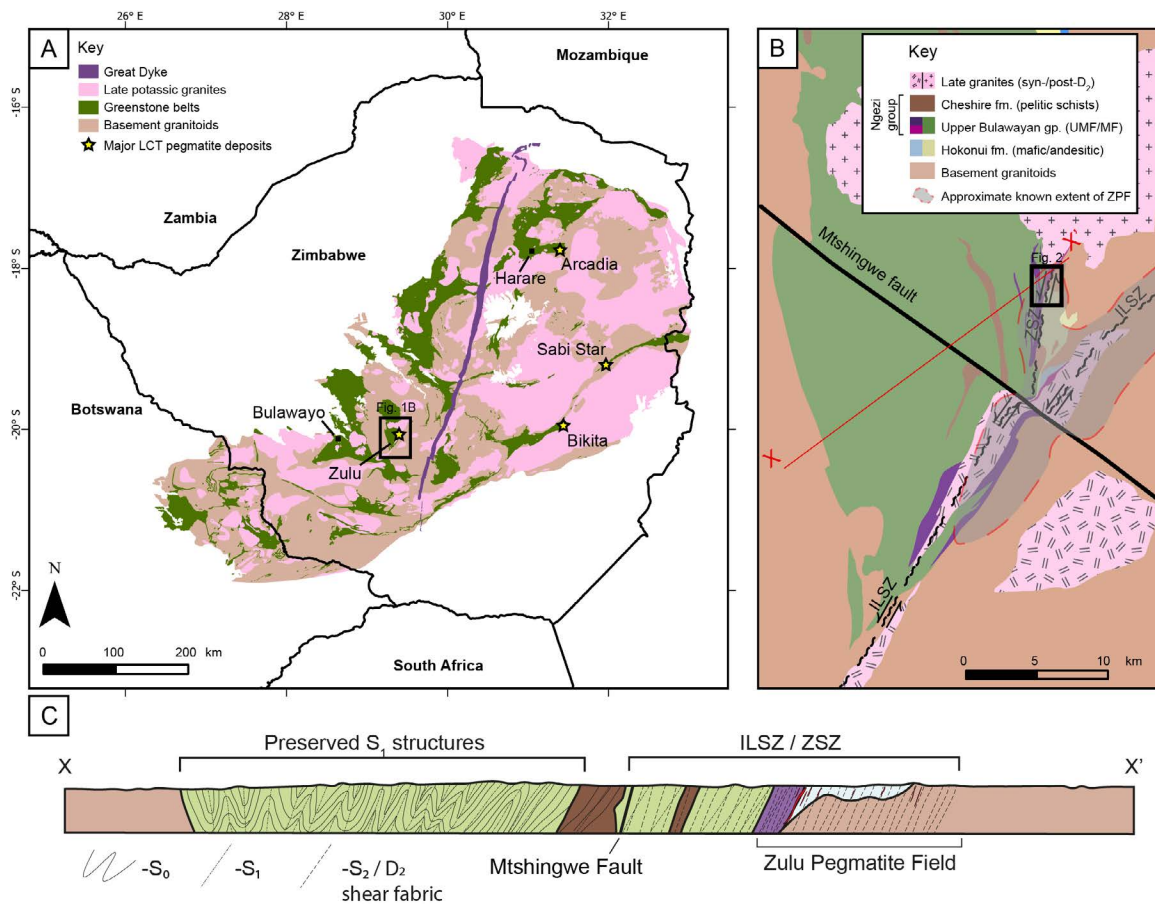


Figure 4.1: A) Simplified geological map of the Zimbabwe Craton, adapted from Ncube (1994). Major lithium deposits of Archean age are shown with yellow stars. Location of Figure 1B noted by black square. B) Regional geological map around the core of the currently known Zulu pegmatite field, adapted from Harrison (1969). Location of Figure 2 noted by black square. C) Schematic cross section X-X' through the Fort Rixon – Shangani greenstone belt, highlighting structural fabrics preserved in the belt. ILSZ: Irisvale – Lancaster Shear Zone, ZSZ: Zulu Shear Zone, ZPF: Zulu Pegmatite Field.

the spatial association with regional granitic bodies, major lithological variations, and crustal-scale fault systems affect the formation and distribution of mineralized lithium-bearing pegmatites. Using geological and structural mapping, field observations, and thin section microscopy, we demonstrate how the stress regime, major lithological competency contrasts, and structural history together influenced the localization and kinematics of shear zones and controlled the geometry, texture, mineralogy and degree of recrystallization of pegmatites during emplacement.

4.2 Geological setting

4.2.1 Geological history

The Zulu Pegmatite Field is situated in the centre of the Zimbabwe Craton, a large cratonic block predominantly exposed in Zimbabwe (Figure 4.1). The Zimbabwe Craton experienced two major magmatic-metamorphic events during the Archaean (Wilson et al., 1995), resulting in a series of folded granite-greenstone belts within a deformed granitoid gneiss terrane (Figure 4.1). Early felsic crust formation occurred at c. 3600–3200 Ma, with magmatic activity preserved in the central Tokwe and Rhodesdale segments (Horstwood et al., 1999; Hofmann et al., 2022), located between the Zulu and Bikita pegmatite fields (Figure 4.1). These segments formed the cratonic nuclei around which the rest of the craton accreted (Jelsma et al., 2021), with the majority of craton growth occurring during c. 3000–2500 Ma. Magmatism culminated during the Neoproterozoic with the emplacement of abundant potassic granites (the Chillimanzi and Razi suites; Wilson et al., 1995; Rollinson, 2022; Chagondah et al., 2023). The formation of several lithium-bearing pegmatite fields (Zulu, Bikita and Arcadia) was contemporaneous with this late felsic magmatism, and a tentative genetic link has been placed between the pegmatite fields and neighbouring granites by some authors based on spatial relationships and geochronological overlap (e.g., Chagondah et al., 2024). Emplacement of the Great Dyke signifies final craton stabilization at c. 2575 Ma (Armstrong and Wilson, 2000; Jelsma et al., 2021).

The c. 3000–2700 Ma Fort Rixon–Shangani greenstone belt (FRSGB) is situated 60 km north-east of Bulawayo (Figure 4.1). It is exposed as a major syncline with a N-S axis and is dominantly comprised of mafic/ultramafic extrusive volcanic units (Harrison, 1969). Large pegmatites of the Zulu Pegmatite Field were emplaced along the eastern margin of the Fort Rixon section of the greenstone belt, where a basal mafic/andesitic tuff succession is unconformably overlain by a mafic/ultramafic extrusive volcanic succession with minor sedimentary input (Figures 4.1 and 4.2). These units can be respectively correlated to the Hokonui and Ngezi Formations (with associated intrusive rocks) of the Belingwe Greenstone Belt (Bickle et al., 1993) to the south-east (Prendergast, 2004a). Following deposition and emplacement, the greenstone belt underwent regional metamorphism to upper greenschist-facies conditions, which overprinted and recrystallized most of the primary magmatic fabrics (Harrison, 1969). The age of this regional metamorphic episode is currently unknown.

A series of potassic granites intruded into and around the greenstone belt; three smaller synkinematic porphyritic intrusions granites (dated to 2653 ± 15 Ma and

2616 ± 21 Ma, Chapter 5), a sheared porphyritic granite along the core of the Irisvale–Lancaster shear zone (2679 ± 18 Ma, Chapter 5), and the postkinematic Nalatale granite (2513 ± 36 Ma, Chapter 5), which almost completely transects the belt north of the Mtshingwe fault, (Figure 4.2, Harrison, 1969; Campbell and Pitfield, 1994). These granites have been correlated with the Chilimanzi suite granites (Chapter 5). There is no clear genetic relationship between the Zulu Pegmatite Field and any of the spatially associated granites.

4.2.2 Structural history

The FRSGB experienced three phases of regional deformation (Table 4.1). The oldest phase (D_1) consisted of folding by WNW-ESE directed shortening, which resulted in a major isoclinal F_1 syncline, and a weak axial-planar S_1 foliation (Campbell and Pitfield 1994). The second phase of deformation (D_2) marked a change in shortening direction from WNW-ESE to NNW-SSE resulting in the formation of large, open, S-shaped warps and tight Z-shaped F_2 kink folds, as well as the formation of a regional-scale transpressive sinistral wrench fault system concentrated along the major NE-SW Irisvale–Lancaster shear zone (ILSZ) and along major lithological boundaries such as the NNE-SSW trending Zulu Shear Zone (ZSZ) (Stowe 1980). This D_2 event locally produced a penetrative shear fabric (S_2) which locally transposed S_1 fabrics into a parallel orientation. Both sinistral and later dextral reactivation shearing has been identified along the ILSZ (Campbell and Pitfield, 1994).

The third phase of deformation (D_3) consisted of NNW-SSE directed shortening, which formed the major NW-SE trending Mtshingwe Fault (Figure 4.2, Harrison, 1969). This fault system drags the earlier S_0/S_1 and D_2 shear zones from striking NNE-SSW to NE-SW. The drag folds and the en-echelon dolerite dyke emplacement filling the fault-system suggest a dextral sense of movement with significant lateral displacement (Campbell and Pitfield, 1994). The Mtshingwe Fault crosscuts the Great Dyke (c. 2575 Ma) further to the south-east.

4.3 The Zulu pegmatite field

4.3.1 The Zulu pegmatite field

The Zulu Pegmatite Field contains >100 individual intrusions that lie along a generally NNE–SSW trend (Figures 4.1, 4.2). Pegmatites of the Zulu Pegmatite Field are LCT-type (after Černý and Ercit, 2005), or Group 1 pegmatites (after Wise

Phase of deformation	Description	Associated structures	Mean attitude of fabrics*
D_1	• WNW–ESE directed shortening	F_1	190/70
	• Bedding parallel to F_1 isoclinal folding	S_1	190/70
	• Weak axial planar S_1 development		
D_2	• NNW–SSE directed shortening	F_2	68/167**
	• S-shaped and Z-shaped F_2 kink folds	S_2	208/72
	• Ductile and sinistral shearing along ILSZ and associated 2 nd order structures		
	• Shear fabric (S_2)		
D_3	• NW–SE directed shortening?	Fault planes	WNW–ESE
	• Brittle dextral faulting along WNW–ESE Mtshingwe fault system	Dolerite-filled tension gashes	NE–SSW

*Attitude is presented as strike/dip using right-hand-rule conventions.

**Plunge and trend of fold axis

Table 4.1: Summary of structural events in the Fort Rixon – Shangani greenstone belt

et al., 2022b), and are significantly enriched in lithium (Harrison, 1969; Goodenough et al., 2025). The largest bodies occur along the intrusive boundary between the Sonop serpentinite and the metamorphosed mafic/andesitic tuffs, although smaller intrusions occur up to 30 km towards the NE within the basement granitoids (Stowe, 1968). The primary mineralogy of bodies in the Zulu Pegmatite Field typically consists of petalite (5–25%), quartz (25–30%), K-feldspar (10–30%), albite (5–30%), and muscovite (5–10%). Abundant spodumene (up to 30%) has also been described in the larger pegmatites, and also occurs within some of the smaller pegmatites hosted within the basement granitoids (Premier African Minerals Limited, 2023).

Individual pegmatite bodies vary in width from 0.5 m to ~50 m, and some strike for up to 2 km. Contacts are sharp with the host rock. Two distinct pegmatite types can be distinguished based on mineralogy and texture (summarized in Table 4.2).

Event	Pegmatites affected	Description and primary mineralogy	Crystal size	Li-aluminosilicate textures
1–Magmatic	Type 1+2	<ul style="list-style-type: none"> Spodumene (Spd1, Type 1 only), Petalite (Type 2 only), K-feldspar, Quartz, Muscovite, Albite 	5–15 cm	Classic SQUI (Type 1), Petalite (Type 2)
2–Dynamic recrystallisation	Type 1	<ul style="list-style-type: none"> Albite, quartz, spodumene, muscovite (purple), \pm spessartine Strong fabric development, strong alignment of albite + muscovite, pre-/syn-kinematic spessartine growth 	1–5 mm	Equant spodumene, fine grained
3–Albitisation	Type 1+2	<ul style="list-style-type: none"> Albite, quartz, muscovite (green), \pm spessartine Weak/absent alignment of albite 	1–5 mm	N/A
4–Late-spodumene	Type 1+2	<ul style="list-style-type: none"> Spodumene, quartz \pm muscovite (green) Along fractures and cleavage planes within petalite (Type 2) 	< 1–5 mm	Equant spodumene (Type 1), SQS (Type 2)
5–Low-T hydrothermal	Type 1+2	<ul style="list-style-type: none"> Fine grained clay assemblages 	< 1 mm	N/A

SQUI: spodumene–quartz intergrowths

SQS: spodumene–quartz symplectites (after Breasley et al., 2025)

Table 4.2: Summary of (re)crystallisation events in the Zulu pegmatite field

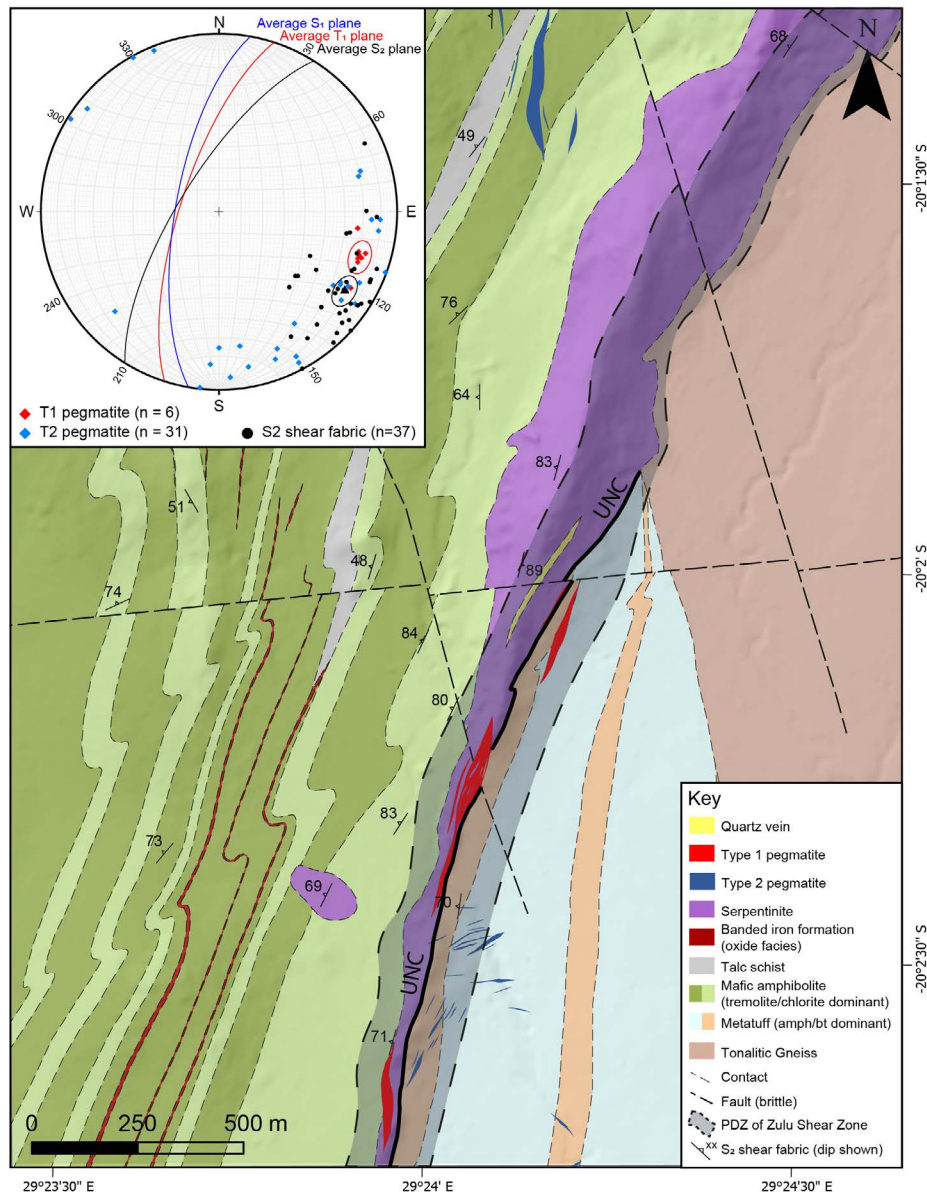


Figure 4.2: Inset: stereonet denoting main structural fabrics within the Zulu pegmatite field. A simplified description of each fabric can be found in Table 1. Main map: geological map of the region surrounding the major pegmatites of the Zulu pegmatite field. PDZ: principal deformation zone, UNC: major lithostratigraphic unconformity between the lower Hokonui and Ngezi formations.

4.3.1.1 Type 1 pegmatites

Type 1 pegmatites attain widths up to 50 m and are therefore the volumetrically dominant pegmatite type in the field. Individual intrusions have asymmetric lenticular shapes with long axes subparallel to the principal deformation zone of their host shear

zone (Figure 4.2). The primary magmatic mineralogy is comprised of spodumene, K-feldspar, albite, quartz, and muscovite. Individual crystals are up to 15 cm long (Figure 4.3C, 4.3D). An early generation of spodumene (Event 1, Table 4.2) in Type 1 pegmatites predominantly consists of crystals (0.1–1 cm) with a strong crystallographic orientation within oikocrystic quartz (Classic SQUI of Breasley et al., 2025, Figure 4.3C, 4.4A).

Primary textures are otherwise poorly preserved. Where magmatic mineralogy is preserved, quartz crystals commonly exhibit chessboard and undulose extinction patterns, whereas plastic deformation of lamellar twins in coarser albite is common (Figure 4.5A, 4.5C). Microcline has well-developed tartan twinning and flame perthite is prevalent (Figure 4.5D). Primary muscovite, where preserved, also exhibits kinked cleavage planes (Figure 4.5B).

An early deformation-driven recrystallization event (Event 2, Table 4.2) is pervasive and largely replaced the primary mineralogy with a fine-grained (1–5 mm) albite + quartz + spodumene + muscovite \pm spessartine assemblage (e.g., Figure 4.3C). Relict magmatic crystals are entrained as porphyroclasts (retaining deformation microstructures described above), and the muscovite has a deep purple colour in outcrop (Figures 4.3C, 4.3D, 4.6A). Individual crystals of spodumene within the replacement mineralogy are equant, although albite and muscovite within these aggregates are commonly strongly aligned to define a fabric within the pegmatites (described in more detail below). Locally, this unit developed a strong schistosity (possibly owing to a horizon of less competent magmatic mineralogy within the pegmatite), and euhedral spessartine is interpreted to be pre- to syn-kinematic with respect to the fabric (Figures 4.4C, 4.6B).

Later albitization (Event 3, Table 4.2) is most common along the hanging wall of the pegmatites, producing zones between 5 cm and 2 m wide of albite + quartz + muscovite and rare spessartine with a weak fabric (Figure 4.4A). Sporadic pockets of albitization also occur internal to the pegmatites. The muscovite in the albitized zones is typically light-green/grey in color. A later, spodumene-rich, replacement unit (Event 4, Table 4.2) is comprised of fine grained (<1 mm) spodumene + quartz + albite + muscovite with no defined orientation. This unit is only locally present and can be seen to overprint all previous crystallization events in the pegmatite (Figure 4.4A). Subsequent low temperature hydrothermal alteration products (Event 5, Table 4.2) are uncommon, though locally spodumene and feldspars are replaced by fine grained clay minerals.

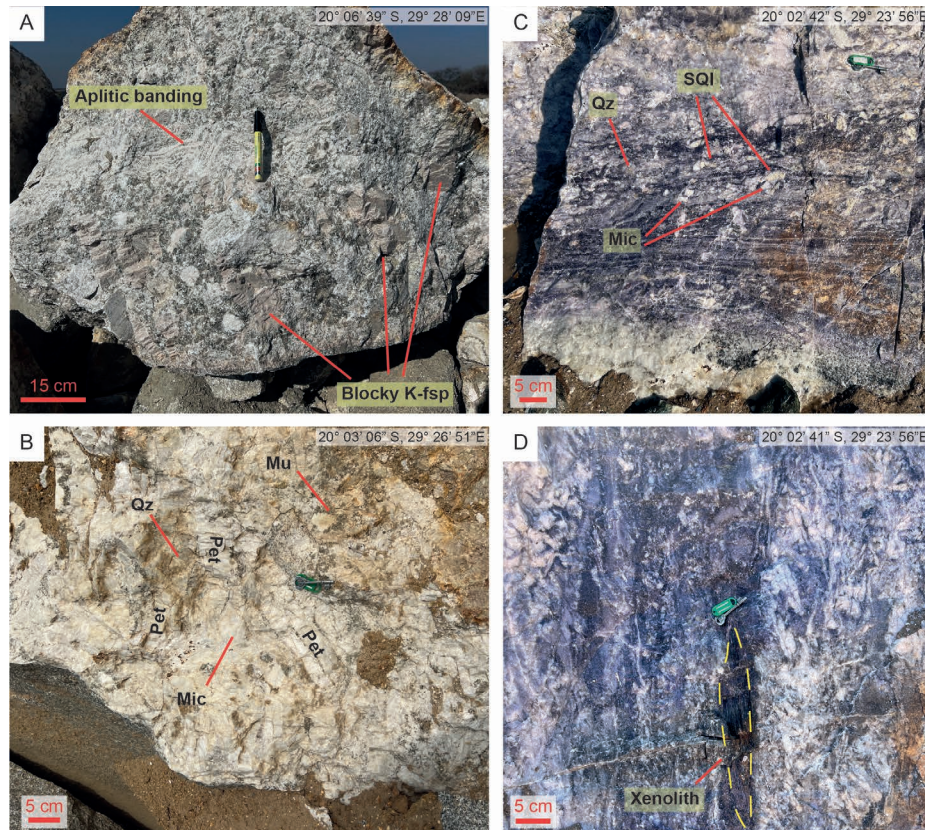


Figure 4.3: Representative photographs of pegmatites in the Zulu Pegmatite Field, with locations of images in top-right corner. A) Typical exposure of barren Type 2 pegmatite near the contact. Note the aplitic banding and unidirectional solidification texture expressed by the K-feldspar megacrysts. B) Mineralized Type 2 pegmatite with large petalite crystals randomly orientated in the intermediate zone. C) Typical exposure of the Type 1 pegmatites, with a strong fabric aligned left-right in the image. D) Rounded xenolith within Type 1 pegmatites within a localized high-strain domain. K-fsp: K-feldspar, Qz: quartz, SQI: spodumene-quartz intergrowths, Mic: microcline, Pet: petalite, Mu: muscovite.

4.3.1.2 Type 2 pegmatites

These are characterized by their relatively simple magmatic mineralogy and textures. Individual intrusions are often limited in width, being no greater than 5 m. A simple pegmatite zonation is common, with a fine-grained border zone (10–30 cm) followed by a zone of irregular aplitic banding (5–30 cm) defined by alternating muscovite-rich and muscovite-poor bands subparallel to the contact (Figure 4.3A). A coarse-grained intermediate zone and a poorly developed quartz core occur towards the centre of the pegmatite bodies (Event 1, Table 4.2). The intermediate zone consists of crystals up to 15 cm in size and can be further subdivided into a mineralized (Figure 4.3B)

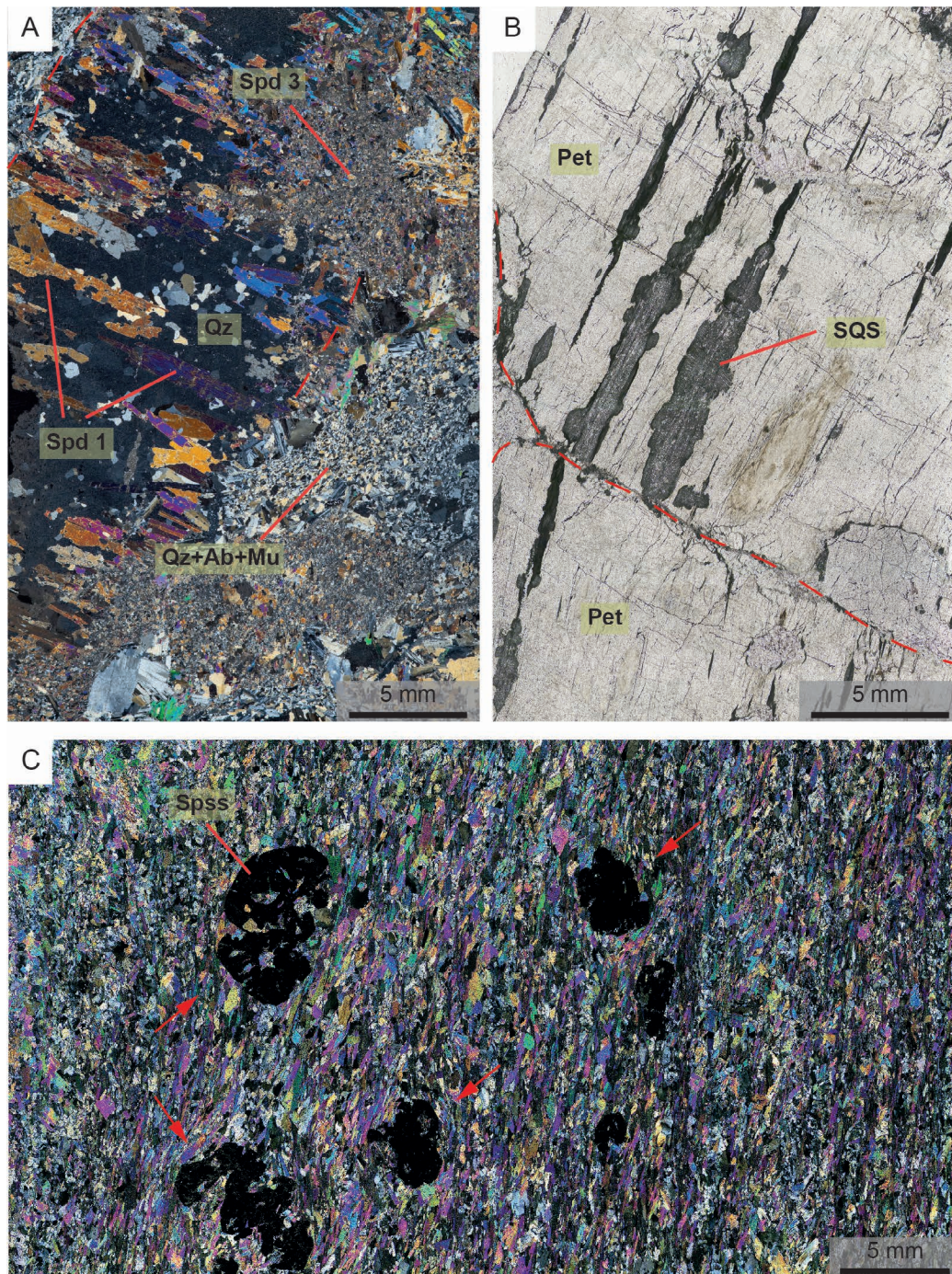


Figure 4.4: Photomicrographs from pegmatites in the Zulu pegmatite field. A) An example of Classic SQUI (Spd 1) within Type 1 pegmatites. Note the overprinting albitization (Event 3, Table 2) and subsequent late spodumene (Spd 3; Event 4, Table 2) events. B) Large petalite crystals in Type 2 pegmatites, with spodumene-quartz symplectites (SQS, Event 4, Table 2) occurring along petalite cleavage planes. C) A strongly foliated zone of a Type 1 pegmatite. Evidence of fabric-wrapping textures around spessartine (opaque mineral) highlighted in red arrows. Qz: quartz, Spd: spodumene, Ab: albite, Mu: muscovite, Pet: petalite, Spss: spessartine.

and barren (Figure 4.3A) subgroup. Petalite (in the mineralized subtype), blocky K-feldspar (in the barren subtype), quartz, albite, and muscovite constitute the major minerals in the intermediate zone. Crystals are commonly euhedral to subhedral. Fabric-forming recrystallization (Event 2, Table 4.2) is not observed within the Type 2 pegmatites.

Locally Type 2 pegmatites have been albitized (Event 3, Table 4.2), leading to replacement of the primary mineralogy by a fine-grained albite + quartz + muscovite assemblage. A muscovite + quartz greisen, predominantly along the margins of blocky K-feldspar and petalite, is sporadically developed. Some petalite in Type 2 pegmatites is partially replaced (Event 4, Table 4.2) by a complex symplectic intergrowth of spodumene + quartz along cleavage planes and fracture faces (Figure 4.4B). Evidence for further low-temperature alteration (Event 5, Table 4.2) is sparse, although feldspars and petalite are patchily replaced with clay-type minerals.

4.3.2 Lithostructural observations on the Zulu pegmatite field

Competency contrasts between lithologies generated distinct high-strain and low-strain domains during D_2 deformation. Across the study area the predominant foliation dips steeply towards the WNW (averaging 212° strike / 72° dip) across a region >5 km wide.

Within the basement granitoids the S_2 fabric is defined by the alignment of biotite and feldspar (Figures 4.7C, D). Strain appears to be heterogeneously distributed through the granitoids, with weakly foliated granitoids gradually obtaining strongly deformed and banded fabrics on the scale of 10s of m (Figure 4.7C). Local megacrystic granitoids within higher strain domains preserve sigma-type clasts indicating a sinistral, SSW oriented, sense of shear (Figure 4.7D).

The basal mafic/andesitic tuff succession on the eastern margin of the FRSGB is locally protomylonitic with the S_2 fabric defined by the alignment of relict amphibole and biotite crystals (Figure 4.8A). Similar mylonitic fabrics are preserved in the Sonop serpentinite, where S/C fabrics and relict phenocrysts also indicate sinistral shearing (Figure 4.8D). Low-strain domains are rarely observed within the Sonop serpentinite. S_2 high strain domains within the overlying mafic/ultramafic succession are defined by the alignment of tremolite/actinolite and chlorite-rich shear bands (Figure 4.8C). Low-strain domains occasionally preserve a S_1 fabric that is locally overprinted by thin, chlorite-rich shear bands (Figure 4.8C). S-shaped kink folds are common within the mafic/ultramafic succession and suggest sinistral NNE-SSW orientated shearing (Figures 4.7A,B).

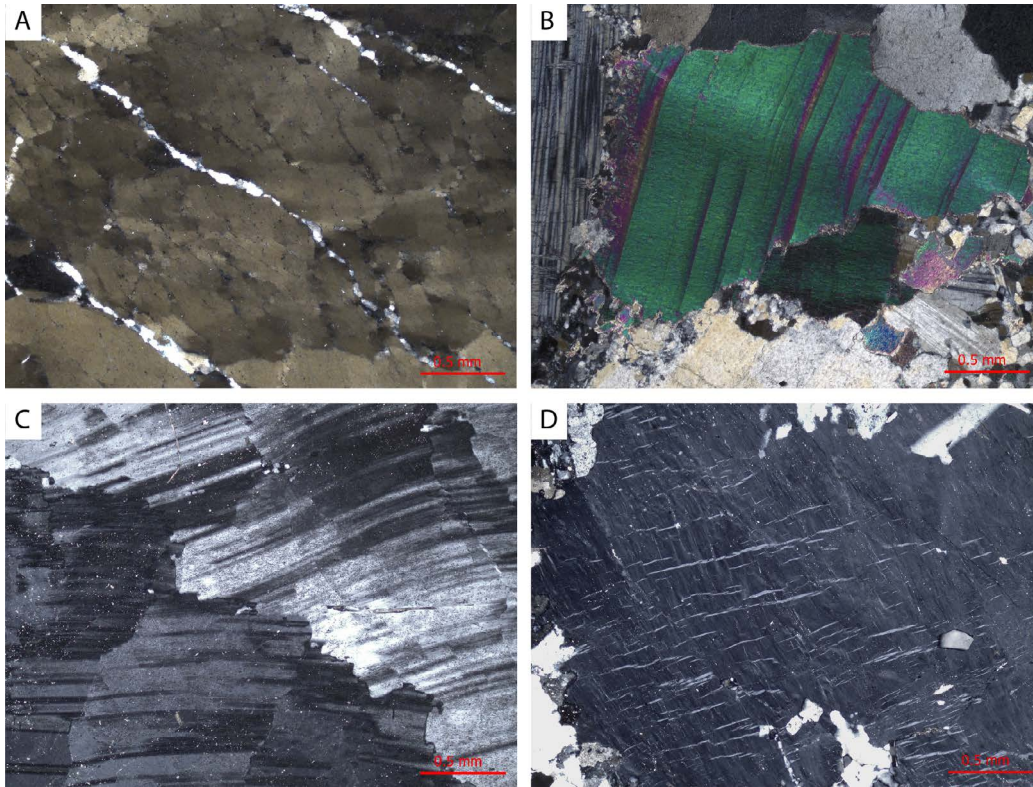


Figure 4.5: Typical deformation microstructures within Type 1 pegmatites. A) Chessboard extinction in relict magmatic quartz. B) Kinked cleavages within a relict muscovite crystal. C) Plastic deformation of early albite twins. D) Orientated flame perthite within early microcline.

Type 1 pegmatites are primarily emplaced within high-strain domains along the contact between the Sonop serpentinite and underlying mafic/andesitic tuff successions. Host rocks to Type 1 pegmatites have been affected by exomorphic wall rock alteration adjacent to the pegmatites which largely destroyed their pre-intrusion mineralogy. A biotite + quartz alteration zone, up to 20 cm thick, is typically observed along the serpentinite hanging wall with a pervasive foliation subparallel to S_2 (Figure 4.9). Asymmetric crenulations are common within the biotite-rich domains. Along the footwall of the pegmatites, a chlorite + biotite \pm holmquistite alteration halo within the mafic/andesitic tuff succession is commonly observed, with the acicular amphibole and platy mica showing a preferred alignment subparallel to S_2 (Figure 4.8B). Fabrics defined by the alignment of albite and mica in the early dynamically recrystallized zones, together with the schistosity preserved in high-strain domains, are all subparallel to the long axis of the pegmatites and S_2 (Figure 4.9).

Type 2 pegmatites are mostly emplaced within the basement granitoids and low-strain domains within the basal mafic/andesitic tuff successions, and often obliquely

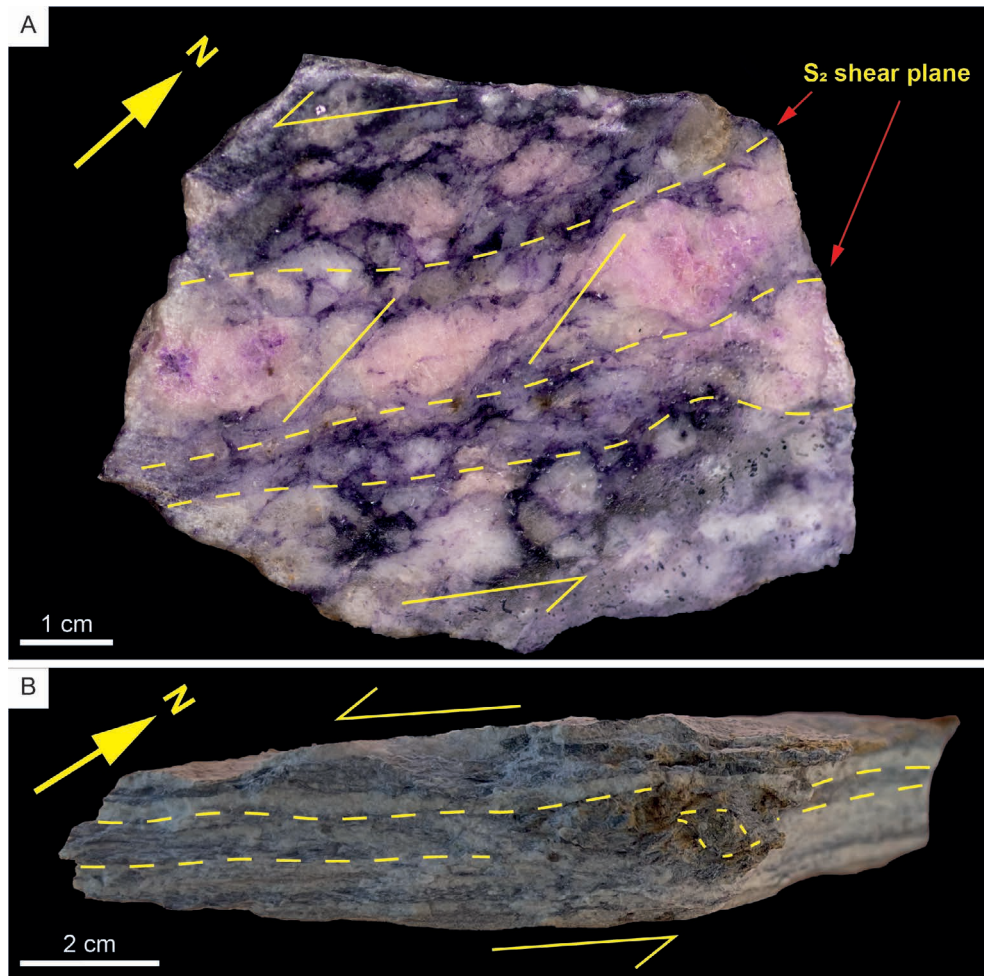


Figure 4.6: Hand samples of Type 1 pegmatites. A) Sheared pegmatite with fabric defined by aligned mica sheets, and antithetic rotation within relict SQI grains. B) Protomylonitic pegmatite with bands defined by albite and lithian muscovite. Spessartine porphyroclasts indicate the sense and direction of shear.

(30–90°) crosscut the primary fabric (Figure 4.10). There are no apparent deformation fabrics within Type 2 pegmatites, and only minor alteration can be observed within the wall rock.

Deformation features related to motion along the Mtshingwe fault (D_3) are absent in regions containing Type 1 pegmatites. Type 2 pegmatites adjacent to D_3 structures are only observed within the granitoid basement and are oriented oblique to the en-echelon dolerite dyke swarm within the Mtshingwe fault zone. In these examples, the primary magmatic mineralogy is pervasively altered to clays.

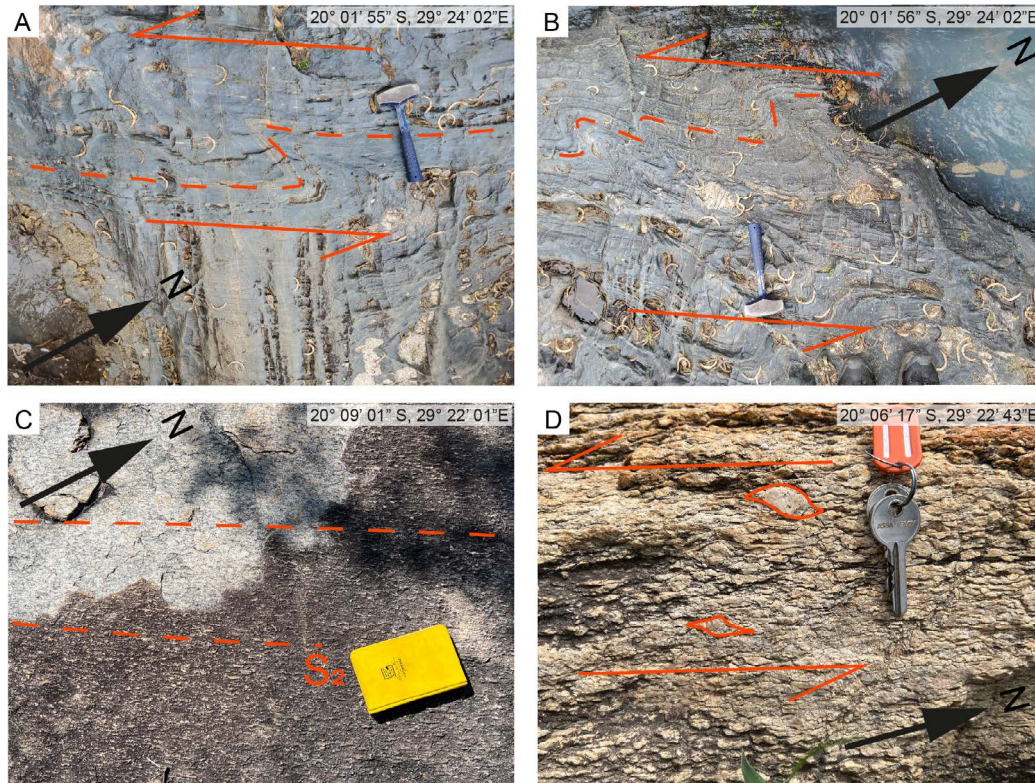


Figure 4.7: Top-down outcrop images of host rocks within the Zulu Pegmatite Field. Location of images in top right. A, B) S-shaped kink folds indicating sinistral sense of shear within mafic/ultramafic extrusives. C) S_2 fabric preserved in moderately deformed basement granitoids highlighted by the alignment of feldspar crystals. D) sigma-type clasts within highly deformed basement granitoids.

4.4 Discussion

4.4.1 Relative timing and structural setting of pegmatite emplacement at Zulu

Transcurrent sinistral shearing along the ISLZ and ZSZ is defined by asymmetric dragging of the pre-existing S_1 foliation along the S_2 shear fabric, asymmetrical sigma clasts, and s-shaped kink folds developed along the margins of the high-strain zones (Figure 4.7).

Type 1 pegmatites are interpreted to have been emplaced during D_2 shearing. Deformation within the pegmatites was accompanied by extensive recrystallization and remobilization of lithium, leading to the formation of lithian muscovite and the replacement of pre-existing lithium silicates with a later fine-grained generation of spodumene (Table 4.2). Critically, the syn-kinematic emplacement during sinistral shearing of the pegmatites is supported by three main lines of evidence: 1) the lentic-

ular

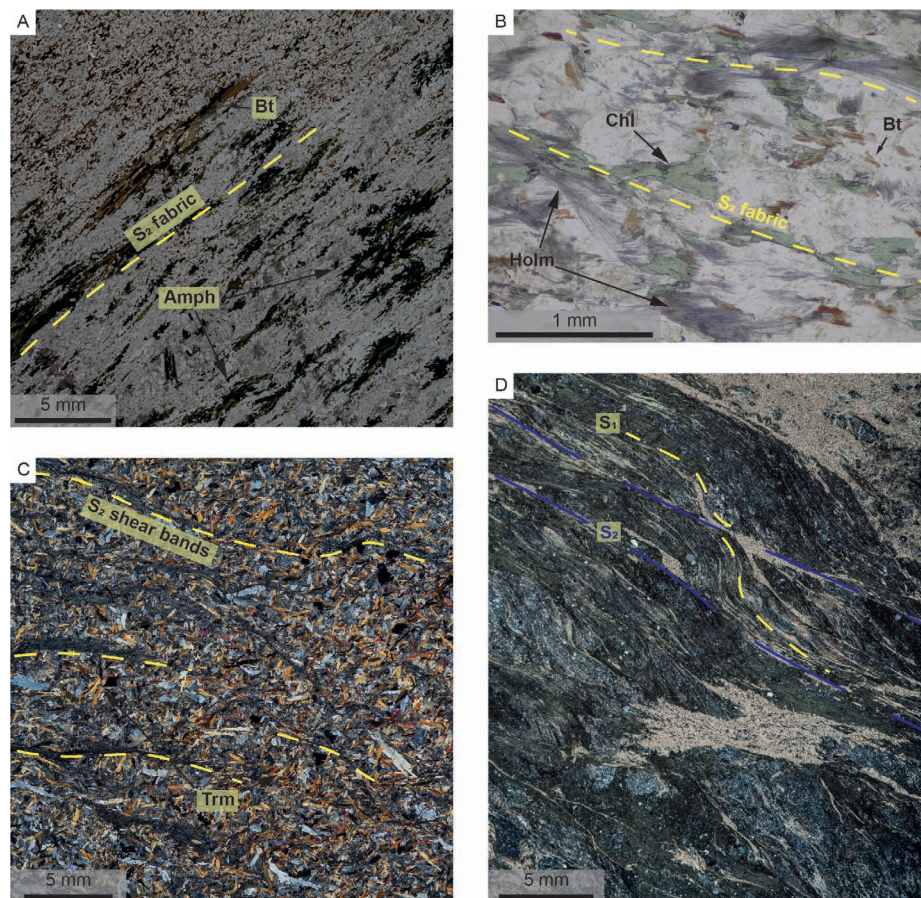


Figure 4.8: Thin section images of host rocks in the Zulu Pegmatite Field. A) Protomylonitic andesitic tuff with the fabric defined by highly attenuated biotite and amphibole crystals. B) Strongly altered andesitic tuff adjacent to a Type 1 pegmatite. Alteration phases are all strongly aligned to S_2 . C) S_2 chlorite shear bands within a mafic/ultramafic extrusive in a low strain domain. D) Highly sheared serpentinite highlighting the relationship between the earlier S_1 fabric (yellow dashed lines) with the well-developed shear fabric (S_2 , blue lines) forming S/C fabrics within the principal deformation zone of the Zulu shear zone. Bt: biotite, Amph: amphibole, Chl: chlorite, Holm: holmquistite, Trm: tremolite.

shape of some Type 1 pegmatites in agreement with emplacement into dilational jogs (in some cases, Type 1 pegmatites present a more contorted shape which may suggest deformation by subsequent shearing), because of which the orientation of the pegmatite-wallrock contacts locally cross-cuts the S_2 fabric (Figures 4.2, 4.10), 2) the internal fabric subparallel to S_2 within pegmatites (Figures 4.3, 4.5, 4.8), and 3) the synkinematic growth of minerals within the exomorphic haloes surrounding Type 1 pegmatites, which indicates that country rock metasomatism must have occurred

during active deformation (Figures 4.8B, 4.11). The significant rheological contrast between the weaker, more ductile Sonop serpentinite and the more competent underlying metamorphosed tuffs (Figure 4.2) likely acted as a concentrating mechanism, localizing melt flow during ascent (Papeschi et al., 2022).

In contrast to Type 1, the Type 2 pegmatites preserve a largely magmatic mineralogy, and have no internal fabric. Cross-cutting relationships with their host foliated granites indicate that these pegmatites were emplaced after the regional fabric was de-

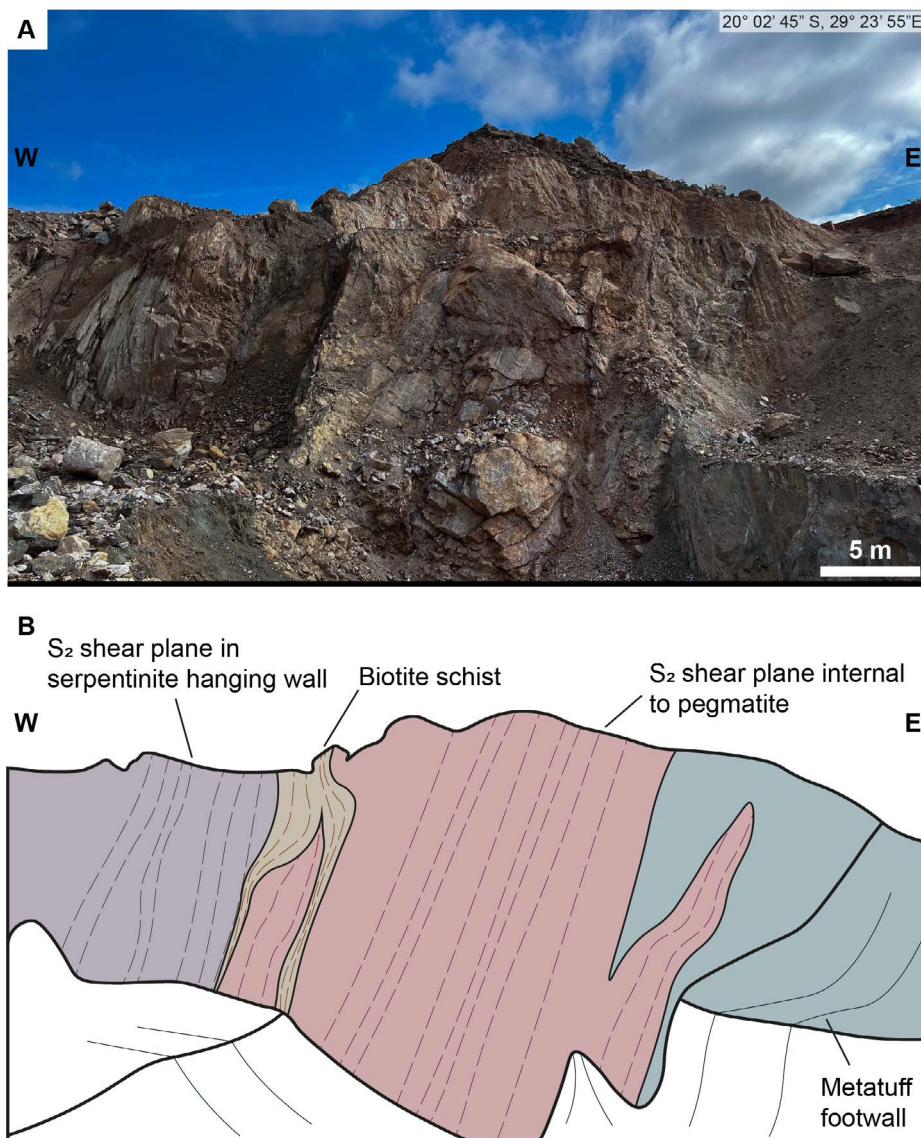


Figure 4.9: A) Outcrop image of a Type 1 pegmatite in an excavated pit. Image traced in B. B) Relationship of the S_2 fabric within the host rocks and the fabric internal to the pegmatite highlighted, with the common biotite schist along the hanging wall of the pegmatite. Location of image in top right.

veloped. Given the systematic orientation of individual bodies, we suggest that Type 2 pegmatites were syn-to-late-kinematic to D_2 and emplaced along typical sinistral strike-slip subordinate fracture sets and/or tension gashes (Figures 4.10 and 4.11). This timing, and emplacement into zones of lower strain, served to truncate the cooling history of Type 2, inhibiting extensive recrystallization and preserving primary magmatic mineralogy and textures.

As well as controlling the emplacement trajectory, host rock lithology also influenced the size of individual pegmatite bodies. As the Zulu pegmatite field extends away from the Fort – Rixon Shangani Greenstone belt into the basement granitoids, individual pegmatites become narrower in width. We interpret this to be driven by the absence of significant rheological contrast within the more competent granitoid host rocks. Whereas the relatively incompetent lithologies and lithological boundaries within the greenstone belt accommodated strain by ductile deformation and concentrated fluid flow along lithological boundaries, the more competent granitoid basement likely accommodated strain through localized brittle fracturing. Narrow mode-1 fractures formed as a result, facilitating pegmatite emplacement within the basement granitoids (Figure 4.11). This behavior would have limited dilation, promoting the formation of several thin, strike-parallel intrusions, instead of thick bodies.

4.4.2 Comparison with other pegmatite fields

Systematic structural analyses of economic pegmatite fields are rare in the literature. However, similar relationships to those observed in the Zulu pegmatite field have been described in other economic pegmatite districts and highlight the importance of regional-scale controlling structures on the size, orientation, and distributions of pegmatites themselves. Four well-studied examples are briefly mentioned here.

The Archaean Greenbushes pegmatite field, Western Australia, is one of the largest hard-rock Li deposits in the world (Bowell et al., 2020, , Figure 4.12A). The Greenbushes pegmatite is emplaced into the 150 km long and 15–20 km wide Donnybrook–Bridgetown shear zone, which preserves a major sinistral component of movement (Partington et al., 1986; Partington, 1988). Detailed fieldwork has shown that mineralized pegmatites of the Greenbushes pegmatite contain relict igneous structures and pegmatitic microstructures indicating syn-kinematic emplacement into structures related to the Donnybrook–Bridgetown shear zone. Individual pegmatite bodies strike subparallel to their controlling structures, as with the Type 1 pegmatites in the Zulu pegmatite field.

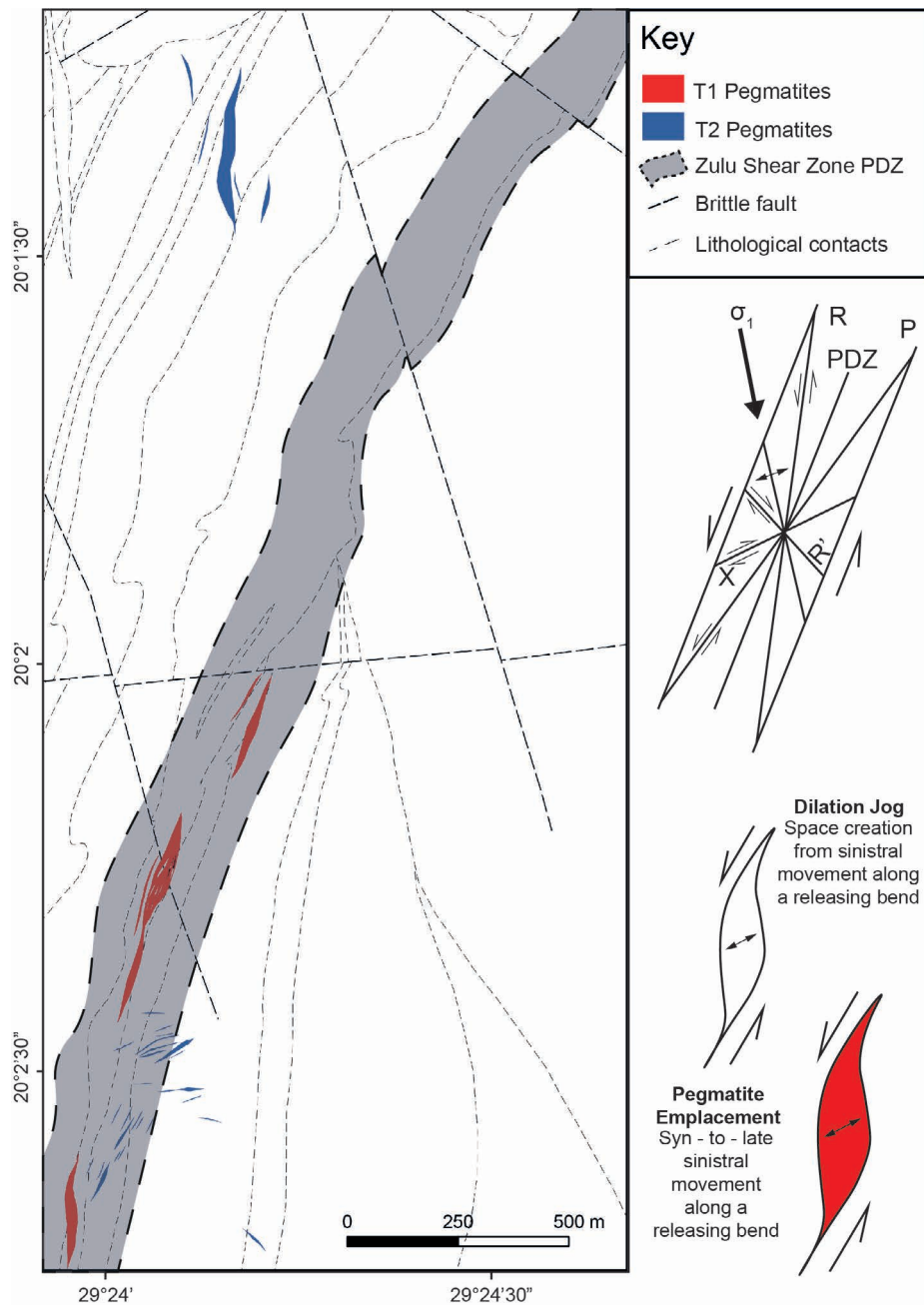


Figure 4.10: Simplified geological map discriminating between group one and two pegmatites. A schematic emplacement model of the Zulu Pegmatite Field with a parallel Riedel schematic is drawn to the right of the map. T1: Type 1, T2: Type 2, PDZ: principal deformation zone.

Pegmatites of the north-eastern Archaean Pilbara craton in Australia are also associated with major regional structures (Figure 4.12B). Several N–S to NNE–SSW trending sinistral shear zones transect the craton. Most large Li-rich pegmatites (e.g., Wodgina, Tabba Tabba) in the Pilbara are hosted within subparallel dilational

sites where greenstone belts are cross-cut by these shear zones (Sweetapple, 2000; Sweetapple and Collins, 2002). The Zulu pegmatite field is similarly located at the transect between a major regional structure and a greenstone belt.

The Brazil Lake pegmatite field was emplaced into the Meguma terrane (540 – 420 Ma) metasedimentary and metavolcanic successions (Figure 4.12D, Culshaw and Reynolds, 1997; Kontak, 2006). Shear zones are common in the area and were repeatedly reactivated during both deformational events. Pegmatites within the Brazil Lake pegmatite field were emplaced at 395 Ma along major lithological boundaries into dilational sites that have orientations sub-parallel to dextral strike-slip movement (Kontak, 2006). Similarly to the Type 1 pegmatites of the Zulu pegmatite field, deformation along the pegmatite–wall rock contact and deformation within primary magmatic phases at Brazil Lake indicate that pegmatite emplacement was syn-kinematic, with shearing having continued post-crystallization.

The Winnipeg River pegmatite district is largely hosted within the Archaean Bird River greenstone belt and hosts several pegmatite fields, including the world-class Tanco deposit (Figure 12C, Černý, 1982; Gilbert et al., 2008). Three major deformation events have been recognized in the region and are summarized by Duguet et al. (2009). Most pegmatites within the Winnipeg River pegmatite district are spatially and temporally associated with reactivated D_3 shear zones, such as the North Bernic Lake shear zone (Baadsgard and Černý, 1993; Kremer, 2010). Both ductile and brittle fabrics formed in response to D_3 , and pegmatites were emplaced into both ductile and brittle dilational sites which in some cases result in pegmatite orientations oblique to the main shear zone (Brisbin and Trueman, 1982). As opposed to the Zulu pegmatite field, where the largest pegmatites formed syn-kinematic to ductile deformation, the largest pegmatites in the Winnipeg River pegmatite district (e.g., the Tanco pegmatite) appear to have occurred late-kinematic during colder, brittle deformation (akin to Type 2 pegmatites at in the Zulu Pegmatite Field, Kremer, 2010). In this way the timing of peak melt flux may be an important factor in determining which structural setting within shear zones are most amenable to large pegmatite deposits.

These examples, together with our observations at Zulu, highlight that major pegmatite fields hosting economic deposits formed by syn-kinematic pegmatite emplacement along major ductile shear zones.

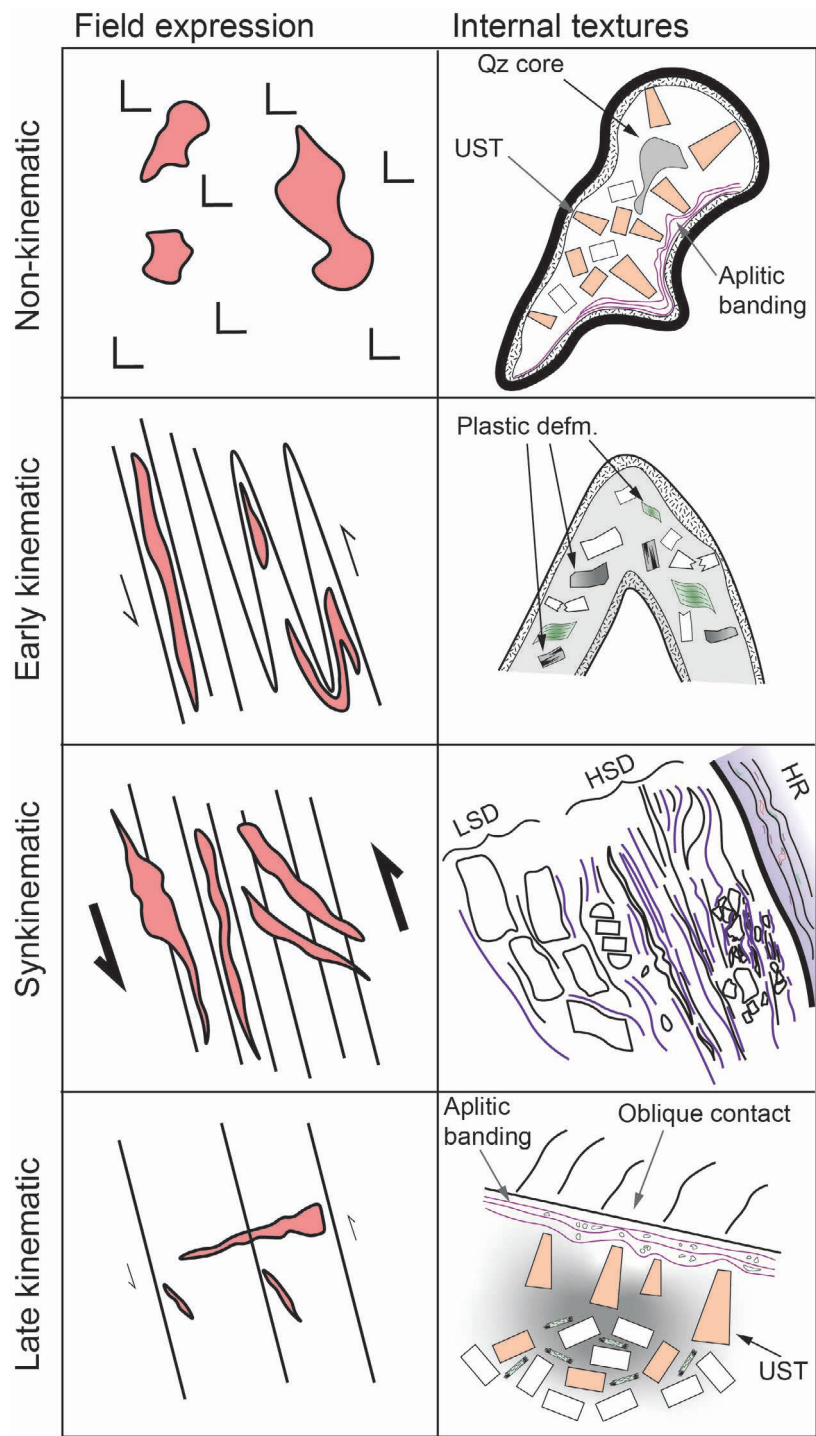


Figure 4.11: Caption on next page.

Figure 4.11 (continued): Schematic diagram highlighting the different styles of emplacement relative to the regional structure for pegmatites emplaced at different times during shear-dominated deformation. Both field expressions (left) and internal textures (right) are drawn. Non-kinematic pegmatites have no relationship to any fabric, nor preserve any dynamic recrystallization textures related to syn-emplacement deformation. Early kinematic pegmatites are likely to become tightly infolded with the shear fabric, and any magmatic phases preserved will exhibit extensive crystal-plastic deformation. In the syn-kinematic case, the fabric internal to the pegmatite is parallel, whilst the orientation of the intrusion locally crosscuts, deformation fabrics in the host rock. Syn-kinematic growth of exomorphic alteration phases provides additional evidence to support syn-kinematic emplacement, as is observed in Type 1 pegmatites in the Zulu pegmatite field. Late kinematic pegmatites will be emplaced into tension gashes and/or mode-1 fractures in the shear zone, and preserve a largely primary mineralogy, as is observed in Type 2 pegmatites in the Zulu pegmatite field. UST: unidirectional solidification textures, HSD: high-strain domain, LSD: low-strain domain, HR: host rock.

4.4.3 The influence of shear zones on pegmatite migration and emplacement

In the mineral systems framework, understanding the formation, migration, and crystallization of pegmatitic melt from source to trap is critical to constrain the expected distribution of such ore bodies in a particular district at the current level of exposure (Wyborn et al., 1994; Hagemann et al., 2016; Gardiner et al., 2024b). Whilst our observations of the Zulu pegmatite field are primarily focused on the trap (given the source is not exposed), the close association between the ILSZ and the Zulu Pegmatite Field inevitably leads to speculation on how such shear zones might be drivers and/or enablers of the localization and migration of pegmatitic melt, leading to the crystallization of economic pegmatites. This is irrespective of the source model (granite fractionation vs. anatectic) invoked, since regardless of melting mechanism, to form a large deposit significant volumes of pegmatitic melt must be extracted and physically accumulated within the mid-upper crust. Although pegmatitic melt migration into pre-existing structures may generate a spatial correlation without having to call upon a genetic association (Deveaud et al., 2013; Lee et al., 2020), here we briefly highlight how active shear zones may serve to promote the accumulation and crystallization of economic (i.e., large) pegmatites.

4.4.3.1 Mechanisms of pegmatite melt extraction

The model of anatectic pegmatite formation requires low degrees of partial melting in a fertile protolith to generate melts that are enriched in incompatible elements (Simmons et al., 1995; Shaw et al., 2016; Müller et al., 2017). The threshold for physical interconnectivity of low-volume silicate melts, is often taken to be 7–10 vol. %, as determined experimentally (Rosenberg and Handy, 2005). However, Etheridge et al. (2021) showed that connectivity may be achieved at lower degrees of melt in actively deforming environments where dehydration melting processes dominate, given the volume increase caused by such reactions may drive sudden brittle failure and allow channelization of melt at proportions as small as 2 vol. %. Given that reducing the melt volume required for extraction enhances the degree of enrichment an anatectic source is able to achieve, active deformation may play an important role in realizing this model.

Similarly, the model of pegmatite melt extraction from highly fractionated granite systems (either as individual plutons or long-lived mid-crustal magma chambers) is also constrained by melt connectivity, in this case limiting the extent of fractionation that can be invoked prior to system lock-up (Vigneresse and Tikoff, 1999, Chapter 3). Although localized overpressure of melt can drive the ejection of melt from granitic source regions through dykes (Rubin, 1995; Baker, 1998), and pegmatitic intrusions formed in this way are frequently seen (e.g., Roda-Robles et al., 2023), these intrusions will be limited in size due to an absence of significant strain anisotropies in the surrounding host rocks (Brisbin, 1986). Alternatively, the presence of tectonic stresses within a melt-present environment facilitates efficient extraction of melt (Brown and Solar, 1998; Sawyer et al., 2011; Brown et al., 2011), and enable large volumes of highly fractionated melts to migrate out of their granite source region (Černý and Ercit, 2005).

It is thus worth noting that irrespective of the source mechanism invoked, the extraction of large amounts of melt is best facilitated in environments of active deformation. This promotes the formation of fewer, large, intrusions (as observed at Zulu), as opposed to several smaller pegmatitic bodies (e.g., pegmatitic pockets internal to –, or radial pegmatite haloes around the source granite).

4.4.3.2 Pegmatite melt migration

Upon extraction of the final pegmatitic melt from the source region, the style and character of pegmatitic melt migration upwards through the crust before reaching a

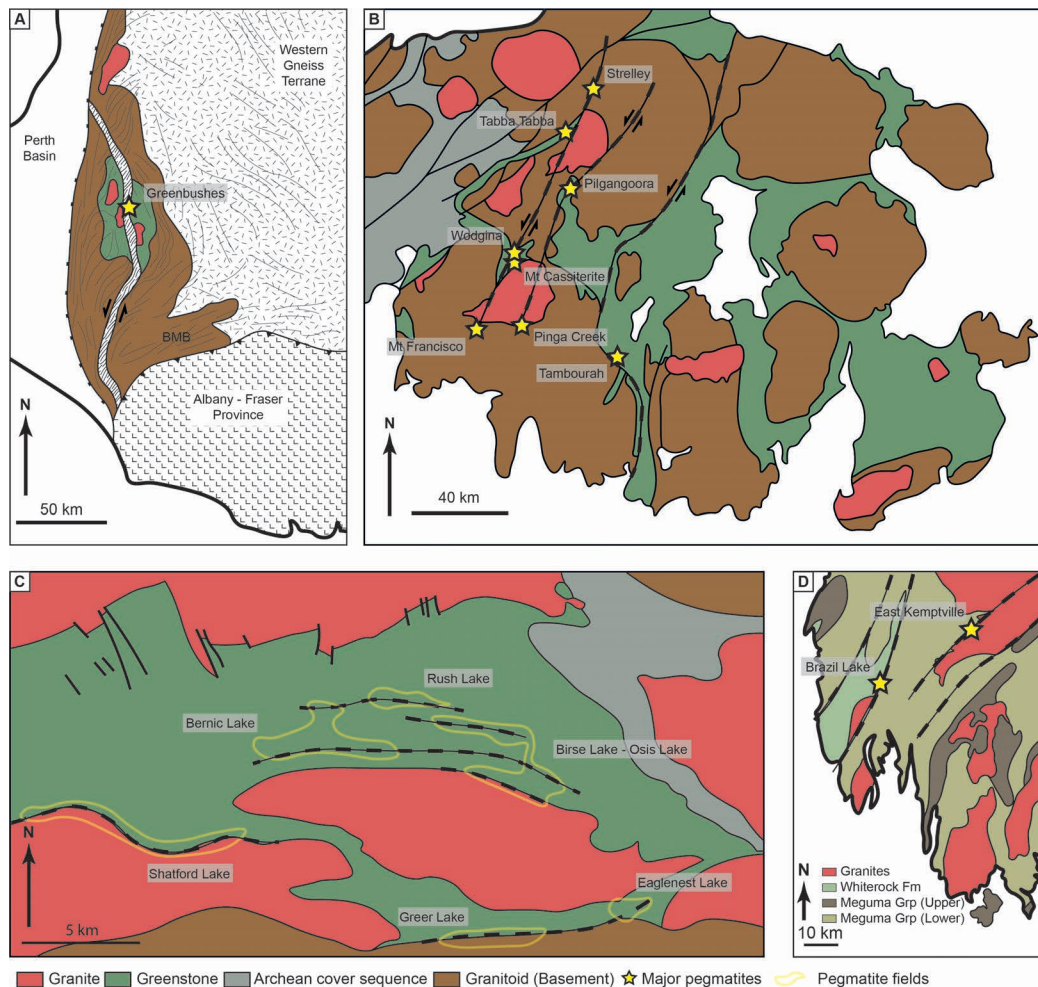


Figure 4.12: Example case studies from the A) Yilgarn Craton (after Partington et al., 1995), B) North-east Pilbara craton (after Sweetapple and Collins, 2002), C) Winniper River pegmatite district (after Kremer, 2010), and D) southwestern Nova Scotia (after Kontak, 2006).

trap depends on host rock rheology, temperature, and strength anisotropies within the crust (Brisbin, 1986). To promote the necessary undercooling to generate characteristic pegmatitic textures (Černý et al., 2005; London, 2005; Simmons and Webber, 2008; McCaffrey and Jowitt, 2023), pegmatitic melts need to be rapidly extracted from their source region and emplaced into cooler host rocks, or chemically quenched during ascent. If melt is present during ductile shearing, then these shear zones can generate regions of increased porosity and dilation, providing a siphon force for melt migration (Sibson et al., 1975; Etheridge et al., 2021), which is an effective way of forcing large volumes of melt to migrate through the crust (Brown, 2013).

The type of shear-zone system (transcurrent, normal, reverse, or oblique) also

has an effect on the vector of melt flow (Figure 4.13, Brisbin, 1986; Brown, 2013; Etheridge et al., 2021). Channelized flow along vertical dilatant conduits preferentially occurs in both transcurrent and reverse shear zones, and are therefore favourable targets for melt migration (Figure 4.13, Lindroos et al., 1996; Brown and Solar, 1998). Furthermore, transcurrent shear zones (such as the ILSZ) will promote a horizontal percolation of flow during shearing, enabling the concentration of melt from a large melt region towards sites of lower pressure prior to vertical extraction, enabling tapping of a large fertile source region (Figure 13, Sibson et al., 1975; Brown and Solar, 1998; Cavalcante et al., 2016). In this way, transcurrent releasing bends are highly favorable sites of pegmatitic melt migration and subsequent emplacement, given a fertile source is intersected, such as at Zulu or the examples presented above.

Importantly, during the waning stages of shearing, pre-existing anisotropies caused by lithological boundaries, fractures, or tensile cracks, will preferentially allow residual melt migration through the system as stress is relieved. This will result in a change of preferred orientation of late-kinematic pegmatites to oblique to the controlling structure, as is seen with Type 2 pegmatites in the Zulu pegmatite field (Brisbin, 1986; Araújo et al., 2001; Demartis et al., 2011; Bhatt et al., 2019). Given sufficiently high melt pressure and vertical flow, late pegmatitic melts may generate large intrusions (e.g., the Tanco deposit) into higher-order late brittle fractures or tension gashes (Brisbin, 1986; Duguet et al., 2009; Kremer, 2010; Bhatt et al., 2019). In this scenario, pegmatites of economic size may not be orientated parallel to the trend of the major shear zone, as subordinate structures can form oblique to the major shear zone strike (Figures 4.10 and 4.11).

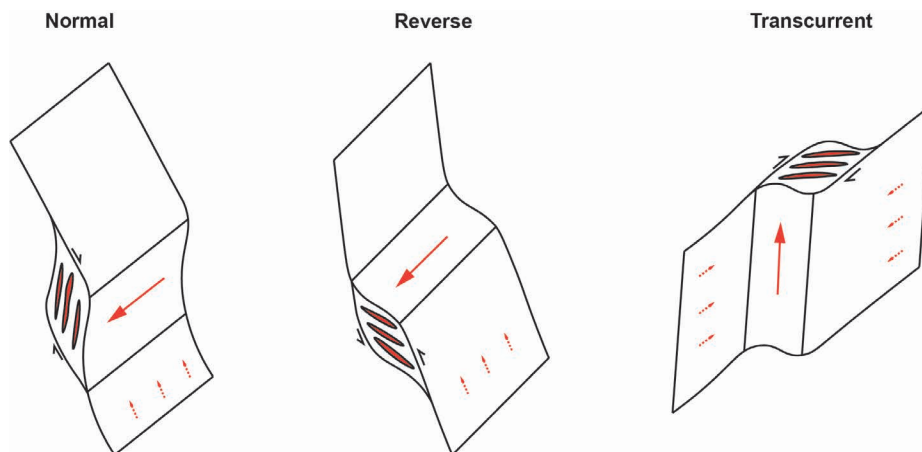


Figure 4.13: Schematic sketches of melt flow orientations relative to three major structural types (modified from Etheridge et al., 2021). Dashed lines represent preferred direction during percolative flow, and solid lines represent preferentially direction of channelized flow. The geometry of dilational jogs is shown for each fault type. Both normal and reverse faults generate subhorizontal zones of dilation, whereas transcurrent systems generate vertical conduits.

4.4.4 Implication for exploration strategies

Irrespective of the source of pegmatites, we suggest that crustal shear zones can play a fundamental role in allowing pegmatitic melt to migrate away from its source region. The localization of melt within shear zones enables the formation of large bodies of evolved melt, in contrast to smaller individual bodies formed around a parental granite in the absence of tectonic stresses (Černý, 1991b). Thus, when present, shear zones act as a first-order control on the spatial distribution of large pegmatite fields, as is seen at several major deposits such as Greenbushes, Tanco, and Zulu.

We propose that desktop studies should focus on delineating potential structural corridors through which pegmatitic melts may have migrated, with a preference for transcurrent shear zones. The presence of granites with fertile characteristics (reduced K/Rb, elevated concentrations of incompatible elements such as Li, Rb, F, Cs, Selway, 2005) along the shear zone may further indicate that the shear zone has intersected a fertile source region, though does not imply that the fertile granites are the source of pegmatites. Such a systematic approach can help delineate the new prospective regions for discovering economic pegmatite bodies.

4.5 Conclusions

The Zulu pegmatite field is an excellent case study with which to highlight the importance of ductile shear-related deformation on both the emplacement mechanisms of pegmatites – where they control the orientation, distribution, and width of individual bodies – alongside controlling the degree of subsolidus recrystallization a pegmatite experiences during deformation. At Zulu, major competency contrasts between lithologies localized high-strain domains during sinistral strike-slip shearing. Pegmatites were emplaced into subparallel dilational jogs syn-kinematic to shearing and subsequently extensively recrystallized during continued deformation, driving the (re)precipitation of fine-grained spodumene and muscovite. By contrast, late-kinematic pegmatites at Zulu are emplaced along late-kinematic subordinate fracture sets and tension gashes oblique to the principal shearing direction and preserve their magmatic mineralogy owing to limited subsolidus deformation.

Shear zones are known to be regions that allow efficient migration of silicate melt through the continental crust. In the presence of a fertile source, specifically transcurrent shear zones enable efficient migration of significant pegmatitic melt from a hot source region to a shallower level, cooler host rock, driving undercooling and enabling the rapid crystallization textures distinctive of pegmatites. Future greenfield exploration should therefore focus on constraining transcurrent shear zones where significant strength anisotropies are present, with indicators of fertility at the melt source (e.g., enriched granites along the shear zone). Constraining the structural history and relative timing of pegmatite emplacement through detailed fieldwork and petrographic analyses can subsequently shed light on the distribution and orientation of individual pegmatite bodies within a pegmatite field. Importantly, a focus on constraining the source (granite vs. anatexis) of a pegmatite district may therefore be of lesser importance, as regional structures have a more important control on the distribution of individual bodies within a district.

Chapter 5

Archean tectonic evolution of the Zimbabwe Craton and the mechanisms of lithium enrichment

5.1 Introduction

5.1.1 Chapter outline

In this study on the Zimbabwe craton, new coupled geochronology and geochemistry is integrated into a compilation of pre-existing data to develop a holistic understanding of cratonic evolution. Using this dataset, and testing it against the results of petrologic modelling, a complete model of craton development is proposed, which includes periods of both vertical and horizontal accretion. Early juvenile continental growth was submarine; later potassic (evolved) magmatism is suggested to have supported continental emergence and the formation of late supracrustal cover sequences. A new model for lithium pegmatite sources is therefore required, which based on the present data is proposed to revolve around repeated recycling of felsic continental crust through melting, rather than erosion, which sufficiently drove lithium enrichment to obviate the need for sedimentary sources in the melting environment.

5.1.2 Sources of TTG melts

Tonalite–trondhjemite–granodiorite (TTG) suite rocks are typically metaluminous, sodic granitoids that constitute >80% of preserved Archean crustal fragments (Bleeker, 2003; Polat, 2012). These juvenile intrusive rocks are characterised by high $\text{Na}_2\text{O}/\text{CaO}$, low $\text{K}_2\text{O}/\text{Na}_2\text{O}$ (TTG *sensu stricto*), and a relatively simple mineralogy of predominantly plagioclase feldspar and quartz, with biotite or amphibole \pm epidote (Martin, 1994; Bédard, 2003; Moyen et al., 2006; Moyen and Martin, 2012). The TTG suite

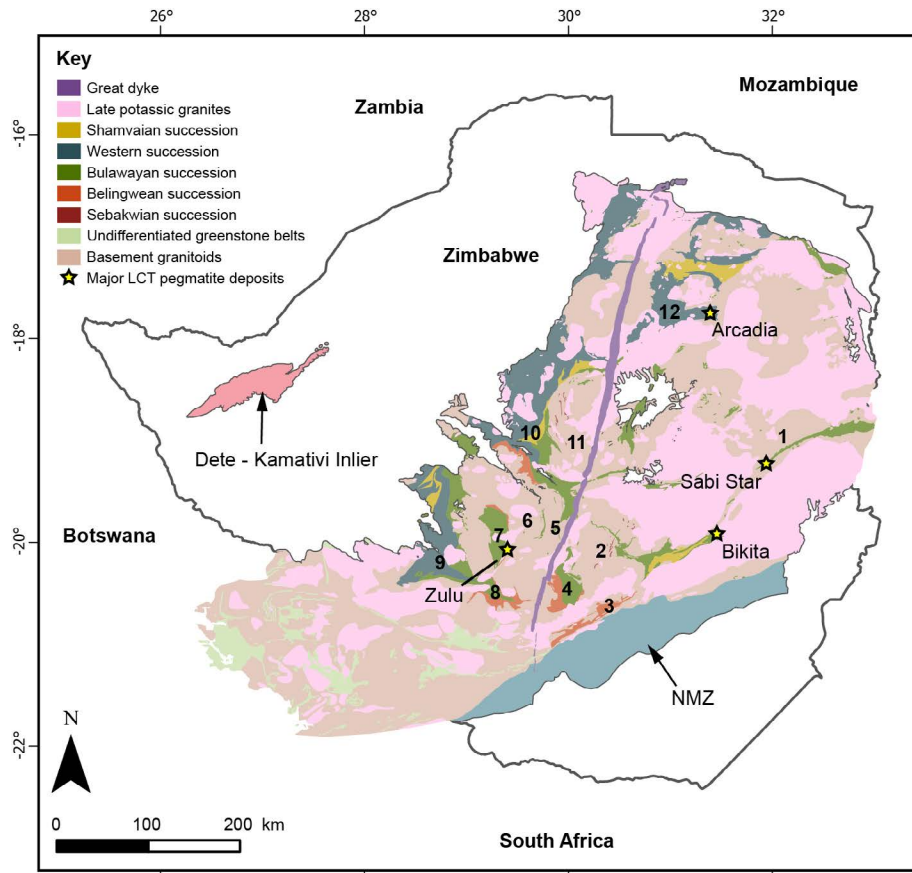


Figure 5.1: Simplified geological map of the Zimbabwe craton, highlighting key locations discussed in this chapter. 1: Murewa–Odzi greenstone belt, 2: Tokwe terrane, 3: Buhwa greenstone belt, 4: Belingwe greenstone belt, 5: Gwenoro dam, 6: Shangani batholith, 7: Fort Rixon–Shangani greenstone belt, 8: Filabusi greenstone belt, 9: Bulawayo greenstone belt, 10: Midlands greenstone belt, 11: Rhodesdale terrane, 12: Harare–Shamva greenstone belt. NMZ: Northern marginal zone.

is often divided into three geochemical groups, which historically have been inferred to represent the depth of melting of their source; 1) a ‘high-pressure’ group characterised by high Al_2O_3 , Na_2O , Sr, and La/Yb, with low Y and Nb/Ta, 2) a ‘low-pressure’ group characterised by low Al_2O_3 , Na_2O , Sr, and La/Yb, as well as high Y and Nb/Ta, and 3) a ‘medium-pressure’ group which displays transitional characteristics between high- and low-pressure groups (Martin and Moyen, 2002; Moyen, 2011; Moyen and Martin, 2012; Laurie et al., 2013). These geochemical characteristics are thought to arise through variations in the restitic mineralogy; high Sr contents in the melt likely indicate a plagioclase-poor residuum, low Nb/Ta corresponds to the presence of rutile, and high La/Yb arises from the presence of garnet in the residuum (Martin and Moyen, 2002; Foley et al., 2002; Rapp et al., 2003). However, varia-

tions in source chemistry, melt mixing, and fractional crystallisation can all impose additional variability in TTG chemistry (Moyen et al., 2006; Getsinger et al., 2009; Hoffmann et al., 2014).

It is generally accepted that TTGs are sourced by melting of a hydrous metabasalt (Drummond and Defant, 1990; Arculus and Ruff, 1990; Rapp et al., 1991; Wolf and Wyllie, 1994; Foley et al., 2002; Moyen and Martin, 2012), though some have argued that fluid absent conditions are preferable (Rapp et al., 2003; Moyen and Stevens, 2006). Experimental studies have shown that Archean TTGs likely formed at temperatures of 750–950 °C (Xiong et al., 2009; Qian and Hermann, 2013), which is supported by phase equilibria modelling (Palin et al., 2016a; Johnson et al., 2017; Kendrick and Yakymchuk, 2020). Generating voluminous melts in this temperature range requires water saturation (e.g., Palin et al., 2016a; Pourteau et al., 2020), therefore supporting a hydrous source. Many modelling studies consider the enriched Archean tholeiite (EAT) of Condie (1981) as a representative metabasaltic protolith, as melts derived from EAT provide a close match to natural Archean TTG compositions from both a major element and trace element perspective (Bédard, 2006; Martin et al., 2014; Palin et al., 2016a).

A key question related to Archean TTG formation is the geological setting in which the melts formed, which provides constraints on the geodynamic character of the early Earth. Given 50–60% of Earth’s continental crust had formed by the Archean (e.g., Taylor and McLennan, 1995; Hawkesworth and Kemp, 2006; Dhuime et al., 2012), the geological processes responsible for generating Archean TTGs must have been highly efficient. Modern day continental crust is primarily formed in continental and island arcs (e.g., Rudnick, 1995; Taylor and McLennan, 1995; Jagoutz and Kelemen, 2015), where dehydration reactions in the subducting slab release water into the overlying mantle wedge, driving fluid-mediated melting of the mantle and the generation of basaltic melts. These basaltic melts either form the basis of arc magmatism (through assimilation and fractional crystallisation), or underplate the lower crust and heat the lower crust to promote fluid-present melting of the lower crust and generating andesitic-dacitic magmas.

A similar tectonic environment has been suggested for the formation of TTGs (e.g., Martin and Moyen, 2002; Rapp et al., 2003), largely based on geochemical similarities between TTGs and modern day adakites (Martin, 1999), the latter of which are derived by directly melting the subducting slab (Castillo, 2012). Experimental studies (see Moyen and Stevens, 2006) show that garnet first appears in metamorphosed EAT-like mafic rocks at pressures > 10 kbar, and so melting at these condi-

tions can generate characteristic 'high-pressure' La/Yb signatures. Proponents of this model where Archean TTGs formed in arc-like environments either invoke modern-day steep subduction zones and thick mantle wedges, or flat-slab subduction with little to no interaction with the asthenospheric mantle, with changes in Ni, Cr, and MgO content of TTGs used to discriminate between these processes (Martin et al., 2005). Given TTGs represent a significant portion of preserved Archean terranes, substantial amounts of hydrated mafic source material needs to be transported to garnet-stability conditions, and subduction zones are an efficient pathway for doing so on the modern Earth.

Whether modern day subduction could have operated on the early Earth is a matter of intense current debate (Stern, 2005), with many researchers concluding that this is an unlikely scenario given the paucity of key geological indicators of convergent plate margin processes in Archean terranes (c.f. paired metamorphic belts, Miyashiro, 1973; blueschists, though see Palin and White, 2016). The Archean mantle was also significantly hotter than today (Korenaga, 2006) and was associated with a thicker and more buoyant primary mafic oceanic crust (Herzberg et al., 2010), which would have reduced the efficiency of subduction (van Hunen and Moyen, 2012; Korenaga, 2013). Further, felsic crustal growth in several Archean terranes appears to have been episodic, which is in contrast to the continuous nature of formation in subduction terranes (Stein and Hofmann, 1994; Rino et al., 2004; Dhuime et al., 2012; Hawkesworth et al., 2016). Archean terranes also show large-scale dome-and-keel structures that are unlike linear fold and thrust belts of younger mountain belts (Macgregor, 1951; Goodwin, 1981; Hamilton, 1998), which has been argued as evidence in favour of a non-uniformitarian style of tectonics. However, others have suggested that these are in fact more typical of mid-crustal structures around diapirs in modern arc terranes (e.g., Kusky et al., 2021).

Opponents to the subduction model have proposed that TTG magma generation may have occurred at the base of thickened mafic crust, such as oceanic plateaux (e.g., Smithies, 2000; Bédard, 2006; Van Kranendonk, 2010). Studies on Archean sediment geochemistry suggest that oceanic plateaux were abundant (Kamber, 2010), and would have attained thicknesses between 40–45 km (Vlaar et al., 1994; Korenaga, 2006). Such thicknesses would have stabilised garnet at the base (corresponding to 12–14 kbar), thereby enabling plateaux-derived sources to impart 'high-pressure' TTG signatures on melts. Furthermore, LILE-enriched basaltic sources (such as EAT, Condie, 1981) were most likely generated by melting of undepleted mantle peridotite, such as those transported from the lower mantle to the surface via plumes

(Guitreau et al., 2012; Martin et al., 2014). Opponents to the plateaux model argue that it is unlikely that the bases of oceanic plateaux are sufficiently hydrated to generate voluminous TTG terranes (Arndt, 2013). However, syn-eruptive hydration during sub-aqueous eruption, followed by repeated burial of lava flows during vertical accretion likely acts as a mechanism to transport hydrated basalts to depths sufficient for partial melting (Kamber, 2015). Alternatively, komatiites may have played an important role in transporting fluid to the lower crust (Hartnady et al., 2022; Tamblyn et al., 2023). This is because the high X_{Mg} in komatiites makes them more susceptible to hydration in low-temperature conditions, generating high-Mg chlorite which then acts as a water carrier during burial until dehydration at 680–800 °C. Given komatiites are a key component of several Archean terranes (e.g., Bickle et al., 1975), they provide a mechanism for water to be brought to suprasolidus conditions (Tamblyn et al., 2023). In either case, a surface water source for TTG melts is supported by recent Si-isotope studies (Murphy et al., 2024).

In reality, it is likely that a combination of these processes were active in the Archean (Bédard et al., 2013), supported by quantitative geodynamic models which support that both models may have been feasible under Archean conditions (e.g., Piccolo et al., 2019). It is therefore pertinent that each individual craton is scrutinised appropriately in order to fully understand its evolution.

In light of this and the newly collected data described below, I outline a model for the evolution of, and crustal formation in, the Zimbabwe craton by integrating the data collected described below (using methods described in Chapter 2) with previously published work.

5.2 Results

5.2.1 New geochronological data from the Zimbabwe craton

Geochronological data collected in this study comprises of zircon, apatite, and titanite U-Pb geochronology, and Rb-Sr geochronology of biotite and white mica samples, and is compiled in Appendix C. Sample locations and field context is provided in appendices A and B. A summary of new data is presented in Figure 5.2 and described in detail below.

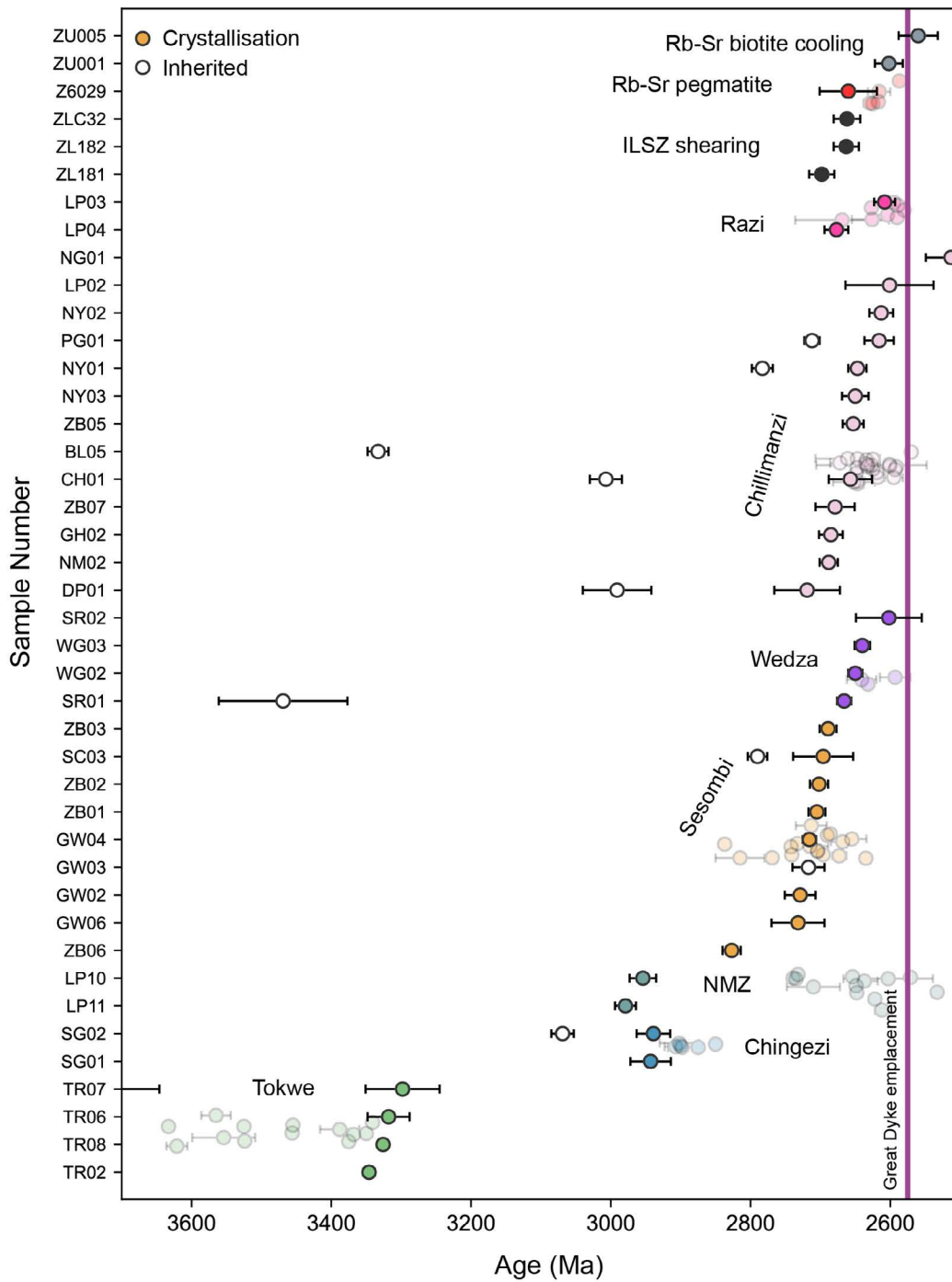


Figure 5.2: Summary diagram of newly collected geochronological data. Data from other studies (Appendix A) is shown as transparent markers, and clustered by the corresponding suites described here. Both crystallisation (full circles), and inherited (open circles) ages are shown.

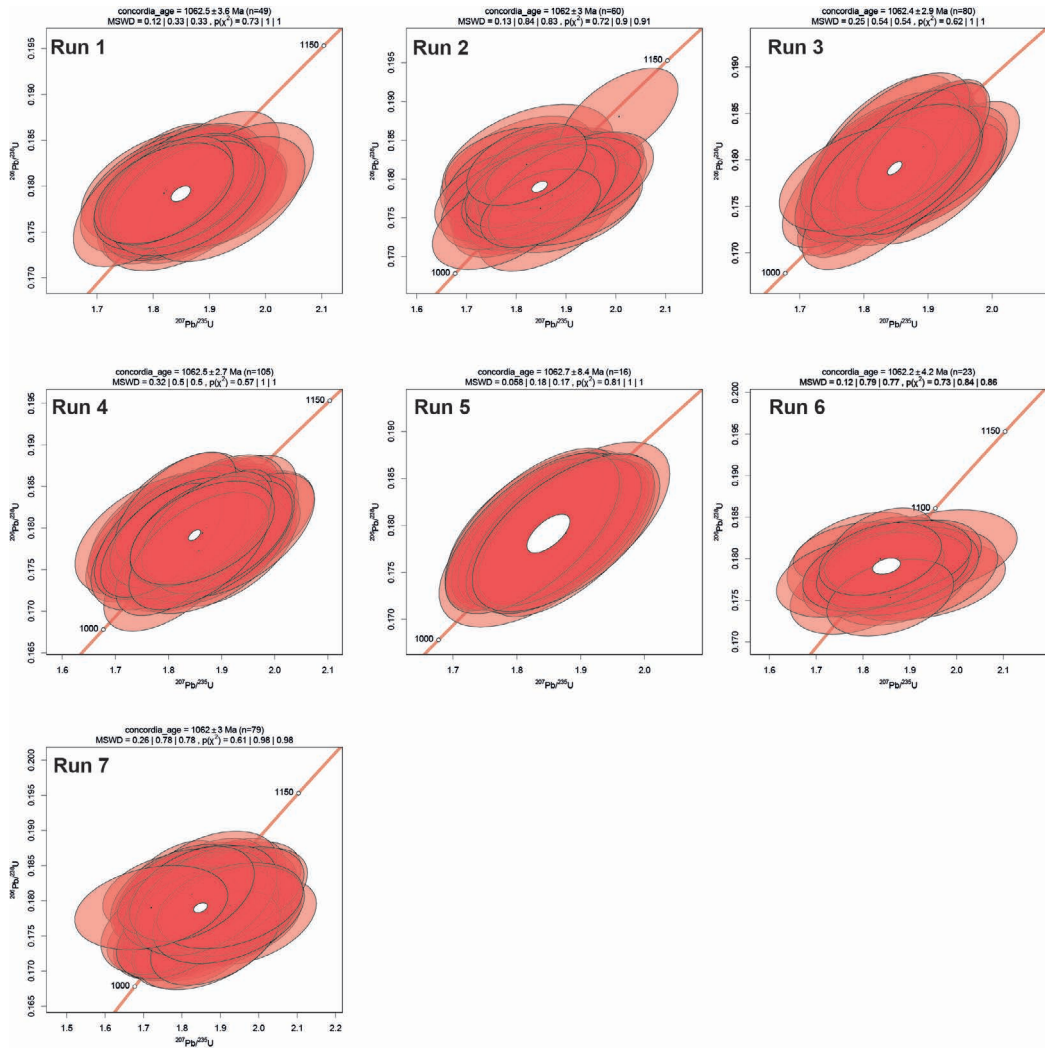


Figure 5.3: Wetherill concordia plots for primary reference standard 91500 grouped by analytical run.

5.2.1.1 U-Pb zircon standards

Primary standard 91500 yielded consistent results across each analytical session (Figure 5.3). All ablated spots were concordant and tightly clustered, with mean square weighted deviates (MSWD) of concordance + equivalence between 0.17 and 0.83. All runs were within error of the accepted age (1063.78 ± 0.65 Ma; (Horstwood et al., 2016; Wiedenbeck et al., 1995)).

Secondary standard Plešovice yielded more variable results, with a range in ages from 325.4–340.1 Ma (Figure 5.4). Runs 2–5 yielded several analyses that were discordant, which may be due to natural variability in the reference standard or ablation through inclusions/cracks in the grains. Excluding discordant analyses, MSWD

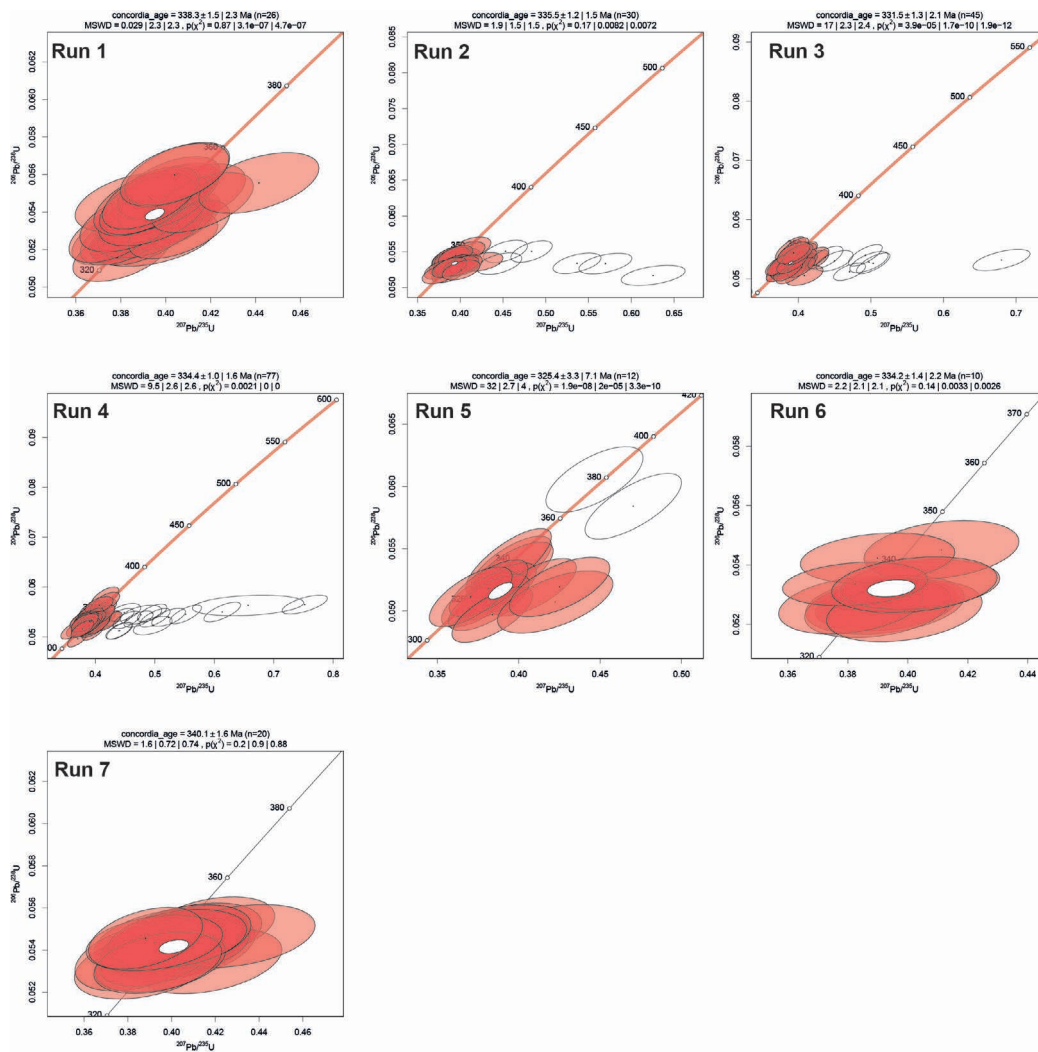


Figure 5.4: Wetherill concordia plots for secondary standard Plešovice grouped by analytical run

ranges from 0.74–4. The results are generally within error of the accepted age (337.13 ± 0.37 Ma; Sláma et al., 2008).

Analyses of Maniitsoq yielded internally consistent results ($\text{MSWD} = 0.28\text{--}1.2$) between each run, with reported ages $^{207}\text{Pb}/^{206}\text{Pb}$ ages ranging between 2990.5 ± 3.7 Ma to 3013 ± 18 Ma (Figure 5.5). Reported ages are all within error of the accepted age (3005 ± 2.6 Ma; Marsh et al., 2019) with the exception of runs 2 and 4 which returned marginally lower ages (2990.5 ± 3.7 Ma and 2992.4 ± 6.3 Ma, respectively).

Secondary standard OG1 was analysed during runs 1–3, and 6–7, and yielded consistent results (Figure 5.6). All $^{207}\text{Pb}/^{206}\text{Pb}$ ages are considered homogeneous ($p(\chi^2) > 0.05$), and are within error of the accepted age (3465.4 ± 0.6 Ma; Stern et al., 2009).

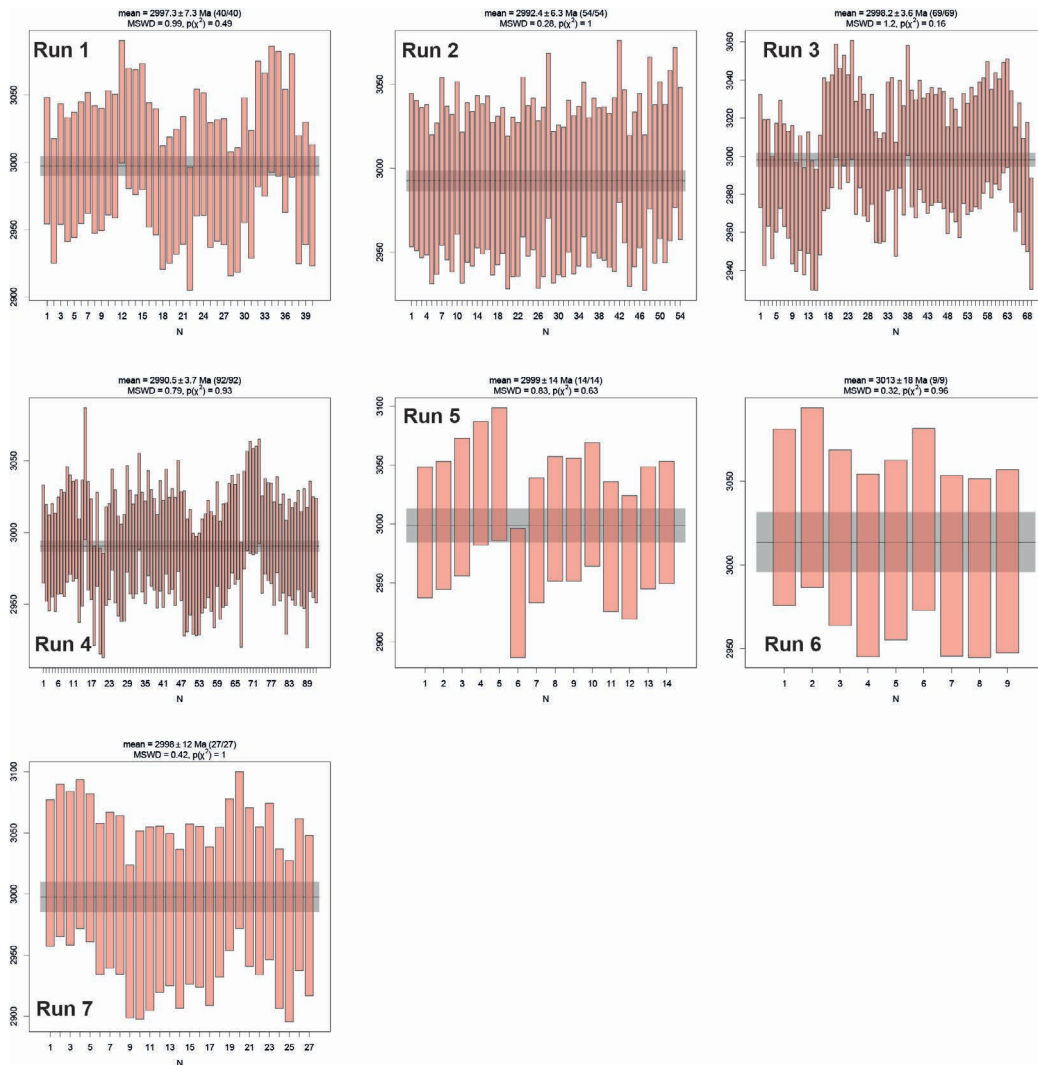


Figure 5.5: $^{207}\text{Pb}/^{206}\text{Pb}$ weighted mean plots for secondary reference material Mani-itsoq across analytical runs.

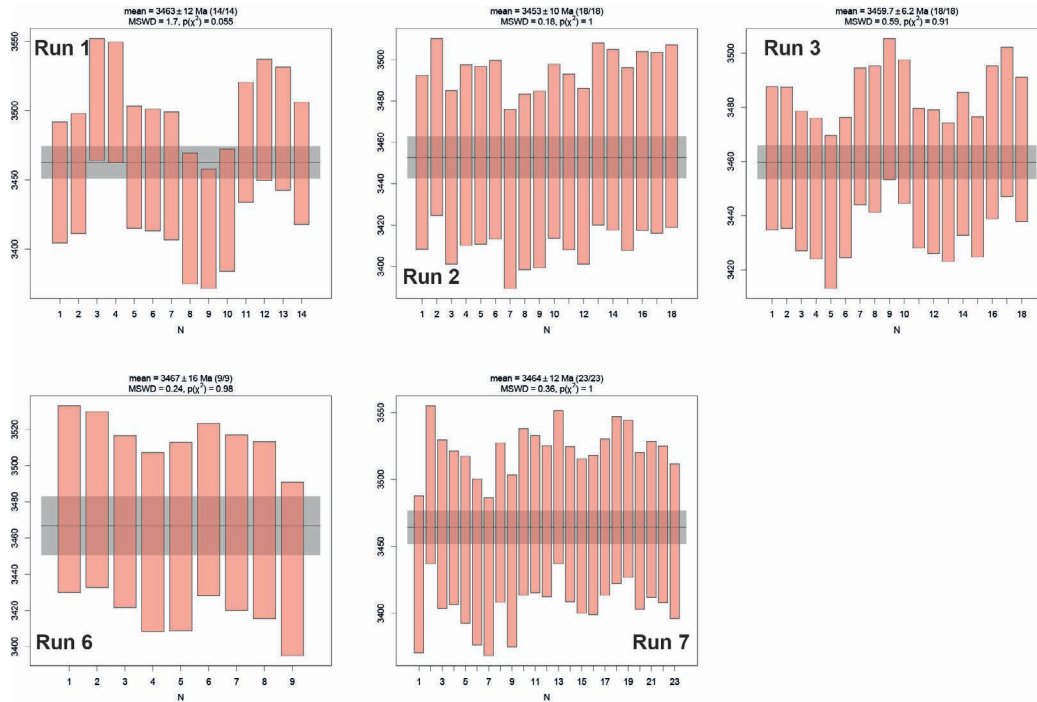


Figure 5.6: $^{207}\text{Pb}/^{206}\text{Pb}$ weighted mean plots for secondary standard OG1, grouped by analytical run.

5.2.1.2 Tokwe segment

TR02 is a complexly folded hornblende-rich tonalitic banded gneiss from the Tokwe river 3 km west of Bwanya. Melanosomes have a well-defined foliation defined by the alignment of hornblende laths, whilst leucosomes present only a weak foliation. Later folding, which post-dates leucosome formation, has caused reoriented some hornblende crystals into an axial-planar alignment. Thirty individual spots from 25 zircons yielded 11 concordant analyses for a concordia age of 3346.0 ± 8.6 Ma (MSWD = 1.2), which is interpreted as a crystallization age of the leucocratic material (Figure 5.7).

TR06 is a strongly banded tonalitic gneiss from the Tokwe river 3 km west of Bwanya. Banding is on the mm–cm scale, with darker bands consisting more abundant well-foliated biotite and chlorite. Thirty individual spots from 21 zircons yielded 18 analyses which define a discordia model age of 3302 ± 40 Ma (MSWD = 4.8) (Figure 5.8). One concordant zircon produced a concordia age of 3318 ± 30 Ma (MSWD = 0.32), which is collectively interpreted with the discordia model age as the age of crystallization.

TR07 is a garnet-bearing trondhjemitic sill which is intrusive to the banded tonalitic gneiss of *TR06*. Contact between the two units is fabric-parallel. The fab-

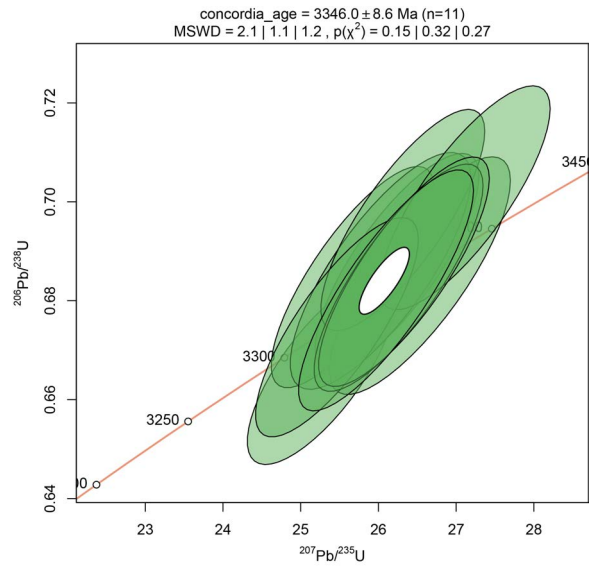


Figure 5.7: Wetherill plot of concordant analyses from sample TR02.

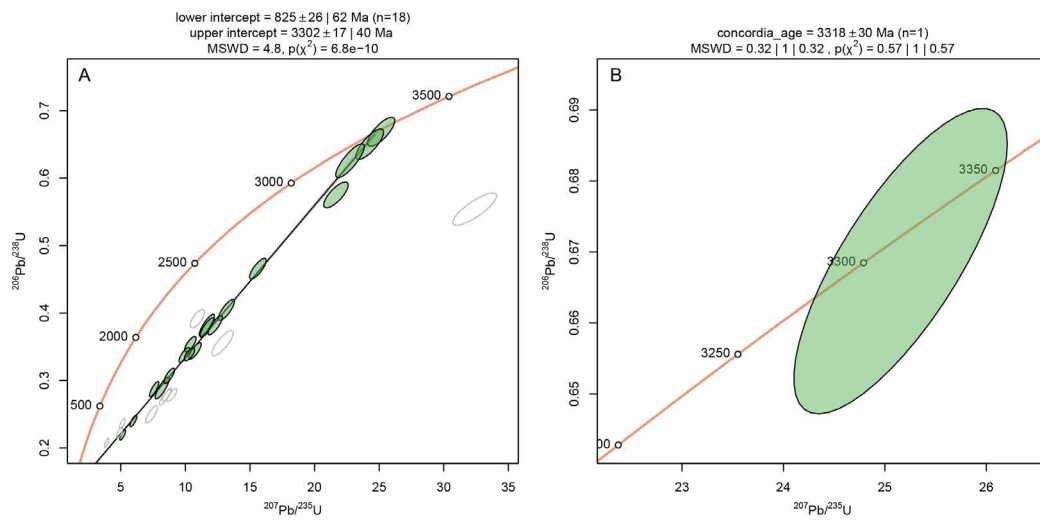


Figure 5.8: Wetherill concordia plots of sample TR06. A) Discordia model, B) zoom in of concordant analysis.

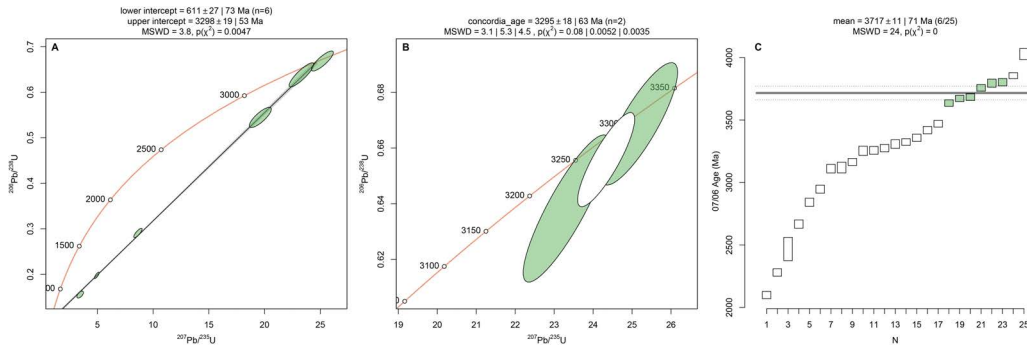


Figure 5.9: Wetherill concordia (A+B) and $^{207}\text{Pb}/^{206}\text{Pb}$ weighted mean (C) plots of sample TR07. A) Discordia model, B) zoom in on concordant analyses, C) inherited analyses cluster.

ric within *TR07* is defined by a strong alignment of muscovite, and garnet appears syn-kinematic to fabric formation. Forty individual spots from 33 zircons yielded 6 zircons which define a discordia age of 3298 ± 53 Ma (MSWD = 3.8, Figure 5.9). Two concordant zircon analyses return a concordia age of 3295 ± 63 Ma (MSWD = 4.5) which in conjunction with the discordia age is interpreted as the age of crystallization. Six analyses of zircon cores yield a $^{207}\text{Pb}/^{206}\text{Pb}$ weighted mean age of 3717 ± 71 Ma (MSWD = 24), which is interpreted as an inherited age.

TR08 is a mafic tonalite collected near *TR02*. Leucosomes within the sample have amalgamated into cm–dm scale fabric-parallel bands along sub-parallel (to the dominant fabric) leucocratic veins. The foliation is defined by strongly aligned hornblende crystals within the melanocratic layers. Thirty analyses of 24 zircon grains returned 16 concordant spots which define a concordia age of 3325.9 ± 6.3 Ma (MSWD = 1.3) and a $^{207}\text{Pb}/^{206}\text{Pb}$ weighted mean age of 3319.9 ± 6.8 Ma (MSWD = 0.42) (Figure 5.10). Collectively, these ages are interpreted as the age of crystallization of the leucocratic material.

5.2.1.3 Shabani gneiss

SG01 is a migmatitic gneiss collected from the Runde river 15 km NE of Zvishavane. A fabric is defined by biotite in the restitic assemblage, that occasionally wraps sericitised feldspar phenocrysts. Radiation haloes are commonly seen within biotite. Thirteen individual analyses from 26 zircon crystals define a discordia with an upper intercept age of 2954 ± 34 Ma (MSWD = 2.2) (Figure 5.11). One concordant analysis yielded a concordia age of 2943 ± 29 Ma (MSWD = 0.32). These ages are collectively interpreted as the age of crystallization of the migmatite. One concordant analysis of an interpreted xenocryst returned a concordia age of 3160 ± 34 Ma (MSWD = 1.3).

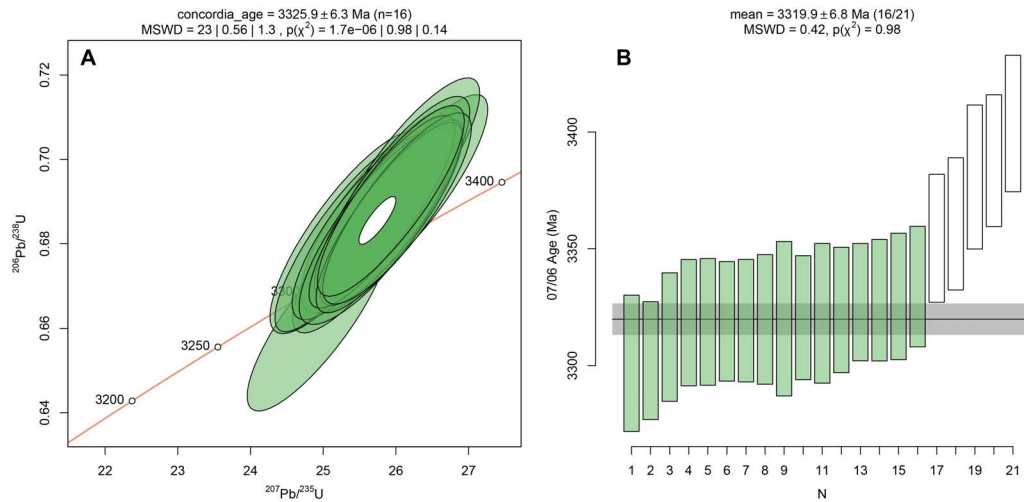


Figure 5.10: A) Wetherill plot of concordant analyses, B) $^{207}\text{Pb}/^{206}\text{Pb}$ weighted mean plot of sample TR08. Statistical outliers excluded in (B).

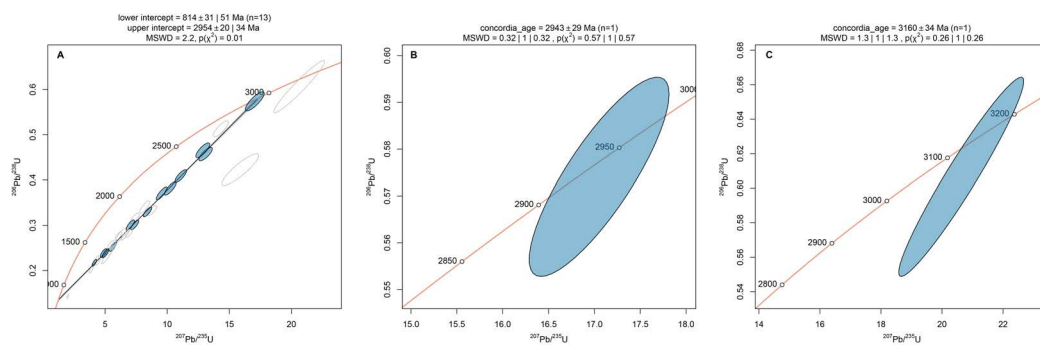


Figure 5.11: Wetherill plots of analyses from SG01. A) Discordia model with spots considered shaded, B) zoom in on concordant analysis of an interpreted crystallisation-age zircon, and C) concordant analysis of inherited grain.

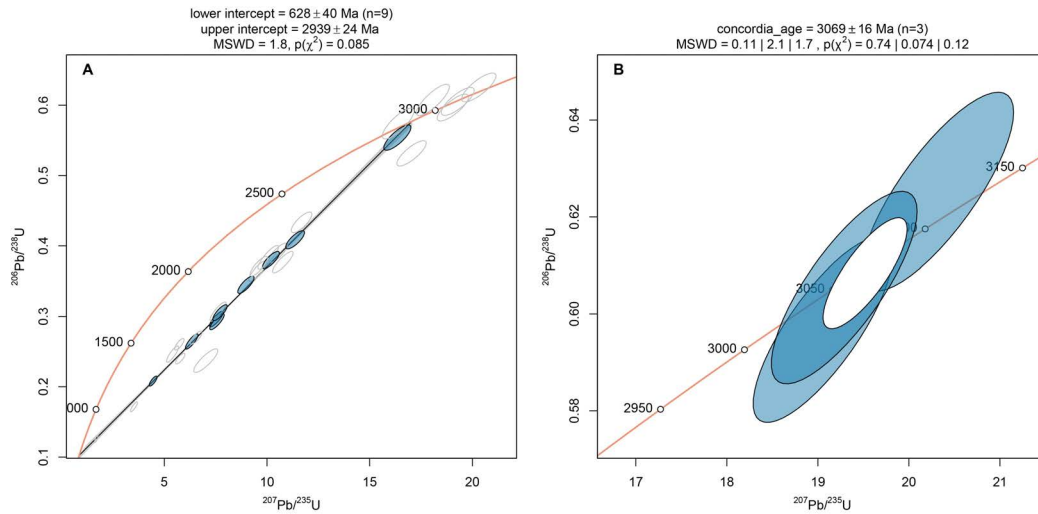


Figure 5.12: Wetherill plots of analyses from SG02. A) Discordia model with considered spots shaded, B) zoom in on concordant analyses of xenocrystic zircon analyses.

SG02 is a foliated tonalitic dyke intrusive to *SG01*. An internal fabric is defined by the alignment of biotite (\pm muscovite) and is parallel to the external fabric in the host gneiss. Contacts between the units are sharp and subparallel to the deformation fabric. Nine individual analyses from 28 zircons returned comparable U/Th ratios (1.5–3.0) and yield an upper intercept age on a discordia model of 2939 ± 24 Ma (MSWD = 1.8), and is interpreted as the crystallization age (Figure 5.12). Three separate analyses returned a concordia age of 3069 ± 16 Ma (MSWD = 1.7), which is interpreted as xenocrystic material inherited from the melting protolith.

5.2.1.4 Gwenoro dam and surroundings

GW02 is a sample of a migmatitic raft hosted within a foliated gneiss on the western edge of the Gwenoro dam spillway. A foliation is defined by the alignment of amphibole and biotite in the melanocratic horizons, which appear in equilibrium. Banding is on the cm scale. Ten analyses from six zircons returned three concordant analyses with a concordia age of 2729 ± 22 Ma (MSWD = 0.73), which is considered to be the crystallization age of the migmatite (Figure 5.13).

GW03 is a sample of a trondhjemitic vein that cross-cuts basement rocks of the Gwenoro dam spillway. The vein is primarily composed of quartz and plagioclase, with minor amounts of chlorite. Ten analyses of three zircons return four analyses that define a discordia model with an upper intercept age of 2717 ± 23 Ma (MSWD = 0.87) (Figure Y.8). All zircons are interpreted to be xenocrysts, and therefore the age is considered an inherited age.

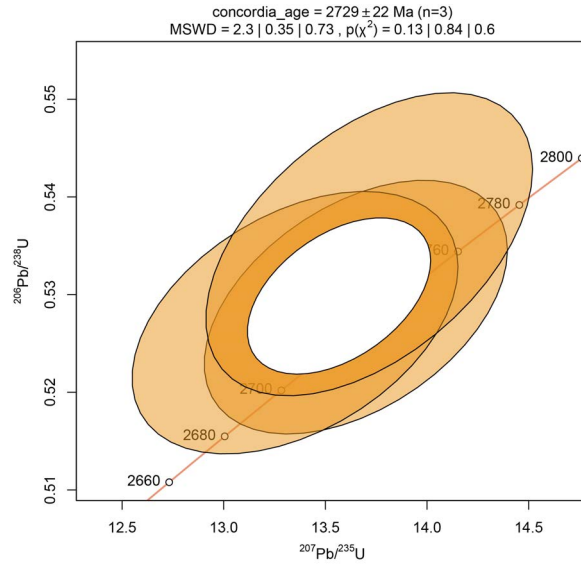


Figure 5.13: Wetherill concordia plot with concordant analyses of GW02 shown.

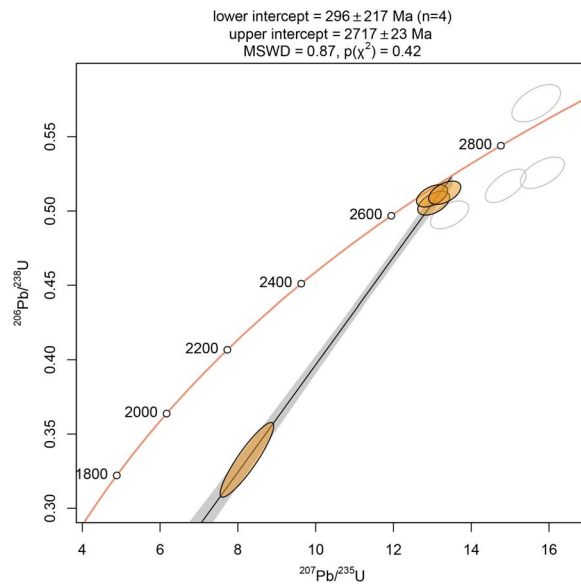


Figure 5.14: Wetherill concordia plot with analyses of GW03 shown. Analyses considered in the discordia model are coloured.

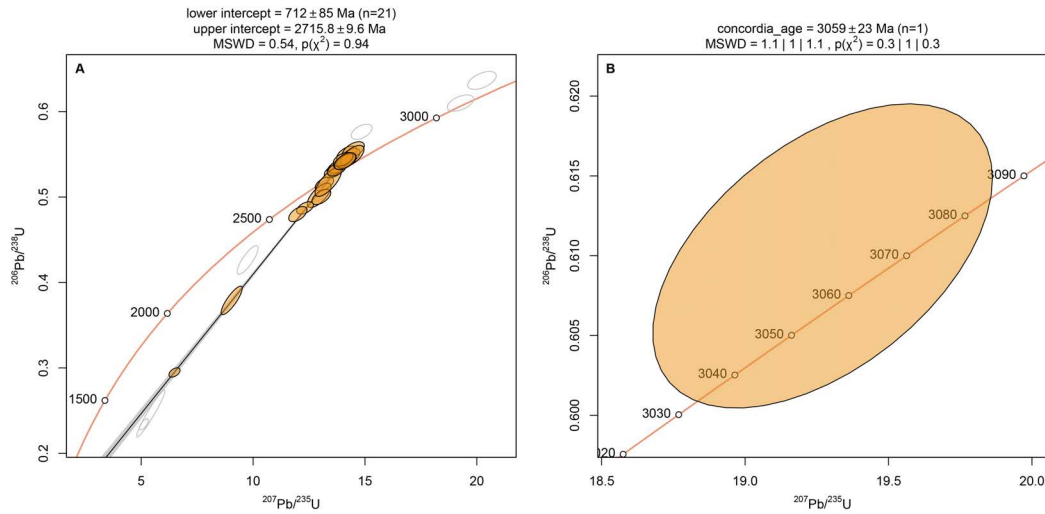


Figure 5.15: Wetherill concordia plots of analyses from GW04. A) Discordia model, with considered analyses shaded, B) zoom in on concordant analyses interpreted to represent the age of the protolith.

GW04 is a sample of a leucosome rich migmatitic xenolith in the Gwenoro dam spillway. A fabric is primarily defined by biotite, with subordinate amphibole and titanite. Thirty analyses of 18 zircons returned 21 analyses that define a discordia model with an upper intercept age of 2715.8 ± 9.6 Ma (MSWD = 0.54), which is interpreted as the crystallisation age of the migmatite (Figure 5.15). One analysis provides a concordant age of 3059 ± 23 Ma (MSWD = 1.1), which is interpreted to represent a xenocryst from the melt source.

NM02 is a granodiorite originally mapped as part of the basement gneisses by Stowe (1968). A weak fabric is defined by the alignment of biotite crystals. The rock has phenocrysts of stubby alkali feldspars, which in many cases preserve strong oscillatory growth zoning. Twenty analyses of 11 zircons yield 16 analyses that define a discordia model with an upper intercept age of 2688 ± 13 Ma (MSWD = 1.3), which is interpreted as the crystallization age of the unit (Figure 5.16).

SC03 is a sample of intensely deformed migmatitic gneiss collected from the Mapongkwe river c. 8 km west of where the river transects the Great Dyke. The fabric is defined by the strong alignment of amphibole in the melanosomes parallel to cm-scale banding. Forty analyses of 35 zircons yield eight analyses that pass an Aitchison (1982) log-ratio discordance filter (Vermeesch, 2021) define a discordia model with an upper intercept age of 2696 ± 43 Ma (MSWD = 0.85), which is interpreted as the crystallization age of the migmatite (Figure 5.17). Three separate concordant analyses define a concordia age of 2790 ± 14 Ma (MSWD = 1.4), which

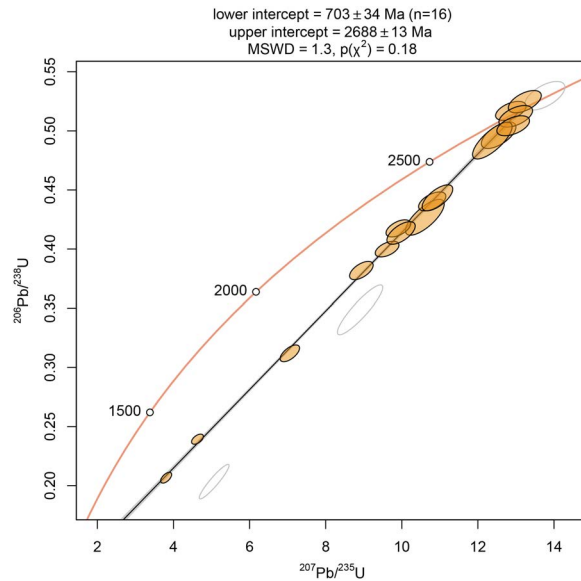


Figure 5.16: Wetherill concordia plot of analyses from NM02. Analyses considered for discordia model are shaded.

is interpreted as an inheritance age from xenocrysts entrained from the melt source.

5.2.1.5 Shangani batholith

ZB01 is a sample from a very weakly foliated tonalite c. 5 km from the eastern margin of the Fort Rixon–Shangani greenstone belt. A weak alignment of biotite laths defines the fabric. Thirty analyses from 22 zircons returned 19 analyses that define a discordia model with an upper intercept age of 2681 ± 11 Ma (MSWD = 1.1), and four concordant analyses which define a concordia age of 2705 ± 12 Ma (MSWD = 1.2, Figure 5.18). Collectively, these ages are interpreted to represent the crystallisation age of the tonalite.

ZB02 is sample of a moderately foliated tonalite collected c. 3 km from the eastern margin of the Fort Rixon–Shangani greenstone belt and along the outer flanks of the Irisvale–Lancaster shear zone (ILSZ). Thirty analyses from 25 zircons returned 13 analyses that define a statistically significant discordia model with an upper intercept age of 2702 ± 13 Ma (MSWD = 1.2), which is interpreted as the crystallisation age of the tonalite (Figure 5.19).

ZB03 is a migmatitic gneiss near the eastern margin of the Fort Rixon – Shangani greenstone belt, about five kilometres east of the Zinjanja monument along the Mt-shingwe river. A fabric is defined by the alignment of hornblende laths in the restitic assemblage, and melt selvages that have been deformed and chloritized by later shearing along the ILSZ. Sixteen of 30 analyses from 18 zircons yielded a discordant

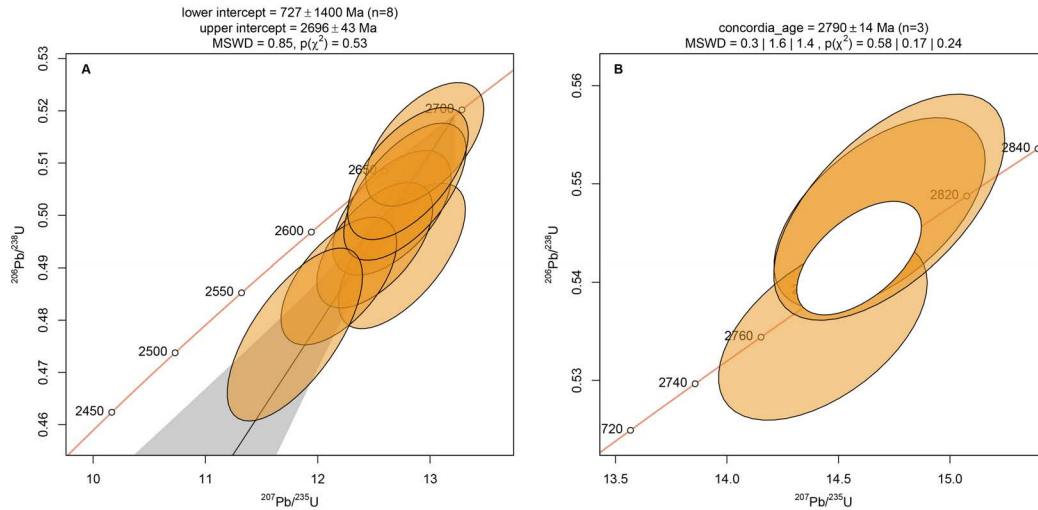


Figure 5.17: Wetherill concordia plots with analyses of SC03. A) Discordia model of analyses interpreted to represent the age of migmatite crystallisation, B) zoom in on concordant analyses of concordant xenocryst analyses.

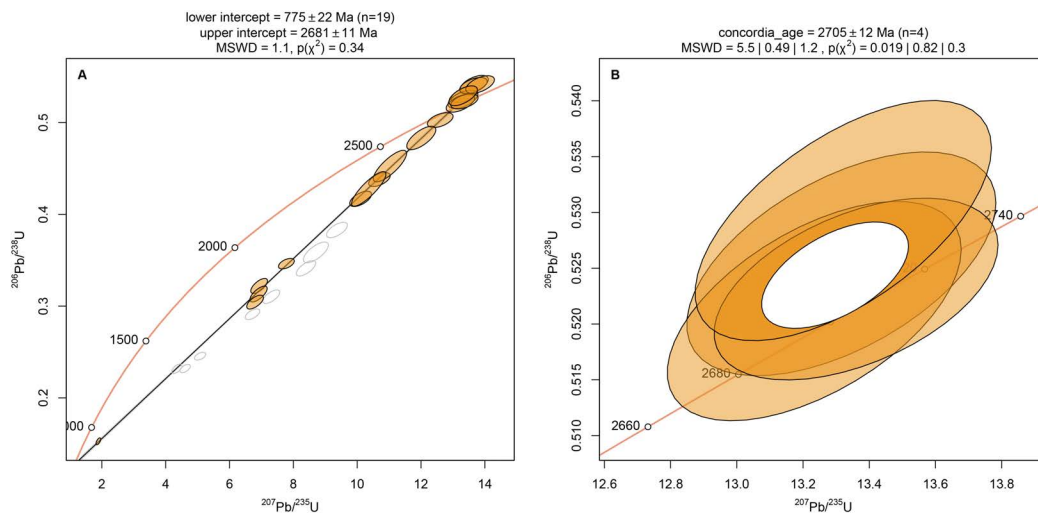


Figure 5.18: Wetherill concordia plots of analyses from sample ZB01. A) Discordia model, with considered analyses coloured, B) zoom in on concordant analyses.

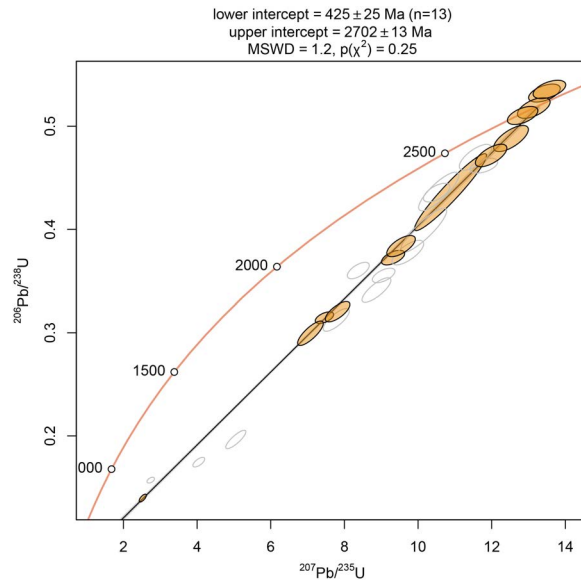


Figure 5.19: Wetherill concordia plot with the discordia model through analyses of ZB02. Analyses considered in the discordia model shaded.

model upper intercept age of 2689 ± 12 Ma (MSWD = 1.5) (Figure 5.20), and three concordant analyses produced a concordia age of 2683 ± 16 Ma (MSWD = 1.4). Both ages are interpreted as the age of crystallization.

ZB06 is a strongly sheared tonalitic gneiss situated 200 m from the eastern margin of the Fort Rixon–Shangani greenstone belt, approximately five kilometres south of the Zinjanja monument. The fabric is defined by a strong alignment of biotite sheets, locally preserving S/C fabrics considered to have formed during shearing along the ILSZ. Twenty-eight analyses from 23 zircons yield a discordia model with an upper intercept age of 2835 ± 14 Ma (MSWD = 1.3) (Figure 5.21). Eight analyses were concordant and return a concordia age of 2827 ± 13 Ma (MSWD = 1.3). The two results are collectively interpreted as the age of crystallization of the original tonalite intrusion.

5.2.1.6 Granitoids around the Odzi greenstone belt

SR01 is a foliated tonalite (part of the Nyashanu Gneissic Belt of Lauerdale, 1988) exposed along the Save river, and sampled c. 1 km NW of the contact between the Odzi greenstone belt and the granitoids. The tonalite has a foliation defined by the alignment of biotite and amphibole laths, and biotite is overgrown over earlier amphibole. Fifteen of 30 analyses from 19 analysed zircons generate a discordia model with an upper intercept age of 2666 ± 10 Ma (MSWD = 1.7) (Figure 5.22). Five

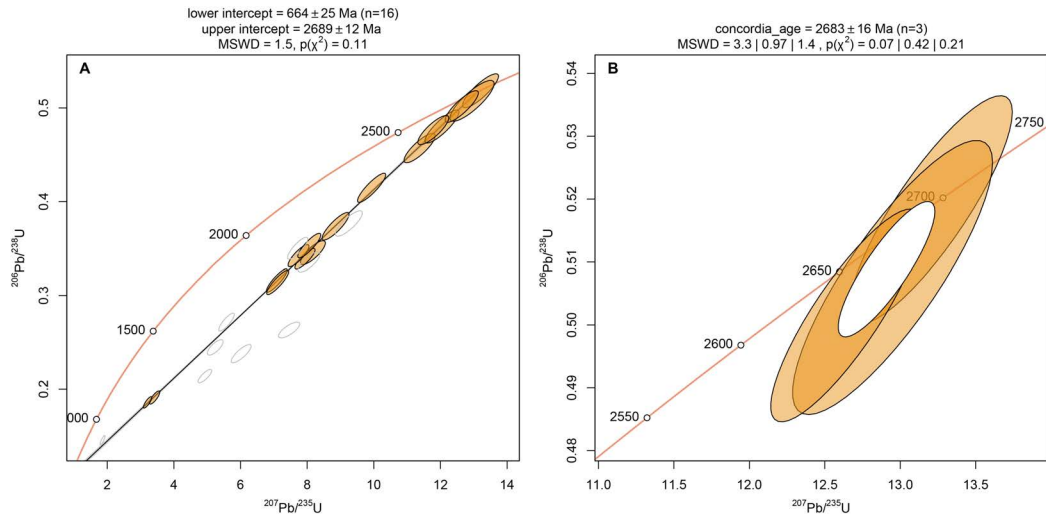


Figure 5.20: Wetherill plots of analyses from ZB03. A) Discordia model with considered analyses shaded, B) zoom in on concordant analyses.

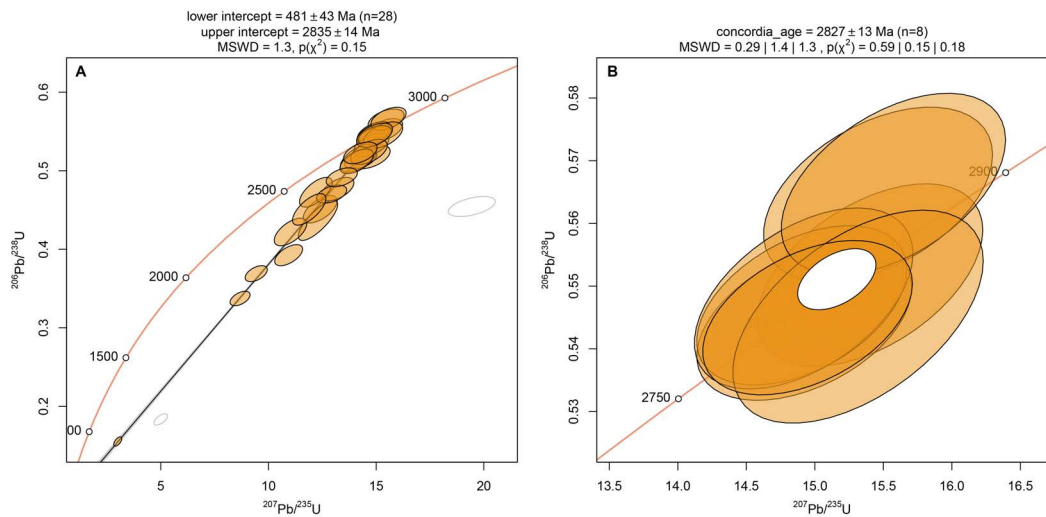


Figure 5.21: Wetherill plots of analyses from ZB06. A) Discordia model with considered analyses shaded, B) zoom in on concordant analyses.

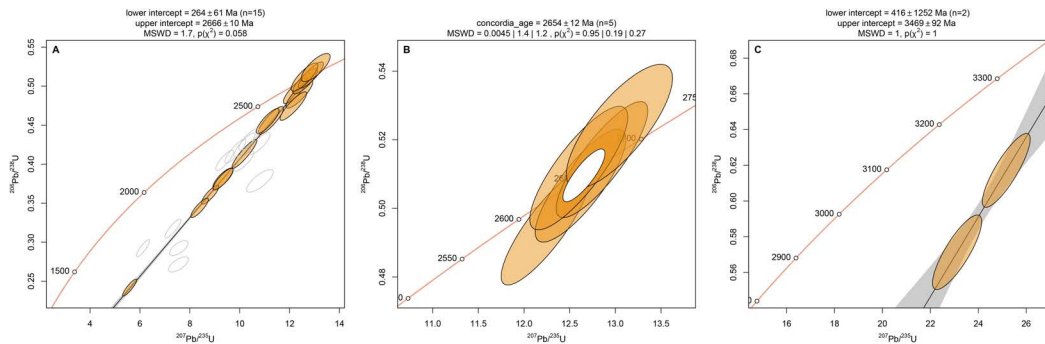


Figure 5.22: Wetherill concordia plots of analyses from SR01. A) Discordia model with considered analyses included, B) zoom in of concordant analyses assumed to represent crystallization, and C) discordia model of two analyses interpreted to represent an inheritance age.

analyses were concordant, and yield a concordia age of 2654 ± 12 Ma (MSWD = 1.2). These ages are collectively interpreted as the crystallization age. Two analyses returned an upper-intercept discordia age of 3469 ± 92 Ma (MSWD = 1), which is interpreted as an inherited age.

SR02 is a strongly foliated tonalite (part of the Nyashanu Gneissic Belt) exposed along the Save river, c. 100 m NW of the contact between the Odzi greenstone belt and the granitoids. The foliation is defined by the alignment of biotite (\pm titanite), which occasionally wraps around feldspar phenocrysts. Three of 30 analyses from 21 zircons were semi-concordant and produce a concordia age of 2602 ± 47 Ma (MSWD = 4.3), which is interpreted as the age of crystallization (Figure 5.23).

WG02 is a weakly foliated tonalite (part of the Romsley Gneiss suite of Lauerdale, 1988) exposed in the Murove river ~ 3 km NW of Nyashanu. A weak fabric is defined by the alignment of stubby amphibole and biotite crystals. Seventeen of 30 analyses from 20 zircons return a discordia model with an upper intercept age of 2650 ± 10 Ma (MSWD = 0.99) (Figure 5.24). Six analyses returned concordant results, defining a concordia age of 2664 ± 17 Ma (MSWD = 1.9). The two dates are collectively interpreted as the crystallisation age.

WG03 is a tonalite that belonging to the Romsley Gneiss suite. The unit has no clear fabric, and is composed of quartz, plagioclase, (retrogressed) biotite, chlorite, epidote \pm calcite, and is considered a retrogressed equivalent of *WG02*. Twenty of 30 analyses from 22 zircons generated a discordia model with an upper intercept at 2640 ± 11 Ma (MSWD = 1.3) (Figure 5.25). Three analyses were concordant, and return a concordia age of 2648 ± 17 Ma (MSWD = 0.47). The two dates are collectively interpreted as the age of crystallisation.

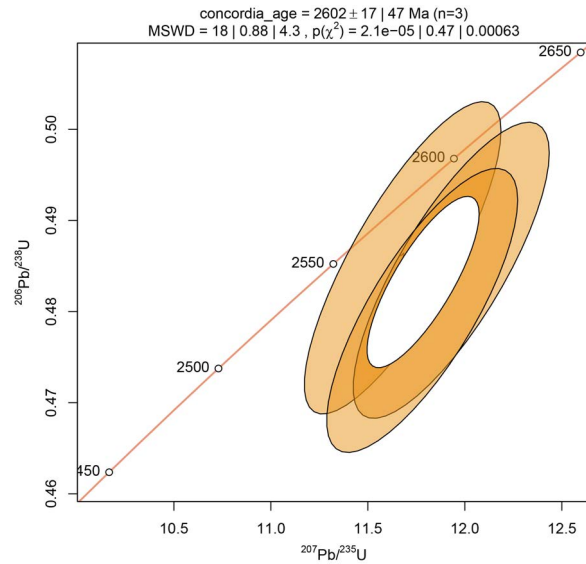


Figure 5.23: Discordia model of analyses from SR02 viewed on a wetherill concordia plot.

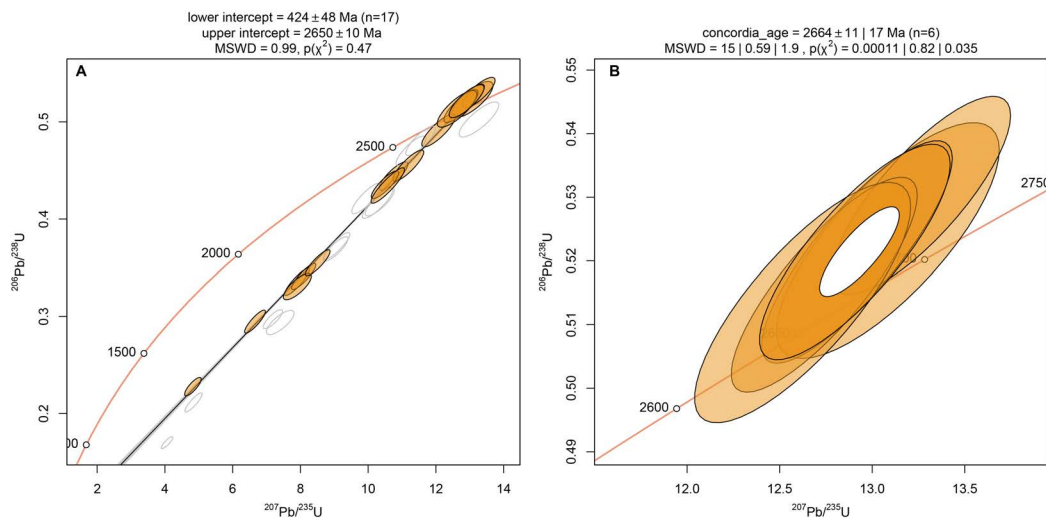


Figure 5.24: Wetherill concordia plots of analyses from WG02. A) Discordia model, with considered analyses shaded, B) Zoom in on concordant analyses.

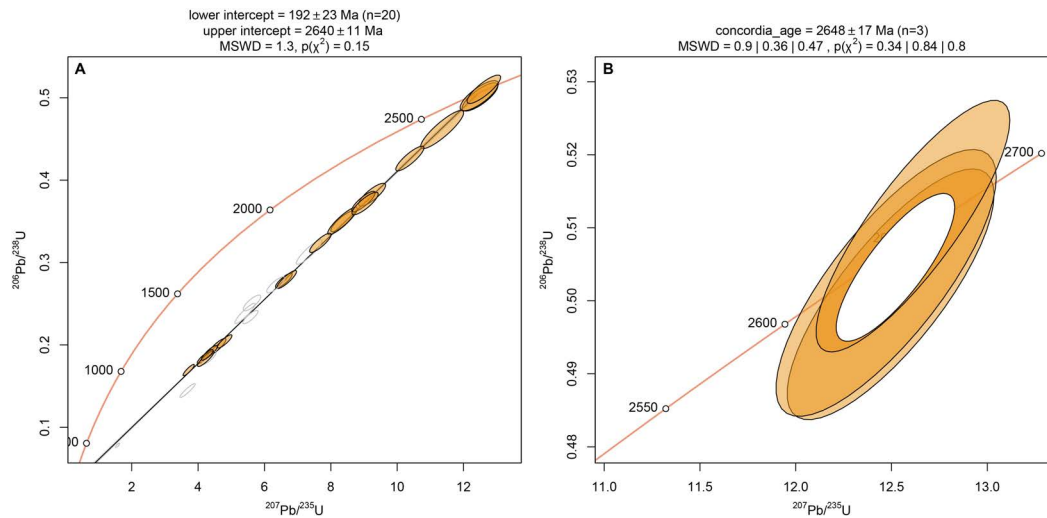


Figure 5.25: Wetherill concordia plots of analyses from sample WG03. A) Discordia model, with considered analyses shaded, and B) zoom in of concordant analyses.

5.2.1.7 Chillimanzi/Razi suite granites

GH02 was collected from a weakly foliated granodiorite that is part of the syntectonic suite of Stowe et al. (1968), collected c. 5 km west of Shurugwi. A weak fabric is defined by the alignment of biotite ± muscovite, and the sample contains more plagioclase than alkali feldspar. Three of 10 analyses from two zircons return a concordia age of 2688 ± 20 Ma (MSWD = 1.3) (Figure 5.26). Five analyses yield a $^{207}\text{Pb}/^{206}\text{Pb}$ weighted mean age of 2685 ± 17 Ma (MSWD = 0.99), and are collectively interpreted as the crystallisation age.

CH01 is a two-mica granite named the Chinoia Stock granite (Phaup, 1973). The sample was collected 7 km north of the Mandamabwe road junction. Thirty analyses from 29 zircons yielded scattered results (Figure 5.27). One concordant analysis returned a concordia age of 2657 ± 31 Ma (MSWD = 2), which is interpreted as the age of crystallisation. Seven analyses define a discordia model with an upper intercept of 3007 ± 23 Ma (MSWD = 0.52), which is interpreted as an inherited age.

BL05 is a two-mica granite sample collected 2 km north of the Mount Vugwe peak along the eastern boundary of the Belingwe greenstone belt. Thirty analyses from 22 zircons yielded 11 points which define a statistically significant discordia model with an upper intercept of 3333 ± 15 Ma (MSWD = 0.83), which is considered to be an inherited age (Figure 5.28).

PG01 is a porphyritic two-mica granite named the Wabai porphyritic granite (Phaup, 1973), and intrudes into the Shangani batholith c. 12 km east of the Fort

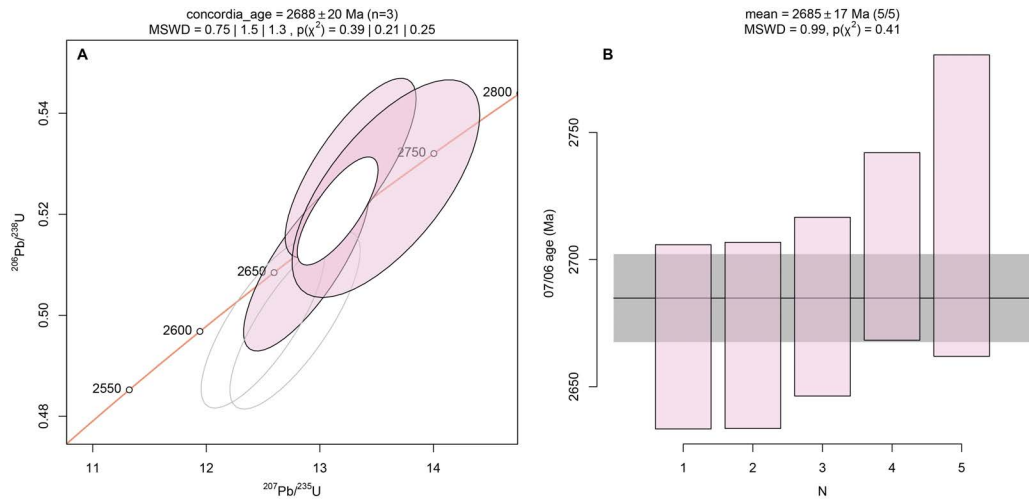


Figure 5.26: A) Wetherill concordia plot of sample GH02. Considered analyses are shaded. B) $^{207}\text{Pb}/^{206}\text{Pb}$ weighted mean plot of sample GH02.

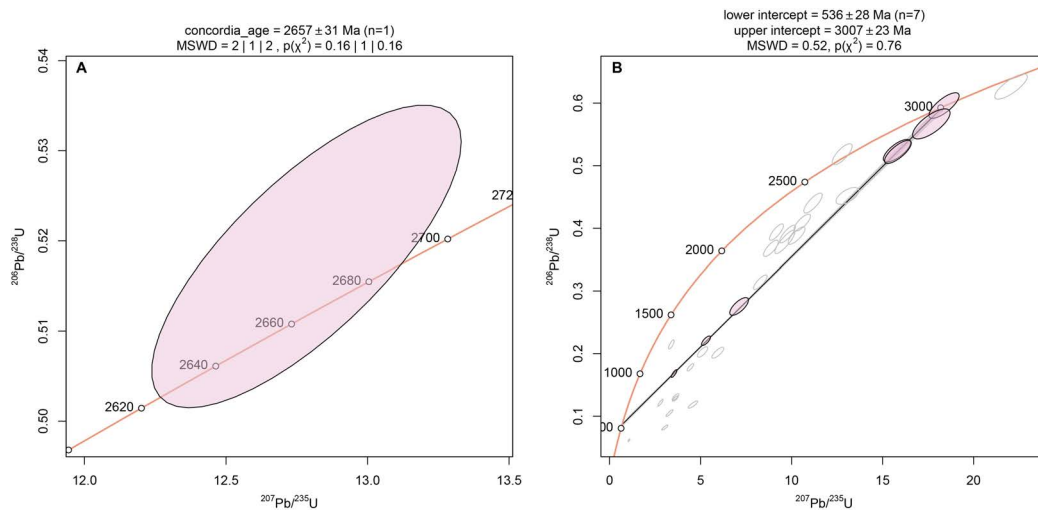


Figure 5.27: Wetherill concordia plots of CH01. A) Zoom in on concordant analysis interpreted to represent crystallisation. B) Discordia model of resolvable inherited population.

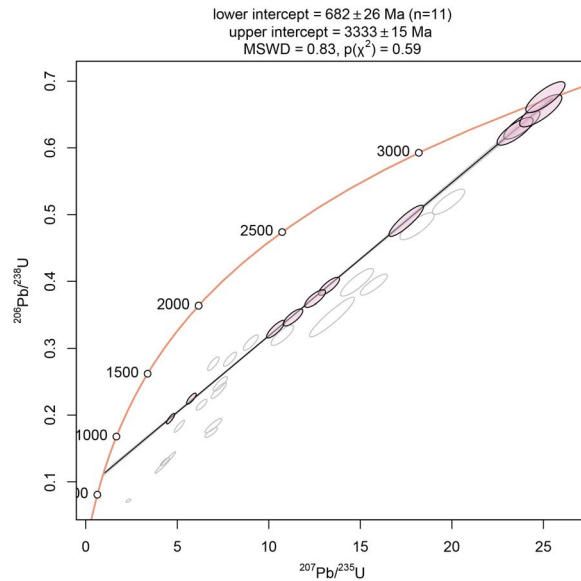


Figure 5.28: Wetherill concordia plot with a discordia model through analyses from BL05. Analyses considered in the model are shaded.

Rixon–Shangani greenstone belt. Forty-five analyses of 22 zircons yielded five analyses forming a statistically significant discordia model with an upper estimate of 2616 ± 21 Ma (MSWD = 0.96) (Figure 5.29). One concordant analysis with a concordia age of 2623 ± 24 Ma (MSWD = 0.48) is part of this cluster, and these ages are collectively interpreted as the crystallisation age of the granite. Seven concordant analyses yielded a concordia age of 2712 ± 11 Ma (MSWD = 7.1), which is interpreted as an inherited age.

ZB05 is a porphyritic two-mica granite named the Three Fingers porphyritic granite (Phaup, 1973), and is intrusive to the Shangani batholith c. 5 km east of the Fort Rixon–Shangani greenstone belt. Thirty analyses of 17 zircons yield 26 analyses that generate a statistically significant discordia model with an upper intercept age of 2653 ± 15 Ma (MSWD = 1.3), interpreted to be the age of crystallization of the granite (Figure 5.30).

ZB07 is a sheared porphyritic two-mica granite that intrudes into and along the eastern margin of the Fort Rixon–Shangani greenstone belt. The sample was collected 2 km south of the Zinjanja monument. Twenty analyses of 11 zircons yield 13 analyses with anomalously low U/Th (<0.02) which were discarded. Six analyses return a discordia model with a lower intercept age in Tera-Wasserburg space of 2693 ± 36 Ma (MSWD = 2.3), which is considered the crystallisation age of the intrusion (Figure 5.31). One concordant analysis returned a concordia age of 3947 ± 52 Ma (MSWD = 1.5) and is considered an inherited age.

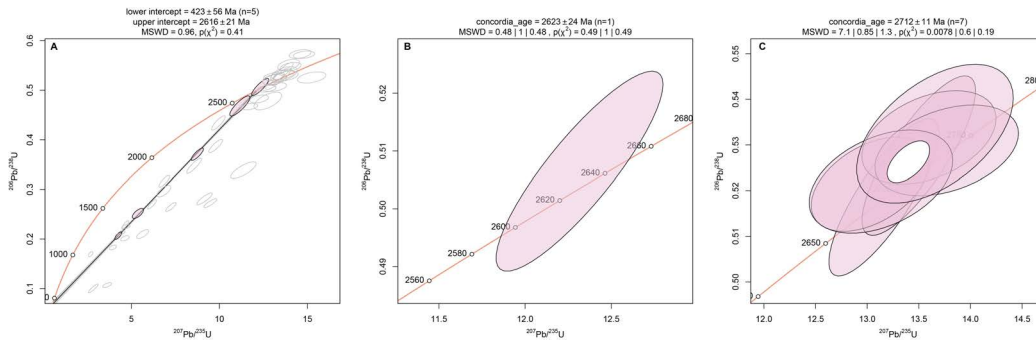


Figure 5.29: Wetherill concordia plots with data from sample PG01. A) Discordia model of interpreted crystallisation; B) zoom in on concordant analysis interpreted to be recording crystallisation; and C) zoom in on concordant analyses interpreted to represent an inheritance age.

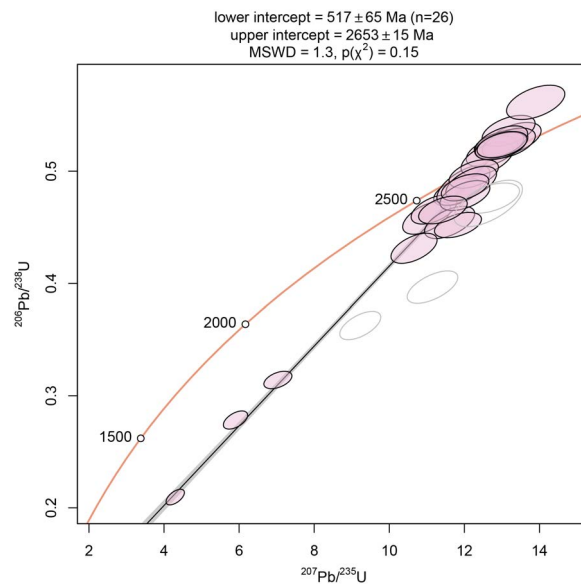


Figure 5.30: Wetherill concordia plot with discordia model through analyses of ZB05.

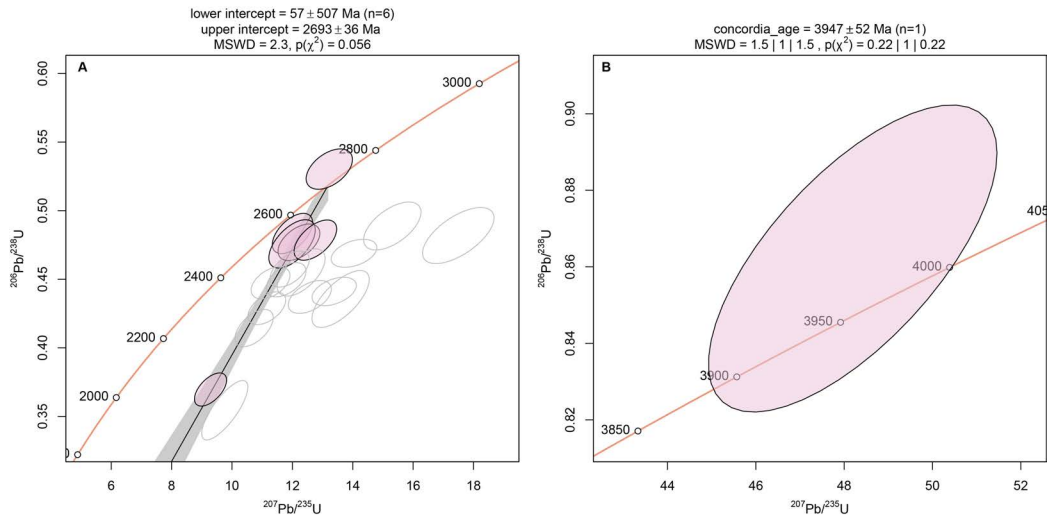


Figure 5.31: A) Wetherill concordia plot with discordia model for analyses of ZB07. Considered analyses shaded. B) Zoom in on an interpreted inheritance concordant analysis.

DP01 is sampled from the same intrusion as ZB07 and collected 5 km NE of the Zinjanja monument. Thirty analyses of 21 zircons yield 5 analyses that return a discordia model with an upper intercept age of 2719 ± 47 Ma (MSWD = 0.89), within error of the age for ZB07 (Figure 5.32). This population is interpreted to record crystallisation. Ten analyses yield a discordia model with an upper intercept age of 2991 ± 49 Ma (MSWD = 1.1), which is interpreted to represent an inherited age.

NY01 is a sample from a granite directly abutting the SW edge of the Odzi greenstone belt, collected c. 2 km NW of Nyashanu village. Thirty analyses of 23 zircons yielded four concordant analyses with a 2647 ± 13 Ma (MSWD = 1.9) concordia age, interpreted as the age of crystallisation of the granite body (Figure 5.33). Eight discordant analyses generate a discordia model with an upper intercept at 2783 ± 15 Ma (MSWD = 1.3), which is interpreted as an inherited age.

NY02 is a sample from a granite on the southern margin of the Odzi greenstone belt, collected c. 3 km south of Nyashanu village. Crystal sizes are smaller than the granite of NY01, and alkali feldspars are sericitized, whilst biotite is commonly chloritized. Thirty analyses from 25 zircons returned 10 analyses that define a discordia model with an upper intercept age of 2613 ± 17 Ma (MSWD = 1.6) (Figure 5.34). Five analyses were concordant, returning a concordia age of 2640 ± 30 Ma (MSWD = 3.3). Collectively, these ages are interpreted to represent the crystallisation age of the granite body.

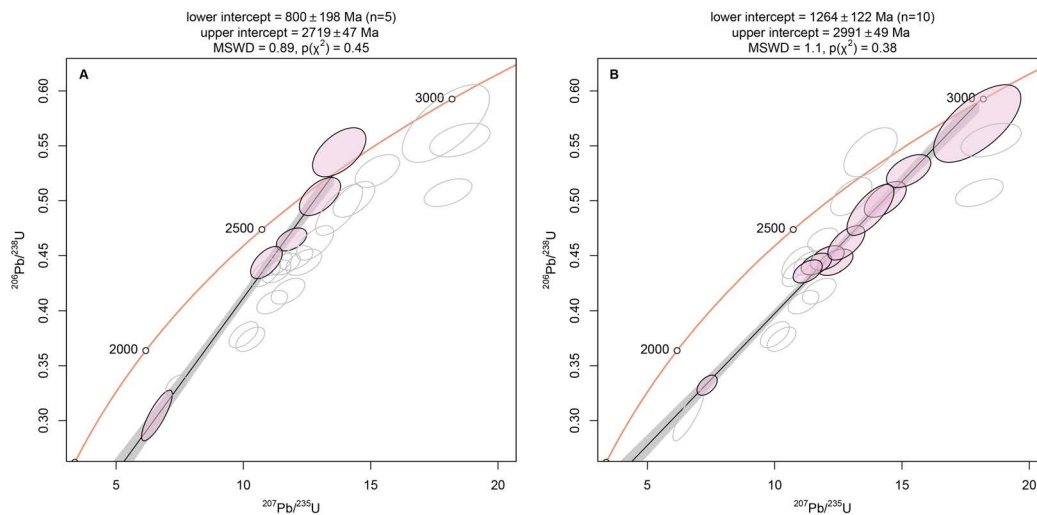


Figure 5.32: Wetherill concordia plots for analyses from DP01. A) Discordia model for analyses considered to represent crystallisation; B) discordia model for analyses considered to represent an inherited population.

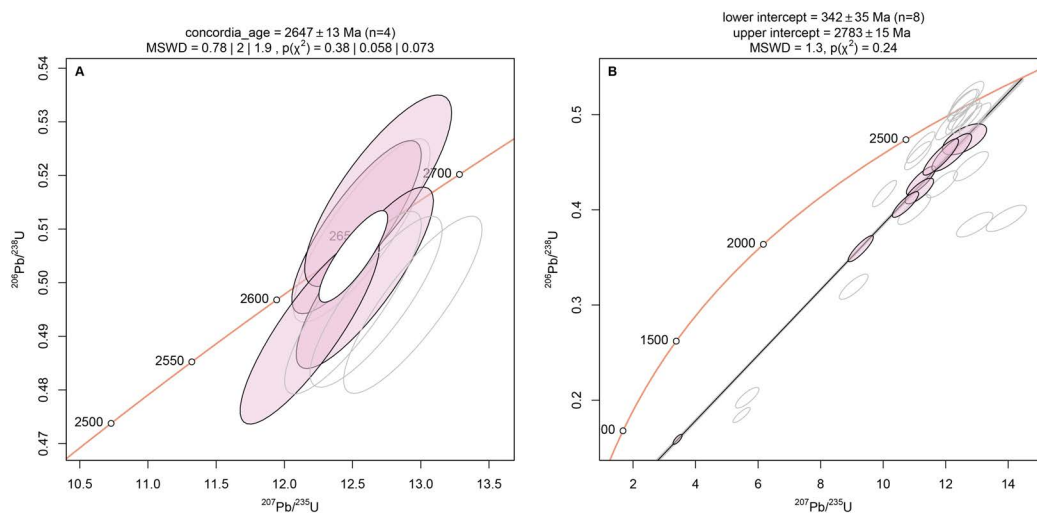


Figure 5.33: Wetherill concordia plots of analyses from sample NY01. A) Concordant analyses of interpreted crystallisation age; B) discordia model of an interpreted inherited population.

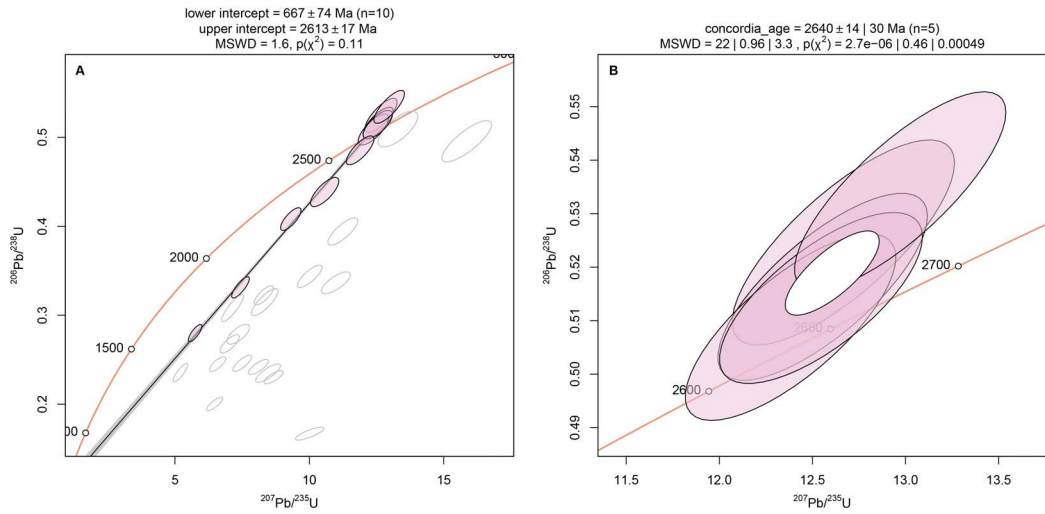


Figure 5.34: Wetherill concordia plots of analyses from sample NY02. A) Discordia model, considered analyses shaded; B) zoom in of concordant analyses.

NY03 is a sample from a granite body c. 10 km north of the Odzi greenstone belt. The body has a foliation defined by the alignment of biotite (\pm muscovite), and alkali feldspar commonly displays growth zoning, tartan twinning, and flame perthitic textures. Thirty analyses from 20 zircons returned seven analyses that define a statistically significant discordia model with an upper intercept age of 2650 ± 19 Ma (MSWD = 1.1), which is interpreted as the crystallization age of the granite body (Figure 5.35).

LP02 is a sample of equigranular granite that is part of the Great Zimbabwe pluton between the Masvingo greenstone belt and the Northern Limpopo thrust. The sample was collected c. 20 km south of Great Zimbabwe, and contains abundant alkali feldspar and biotite. Thirty analyses of 19 zircons yield a cluster of six analyses with comparable U/Th ratios (3.5–5.2) that define a discordia model with an upper intercept age of 2601 ± 63 Ma (MSWD = 2.6), which is interpreted as the age of crystallisation of the intrusion (Figure 5.36).

LP03 is a sample of porphyritic granite along the southern margin of the Great Zimbabwe pluton, 10 km north of the major Northern Limpopo thrust zone (NLTZ). A strong fabric is present in the sample, defined by the alignment of biotite, amphibole, and titanite crystals. Thirty analyses of 25 zircons returned 19 analyses that define a discordia model with an upper intercept age of 2608 ± 15 Ma (MSWD = 1.5) (Figure 5.37. Three analyses were concordant, with a concordia age of 2610 ± 17 Ma (MSWD = 1.3) (Figure 5.37B). Collectively, these are interpreted to represent the crystallisation age of the granite body.

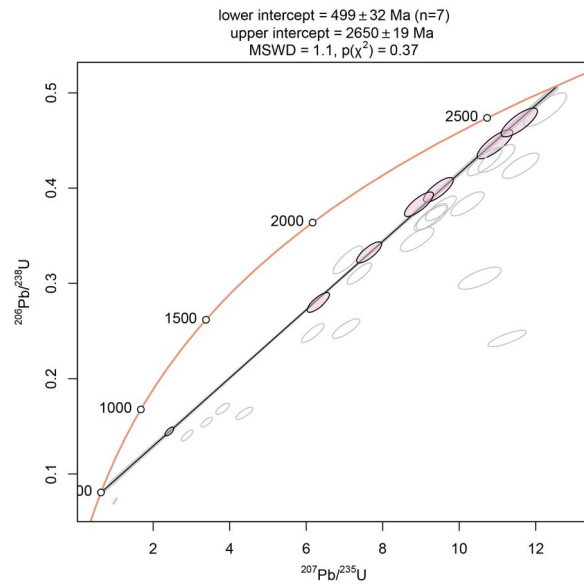


Figure 5.35: Wetherill concordia plot with the discordia model through analyses of NY03. Considered analyses are shaded.

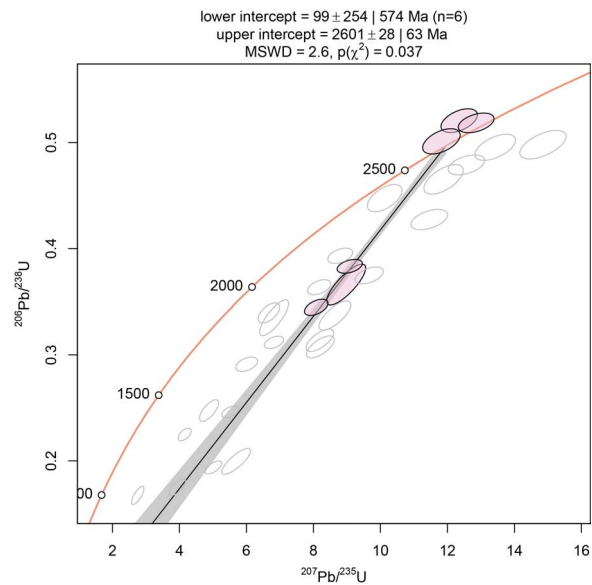


Figure 5.36: Wetherill concordia plot with the discordia model through analyses of LP02. Considered analyses are shaded.

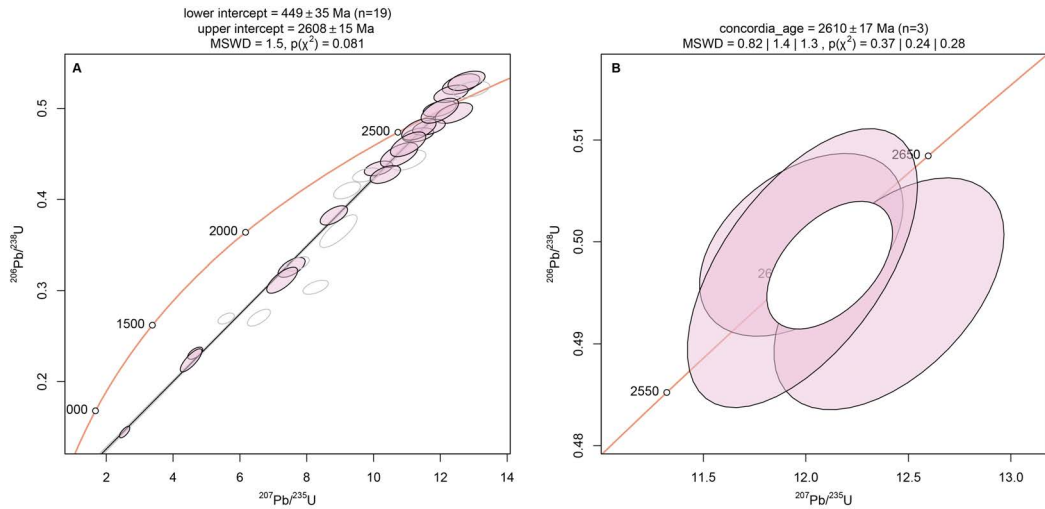


Figure 5.37: Wetherill concordia plots of analyses from LP03. A) Discordia model, considered analyses are shaded; B) zoom in on concordant analyses.

LP04 is an amphibole-bearing granite collected 200 m south of LP03, where the granitic intrusion contains abundant amphibolite xenoliths. The fabric within the xenolith is parallel to that of the host granite. Twenty analyses of 12 zircons returned a statistically significant discordia age through nine analyses, with an upper intercept age of 2677 ± 17 Ma (MSWD = 1.6), which is considered to be the crystallisation age of the intrusion (Figure 5.38).

5.2.1.8 Northern Marginal Zone

LP10 is a sample of tonalitic gneiss from the basement of the Northern Marginal Zone (NMZ) and appears to intrude through *LP11* (described below). *LP10* was collected 1 km north of where the Nyamawanga river joins the Mtilikwe river, c. 2 km east of Renco. The gneiss consists of garnet, biotite, quartz, plagioclase, and trace alkali feldspar, and the fabric is parallel to the NLTZ. Thirty analyses of 26 zircons returned 22 analyses that define a discordia model with an upper intercept age of 2954 ± 19 Ma (MSWD = 1.5), which is considered to be the crystallisation age of the tonalitic body (Figure 5.39).

LP11 is a sample of the felsic basement that *LP10* intrudes through. It is strongly deformed, with a fabric defined by biotite (wrapping around synkinematic garnet) parallel to the NLTZ. Thirty analyses of 20 zircons enabled a discordia model through 19 analyses that yields an upper intercept age of 2979 ± 15 Ma (MSWD = 1.5), which is considered the age of crystallisation of the intrusive protolith (Figure 5.40).

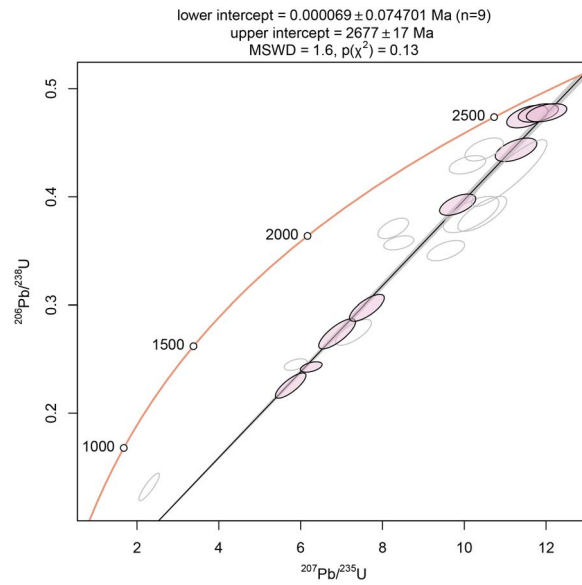


Figure 5.38: Wetherill concordia plot with the discordia model through considered analyses (shaded) of sample LP04.

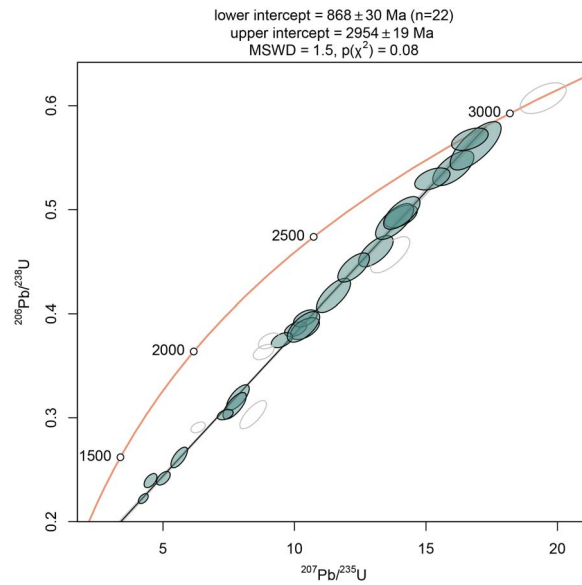


Figure 5.39: Wetherill concordia plot with analyses of LP10 shown. Considered analyses for plotted discordia model coloured.

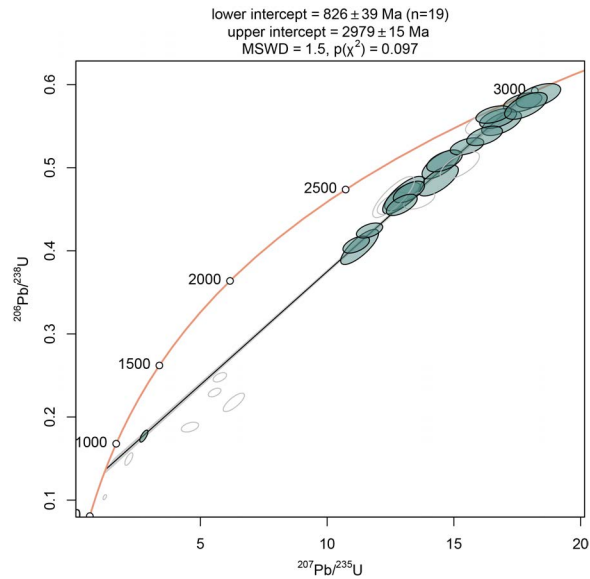


Figure 5.40: Wetherill concordia plot with analyses of LP11 shown. A discordia model is plotted, with considered analyses coloured.

5.2.1.9 Apatite geochronology

GW06 is a foliated tonalite from the Gwenoro dam spillway, which includes rafts of *GW02* and *GW04*. The fabric in *GW06* is defined by the alignment of biotite. Forty-six analyses of apatites yield a lower intercept age of 2732 ± 38 Ma (MSWD = 0.09), which is interpreted to represent the crystallisation age of the tonalite (Figure 5.41).

NG01 is a sample of the Nalatale granite (Phaup, 1973) which almost completely bisects the Fort Rixon–Shangani greenstone belt. The granite contains both biotite and muscovite, and also accessory topaz. Sixty-two analyses of apatites yield a lower intercept age of 2512 ± 36 Ma (MSWD = 0.98), which is interpreted as the age of crystallisation of the granite (Figure 5.42).

5.2.1.10 Titanite geochronology in the Irisvale-Lancaster shear zone

Titanite was analysed for U-Pb geochronology from three samples within the ILSZ. Sample *ZLC32* is a metamorphosed komatiitic basalt of the Reliance/Zeederbergs formation in the Fort Rixon–Shangani greenstone belt. The unit has a strong fabric parallel to the major trend of the ILSZ (Chapter 4). Titanite crystals form two textural populations: 1) mm-scale porphyroblasts within quartz knots, and 2) grains complexly intergrown with syn-kinematic tremolite. Seventy-two analyses of titanite grains returned a lower intercept age of 2662 ± 19 Ma (MSWD = 1.4) (Figure 5.43).

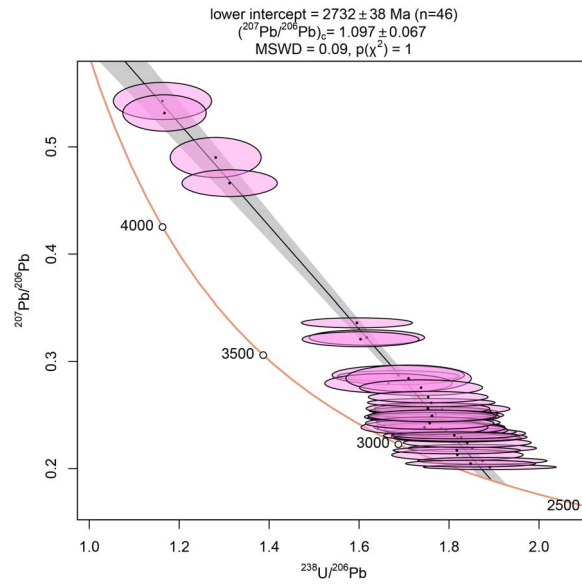


Figure 5.41: Tera-Wasserberg plot of apatite analyses from sample GW06.

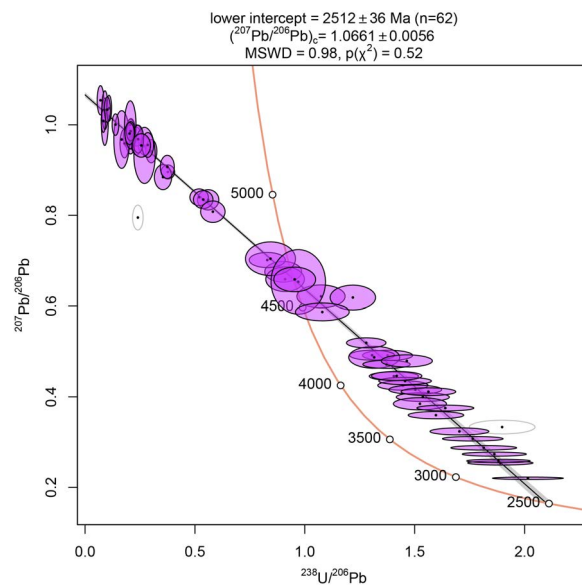


Figure 5.42: Tera-Wasserberg plot of apatite analyses from sample NG01.

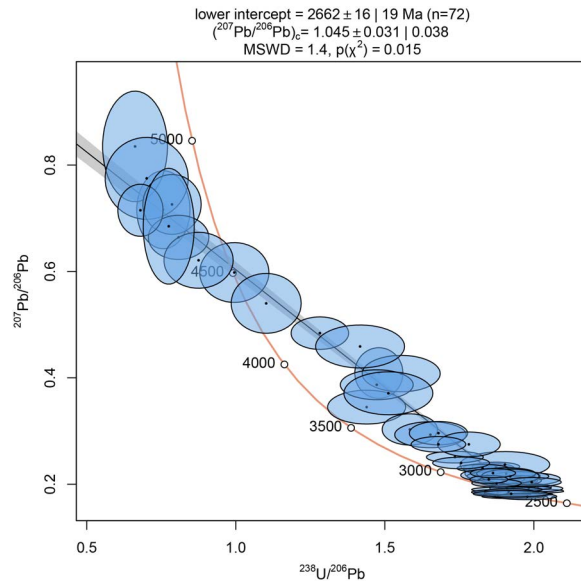


Figure 5.43: Tera-Wasserburg plot with analyses from titanite in sample ZLC32.

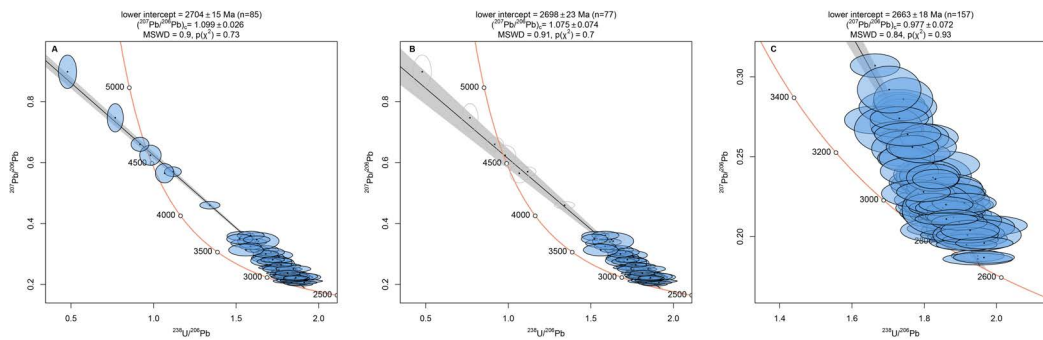


Figure 5.44: Tera-Wasserburg plots of titanite analyses from A) ZL181, B) ZL181, with the removal of extreme analyses, C) ZL182.

Samples *ZL181* and *ZL182* are Belingwean group metamorphosed tuffs that occur at the base of the Fort Rixon–Shangani greenstone belt. Titanites in both samples form porphyroblasts up to 1 cm wide within pressure shadows of ILSZ-related shearing. Within *ZL182*, titanite is also present as fine crystals parallel to the fabric defined by amphibole crystals. Considering all 85 analyses of titanites in *ZL181* (Figure 5.44A), produces a lower intercept age of 2704 ± 15 Ma (MSWD = 0.9). Removing the outliers, the lower intercept age becomes 2698 ± 23 Ma (MSWD = 0.91). Sample *ZL182* returns a lower intercept age of 2663 ± 18 Ma (MSWD = 0.84) (Figure 5.44C). Collectively, these ages are interpreted to bracket the age of deformation along the ILSZ.

5.2.1.11 Rb-Sr mica geochronology around the Zulu pegmatite field

White micas were analysed from four samples from within the Zulu pegmatite field, which is described in Chapter 4, and all samples are from Type 1 pegmatites. Sample *Z6007* contains micas from both magmatic and albitisation events (Table 4.2), and returns an isochron age of 2763 ± 27 Ma (Figure 5.45A). Sample *Z6029* contains micas exclusively derived from dynamic recrystallisation (Table 4.2, and returns an isochron age of 2660 ± 41 Ma (Figure 5.45B). Sample *Z6035* contains micas crystallised in the magmatic stage (Table 4.2), and returns an isochron age of 2580 ± 19 Ma (Figure 5.45C). Sample *ZU008* contains micas derived in both magmatic and dynamic recrystallisation events (Table 4.2), and returns an isochron age of 2738 ± 22 Ma (Figure 5.45D).

Samples *Z6007* and *ZU008* are considered to reflect mixing “pseudo-chrons” (Siebel, 1994; Nebel, 2014) between distinct magmatic and metasomatic isotopic reservoirs. Sample *Z6029* is interpreted to reflect the crystallisation age of the pegmatite during deformation (see Chapter 4). Sample *Z6035* is interpreted to reflect an isotopic resetting event overlapping with the age of the Great Dyke (2575.4 ± 0.7 Ma, Oberthür et al., 2002).

Biotite was analysed in samples *ZU001* and *ZU005* from the ILSZ, which are Belingwean group metamorphosed tuffs located at the eastern margin of the Fort Rixon–Shangani greenstone belt (south of the Nalatale granite). Biotite is interpreted to have formed through metasomatic alteration driven by fluids expelled from the pegmatite bodies (Chapter 4). Sample *ZU001* returns an isochron age of 2602 ± 20 Ma (Figure 5.46A), and *ZU005* returns an isochron age of 2560 ± 28 Ma (Figure 5.46B), which are interpreted to represent cooling ages post-metasomatism.

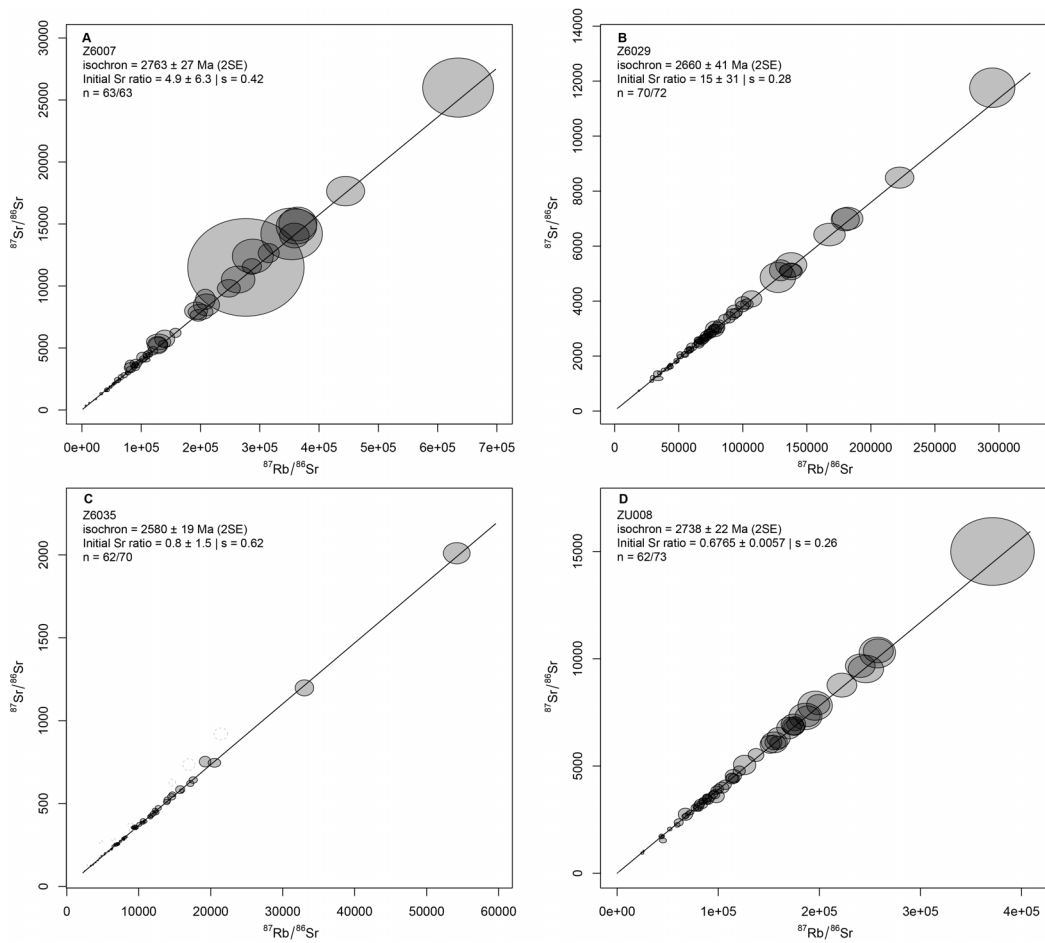


Figure 5.45: Rb–Sr isochron plots for A) Z6007, B) Z6029, C) Z6035, D) ZU008

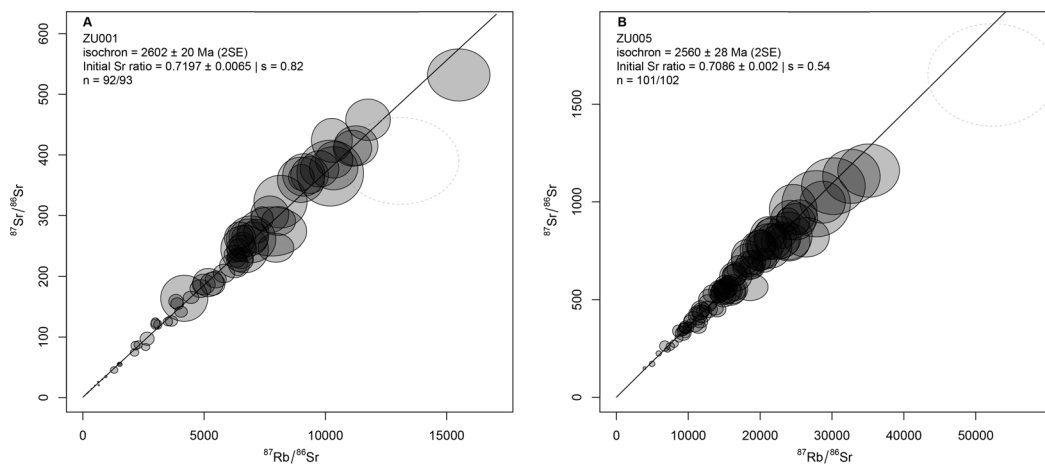


Figure 5.46: Rb-Sr isochrons of A) ZU001, and B) ZU005.

5.2.2 Geochemistry of the Zimbabwe craton

This study reports 76 new geochemical analyses of rocks (described in appendices A and B) from the Zimbabwe craton and the Northern Marginal Zone (Appendix D). Most analyses are of intrusive rocks and have been incorporated into the geochemical database described in section 2.1. In the following sections, granitoids are separated by major temporal suites determined by geochronology. The TTG suites are named in order of younging as follows; the Tokwe suite (3.7–3.2 Ga), the Chingezi suite (3.0–2.8 Ga), The Sesombi suite (2.8–2.7 Ga), the Wedza suite (2.7–2.6 Ga). All rocks of the NMZ are clustered, as the geochemical data are available in the literature lacks sufficient detail to segregate the basement gneisses from later intrusive charnockitic/enderbitic suites. The Late Archean granites are grouped into Chillimanzi and Razi suites, depending on whether they intrude into the ZC (Chillimanzi) or the NMZ and along the NLTZ (Razi).

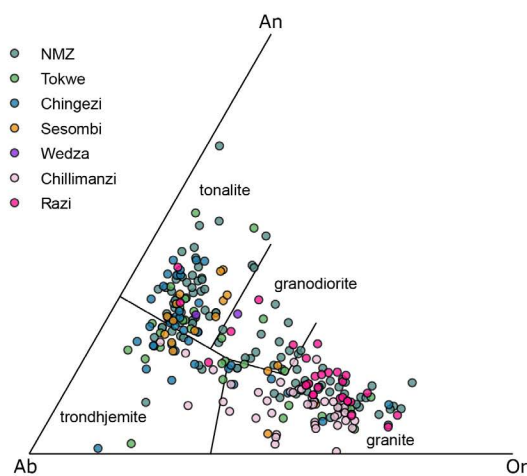


Figure 5.47: Normative Ab–An–Or diagram with granitoid discrimination lines of Barker et al (1979).

5.2.3 Intrusive rocks

Granitoids from the Zimbabwe Craton (ZC) and Northern Marginal Zone (NMZ) largely overlap in normative Ab–An–Or space, spanning all fields (Figure 5.47). Trondhjemitic rocks are rare in the NMZ, whereas granodioritic rocks are absent in the

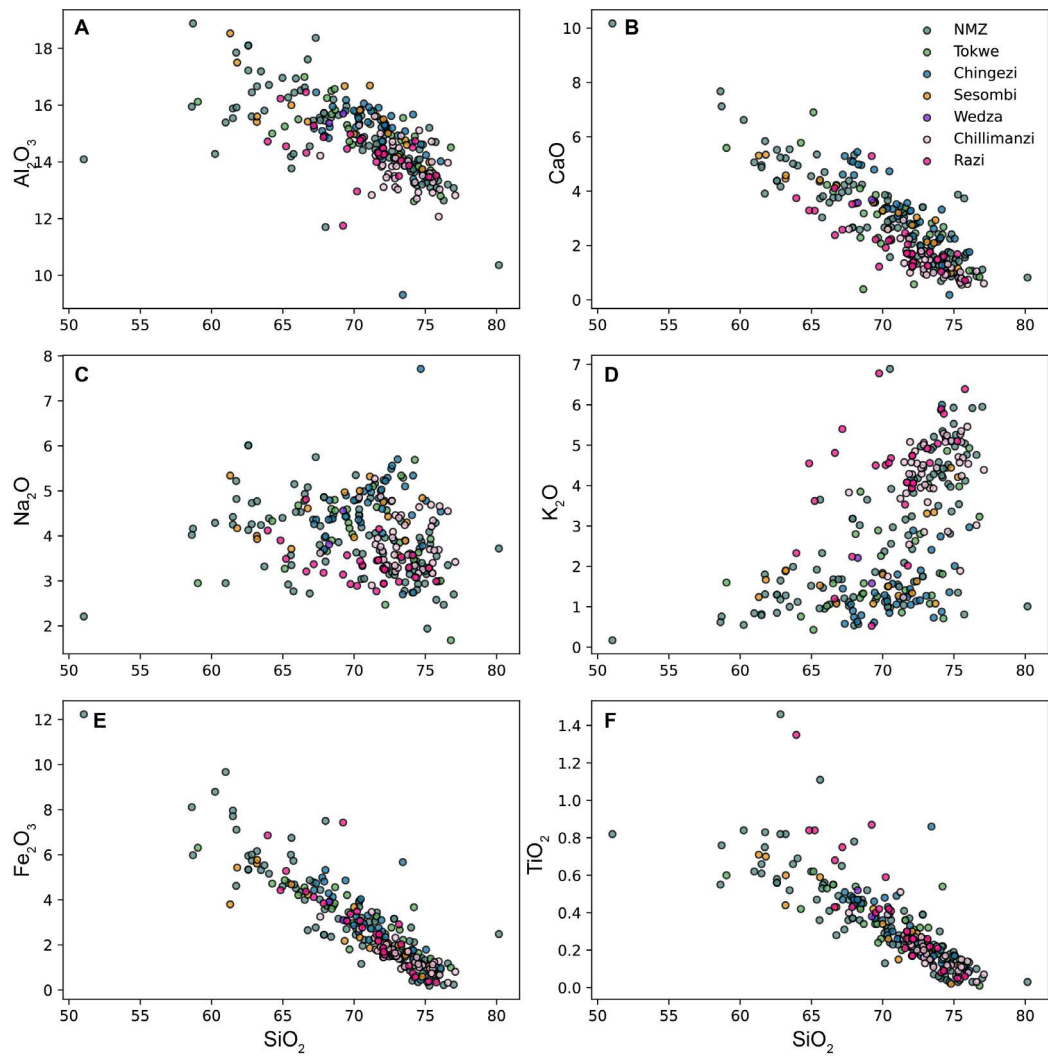


Figure 5.48: Harker diagrams for TTG and granite samples from the Zimbabwe Craton and the Northern Marginal Zone.

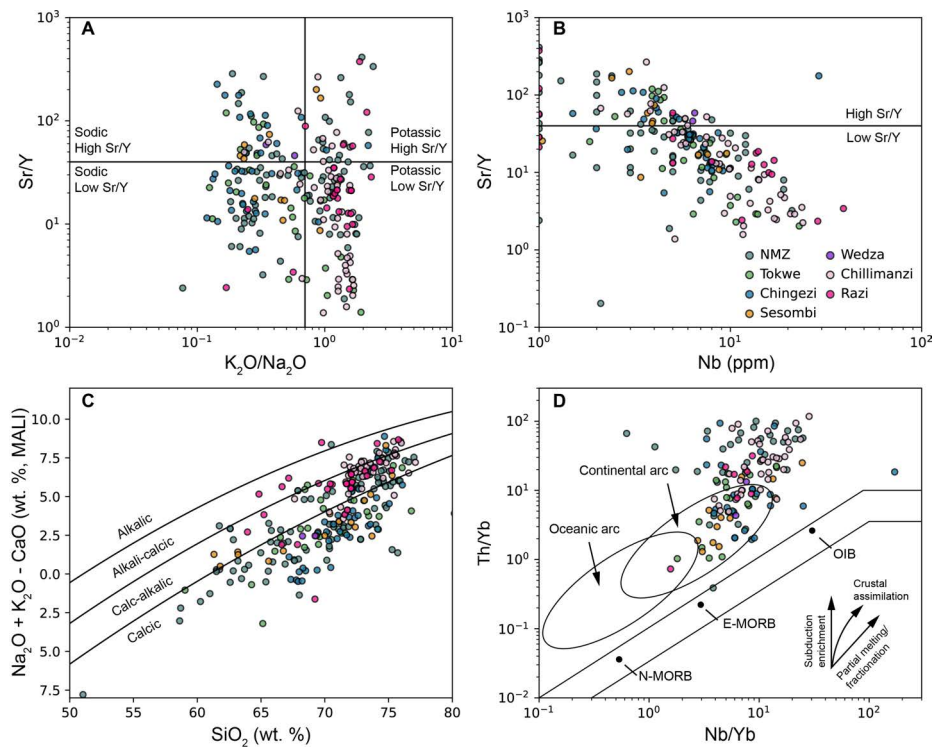


Figure 5.49: Geochemical discrimination diagrams for granitoids. A) Sr/Y vs. K_2O / Na_2O diagram with discriminating principles from Moyen and Martin (2012) and Smithies et al. (2018); B) Sr/Y vs. Nb for discrimination of adakites and TTGs against typical arc rocks after Drummond and Defant (1990) and Nymoen et al. (2025); C) Modified alkali-lime index (MALI) plot after Frost (2001); and D) Th/Yb vs Nb/Yb after Pearce (2008, with vectors for subduction enrichment, crustal assimilation, and partial melting/fractional crystallisation after Szilas et al. (2016).

Chingezi suite. Broadly, Razi suite granites exhibit more calcic (An-rich) affinities relative to the Chillimanzi suite.

Harker diagrams are plotted in Figure 5.48. Granitoids from the NMZ exhibit a wider range of SiO₂ content (c. 51–80 wt. %) relative to the ZC (c. 58–77 wt. %). All TTG and granite suites exhibit clear and parallel fractionation trends in Al₂O₃, Fe₂O₃, CaO, and TiO₂ (Figure 5.48A,B,E,F). Razi suite granites cover a wider potassic compositional space (c. 0.3–6.8 wt. %), whereas Chillimanzi suite granite are generally more constrained (K₂O between c. 1.8–5.5 wt. %), and are both predominantly within the potassic field (Figure 5.49A). An inflection point in K₂O at c. 70 wt. % SiO₂ is observed in both NMZ and ZC granitoids (Figure 5.48D), though not all samples follow this trend. Similarly, Na₂O exhibits a weak negative correlation to SiO₂, and Razi suite granites are consistently less sodic than Chillimanzi suite granites. All TTG suites in the NMZ and ZC show sodic and potassic affinities, and both high and low Sr/Y values are observed (Figure 5.49A). A negative trend is evident in log-log space for Sr/Y vs. Nb in all suites of the NMZ and ZC (Figure 5.49B). The majority of less silicic TTG analyses fall within the calcic trend of Frost (2001), and trend towards being calc-alkaline with increasing SiO₂ (Figure 5.49C). Chillimanzi and Razi suite granites primarily straddle the calc-alkalic trend (Figure 5.49C). In Th/Yb vs. Nb/Yb space, all suites follow partial melting/fractionation and/or crustal assimilation trends away from continental arc compositions (Figure 5.49D).

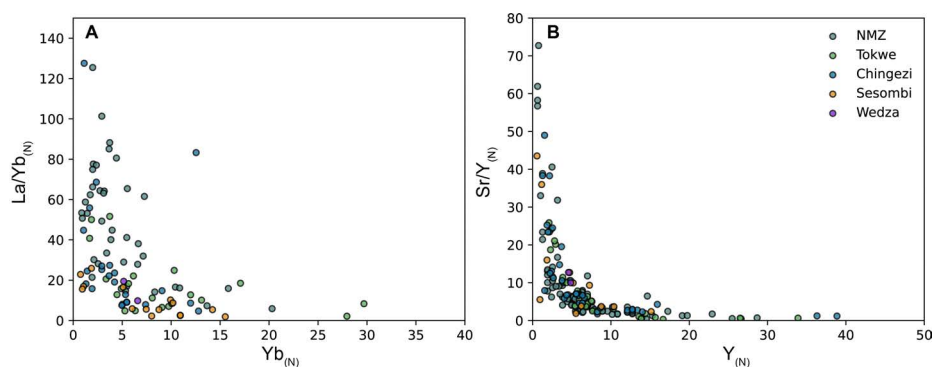


Figure 5.50: A) La/Yb_N vs Yb_N; and B) Sr/Y_N vs Y_N for all TTG suites in the Zimbabwe craton and northern marginal zone.

Each suite has distinct REE and compatibility trends, described individually below.

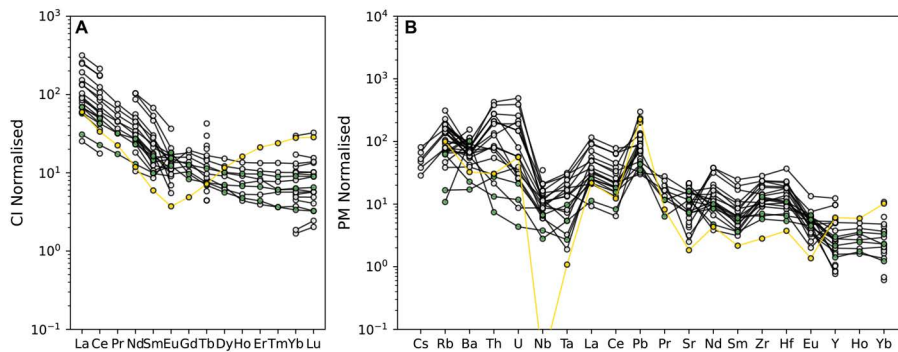


Figure 5.51: Chondrite-normalised REE plot (A), and primitive mantle-normalised spidergram (B) of samples from the Tokwe suite. Data collected during this study coloured in. Yellow line represents the early intrusive sample TR07. Elements are ordered in order of compatibility within a mantle-derived melt. Normalisation values from McDonough and Sun (1995).

5.2.3.1 Tokwe suite

Samples from the Tokwe suite exhibit comparable REE and trace element profiles (with the exception of *TR07*, Figure 5.51). All REE profiles are negative towards the HREEs, with a minimal or absent Eu anomaly (Figure 5.51A). The slope of the profiles shallows at Eu in all samples. Incompatible elements are significantly enriched relative to primitive mantle ($2\text{--}800 \times \text{PM}$), with notable positive Pb, and negative Nb-Ta anomalies (Figure 5.51B). Sample *TR07* exhibits a negative trend towards Eu, coupled with a strongly positive HREE trend (Figure 5.51A). Sample *TR07* additionally has a strongly negative Nb anomaly.

5.2.3.2 Chingezi suite

Samples from the Chingezi suite display three distinct REE trends (Figure 5.52A); a very strong negative REE trend coupled with a moderate Eu anomaly, a moderately negative REE trend with no strong Eu anomaly, and a moderately negative REE trend with a minor positive Eu anomaly. The three groups have overlapping spidergrams, with a prominent positive Pb anomaly, and a more developed negative Nb-Ta anomaly when compared to the Tokwe suite (Figure 5.52B).

5.2.3.3 Sesombi and Wedza suite

Samples from the Sesombi and Wedza suite overlap in both REE and spidergram patterns (Figure 5.53). La/Yb_N values are consistently < 30 and $\text{Sr}/\text{Y}_N < 45$ (Figure 5.50). All samples display a negative REE slope, and Eu anomalies are weak (Figure

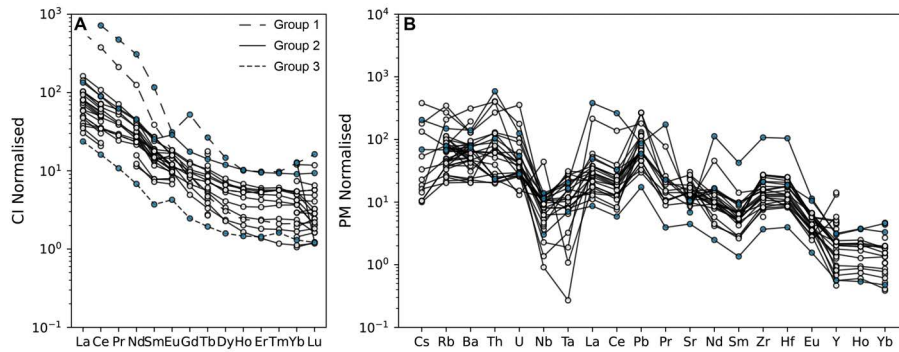


Figure 5.52: Chondrite-normalised REE plot (A), and primitive mantle-normalised spidergram (B) of samples from the Chingezi suite. Data collected during this study coloured in. Elements are ordered in order of compatibility within a mantle-derived melt. Normalisation values from McDonough and Sun (1995).

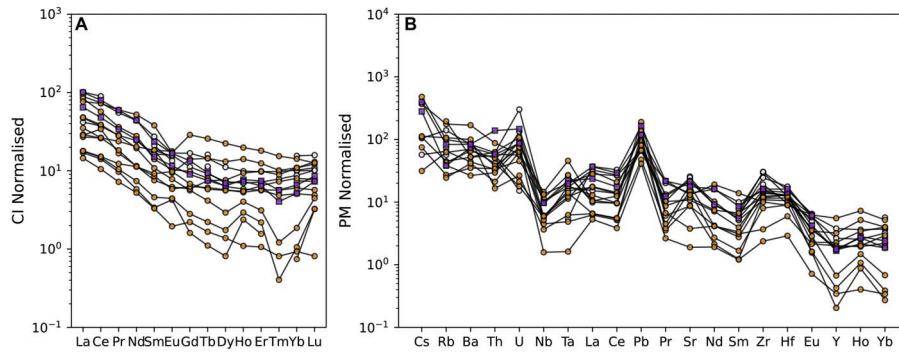


Figure 5.53: Chondrite-normalised REE plot (A), and primitive mantle-normalised spidergram (B) of samples from the Sesombi (orange circles) and Wedza (purple squares) suites. Data collected during this study coloured in. Elements are ordered in order of compatibility within a mantle-derived melt. Normalisation values from McDonough and Sun (1995).

5.53). Four samples have a positive trend from Tm to Lu. All samples exhibit a positive Pb anomaly, and Nb-Ta anomalies are absent. The four least evolved samples display sub-PM Y-Ho-Yb concentrations (Figure 5.53).

5.2.3.4 Chillimanzi and Razi suites

Samples from both the Chillimanzi and Razi suites display overlapping trends in REE and spidergram patterns (Figure 5.54). All samples display a negative Eu anomaly and HREE are more depleted than LREE (Figure 5.54). Razi suite samples are systematically more enriched ($20\text{--}800 \times \text{CI}$), than Chillimanzi suite samples ($1.5\text{--}400 \times \text{CI}$). All samples display a positive Pb anomaly, strongly elevated Th-U values, and

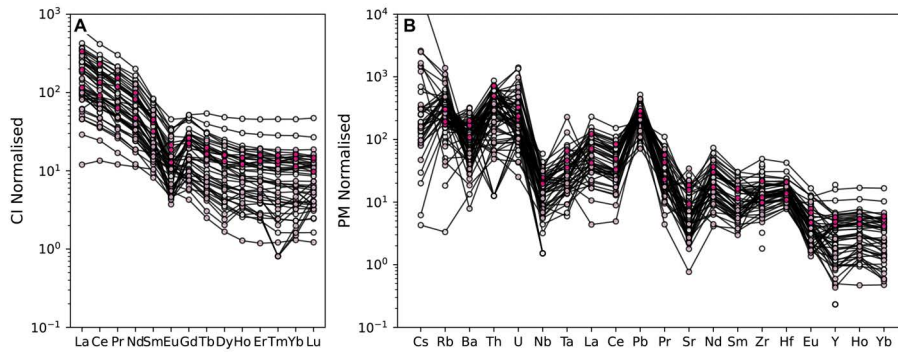


Figure 5.54: Chondrite-normalised REE plot (A), and primitive mantle-normalised spidergram (B) of samples from the Chillimanzi (light pink) and Razi (dark pink) suites. Data collected during this study coloured in. Elements are ordered in order of compatibility within a mantle-derived melt. Normalisation values from McDonough and Sun (1995).

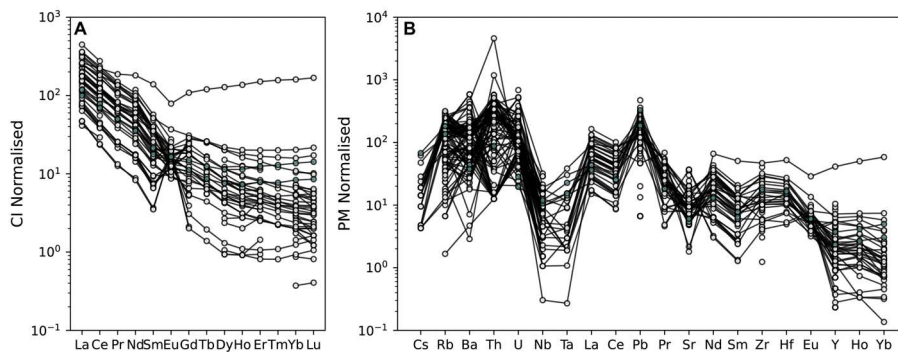


Figure 5.55: Chondrite-normalised REE plot (A), and primitive mantle-normalised spidergram (B) of samples from the Northern Marginal Zone. Data collected during this study coloured in. Elements are ordered in order of compatibility within a mantle-derived melt. Normalisation values from McDonough and Sun (1995).

negative Sr anomalies (Figure 5.54). Several samples display significantly elevated Cs and Rb (Figure Z.8B).

5.2.3.5 Northern Marginal Zone

With one exception, all samples from the Northern Marginal Zone (NMZ) exhibit strongly depleted HREE patterns, and both positive and negative Eu anomalies (Figure 5.55). Samples display limited Rb-Cs enrichment, a negative Nb-Ta anomaly, and a positive Pb anomaly (Figure 5.55). Thorium and U values are commonly significantly elevated ($10\text{--}1000 \times \text{PM}$, with one outlier at $5000 \times \text{PM}$).

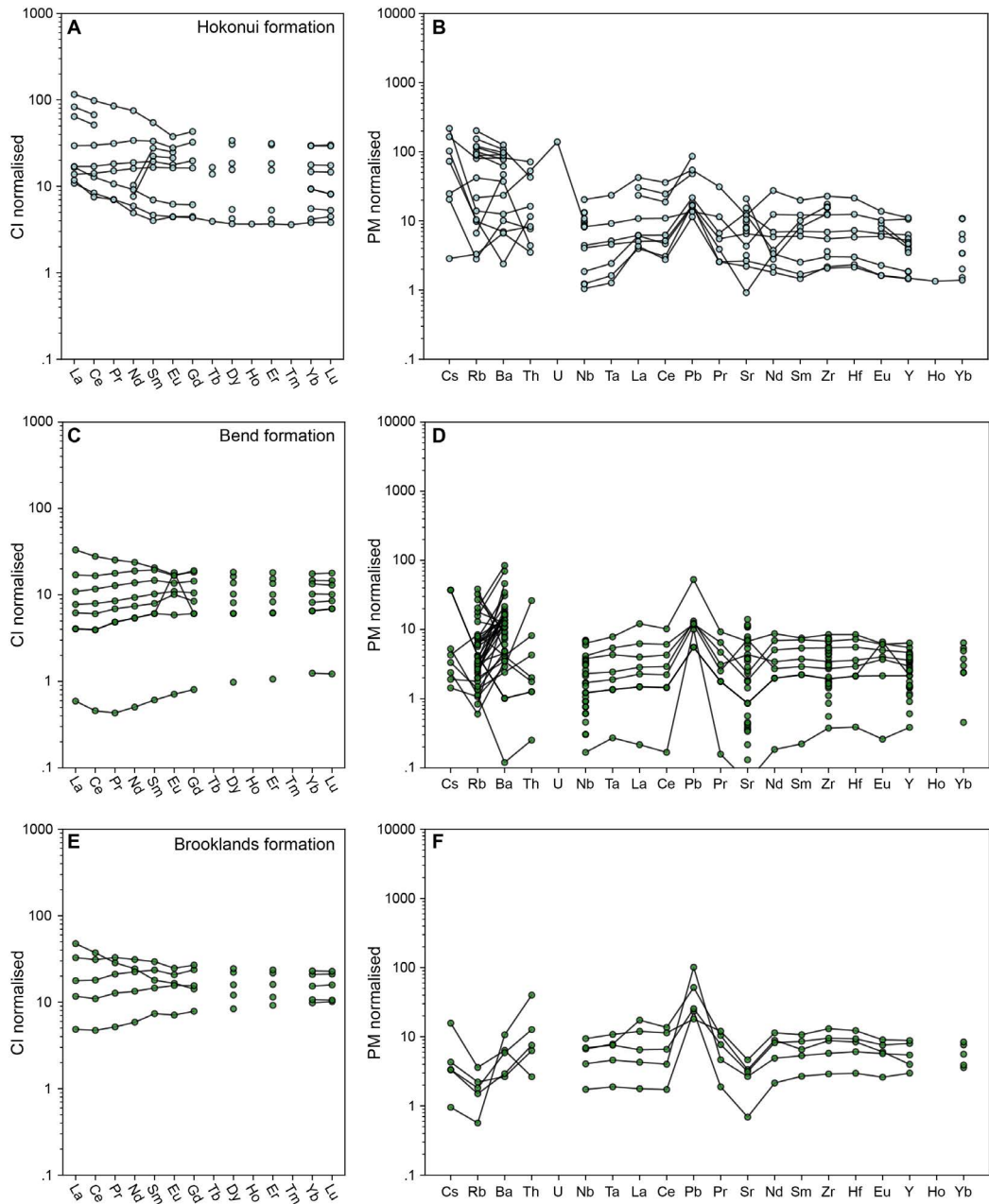


Figure 5.56: Chondrite-normalised REE diagrams (A,C,E), and primitive mantle-normalized spidergrams (B,D,F) for the andesitic Hokonui (A,B), (ultra)mafic Bend (C,D), and mafic Brooklands (E,F) formations.

5.2.4 Extrusive rocks

5.2.4.1 Belingwean group

The Belingwean group comprises the Bvute, Hokonui, Bend, and Brooklands formation at the type locality in the Belingwe greenstone belt. No samples with trace element geochemistry from the Bvute formation are recorded in the literature. The Hokonui formation samples exhibit two trends – mafic lithologies display flat REE patterns, whilst andesitic successions have weakly depleted HREE patterns (Figure 5.56). A Pb anomaly is present in all samples, and trace elements are moderately to strongly enriched ($1\text{--}200 \times \text{PM}$).

Samples from the Bend formation all display a relatively flat REE profile at $5\text{--}20 \times \text{CI}$. Trace elements are moderately enriched to sub-PM concentrations ($0.1\text{--}100 \times \text{PM}$), with a flat spidergram profile with the exception of a positive Pb anomaly (Figure 5.56).

Samples from the Brooklands formation exhibit flat REE and spidergram profiles (Figure 5.56). Trace elements are weakly–moderately enriched ($0.7\text{--}100 \times \text{PM}$), and a positive Pb anomaly is present in all samples.

5.2.4.2 Lower Bulawayan

The lower Bulawayan Koodoovale formation has limited geochemical data. One sample with REE data exhibits strongly depleted HREE and negative spidergram profiles (Figure 5.57).

5.2.4.3 Upper Bulawayan

The upper Bulawayan volcanics in the eastern greenstone belts (Murewa–Odzi, Masvingo, Belingwe, Gweru, Fort Rixon–Shangani, Gwanda, and the eastern portions of the Filabusi and Bulawayo belts) are predominantly komatiite–komatiitic basalt–basaltic successions of the Reliance and Zeederberg formations.

Samples of the Reliance formation have weakly depleted HREE to flat REE profiles (Figure 5.57). Samples are moderately to poorly enriched in trace elements relative to Primitive Mantle ($0.2\text{--}200 \times \text{PM}$), with minor Pb anomalies. Normalised trace element profiles are flat (Figure 5.57).

Samples of the Zeederbergs formation similarly display a weakly HREE depleted to flat REE profiles (Figure 5.57). Trace element profiles are flat, and enrichment is weak–moderate ($0.1\text{--}300 \times \text{PM}$). On average, samples are more enriched than the Reliance formation volcanics.

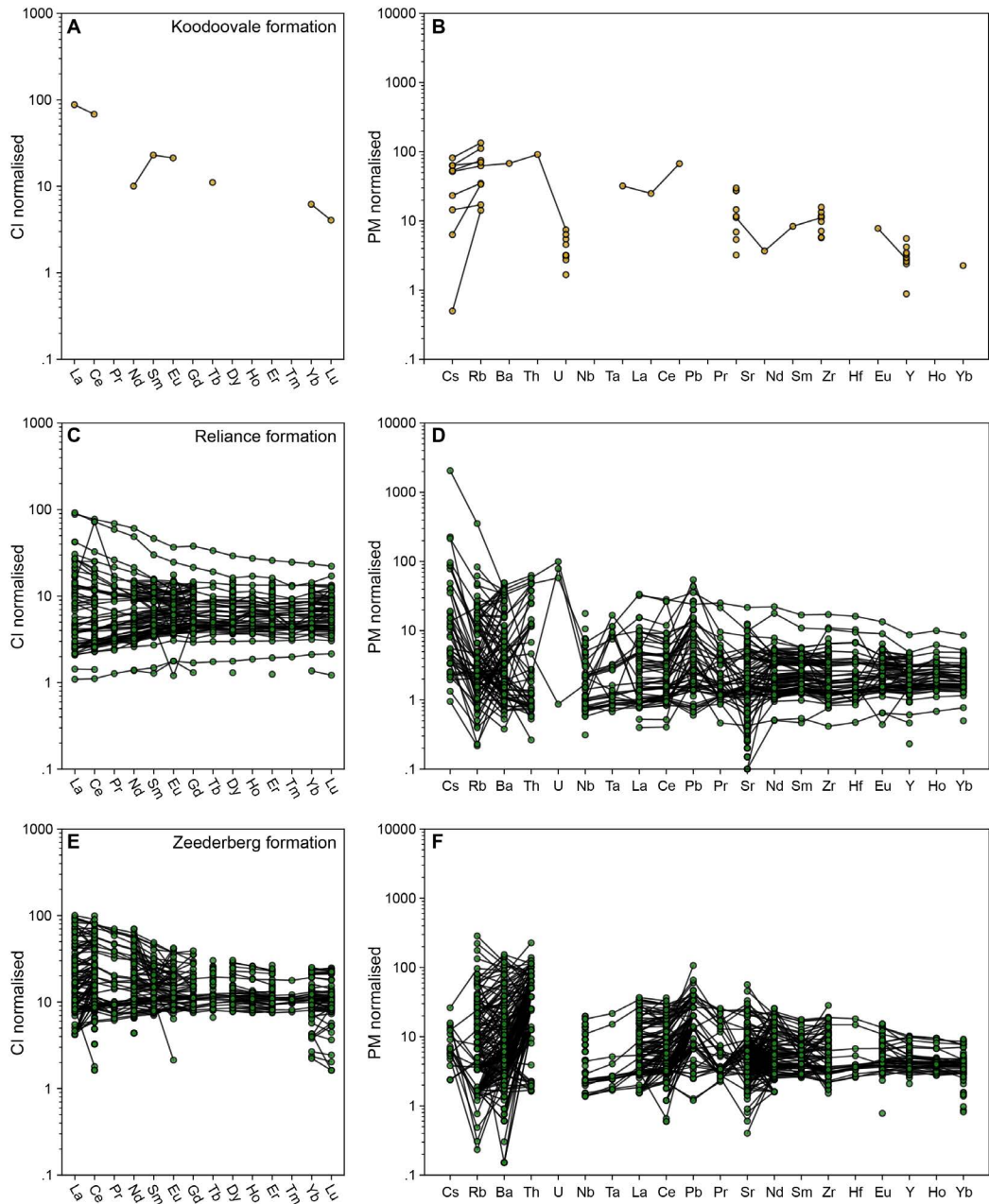


Figure 5.57: Chondrite-normalised REE diagrams (A,C,E), and primitive mantle-normalized spidergrams (B,D,F) for the Koodoovale (A,B), Reliance (C,D), and Zeederberg (E,F) formations.

5.2.4.4 Shamvaian

Limited geochemistry on Shamvaian volcanics exhibit depleted HREE profiles (Figure 5.58). Trace element profiles also show a negative trend, with notably low U contents, and a general enrichment of $4\text{--}300 \times \text{PM}$.

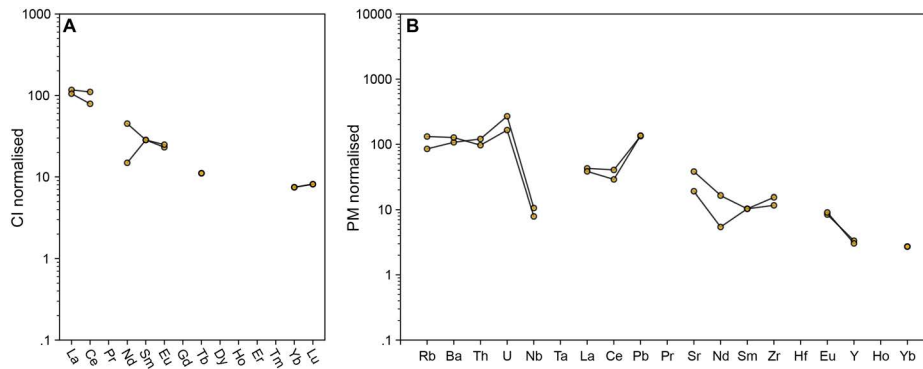


Figure 5.58: Chondrite-normalised REE plot (A), and primitive mantle-normalised spidergram (B) of samples from the Shamvaian group.

5.3 Discussion

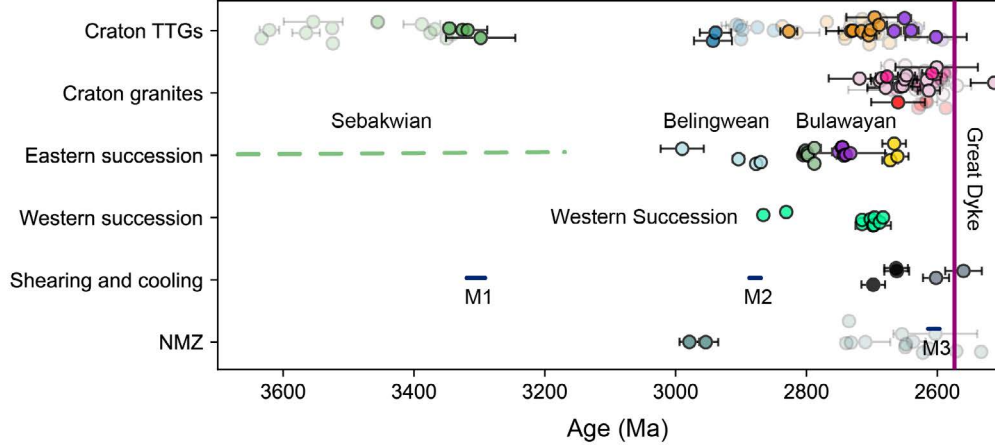


Figure 5.59: Summary of events in the Zimbabwe craton. Colours correspond to those used for suites for figures 5.47, with the addition of red circles (pegmatite crystallisation), and dark blue (metamorphic events).

5.3.1 Early cratonic nuclei in the Zimbabwe Craton (3.8–3.2 Ga)

The earliest *in-situ* remnants of felsic magmatism in the Zimbabwe craton are preserved in two cratonic nuclei named the Tokwe and Rhodesdale terranes (Figure 5.1), with oldest ages of 3633 ± 3 Ma (Hofmann et al., 2022) and 3565 ± 21 Ma (Horstwood et al., 1999), respectively (Stowe, 1979; Wilson et al., 1995; ; Figure 5.59). The presence of earlier felsic magmatism is suggested by detrital (Bolhar et al., 2017; Dodson et al., 1988) and inherited zircons (e.g., Hoffman et al., 2022 and this study), though these have not yet been identified. It is unclear whether the Rhodesdale and Tokwe terranes were adjacent to each other during the Paleoproterozoic, however this early magmatic episode appears to have generated TTGs sporadically from at least 3633–3340 Ma with remarkably similar geochemistry (Figures 5.60 and 5.61). Tokwe suite TTGs are characterised by high $\text{Na}_2\text{O}/\text{K}_2\text{O}$, low La/Yb , Sr/Y , and shallow negative REE patterns, which suggests an absence of rutile and garnet in the source (Martin and Moyen, 2002; Foley et al., 2002; Rapp et al., 2003). The Tokwe suite was therefore likely sourced by melting of relatively shallow (< 30 km) melting of a mafic hydrated crust, arguing against formation in oceanic plateaux or subduction zone environments during this period (Figure 5.62). Both Re-Os and Lu-Hf isotope systematics suggest that the mafic proto-crust precursor to the Tokwe suite likely separated from the

primitive mantle by 3.8 Ga (Bolhar et al., 2017; Hofmann et al., 2022; Nagler et al., 1997).

Early crustal recycling related to an amphibolite facies metamorphic episode generated potassic/granodioritic melts – the Mushandike and Mont D’Or granites – at c. 3300–3375 Ma (Dodson et al., 2001; Hickman, 1974; Horstwood et al., 1999; Moorbath et al., 1976). These early ‘potassic’ granitoids have Sm-Nd isotopic ages of 3540–3670 Ma, suggesting the incorporation of older felsic crustal material (Taylor et al., 1991).

Sample *TR07* from this study is intruded into c. 3.4 Ga crust that has a remarkably distinct geochemical signature (yellow dataset in Figure 5.60A). Positive HREE, negative Eu, and \ll Nb/Ta anomalies suggest the presence of rutile/amphibole and plagioclase in the melt source, with magmatic garnet retaining the HREE. The presence of Eoarchean inherited zircons within error of the earliest detrital zircon data (Figure 5.59, Bolhar et al., 2017) suggest that similar recycling was required to form these early leucogranitic sills.

Remnants of early extrusive volcanism and surface sedimentation are limited to tightly infolded keels within the Paleoproterozoic gneisses. Little is known about their origin, although ultra-mafic talc-tremolite schist and serpentinites, and metamorphosed cherts and ironstones suggest that mafic magmatism was coupled with submarine chemical sedimentation during this time period.

5.3.2 Quiescence, followed by renewed Mesoarchean magmatism (3.0–2.85 Ga)

Following Paleoproterozoic magmatism, the Zimbabwe craton experienced a period of relative quiescence and stability. Along the southern margin of the Tokwe terrane a stable passive margin developed now preserved in the Buhwa greenstone belt (Figure 5.1), suggesting a period of (limited) crustal emergence must have occurred to provide Tokwe-age detritus that was largely deposited under submarine conditions (Breneman et al., 2021; Fedo et al., 1995; Fedo and Eriksson, 1996; Fedo et al., 1996).

Following a period of magmatic quiescence, activity resumed with the generation of the Chingizi TTG suite between 2.95–2.85 Ga (Figure 5.59). The type area of the Chingizi suite lies on the western and southern edges of the Belingwe greenstone belt (Martin et al., 1993; Wilson et al., 1995, Figure 5.1). Extensions of the Chingizi suite have been found in the Mashaba tonalite, which transects the Tokwe terrane along a N-S axis (Wilson et al., 1995), the Chipinda batholith around the Buhwa greenstone belt, which was emplaced along major ductile oblique-slip dextral shear zones (Fedo et al., 1995), and along the eastern edge of the Mutare-Odzi greenstone

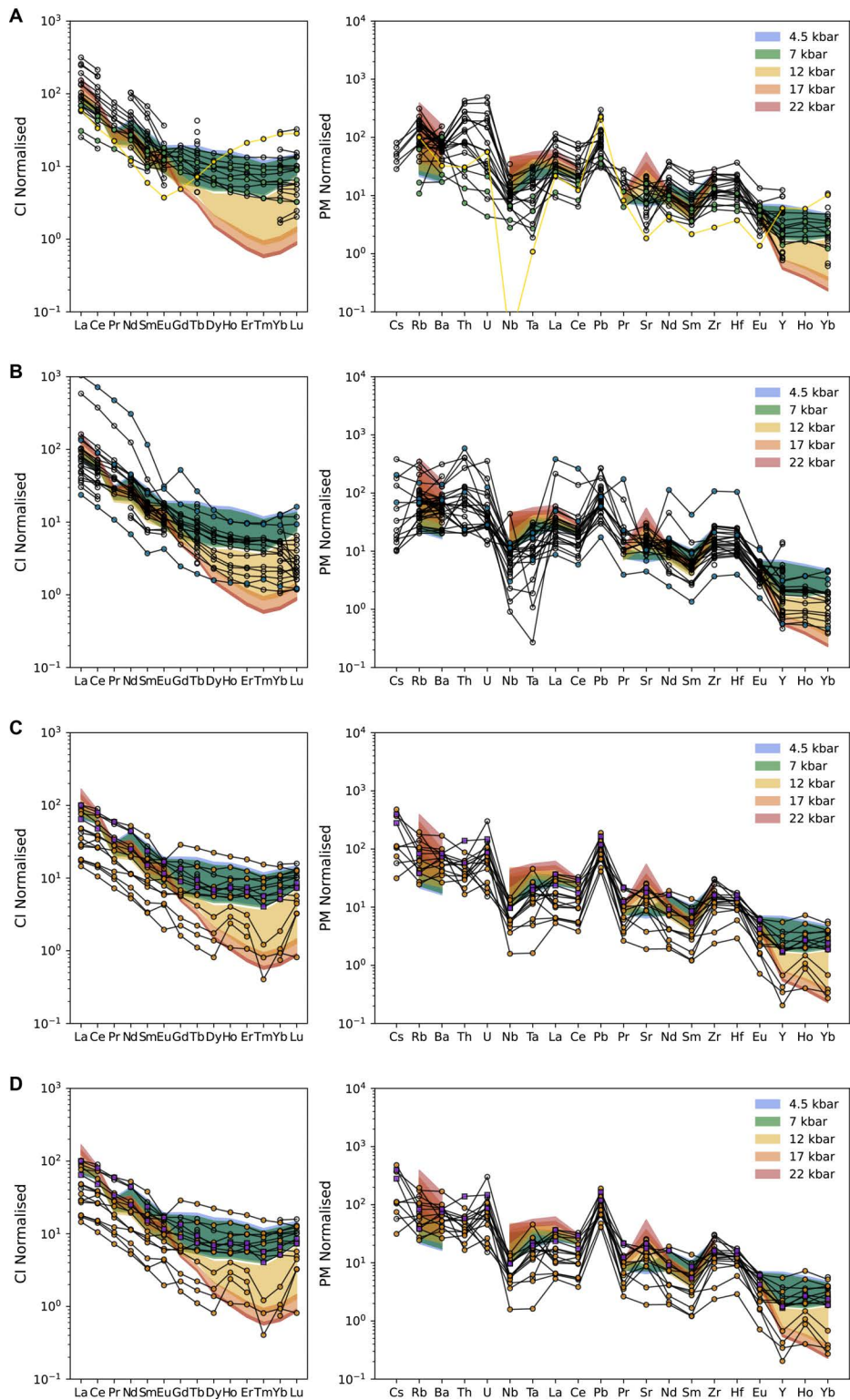


Figure 5.60: Trace element spidergrams (right) and REE plots (left) for all TTG suites in the Zimbabwe craton, with samples collected in this study coloured in. A) Tokwe suite, B) Chingezi suite, C) Sesombi/Wedza suites, D) Northern marginal zone. All plots are underlain by modelled EAT (Condie, 1981) melt compositions, coloured by pressure of the isobaric melt path.

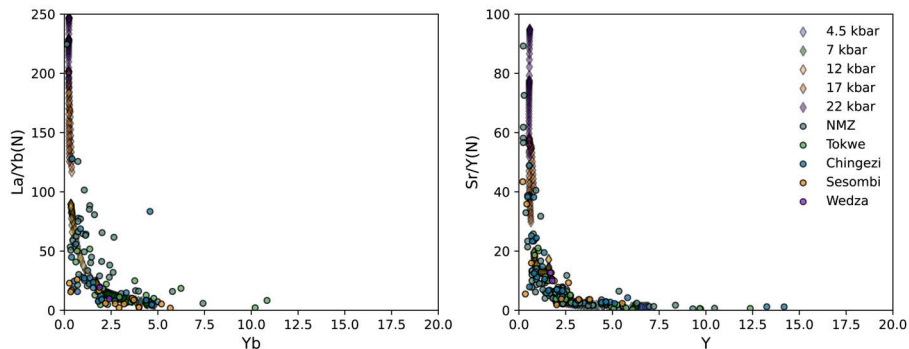


Figure 5.61: La/Yb vs Yb (left) and Sr/Y vs. Y (right), normalised to Primitive Mantle, (McDonough and Sun, 1995), diagrams of all TTG suites in the Zimbabwe craton. The data are underlain by modelled melts by melting EAT (Condie, 1981), colour-coded by pressure at which melting is modelled.

belt in Mozambique (Mantäri, 2008). The Chingezi suite presents elevated La/Yb, Sr/Y and more steeply negative REE plots, suggestive of a source melted at greater depth (>10 kbar, Figures 5.60 and 5.61, Luais and Hawkesworth, 1994), implying the crust substantially thickened prior to magmatism during time period.

This study also reports the first evidence of Chingezi-age magmatism in the Northern Marginal Zone (NMZ), which is preserved as highly deformed and migmatized granulitic basement into which later charnockitic and enderbitic magmas intruded. This discovery lends credence to the interpretation that the NMZ is a deep-crustal equivalent of the Zimbabwe craton (cf. Rollinson, 2022), and extends its magmatic history by 200 Myr. It is highly likely that pre-3.0 Ga felsic magmatism is not preserved in the NMZ, given that early Tokwe-suite melts were derived from sources shallower than reported depths of the currently exposed NMZ (Kamber and Biino, 1995; Rollinson, 1989).

Belingwean group volcanics formed coevally with intrusive magmatism members of the Chingezi suite, and are best exposed in the Belingwe (as the Hokonui, Brooklands, and Bend formations), Filabusi (as the Shamba-Kudu formations), Fort Rixon–Shangani (the “ycc” series of Harrison, 1969), and Gweru greenstone belts (Ranche formation) (Wilson et al., 1995, Figure 5.1). This group is primarily composed of a bimodal mafic extrusive and andesitic-dacitic agglomerate/tuff succession that document multiple major explosive volcanic episodes (Bickle et al., 1993; Martin et al., 1993). Minor chemical sedimentation forming thin banded iron formations occurred during this period, and textures within the agglomerate and tuff horizons suggest the majority of preserved deposition occurred in shallow marine environments

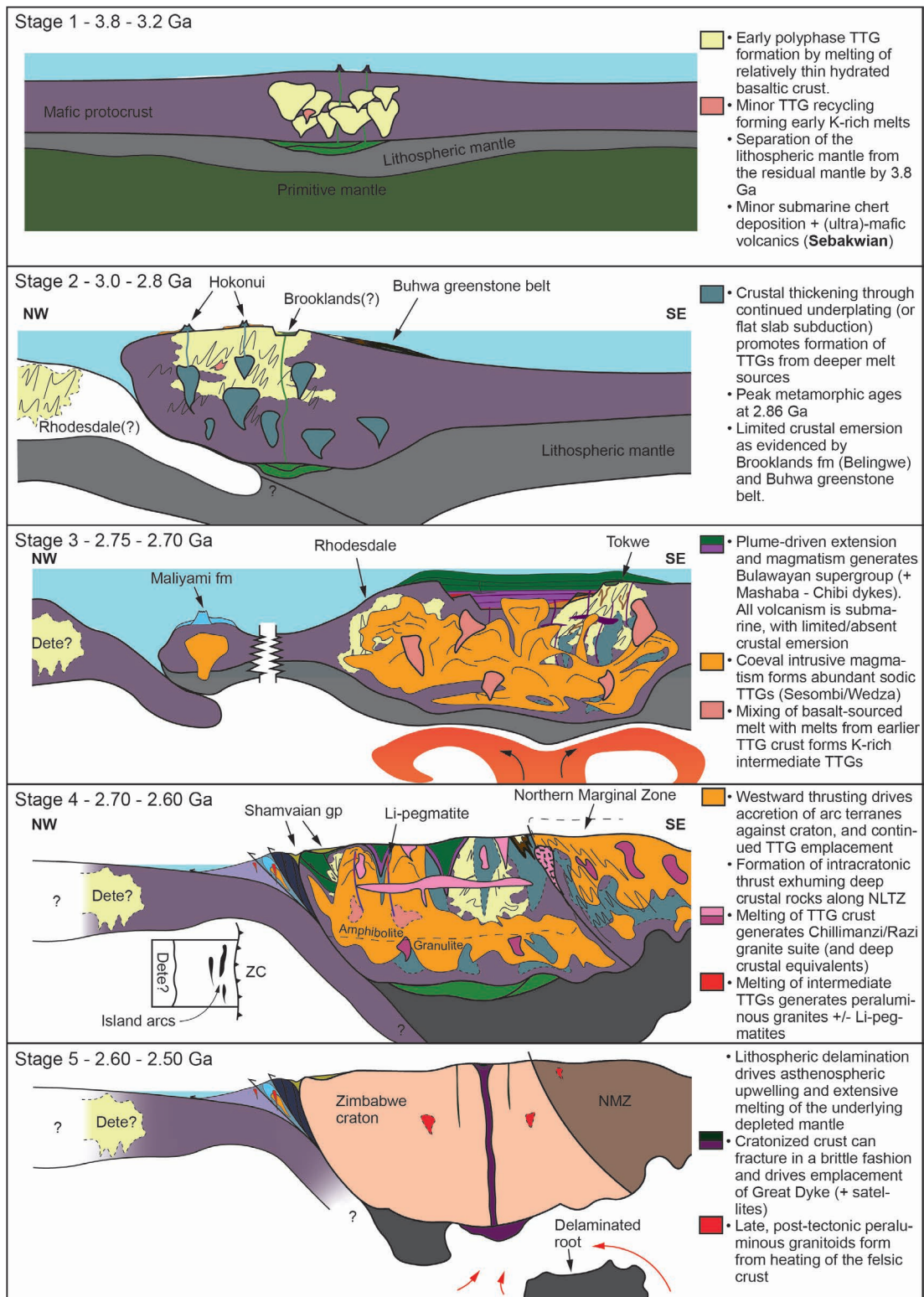


Figure 5.62: Schematic diagram summarizing the major stages of Zimbabwe craton evolution.

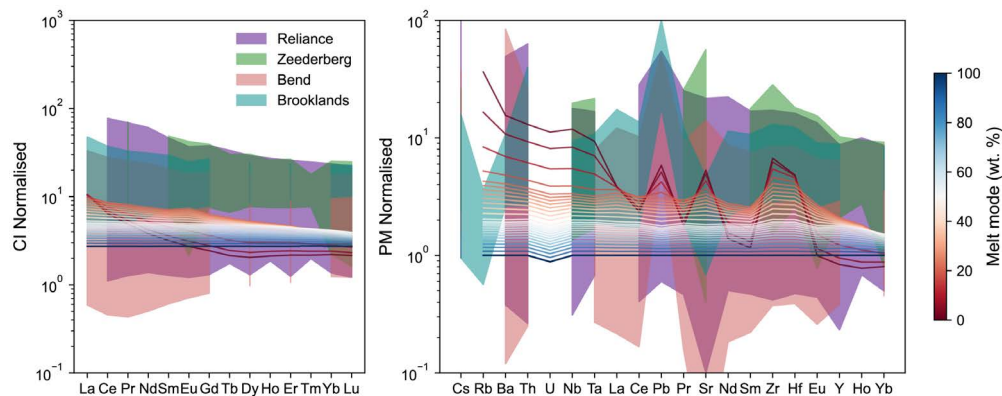


Figure 5.63: Chondrite-normalised REE diagram (left) and trace element spiderplot (right) showing the average composition of mafic volcanics in the central Zimbabwe craton. A model melt of relatively shallow (22 kbar) melting of primitive mantle is overlain, colour-coded by melt mode. Normalisation values are from McDonough and Sun (1995).

(Martin et al., 1993). The andesitic compositions and elevated LILE signatures (Figure 5.56) suggest fluid-rich melting environment, perhaps akin to modern-day mantle wedges situated above dehydrating subducting slabs.

Crustal thickening during this time period coincides with a second amphibolite-facies metamorphic event dated to c. 2.86 Ga (Figure 5.59; Nagler et al., 1997), and emplacement of the Chingezi suite was, at least in part, along major active shear zones (Fedo et al., 1995). These events may have coincided with the accretion of the Rhodesdale terrane onto the Tokwe terrane, with flat slab subduction suggested to promote substantial thickening via underthrusting (Figure 5.62). These processes are thought to have led to the formation of the Sebakwe protocraton (Jelsma et al., 2021), which subsequently evolved as a single cratonic block. In this scenario, the Brooklands formation preserved along the south-eastern margin of the Belingwe greenstone belt may represent a back-arc type rift basin within the Tokwe segment (e.g., Nisbet et al., 1993). In any case, both the Brooklands and Buhwa belt sedimentation indicate that there is some component of crustal emergence present during this time period, both preserving Tokwe-age detritus, however only generating limited sedimentary basins.

5.3.3 Plume-related growth of the craton (2.8–2.7 Ga)

Following a short period of quiescence, the last major period of crustal growth in the Zimbabwe craton began at c. 2.75 Ga. Large swaths of the central Zimbabwe craton preserve a coherent Bulawayan group stratigraphy in the greenstone belts, which

is subdivided into the felsic volcanic Koodoovale, shallow sedimentary Manjeri, ultramafic Reliance, mafic Zeederberg, and sedimentary Cheshire formations (Bickle et al., 1993; Jelsma et al., 2021; Wilson et al., 1995). The post-Koodoovale formation succession has been variously suggested to have an autochthonous origin (e.g., Blenkinsop et al., 1993; Hunter, 1998; Wilson et al., 1995), or an allochthonous origin (e.g., Hofmann and Kusky, 2004; Kusky, 1998; Kusky and Winsky, 1995).

Proponents for the allochthonous origin argue that the volcanic pile of the Bulawayan succession was tectonically emplaced along a major décollement onto a stable shelf — now preserved as the Manjeri formation — overlying basement gneisses (Hofmann and Dirks, 2003; Kusky and Kidd, 1992). The contact between the Manjeri formation and overlying volcanics commonly consists of a sulphidic banded iron formation (the Jimmy member of the Manjeri formation), which has been interpreted as a major mylonitic horizon (Hofmann and Dirks, 2003; Hofmann and Kusky, 2004; Kusky and Winsky, 1995). Hofmann et al. (2001; 2024) further argued that the top-most Cheshire formation experienced deformation prior to folding into the current synformal geometry, which they interpret to show that the Belingwe belt underwent an earlier stage of tectonic activity. To these authors, the upper Bulawayan group (Reliance fm. and up) is therefore a segment of an oceanic plateau that was thrust onto the Sebakwe protocraton.

This scenario conflicts with the model (e.g., Blenkinsop et al., 1993) for an autochthonous origin for the Bulawayan group. In the autochthonous model, the Manjeri formation represents a cyclical shallow fluvial–marine transgressive sequence within rifted grabens, capped by banded iron formations formed during sediment-starved deposition (Blenkinsop et al., 1993; Hunter, 1998). A clear nonconformity between the Manjeri formation and 3.35 Ga granitoid gneisses (Hofmann et al., 2022) is exposed at the Basement Unconformity Exposure national monument south of Zvishavane (Blenkinsop et al., 1993; Hunter, 1998; Martin et al., 1993; Nisbet et al., 1977), and similarly as an angular unconformity against the Hokonui formation along the Dohwe river along the western edge of the Belingwe greenstone belt (Martin et al., 1993). Continuous deposition and a transition to volcanic activity is inferred by the presence of isolated quartzose horizons between the basal Reliance formation volcanics (Blenkinsop et al., 1993), and Prendergast (2004b) further suggests that interdigitation between the Koodoovale, Manjeri, and Reliance formation occurs at various points across the central craton. This would eliminate the possibility of an allochthonous origin for the volcanic pile. In addition, trace element (e.g., elevated

Th) and Nd-isotope geochemical evidence suggests that the Bulawayan volcanics experienced a component of crustal assimilation prior to eruption (Bolhar et al., 2003; Chauvel et al., 1993; Shimizu et al., 2005), as also supported by inherited zircons found in felsic volcanic horizons (Wilson et al., 1995). Arguably most convincing is the temporal overlap between the Mashaba-Chibi dyke swarm and associated ultramafic intrusive bodies (e.g., the Shabani and Mashaba sills), interpreted to be the intrusive feeders to the Bulawayan volcanics (Prendergast, 2004a; Prendergast and Wingate, 2007), which are emplaced into the Tokwe segment (Prendergast and Wingate, 2013; Wilson et al., 1995).

It is therefore most likely that the Bulawayan group was deposited onto the Sebakwe proton, and not tectonically juxtaposed. The volcanic succession is generally considered to have formed by plume-driven mantle melting, forming a large igneous province (c. 250 000 km², Prendergast, 2004a) atop a felsic protocrust, starting at c. 2.75 Ga (Hofmann et al., 2024; Prendergast and Wingate, 2013). A mantle plume origin is invoked to generate the high temperatures required to form the komatiitic successions (Shimizu et al., 2005), and supported by melt inclusion studies by McDonough and Ireland (1993), who showed that incompatible/compatible element ratios in the komatiitic melts are comparable to modern plume-related magmas. Models of primitive mantle melting closely align with the generally observed flat REE profiles of the Reliance and Zeederberg formations, with progressive enrichment inferred to be caused by fractional crystallisation within a magma chamber prior to eruption (Figure 5.63, Bolhar et al., 2003). A very deep (> 150 km) melt origin, as suggested by Shimizu et al. (2005), is not required (Figure 5.63), though the absence of depletion in incompatible elements suggests that the mantle was not extensively depleted prior to this event (e.g., Salters and Stracke, 2004). The upper Cheshire formation, which consists primarily of a basal carbonate ramp grading upwards into a deeper water siliciclastic facies formed by turbidity currents (Hofmann et al., 2001; Martin et al., 1993), must have formed prior to 2.71 Ga, thereby constraining the (ultra)-mafic volcanism to a <40 Myr period (Hofmann et al., 2024). This encompasses sufficient time to generate comparable plume-related large-igneous provinces (e.g., Columbia River flood basalts, 210 000 km² in c. 11 Myr, Reidel et al., 2013; Deccan Traps, 500 000 km² in c. 3 Myr, Jay and Widdowson, 2008).

Initial plume impingement is expected to have driven partial melting of the lower crust, generating the intermediate-felsic volcanic Koodoovale formation (in the Belingwe greenstone belt) early in the Bulawayan succession (Prendergast, 2004a; Prendergast and Wingate, 2013). Similar stratigraphy is observed in the Filabusi (Eldo-

rado formation), Bulawayo (Lonsdale formation), Gweru (Arizona formation), and Fort Rixon–Shangani (Series formation). Associated with the volcanic activity is extensive intrusive magmatism, generating the Sesombi suite (Figure 5.62). The presence of both high and low La/Yb, Sr/Y, and variably negative REE patterns all suggest that melting of a mafic protolith would have occurred at varying depths in the crust during this time. Some of the protolith may have been underplated equivalents of the Bulawayan volcanic succession. The relative increase in proportion of TTGs with elevated K_2O/Na_2O values in TTGs of this time period (Figure 5.49) in both the NMZ and the Zimbabwe craton suggests that earlier (Tokwe/Chingezi suite) TTGs were also incorporated in the melting environment and mixed with primitive EAT-sourced melts (cf. Rollinson et al., 2024).

5.3.4 Accretion of the Western Successions and Shamvaian deposition (2.71–2.60 Ga)

The Western Succession comprises an almost continuous greenstone belt from southwest of Bulawayo towards Shamva in the north-east of Zimbabwe (Figure 5.1). Typifying the succession is extensive calc-alkaline felsic volcanism, and a relatively high proportion of felsic and intermediate volcanics (c. 45% of the stratigraphy) when compared to greenstones of the central craton (Jelsma and Dirks, 2002). Intermediate–felsic volcanism is dated at 2715–2683, with later volcanic activity found in the Harare greenstone belt at 2648–2643 Ma (Jelsma et al., 2021; Wilson et al., 1995).

The stratigraphy within the western succession is often complicated and obscured by later tectonic repetition and shearing. In the Midlands greenstone belt, the Mafic formation (2880 ± 8 Ma, Horstwood, 1998) and What Cheer formation (2683 ± 8 Ma, Wilson et al., 1995) are interdigitated basalt, quartz grits, and siltstones (Jelsma and Dirks, 2002), and are abutted by a major transpressive shear zone against the Maliyami formation; a thick sequence of calc-alkaline andesites and dacites (2702 ± 6 Ma, Wilson et al., 1995). All are intruded by several late TTGs dated to c. 2670 Ma (Jelsma et al., 2021). The Midlands greenstone belt has been shown to have undergone significant west-directed thrusting, forming complex compressional duplexes which were later reactivated as (minor) strike-slip shear zones (Campbell and Pitfield, 1994; Dirks et al., 2002). This deformational event incorporated slivers of Paleoproterozoic TTGs of the Rhodesdale terrane (Jelsma and Dirks, 2002). Similarly, the Bindura–Shamva and Harare greenstone belts (Figure 5.1) experienced significant west-directed thrusting and imbricate stacking above major layer-parallel shear zones, stacking a 6 km thick tholeiitic basalt succession (Arcturus formation, 2697 ± 9 Ma, Wilson et al.,

1995) on top of a basal andesitic Iron Mask formation (2712 ± 7 Ma, Jelsma et al., 2021). Structurally overlying the Arcturus formation is the andesitic/dacitic volcanic Passaford/Lower Shamva and Mt Hampden (graphitic argillite) formations, dated at 2643 ± 8 Ma (Wilson et al., 1995). The felsic volcanics are closely associated with porphyry stocks and temporally overlaps with late tectonic TTGs of the Wedza suite (2648 ± 6 Ma, Jelsma et al., 1996), and SR01-2, WG02-3 in this study). The felsic volcanic rocks in the Western Succession have relatively unevolved isotopic signatures (e.g., initial Sr-ratios of 0.701 and 0.704, Baldock and Evans, 1988; Jelsma et al., 2021, $\epsilon\text{Nd } 2.5 \pm 0.4$, Horstwood, 1998), and μ_1 values between 7.7 and 8.6 (Taylor et al., 1991), suggesting a significant juvenile component in the source.

In the Bulawayo greenstone belt (Figure 5.1), the Avalon and Kensington formations are infolded and thrust mafic volcanic and calc-alkaline successions. The top of the sequence (Umzingwane formation) is a thick volcanic laharic type breccia overlain by volcanoclastic turbidites and containing intercalations of rhyodacitic lava (Garson, 1995). Interestingly, the calc-alkaline succession in the Bulawayo greenstone belt structurally overlies the Westacre and Umganin formations, which can be correlated to the Bulawayan volcanic successions in the central craton. This succession appears to be imbricated in a series of W-NW directed thrust stacks into the Kensington and Avalon formations (2683 ± 8 Ma, Jelsma et al., 2021).

Throughout the Western succession greenstones are a series of fault-bounded siliciclastic successions which are defined as the Shamvaian group. These are primarily preserved in the Midlands and Harare–Shamva greenstone belts, and correlative successions are also found in the Masvingo greenstone belt. The Shamvaian group are predominantly coarse grained fluvial-deltaic successions composed of conglomerates, grits, and poorly sorted sandstones and minor shales (Dirks et al., 2002; Hofmann et al., 2002). Sporadic lenses of intermediate-felsic volcanism exist, which alongside cross-cutting intrusive rocks date Shamvaian deposition to 2.68–2.61 Ga (Jelsma et al., 2021). Importantly, the Shamvaian succession presents the only significant siliciclastic sedimentary basins documented in the craton.

Combined deformation and shear-sense observation suggests that the Western Succession experienced several generations of W-NW directed thrusting (Garson, 1995; Jelsma and Dirks, 2002), and later strike-slip shearing and reactivation (Campbell and Pitfield, 1994; Dirks et al., 2002). The fault-bounded Shamvaian successions at c. 2.68–2.61 Ga coincide with U-Pb titanite ages dating shearing along the Irisvale-Lancaster shear zone (Figures 5.43 and 5.44), suggesting that the major crustal shear

zones (Campbell and Pitfield, 1994; Stowe, 1980) that transect the craton likely developed during the same deformational event.

The calc-alkaline signatures, along with some overlapping ages with the Bulawayan succession of the central craton indicates that the Western Succession volcanics are as a series of accreted volcanic arcs, possibly incorporating small oceanic plateaux (e.g., Mafic formation, Condie and Harrison, 1976; Horstwood, 1998). Lithologies to the west may be a different Archean cratonic block, which today are exposed in the Dete-Kamativi inlier (Glynn et al., 2020, Figure 5.1). If true, this juxtaposition may mark the closure of an early ocean that represents a valuable opportunity for future investigation. Accretion onto the central Sebakwe protocraton (Jelsma et al., 2021) coincided with the continued emplacement of TTGs, and likely aided by the significant density contrast with the overlying (ultra)-mafic volcanics, promoted (at least partially) solid-state diapirism, generating the general dome-and-keel pattern observed throughout the craton (e.g., Copley et al., 2023; Kusky et al., 2021; Weinberg and Podladchikov, 1995). This presents a unique scenario where geological features conventionally attributed to vertical-style tectonic processes may have formed during a horizontal tectonic event. Similar relationships have recently been proposed in the Wabigoon superterrane, Canada (Tóth et al., 2025). A metamorphic event, coupled with the development of major crustal shear zones accommodated strain in the cratonic block c. 2.68–2.61 Ga (Figure 5.62). The onset of exhumation of the NMZ along the Northern Limpopo Thrust Zone, which overlaps in age with this event (Blenkinsop, 2004; Rollinson and Blenkinsop, 1995) may be related to intracratonic basement uplift related to this accretionary event. The NMZ may therefore be similar to the Kapuskasing zone in the Superior Craton, Canada (c. 2.0 Ga, Percival and Card, 1983; Percival and West, 1994).

5.3.5 Emplacement of potassic granites (2.68–2.51 Ga)

Neoproterozoic potassic granites occur throughout the craton and occupy c. 50% of the current level of exposure (Blenkinsop and Treloar, 2001; Chagondah et al., 2023; Jelsma et al., 2021). Commonly characterised by K-feldspar megacrysts, the late granites have historically been grouped into the Chillimanzi (cratonic) and Razi (NMZ and NMZ-craton margin) suites. Most large-scale Chillimanzi suite intrusions were likely fed by thin vertical dykes and emplaced as thick horizontal sheets making them appear voluminous (Blenkinsop and Treloar, 2001). A general younging eastward trend has been noted previously in the granitic suite, from c. 2680 Ma near the Western Succession towards 2517 Ma along the eastern extent of the NMZ (Frei

et al., 1999; Horstwood, 1998; Jelsma et al., 1996; Mkweli et al., 1995). However, the first age of the Nalatale granite (2512 ± 36 Ma, apatite U-Pb, Figure 5.42), which partially bisects the Fort Rixon–Shangani greenstone belt (Figure 5.1), indicates that late Archean felsic magmatism occurs near the Western Succession as well. The older Rb-Sr biotite cooling ages (2602 ± 20 Ma and 2560 ± 28 Ma, Figure 5.46) supports the fact that this is a crystallisation age and not a later thermal resetting.

Compositionally, most Chillimanzi and Razi suite granites are metaluminous to slightly peraluminous (I-type, Figure 5.65), with some highly peraluminous ($A/CNK > 1.1$), and plot along the calc-alkaline trend of Frost et al. (2001, Figure 5.49C). Evolved geochemical (elevated K/Na, Th, U, and ΣREE , Figures 5.49, 5.54) and isotopic ($\epsilon Nd = -1.6$, Horstwood, 1998) characteristics strongly support their formation by melting of pre-existing felsic lower crust. Modelling of potential felsic sources suggest the best match occurs when melting both sodic and potassic Sesombi suite TTGs at mid-lower crustal levels (Figure 5.64). Importantly, the most evolved units, characterised by elevated Th and U, and the lowest Mg #, best fit melting intermediate TTGs (Figure 5.65B), in agreement with observations made by Rollinson et al. (2024).

The wide range of ages (2.68–2.51 Ga) during which the Chillimanzi and Razi suite granites were emplaced suggests that multiple processes may have generated the collective suites. Early granites (both metaluminous and peraluminous) may be associated with final accretion of the Western Succession, followed by late, peraluminous granites (e.g., the Nalatale granite) possibly driven by mantle heating during emplacement of the Great Dyke and final cratonic stabilisation (discussed below).

5.3.6 Emplacement of the Great Dyke (2.575 Ga)

The Great Dyke remains a feature unique to the Zimbabwe craton. The main intrusion is intruded as a 550 km long, and up to 11 km wide dyke (Worst, 1960; Wilson, 1990), and is dated to 2574 ± 2 Ma (Armstrong and Wilson, 2000; Wingate, 2000; Oberthür et al., 2002). Two main satellite dykes occur both west (Umvimeela Dyke) and east (East Dyke) of the main intrusion, and are related to the final pulses of magmatism of the suite (Harrison, 1969; Wingate, 2000). The Great Dyke is emplaced along a NNE trending fault set (Worst, 1960; Wilson, 1990), which may have formed by extensional relaxation following the accretion of the Western Succession. The ultramafic compositions of the Great Dyke imply a significant amount of mantle melting must have occurred (Oberthür et al., 2002; Wilson, 1996). Compositionally, there is no evidence of a deep mantle source, and extensive melting of the depleted

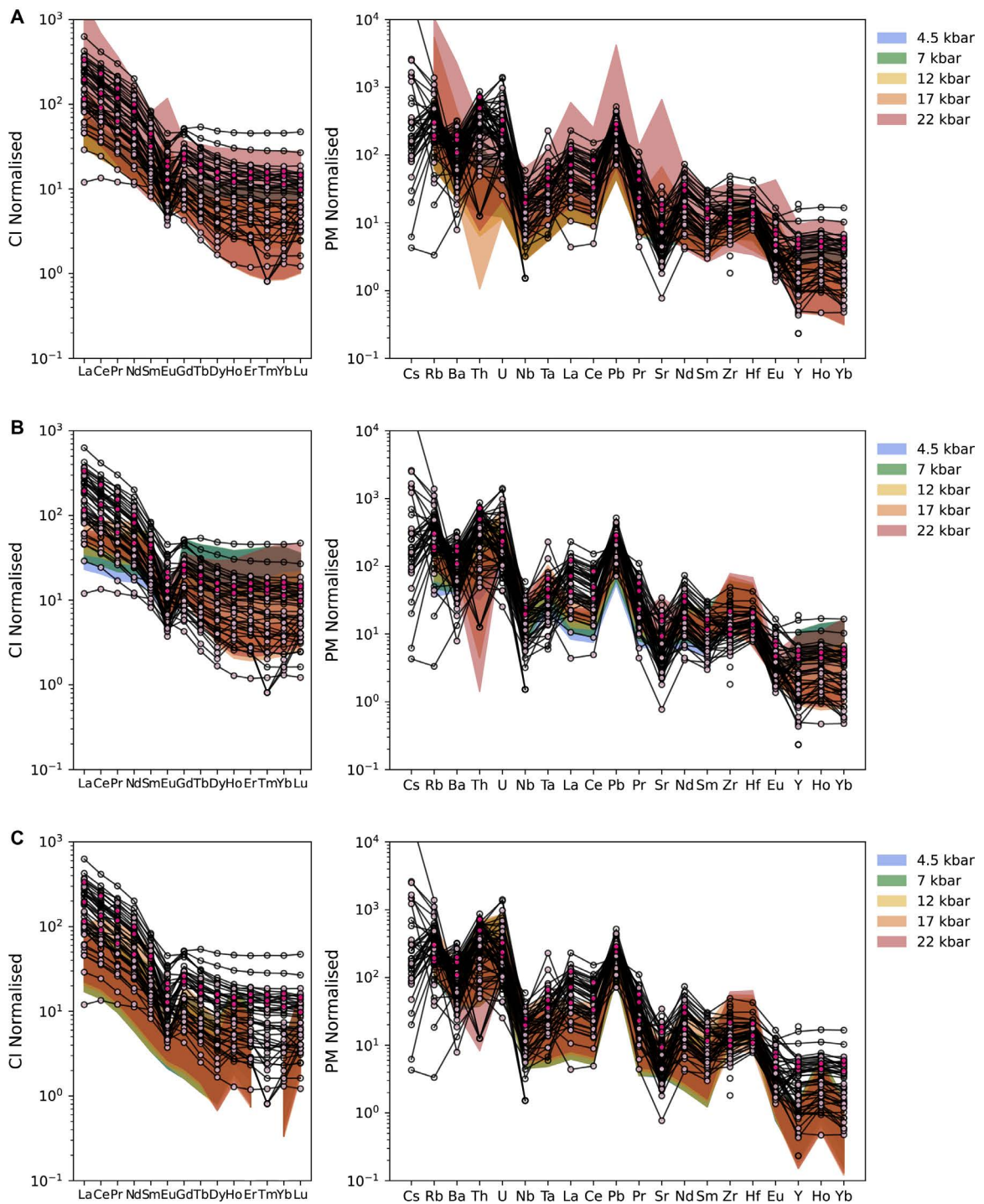


Figure 5.64: Chondrite normalised REE plots (Left) and primitive mantle normalised trace element spidergrams (right) of the Chillimanzi and Razi suite granites overlain onto modelled compositions of direct partial melts of Tokwe suite (TR02, A), Sodic Sesombi suite (GW04, B), and intermediate TTG Sesombi suite (ZB02, C).

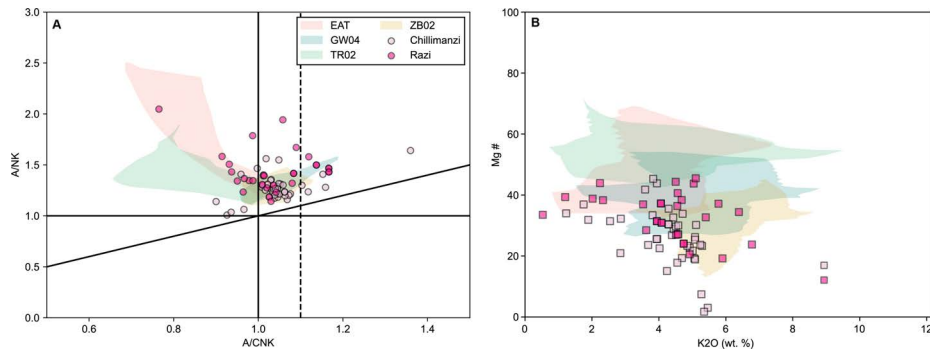


Figure 5.65: A) $\text{Al}_2\text{O}_3/(\text{Na}_2\text{O}+\text{K}_2\text{O})$ (A/NK) vs. $\text{Al}_2\text{O}_3/(\text{CaO}+\text{Na}_2\text{O}+\text{K}_2\text{O})$ (A/CNK) (molar proportions) diagram, after Maniar and Piccoli (1989). B) Mg # ($\text{Mg}/(\text{Mg}+\text{Fe})\times 100$, molar proportions) vs. K_2O diagram. Both present data from the Chillimanzi and Razi suites overlain on modelled compositions of direct melting of EAT (Condie, 1981), TR02 (Tokwe Suite), GW04 (sodic Sesombi suite), ZB02 (intermediate TTG, Sesombi suite).

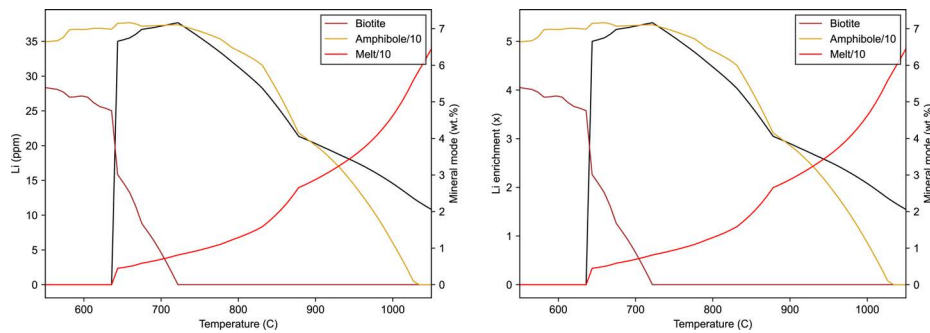


Figure 5.66: Li concentrations (left) and maximum melt enrichment (right) during modelled melting of EAT (Condie, 1981). Modal proportions of melt, biotite, and amphibole are shown.

mantle below the craton is invoked (Oberthür et al., 2002). It is feasible that mantle depletion would have occurred during the Bulawayan volcanic event, and that subsequent melting was driven by significant delamination of the lower crust (Figure 5.62).

Exactly how such a large, effectively singular intrusion was able to form in the craton is unclear, however it is worth noting that this must have been able to occur in the midst of extensive granitic magmatism, given Chillimanzi and Razi suite crystallisation ages straddle the emplacement of the Great Dyke.

5.3.7 Lithium fertility during crustal evolution

Compared to other Archean cratons worldwide, the Zimbabwe craton is uniquely endowed in lithium-rich LCT/Group 1 pegmatites (Goodenough et al., 2025). However, unlike other Li pegmatite-rich domains, the Zimbabwe craton lacks the sedimentary successions commonly assumed to be protoliths for such melts (Černý, 1991b; Černý et al., 2012; London, 2008). It is therefore unclear which geological mechanisms allowed such enrichment in Zimbabwe.

The assumption that sedimentary material is needed within a source region undergoing partial melting arises from the compositional affinity that LCT/Group 1-type signatures have with S-type granites (Černý et al., 2012; Chappell and White, 1992, 2001). Enrichment in lithium is a key characteristic of LCT pegmatites (Černý et al., 2012), and is considered to arise from melting a previously unmelted, mica-bearing metapelite or metagreywacke (Acosta-Vigil et al., 2012; Černý et al., 2012; London and Morgan, 2012). Similarly, elevated P contents commonly seen in LCT-type pegmatites are considered diagnostic of a metapelitic protolith (London et al., 1999; Martin and De Vito, 2005). Initial enrichment in Li, Cs, Rb, P, etc. in the metasedimentary source is a product of the relative incompatibility of these elements in granitic minerals (i.e. quartz and two feldspars), thereby promoting their release during weathering, which thus acts as a pre-concentrating step. In this model, the source for LCT/Group 1 pegmatites must form via crustal recycling prior to attaining the required enrichments to form key LCT ore minerals (see also Chapter 3).

If recycling of felsic continental crust is required to generate a fertile source rock, this can be achieved simply through crustal melting–remelting processes. In the Zimbabwe craton, and Archean cratons more generally, the first stage of differentiation occurs during the formation of primitive (sodic) TTGs from a hydrated metabasaltic source. Assuming average EAT concentrations for lithium (c. 7 ppm, Condie, 1981), mid-lower crustal melting of hydrated EAT can generate a melt of TTG composition with Li contents of up to 36 ppm (Figure 5.66). The release of lithium into the melt is primarily controlled by the breakdown of biotite, with minor influence of amphibole. Lithium concentrations of ≤ 36 ppm is within the range of Li concentrations measured in Tokwe suite TTGs.

In the Zimbabwe craton, there is evidence for the mixing of primitive TTG melts (sourced from hydrated metabasalt) and those derived by melting of pre-existing TTG crust, generating intermediate-stage TTGs with elevated K/Na throughout craton evolution (Figure 5.49). Taking *TR02* (Li = 32 ppm) as a typical Tokwe suite TTG, melting of early, sodic TTGs can generate melts with Li up to c. 300 ppm (Figure

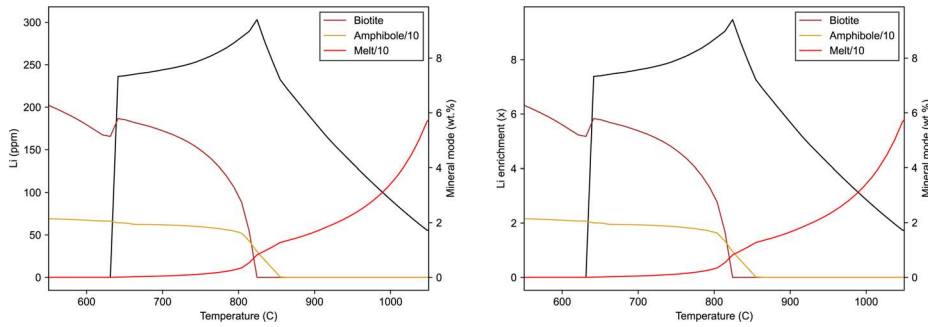


Figure 5.67: Li concentrations (left) and maximum melt enrichment (right) during modelled melting of TR02, as a representative TTG of the Tokwe suite. Modal proportions of melt, biotite, and amphibole are shown.

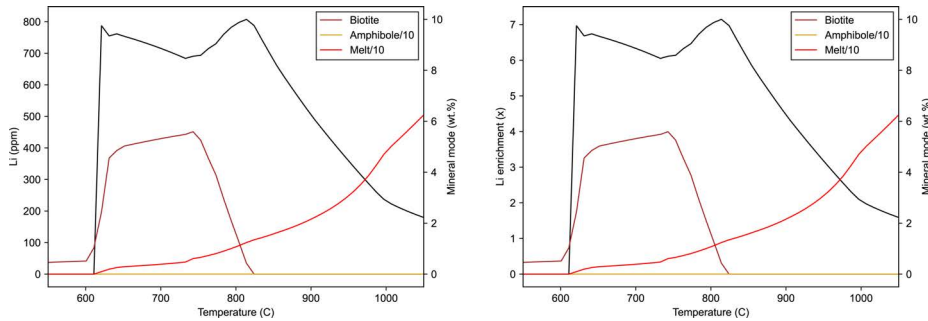


Figure 5.68: Li concentrations (left) and maximum melt enrichment (right) during modelled melting of ZB02, as a representative intermediate TTG of the Sesombi suite. Modal proportions of melt, biotite, and amphibole are shown.

5.67). Mixing of this melt (25%) with a primitive melt (75%) sourced from a hydrated mafic source would therefore generate an intermediate stage TTG ($K_2O/Na_2O \approx 1$) with Li concentrations of ~ 100 ppm, which is in line with what is observed in the craton.

Finally, melting of intermediate TTGs (taking *ZB02* as an example, Li = 113 ppm) can generate enriched melts with concentrations of c. 800 ppm Li, with a maximum enrichment factor of c. 7 (Figure 5.68). Whilst this is insufficient to saturate in spodumene or petalite (>5000 ppm Li), subsequent fractional crystallisation of the granitic melt can further enrich Li, either within granitic staging bodies or directly within pegmatites generating strongly zoned pegmatitic bodies (London, 2014). Such strongly zoned bodies are seen in the Zimbabwe craton (e.g., Arcadia, Goodenough et al., 2025; Mistress, Baldock, 1991). Later melting of early formed fertile Chillimanzi/Razi suite granites (cf. Chapter 3) can, alongside fractional crystallisation (owing to the higher starting concentrations), directly form melts with Li contents >5000 ppm, enabling the formation of unzoned pegmatites.

Melting of intermediate TTGs generates peraluminous–highly peraluminous melt compositions (Figure 5.65A), low Mg # (Figure 5.65B), and low degrees of partial melting promotes the formation of melts with distinct fertile characteristics (Figure 5.69). Whilst compositions comparable to Chillimanzi and Razi suite granites can form melting of both primitive (sodic) and intermediate (potassic) TTGs, regions with a prevalence of intermediate TTGs may prove fertile for future exploration for Li pegmatites.

5.3.8 Conclusions

The Zimbabwe craton preserves an extensive and complicated history spanning the Eo- to Neoproterozoic. Detailed geochronology, geochemistry, and modelling suggests that the earliest TTGs of the Tokwe and Rhodesdale segments likely formed through relatively shallow (<10 kbar) melting of a hydrated mafic protocrust. Intrusive TTGs of the Tokwe suite became complexly infolded with small remnants of the early Sebakwian greenstone cycle during an amphibolite-facies metamorphic event which initiated the first cycle of crustal recycling within the craton, forming intrusive sills and larger

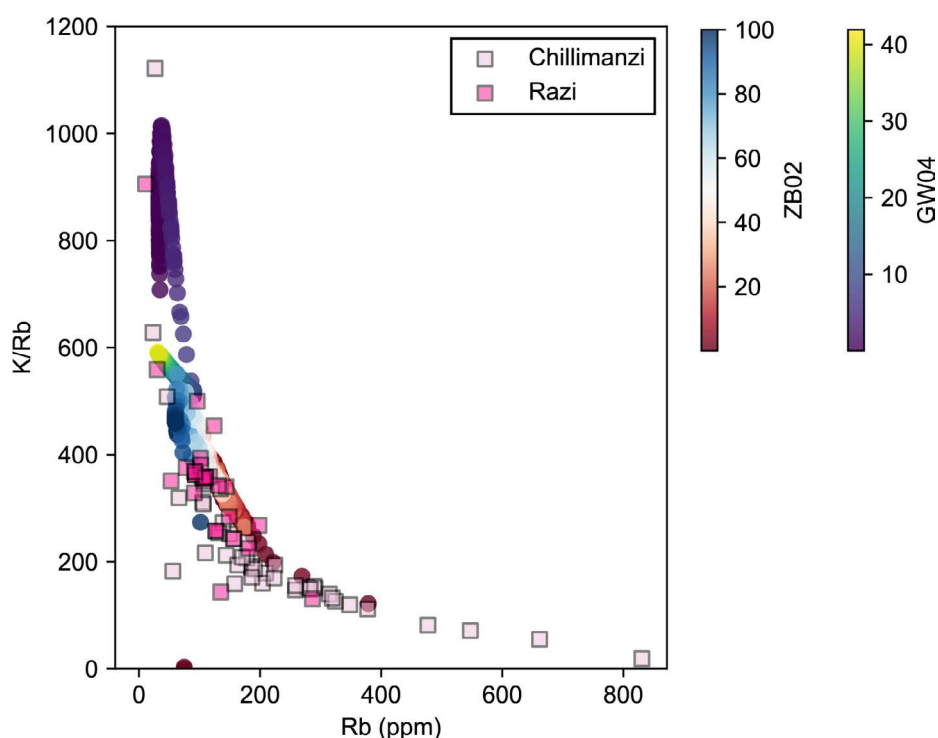


Figure 5.69: K/Rb vs. Rb plot of Chillimanzi and Razi suite granite samples, overlain by model melt compositions of melting intermediate TTG (ZB02) and primitive TTG (GW04) of the Sesombi suite in mid-lower crustal levels (8-10 kbar).

intrusive bodies (e.g., the Mushandike granite).

Following a period of quiescence and relative stability, during which the southern margin of the Tokwe segment developed a passive margin, deeper-seated melting of mafic crust (forming the Chingezi TTG suite) was coupled with extensive bimodal volcanism in the first major generation of greenstone belts (Belingwean) in the craton. This may have been driven by the amalgamation of the Tokwe and Rhodesdale terranes through underthrusting or flat-slab subduction underneath the Tokwe terrane, generating the Sebakwe protocraton during a regional metamorphic event at c. 2.86 Ga (Figure 5.62).

Plume impingement underneath the protocraton first generated felsic volcanism and the initiation of extension, prior to development of a major (c. 250 000 km²) large igneous province (Bulawayan) which covered the majority of the central Zimbabwe craton. Coupled with the extensive volcanism was the emplacement of the Mashaba-Chibi dyke swarm and ultramafic sills as intrusive feeders for the flood-basalt type volcanism under shallow submarine conditions. Heating of the lower crust generated abundant melting of underplated basalts and relatively mafic lower crust, generating the Sesombi TTG suite. Through assimilation and mixing, earlier formed TTG suites were incorporated into the primitive TTG melts to form a suite of intermediate (elevated K/Na) TTGs (Figure 5.62).

Immature and mature volcanic arcs and oceanic plateaux accreted onto the western margin of the Sebakwe protocraton during a major accretionary event c. 2.68–2.61 Ga, which formed the spatially extensive and complexly deformed Western Succession greenstone belts. Coupled with this orogenic event was the doming of TTGs, generating the dome-and-keel style structures seen in the craton today, emplacement of the major Chillimanzi and Razi suite granites, and the thrusting of the deep-crustal Northern Marginal Zone onto the craton (Figure 5.62). Final craton stabilisation was marked by the emplacement of the Great Dyke at c. 2574 Ma during a major delamination event and coincided with continued crustal melting and granitic magmatism. By 2.5 Ga, the Zimbabwe craton became a fully stable cratonic block.

The Zimbabwe craton presents an intriguing case study for the study of LCT pegmatite petrogenesis, as the previously assumed melting protoliths (metasedimentary rocks) are scarce, and therefore the craton presents a markedly different example to Archean pegmatite formation when compared to well-studied cratons such as the Superior. In the Zimbabwe craton, repeated recycling of the early formed crust during repeated crustal melting events likely provided the necessary enrichment of lithium, where models suggest that three to four crustal melting/recycling events is sufficient

to generate the concentrations required to form potentially economic LCT type pegmatites. This can also explain an Archean enigma: lithium-rich pegmatites only appear towards the end of cratonisation in any Archean craton as only at this point has the crust been sufficiently recycled to generate the required lithium enrichment.

Chapter 6

Conclusions

6.1 Thesis conclusions

This thesis explored the source and structural controls of lithium-rich, LCT/Group 1 type pegmatites, and integrated this understanding into the broader context of the evolution of the Zimbabwe craton. First, I conducted a modelling based study to test whether the two main formation models of lithium pegmatites can generate the necessary lithium enrichment (>5000 ppm Li for spodumene saturation) to form unzoned, albite-spodumene, pegmatites which characterise the majority of economic deposits (Chapter 3). Next, I presented a structural case study on the Zulu pegmatite field to determine the relative timing of emplacement and potential structural controls on the distribution of the pegmatite field (Chapter 4). Finally, I integrated new geochronological and geochemical data from the Zimbabwe craton to propose a refined holistic model of the evolution of the craton during the Archean (Chapter 5).

6.1.1 Modelling of pegmatite sources

In Chapter 3, I showed through thermodynamic modelling that the lithium content of a melt generated by anatexis of metasedimentary protoliths is primarily controlled by the breakdown of biotite. When comparing a typical pelite and greywacke as protoliths, greywackes were found to be more fertile given lower mica content. Direct anatexis of metasedimentary rocks was found to have a maximum enrichment factor of c. 4.3. Flux melting (by introduction of external fluids to the system), which has historically been considered a potential source mechanism for pegmatite formation (e.g., Müller et al., 2017), served to depress lithium enrichment in the melt due to persistent mica stability. Given sedimentary rocks with lithium contents >1000 ppm

are exceptionally rare, it is unlikely that direct anatexis of metasedimentary protoliths is a viable mechanism to generate economic albite-spodumene pegmatites.

By comparison, fractional crystallisation was simulated to test the granite fractionation model. Fractionation can enrich lithium in a granitic melt by a factor ≤ 6 at 90% fractionation (above the generally accepted particle locking threshold of Vigneresse et al., 1996, at which separation of crystals from remaining melt becomes untenable). Therefore elevated initial lithium contents of c. 600 ppm (a melt that formed a result of direct anatexis of enriched metasedimentary rocks tested above) generated melts with a maximum lithium of up to 3400 ppm, still below the minimum saturation point for spodumene.

Forming economic albite-spodumene pegmatites through either generally accepted mechanism is therefore considered unlikely. In this study, I proposed a new petrogenetic model, whereby an initial melting event of metasedimentary rocks generated an early generation of granites. Subsequent reheating of the system, either within the same or during a later orogenic event, promotes partial melting of the earlier granitic body, which modelling suggest is an extremely efficient mechanism at concentrating lithium. The resultant melts readily pass the required 5000 ppm threshold, thereby providing a viable mechanism to generate spodumene saturation at initial crystallisation of a pegmatitic melt.

This model accounts for two features observed in several pegmatite fields: 1) economic pegmatites are commonly unzoned, and therefore spodumene saturation must be achieved at the onset of crystallisation, and 2) pegmatites are often significantly younger than spatially related granites (e.g., Brou et al., 2022; Goodenough et al., 2014). It should be noted that the classic fractionation and direct anatectic models can potentially explain the genesis of some (zoned) deposits worldwide.

6.1.2 Structural controls on economic deposits

In chapter 4, I presented a field-based case study on the Zulu pegmatite field in the Zimbabwe craton, within which two main pegmatite types were delineated.

Type 1 pegmatites were interpreted to have been emplaced synkinematic to transcurrent shearing (D_2) along the Irisvale–Lancaster and Zulu shear zones. This was highlighted by internal deformation within the pegmatites, their lenticular shape subparallel to the main shear fabrics, and the synkinematic growth of exomorphic alteration phases surrounding the pegmatite bodies. The significant rheological contrast between the weaker serpentinite hanging wall and amphibolitic footwall likely acted as a concentrating mechanism to generate large individual pegmatite intrusions.

Type 2 pegmatites, on the other hand, primarily preserved magmatic textures and were emplaced obliquely to the main deformation fabric. Type 2 pegmatites were therefore interpreted to be syn-to-late kinematic to D_2 , and emplaced along typical conjugate strike-slip fracture sets and tension gashes.

The study also highlighted that the rheology of the country rock has a strong control on the potential dimensions of individual pegmatite bodies, which provided an explanation for why pegmatites within the Fort Rixon–Shangani greenstone belt are larger than those in the surrounding (more rheologically competent) granitoid basement.

By comparing the Zulu pegmatite field to other major pegmatite fields globally, shear zones are shown to be a critical component in developing large economic deposits, and can enable migration of pegmatitic melt far from the source region. Constraining the structural history surrounding a pegmatite field is therefore shown to be a highly effective method to identify new greenfield and brownfield targets.

6.1.3 Tectonic evolution of the Zimbabwe craton

Chapter 5 presented an integrated geochemical, geochronological, and petrological modelling approach to developing a holistic tectonic evolution model of the Zimbabwe craton and summarised below.

Early cratonic nuclei (Tokwe and Rhodesdale terranes) formed through relatively shallow (<10 kbar) partial melting of an early mafic protocrust as a result of an elevated geothermal gradient. This likely coincided with the formation of the first extrusive ultramafic volcanic succession (the Sebakwian group). Early (felsic) crustal recycling is evidenced by the formation of the c. 3300–3375 Ma Mont D’Or and Mushandike granitoids, as well as the associated intrusive sills found cross-cutting the intrusive basement.

Following a period of relative quiescence and limited continental emergence (forming shelf successions of the Buhwa greenstone belt), A second major TTG-forming event (the Chingezi suite, 3.0–2.85 Ga) possibly resulted from the accretion of the Tokwe and Rhodesdale terranes to form the Sebakwe protocraton, and drove andesitic extrusive volcanism to form the Belingwean group along the western and northern margin of the Tokwe terrane. Crustal thickening during this period is suggested from trace element signatures which overlap with models of deeper (>10 kbar) melting of mafic protocrust and coincided with a major metamorphic event and ductile shearing.

The impingement of a plume at the base of the protocraton initiated a period of major extrusive volcanism, generating the Bulawayan large igneous province that

typifies the majority of central cratonic greenstone belts between 2.75–2.71 Ga. Associated melting of the (still predominantly) mafic lower crust formed the Sesombi TTG suite, though earlier felsic crustal melts were locally mixed to generate a series of more potassic ($K/Na \approx 1$) intermediate TTGs within this suite.

During a major tectonic event c. 2.68–2.61 Ga, a series of volcanic arcs and oceanic plateaux are proposed to have accreted onto the Sebakwe protocraton along its western margin to form the Western Successions. This generated a mixture of more primitive TTG type magmatism (the Wedza suite), and extensive crustal recycling to form the Chillimanzi and Razi suite by melting both sodic (generating metaluminous granites) and intermediate (generating peraluminous granites) TTGs formed previously. Major shear zones developed, and the Northern Marginal Zone – a deep crustal block of the Zimbabwe craton – was thrust onto the central craton during this event. The Shamvaian group, the major siliciclastic supracrustal succession in the Zimbabwe craton, was deposited into major syntectonic basins during this episode, marking extensive continental emergence.

Final craton stabilisation was marked by the emplacement of the Great Dyke at 2574 Ma, which likely formed by melting of a depleted mantle source during delamination of part of the lower lithosphere underneath the craton. Late state granitic melts, probably formed by coincident elevated geotherms melting the lower crust, continued to be emplaced until c. 2.51 Ga.

6.1.4 Implications for lithium enrichment

The proposed tectonic model for Zimbabwe craton highlights its unique position in testing how LCT/Group 1 pegmatites form, as siliciclastic sedimentary rocks are unrealistic source rocks given their concurrent formation with pegmatites. In this way, the lithium-rich pegmatites in the craton are good evidence to indicate that alternative formation mechanisms must exist.

The tectonic model of the Zimbabwe craton highlights the episodic growth of the craton, and indicates felsic crustal recycling initiated early in the cratonic history. Modelling of these recycling events suggest that three to four crustal melting events is sufficient to generate highly peraluminous melts with economic concentrations of lithium. Importantly, this implies that recycling, which can either occur through erosion and sedimentation, or repeated partial melting, is the key factor required to generate lithium-rich melts, obviating the assumed requirement of sedimentary material in the source.

Put differently, the presence of abundant hard rock lithium deposits in the Archean suggests that the enrichment mechanisms driving lithium enrichment must have been active early in Earth history. If sedimentary recycling is a necessary component for LCT/Group 1 pegmatites, we should expect to see major pegmatite deposits later in Earth history as sedimentary processes become more effective (e.g., Bucholz, 2022; Bucholz and Hernández-Montenegro, 2025). The propensity for major hard rock lithium deposits to be Archean in age suggests partial melting, rather than sedimentary recycling, is a more efficient lithium enrichment mechanism. Only in post-Archean terranes which rework significant components of Archean crust, such as the Dete–Kamativi inlier in Zimbabwe (Shaw et al., 2022) or the Kibaran belt (De-waele et al., 2016) during later orogenic events do we see significant lithium deposits post-Archean.

6.1.5 Implications for exploration

The conclusions made above have several key implications for exploration for new lithium deposits. First, desktop studies should not focus on constraining potential sources, as regional structures place a more significant control on the distribution pegmatite fields. Second, transcurrent shear zones are very effective at extracting, vertically migrating, and subsequently emplacing pegmatitic melts, and should therefore be targeted. Finally, Archean terranes, or terranes which significantly rework Archean protoliths, are most likely to host major lithium deposits.

6.1.6 Summary

In summary, this thesis has shown: 1) Whilst pegmatites can form through both anatexis and granitic fractionation, a simple single-stage model is unlikely to account for the necessary enrichment required to generate economic hard rock lithium deposits, 2) The Zimbabwe craton grew through episodic periods of growth through various geodynamic mechanisms, and repeated melting events provided a mechanism by which to progressively enrich a source region in lithium through the formation of intermediate-stage TTGs, 3) The absence of siliciclastic sedimentary sources for lithium pegmatites in Zimbabwe, and the feasibility of a melt-driven recycling mechanism implies sedimentary sources are not an essential component of LCT/Group 1 pegmatites, and 4) From an exploration perspective, a focus should therefore be put on delineating structural corridors across Archean or reworked Archean terranes which may have enabled pegmatitic melt migration and emplacement.

6.2 Future Work

The results of this thesis present several exciting avenues for further research, some of which are highlighted below.

6.2.1 Archean evolution of Zimbabwe

Whilst Chapter 5 presents a reasonable attempt at synthesizing the complex geological history of the Zimbabwe craton, several questions remain and would benefit from further work:

- The recognition of Archean protoliths in the Dete–Kamativi inlier (Glynn et al., 2020) presents an interesting opportunity to test the accretion model for the Western Succession. If the Western Succession is allochthonous, then the Dete–Kamativi block could present the cratonic block which, upon collision with the Sebakwe protocraton, marked the closure of an Archean ocean. Further geochronological and isotopic studies on the Dete–Kamativi block to determine whether it evolved independently to the rest of the Zimbabwe craton is highly recommended.
- Additional structural studies on the Midlands and Bulawayo greenstone belts will help better unravel the complex history of the Western Succession, and better delineate the potential geometries of accretion.
- Our understanding of the evolution of the Zimbabwe craton as a whole will benefit from significantly more coupled geochemical and geochronological analyses of samples across the craton, and extending the database towards Botswana and towards the north-east extent of the craton will help potentially delineate other early protocratonic blocks.

6.2.2 Testing source models

Since the publication of Chapter 3 (as Koopmans et al., 2024), several researchers have proposed areas in which the multi-stage mechanism appears to have formed lithium pegmatites. Additional studies are needed to further test and refine our understanding of the exact source requirements for economic lithium pegmatites.

- Modelling of lithium during anatexis will benefit from additional experimental work to better constrain the partition coefficients for lithium across $P-T$ space, and may reveal additional caveats to the presented source models.

- Further field based studies in terranes which expose a larger cross-section of the crust, in order to capture both the source and trap, are recommended to validate the proposed source models.
- Delineating source models would benefit from developing a technique by which to differentiate the various proposed processes. Whilst the source mechanisms proposed result in comparable geochemical signatures, isotopic fractionation may differ, and thus may prove a fruitful avenue for future research.

6.2.3 Refining structural model

Detailed structural studies around economic pegmatite fields remain sparse. Our understanding of the structural controls of pegmatite systems would greatly benefit from additional field-based investigations, which may reveal additional factors which will help delineate new exploration targets.

6.3 The lithium mineral system model

Whilst this thesis has contributed to a solid foundation for developing a robust mineral system model for hard rock lithium deposits, it is evident more work is needed. At present, the following simple model can be proposed: for the source, the key component appears to be a sufficiently recycled felsic crust, with or without sedimentary processes to drive recycling. Whilst precise mechanisms for how to generate the necessary lithium enrichment have been proposed, these will need to be tested and validated in field-based studies. For the migration and trap, (transcurrent) crustal-scale shear zones appear to be pivotal in initial extraction, migration, and subsequent emplacement of the pegmatitic melt to form large intrusions. Further structural studies on additional pegmatite fields may shed light on additional structural controls which promote the formation of large pegmatitic bodies. Post emplacement, mechanisms must be present to promote late-stage metasomatic fluid migration so as to minimise the effects of albitisation and greisenisation (Goodenough et al., 2025), preserving the magmatic lithium ore-forming minerals for subsequent extraction. The exact controls on what promotes fluid expulsion remains elusive (Sweetapple et al., 2024).

Finding new lithium deposits is critical to meet and supply the growing demand for lithium, therefore detailed geological research to refine the mineral system model remains vital to realize a renewable-based future.

Appendix A

List of sample locations

The following tables list and describe the samples that have been referenced in Chapter 5.

A.1 Surface samples

Sample	UTM	X	Y	Rock suite	Description
BL005	WGS84, 36S	201758	7735901	Chillimanzi	30-35% qz, 30-35% ksp (tartan twinning, perthitic), 15-20% plg, 5-10% bt, 5% mu.
BL017	WGS84, 36S	187143	7755960	Chillimanzi	30-35% qz, 20-25% ksp, 30-35% plg, 5% bt, <5% mu.
CH001	WGS84, 35S	849026	7778164	Chillimanzi	30-35% qz, 25-30% kps (perthitic), 20-25% plg, 5-10% bt, <5% mu.
DP01	WGS84, 35S	754025	7778901	Chillimanzi	30-35% qz (undulose extinction), 25-30% ksp (partly sericitised), 20-25% plg (sericitised), 5-10% bt (fabric defining), 5% mu.
GH002	WGS84, 35S	805135	7819670	Chillimanzi	40-45% qz, 15-20% ksp (partly sericitised), 25-30% plg (partly sericitised), 5-15% bt (fabric defining), <5% chl (replacing bt).
GH003	WGS84, 35S	807590	7821208	Chillimanzi	30-35% qz, 25-30% ksp (perthitic), 20-25% plg, 5-10% bt, <5% tit, <5% opq.
GW001	WGS84, 35S	801104	7811636	Sesombi	<5% qz, 25-30% plg (partly sericitised), 65-70% amph.
GW002	WGS84, 35S	801104	7811636	Sesombi	25-30% qz, 30-35% plg (partly sericitised), 20% amph (fabric defining), 15-20% bt (fabric defining), <5% tit, <5% zrc.
GW003	WGS84, 35S	801104	7811636	Chillimanzi	40-45% qz, 35-40% plg (partly sericitised), 5-10% bt (fabric defining), 5% opq, <5% grt.
GW004	WGS84, 35S	801104	7811636	Sesombi	30-35% qz, 35-40% plg (partly sericitised), 15-20% bt, 5% tit, <5% opq,
GW006	WGS84, 35S	801104	7811636	Sesombi	30-35% qz, 40-45% plg (partly sericitised), 10-15% amph, 5-10% bt,
LP002	WGS84, 36S	285702	7748398	Chillimanzi	30-35% qz, 25-30% ksp, 25-30% plg, 5% bt, <5% opq.
LP003	WGS84, 36S	289653	7736345	Chillimanzi	30% qz, 25-30% ksp, 20-25% plg, 10-15% bt, <5% opq.
LP004	WGS84, 36S	289542	7736329	Razi	30% qz, 20% ksp (tartan twinning), 15-20% plg (partly sericitised), 15-20% bt (fabric defining), 10-15% amph (fabric defining), 5% tit.
LP010	WGS84, 36S	3144489	7718183	NMZ	20% qz, 15% ksp (partly sericitised), 30-40% plg (partly sericitised), 5-10% bt (fabric defining), 5% grt.
LP011	WGS84, 36S	314501	7718207	NMZ	30% qz, 5-10% ksp (partly sericitised), 25-30% plg (partly sericitised), 25-30% bt (fabric defining, partly altered to chl+mt), 5% grt, <5% opq.
MA001 A+B	WGS84, 36S	218729	7783769	Chingezi	35-40% qz, 5% ksp (occasionally phenocrystic), 40-45% plg (sericitised), <5% bt, <5% chl (replacment) <5% opq.
MO01	WGS84, 36S	403329	7785110	Bulawayan	80-85% qz, <5% plg, 5-10% bt, 5% amph.
MO03	WGS84, 36S	403329	7785110	Bulawayan	70-75% cpx, 10-15% plg, 10-15% amph.
MO04	WGS84, 36S	403329	7785120	Bulawayan	<5% qz, 20-25% plg, 55-60% trm, 10-15% chl, 5% chr, 5% opq.
MO05	WGS84, 36S	403329	7785137	Bulawayan	5-10% qz, 60-65% clc, 25-30% chl, 5-10% opq.
MU001 A+B	WGS84, 35S	886715	7797507	Tokwe	20-25% qz, 50-55% plg, 10-15% bt, 5% opq.
MU003	WGS84, 35S	886715	7797507	Tokwe	35-40% qz (some polycrystalline), 40-45% plg (sericitised, some porphyroblastic), 10-15% bt (fabric defining), <5% gt, <5% opq.
NG01	WGS84, 35S	743951	7792632	Chillimanzi	40% qz, 20-25% ksp, 25-30% plg (partly sericitised), 5% bt (partly chloritised), 5% mu, <5% opq.
NG02	WGS84, 35S	751515	7784631	Chillimanzi	35-40% qz, 20-25% ksp, 30-35% plg (partly sericitised), <5% bt (partly chloritised), 5% mu, <5% opq.
NM002	WGS84, 35S	799792	7819146	Chillimanzi	25-30% qz, 20-25% ksp (partly sericitised, oscillatory zoning preserved), 30-35% plg (partly sericitised), 15-20% bt (defining weak fabric), <5% chl (replacing bt), <5% opq.
NM003A+B	WGS84, 35S	799783	7819130	Amphibolite	5% qz, 10% plg, 75% amph (fabric defining), 5% altered grt (to opq+amp+qz).
NY01	WGS84, 36S	379519	7870697	Chillimanzi	25-30% qz, 30% ksp (partly sericitised), 25-30% plg (sericitised), 5% chlorite, <5% bt (relict), <5% opq.
NY02	WGS84, 36S	380107	7865946	Chillimanzi	35-40% qz, 30% ksp (tartan twinning), 20-25% plg (partly sericitised), 5% bt, <5% mu, <5% opq.
NY03	WGS84, 36S	376339	7877660	Chillimanzi	25-30% qz, 20-25% ksp (oscillatory zoning, partly sericitised), 15-20% plg (partly sericitised), 5-10% bt (fabric defining), 5% mu (fabric defining), 5% opq.
PG01	WGS84, 35S	765196	7778059	Chillimanzi	30-35% qz, 30-40% ksp (megacrystic (<5cm), tartan twinning, oscillatory zoning), 15-20% plg, 5% bt, 5% mu, <5% opq.
SC003 A+B+C	WGS84, 35S	802312	7794971	Sesombi	20-25% qz, 35-40% plg, 20-25% amph (fabric defining), 10-15% bt (fabric defining), <5% opq.
SC004	WGS84, 35S	802312	7794971	Chillimanzi	30-35% qz, 25-30% ksp (sericitised), 30-35% plg (sericitised), <5% bt.
SG01	WGS84, 36S	201881	7759606	Chingezi	30-35% qz, <5% ksp, 15-20% plg, 20-30% bt (fabric defining), 5% chl, 5% zrc, <5% opq.
SG02	WGS84, 36S	201881	7759606	Chingezi	30-35% qz, <5% ksp, 25-30% plg, 15-20% bt (fabric defining), 5% mu, <5% tit, 5% opq.
SR01	WGS84, 36S	394943	7881998	Wedza	30% qz, 15-20% ksp, 30-35% plg, 10% bt (fabric defining), 5% amph, <5% chl, <5% opq.
SR02	WGS84, 36S	395955	7881592	Wedza	30% qz, 10-15% ksp, 25-30% plg, 20-25% bt (fabric defining), <5% chl, 5% tit, 5% opq.
TR002	WGS84, 35S	856534	7781890	Tokwe	25-30% qz, 20-25% plg (partly sericitised), 40-45% amph (fabric defining), <5% opq.
TR004	WGS84, 35S	856534	7781890	Tokwe	30-35% qz, 25-30% plg (partly sericitised), 35-40% amph (fabric defining), <5% opq.
TR006	WGS84, 35S	856534	7781890	Tokwe	30-35% qz, 45-50% plg (partly sericitised), 10-15% bt (fabric defining), 5% chl, <5% opq.
TR007	WGS84, 35S	856505	7781948	Tokwe	20-25% qz, 5% ksp, 25-30% plg, 35-40% mu, 5-10% gt, <5% opq.

Continued on next page

Sample	UTM	X	Y	Rock suite	Description
TR008	WGS84, 35S	856505	7781948	Tokwe	25-30% qz, 25-30% plg (partly sericitised), 35-40% amph (fabric defining), <5% opq.
WG01	WGS84, 36S	377196	7871516	Amphibolite	90% amph (fabric defining), 5% opq, <5% plg, <5% qz.
WG02	WGS84, 36S	377196	7871516	Wedza	25-30% qz, 5-15% ksp, 25-30% plg (partly sericitised), 20% amph, 5-10% bt, <5% opq.
WG03	WGS84, 36S	377203	7871429	Wedza	25-30% qz, 5-10% ksp, 25-30% plg (partly sericitised), 5-10% bt, 20-25% chl (fabric defining), <5% opq.
ZB01	WGS84, 35S	756648	7774436	Sesombi	35-40% qz (some polycrystalline), <5% ksp, 30-35% plg (partly sericitised), 10-15% bt, 5-10% opq.
ZB02	WGS84, 35S	757278	7772204	Sesombi	40-45% qz (some polycrystalline), <5% ksp, 35-40% plg (partly sericitised), 5-10% bt (fabric defining), 5% opq.
ZB03	WGS84, 35S	752669	7770703	Sesombi	15-20% qz, 25-30% plg, 30-35% amph (fabric defining), 15-20% chl (fabric defining).
ZB04	WGS84, 35S	757629	7771363	Sesombi	30-35% qz, 35-40% plg (partly sericitised), 15-20% bt, 5% tit, <5% opq.
ZB05	WGS84, 35S	753762	7767205	Chillimanzi	30-35% qz (undulose extinction), 25-30% ksp (partly sericitised, megacrystic), 20-25% plg (sericitised), 5-10% bt (fabric defining), <5% chl (replacing bt).
ZB06	WGS84, 35S	747776	7769137	Sesombi	35-40% qz (undulose extinction), <5% ksp (partly sericitised), 25-30% plg (partly sericitised), 10-15% bt (fabric defining), 5-10% chl (fabric defining, replacing bt), <5% opq.
ZB07	WGS84, 35S	747441	7770096	Chillimanzi	40-45% qz, 20-25% ksp (partly sericitised, perthitic), 15-20% plg (partly sericitised), 10% bt (weak fabric defining), 5% mu (weak fabric defining).
ZL01	WGS84, 35S	751318	7780779	Belingwean	20-25% qz, 55-60% plg, 5-10% bt (fabric defining), 5-10% amph (fabric defining), <5% opq.
ZL09	WGS84, 35S	751007	7784869	Bulawayan	<5% qz, 10-15% plg, 70-75% trm, <5% chl, <5% chr.
ZL11	WGS84, 35S	750368	7782271	Bulawayan	<5% qz, 15-20% plg, 65-70% trm, 5%chl, <5% opq.
ZL15	WGS84, 35S	751065	7783893	Bulawayan	10-15% qz, 30-35% plg, 5-10% chl, 35-40% talc, <5% opq.
ZL16	WGS84, 35S	750515	7784006	Bulawayan	<5% qz, 15-20% plg, 50-55% trm, 20-25% chl, <5% opq.
ZL20	WGS84, 35S	751072	7782166	Bulawayan	<5% qz, 15-20% plg, 60-65% trm, 10-15% chl (defining shear bands), <5% chr.

A.2 Drill core samples

Core samples are collected from core drilled by Premier African Minerals around their Zulu project.

Sample	Rock suite	Core	UTM	Hole X	Hole Y	Depth	Description
ZLC06	Chillimanzi	ZDD18	WGS84, 35S	751222	7784390	100.68 - 100.87	30-35% qz, 25-30%ksp (tartan twinning), 20-25% plg (partly sericitised), 10-15% bt (partially replaced with clay), 5% mu.
ZLC07	Chillimanzi	ZDD18	WGS84, 35S	751222	7784390	102.17 - 102.33	25-30% qz, 30-35% ksp (tartan twinning, perthitic), 25-30% plg (partly sericitised), <5% bt, 5% mu.
ZLC15	Belingwean	ZDD64	WGS84, 35S	751792	7783148	63.63 - 63.80	35-40% qz, 30-35% plg (partly sericitised), 15-20% amph (fabric defining), <5% bt (fabric defining), 5-10% tit.
ZLC17/15	Belingwean	ZDD64	WGS84, 35S	751792	7783148	85.20 - 85.40	20-25% qz, 30-35% plg (partly sericitised), 20-25% amph (fabric defining), 15-20% tit. Melt lens in ZLC15.
ZLC20	Belingwean	ZDD64	WGS84, 35S	751792	7783148	106.57 - 106.67	35-40% qz, 30-35% plg (partly sericitised), 20-25% amph (fabric defining), 5-10% tit.
ZLC32	Bulawayan	ZDD142	WGS84, 35S	751346	7784116	159.10 - 159.23	5% qz, 20-25% plg, 65-70% trm (fabric defining), 5% tit.
ZLD01	Chillimanzi	ZDD65	WGS84, 35S	751889	7782800	175.60 - 175.80	30-35% qz, 15-20% ksp, 30-35% plg (partly sericitised), 10-15% bt, <5% mu.
ZLD02	Chillimanzi	ZDD65	WGS84, 35S	751889	7782800	130.10 - 130.25	35-40% qz, 15-20% ksp, 25-30% plg (partly sericitised), 10-15% bt (fabric defining), <5% mu.
ZLD03	Bulawayan	ZDD89	WGS84, 35S	751373	7783468	186.52 - 186.62	10% plg, 70-75% serp (atg), 10% cc, 5-10% mt.
ZLD05	Bulawayan	ZDD136	WGS84, 35S	751099	7782841	336.75 - 336.85	<5% plg, 40-45% serp, 25-30% tlc (rhombic, fabric defining), <5% trm, 5-15% mt, <5% pyr.
ZLD06	Belingwean	MDGT06	WGS84, 35S	751070	7782109	32.43 - 32.53	35-40% qz, 30-35% plg (partly sericitised), 10-15% bt (fabric defining), 20-25% chl (fabric defining), 5% gt, <5% opq. Metasomatically altered by nearby pegmatite, protolith equivalent to ZLC15
ZLD07	Belingwean	ZDD122	WGS84, 35S	751128	7782511	104.20 - 104.30	30-35% qz, 35-40% plg (partly sericitised), 15-20% amph, 5-10% bt, <5% tit.
ZLD08	Bulawayan	ZDD102	WGS84, 35S	750930	7782260	50.05 - 50.15	<5% plg, 65-70% serp (atg dominant), 5% tlc, 15-20% cc, 5-10% mt.
ZU06	Bulawayan	ZDD134	WGS84, 35S	750825	7781918	205.0 - 205.12	40-45% qz, 30-35%plg (matrix), 5% plg (phenocrystic), 10-15% bt (fabric defining), <5% chl (replacement of bt).

Appendix B

Field photographs of sample locations

The following images are of sample localities that have been referenced in this study in Chapter 5, and described in Appendix A. Where multiple samples were collected in close proximity, the sampled lithologies are labelled.



Figure B.1: Locality BL17, a quarry north of Zvishavane along the Zvishavane – Gweru road.



Figure B.2: Locality CH01, A small granite plug between Masvingo and Zvishavane.



Figure B.3: Locality DP01, sampled near the Deep Purple deposit east of the Fort Rixon – Shangani belt.



Figure B.4: Locality GH02, sampled from a small plug along a side road between Gwenoro dam and Shurugwi.



Figure B.5: Locality GH03, sampled from a small plug east of GH02.

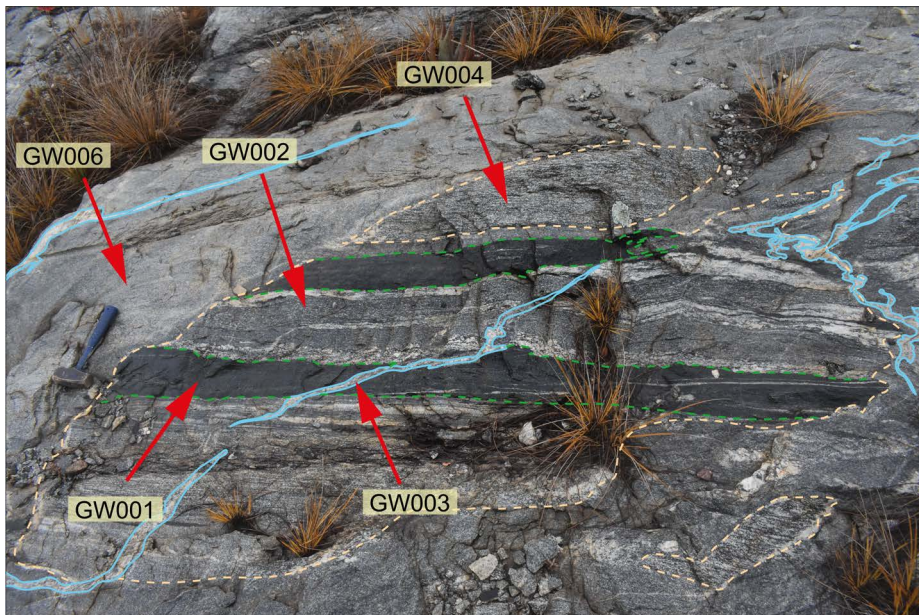


Figure B.6: Gwenoro dam (GW) locality, along the western flank of the Gwenoro Dam spillway.



Figure B.7: Locality LP02, collected south of the Masvingo greenstone belt along the road to Renco.



Figure B.8: Locality LP03, collected south of LP02.



Figure B.9: Locality LP04, collected further south of LP03 along the Renco road.



Figure B.10: Locality LP10, in dry river bed east of Renco.



Figure B.11: Locality LP11, in a dry river bed east of Renco.



Figure B.12: Trench from which MO01-5 were collected, trenched by Li3 in the Odzi greenstone belt.



Figure B.13: Locality MU01, a dry river bed c. 20 km north of Masvingo along the Masvingo – Mvume road.



Figure B.14: Locality MU03, a dry river bed c. 20 km north of Masvingo along the Masvingo – Mvume road.



Figure B.15: Locality NG01, an exposure of the Nalatale granite along the Perlycross road north of Fort Rixon.



Figure B.16: Locality NG02, an exposure of the Nalatale granite in a quarry north of the Zulu mine.

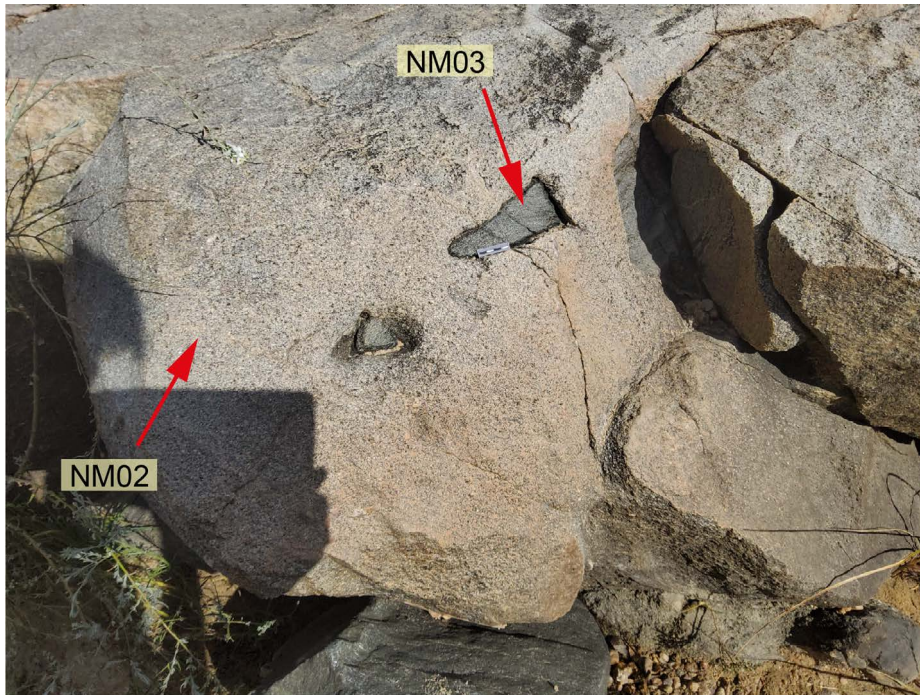


Figure B.17: Locality of NM samples, along the Nyamaziva river at the norther end of Gwenero dam.



Figure B.18: Locality NY01, a granitic plug south of Nyashanu.



Figure B.19: Locality NY02, a granitic plug north of Nyashanu.



Figure B.20: Locality NY03, a granitic plug north of Nyashanu.



Figure B.21: Locality PG01, a large granite koppie east of the Fort Rixon – Shangani greenstone belt.



Figure B.22: Locality SC03, a dry river bed along the Mapongokwe river south of Gwenoro dam.

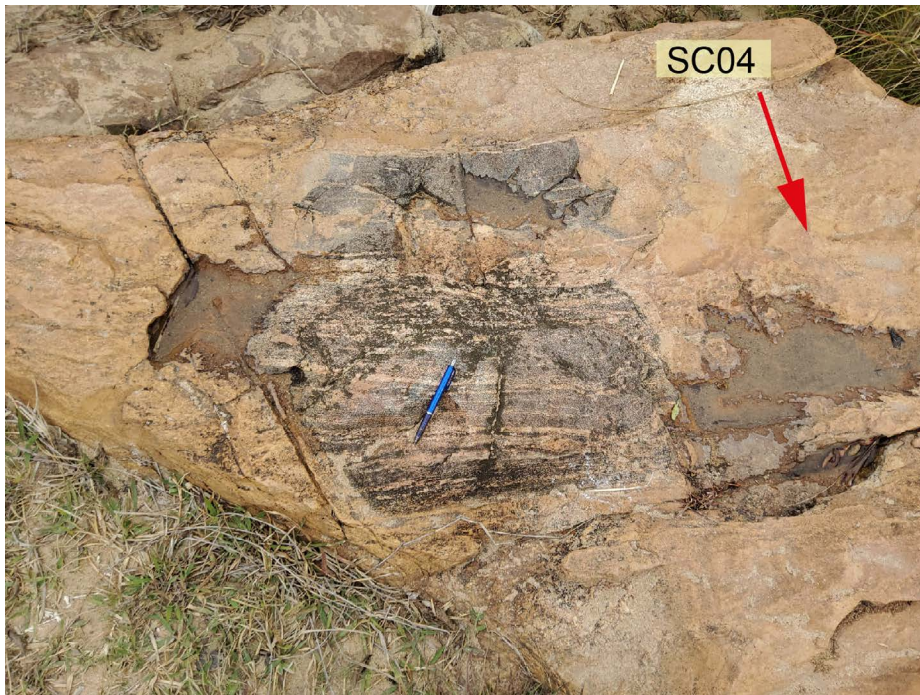


Figure B.23: Locality SC04, a dry river bed along the Mapongokwe river, west of SC03.

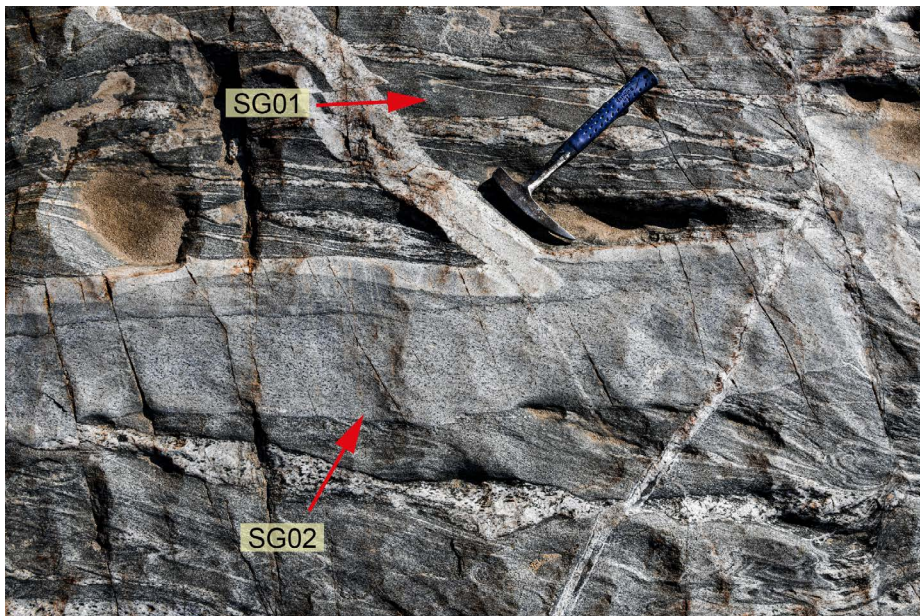


Figure B.24: Locality for SG samples, A dried river bed of the Runde river north of the Zvishavane – Masvingo road.



Figure B.25: Locality SR01, an outcrop along a dried river bed of the Save river where it crosscuts the Odzi greenstone belt.



Figure B.26: Locality SR02, an outcrop southeast of SR01 along a dried river bed of the Save river where it crosscuts the Odzi greenstone belt.



Figure B.27: Locality TR02, an outcrop along the Tokwe river north of the Zvishavane – Masvingo road.



Figure B.28: Locality TR04, sampled near TR02.



Figure B.29: Locality TR06, sampled near TR02.



Figure B.30: Locality TR07, sampled near TR02.



Figure B.31: Locality TR08, sampled near TR02.

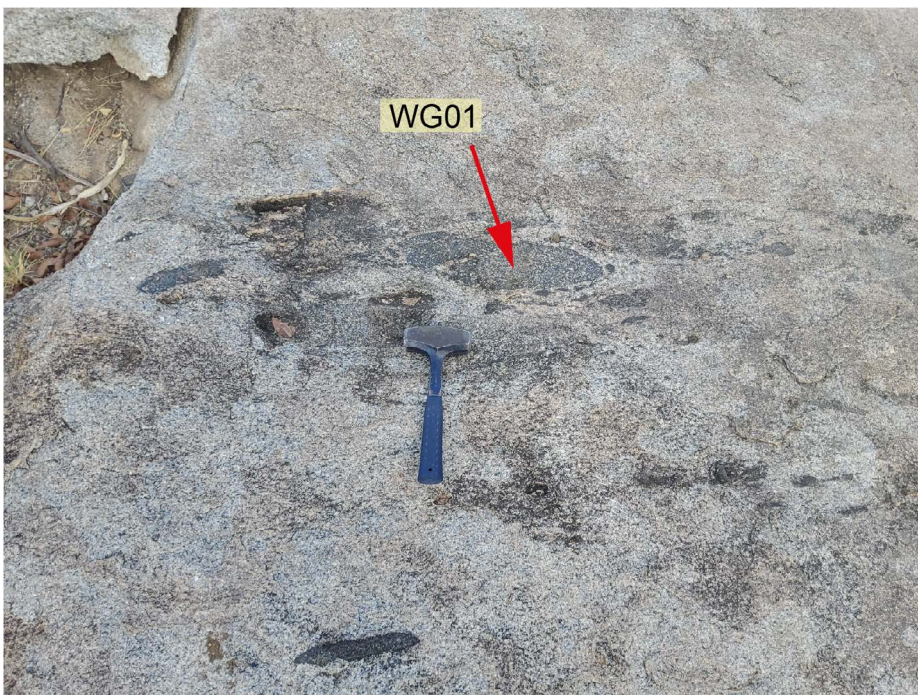


Figure B.32: Locality WG01, an outcrop along the bank of the Murove river.



Figure B.33: Locality WG02, sampled near WG01.



Figure B.34: Locality WG03, sampled 100 m south of WG01 along the Murove river.



Figure B.35: Locality ZB01, an outcropping hill east of the Fort Rixon – Shangani greenstone belt.



Figure B.36: Locality ZB02, an outcropping hill east of the Fort Rixon – Shangani greenstone belt.



Figure B.37: Locality ZB03, a small outcrop in a river bank of the Machakwe river, east of the Fort Rixon – Shangani greenstone belt.



Figure B.38: Locality ZB04, an outcrop along the south-eastern margin of the Fort Rixon – Shangani greenstone belt.



Figure B.39: Locality ZB05, a washed outcrop in the Three Fingers ranch.

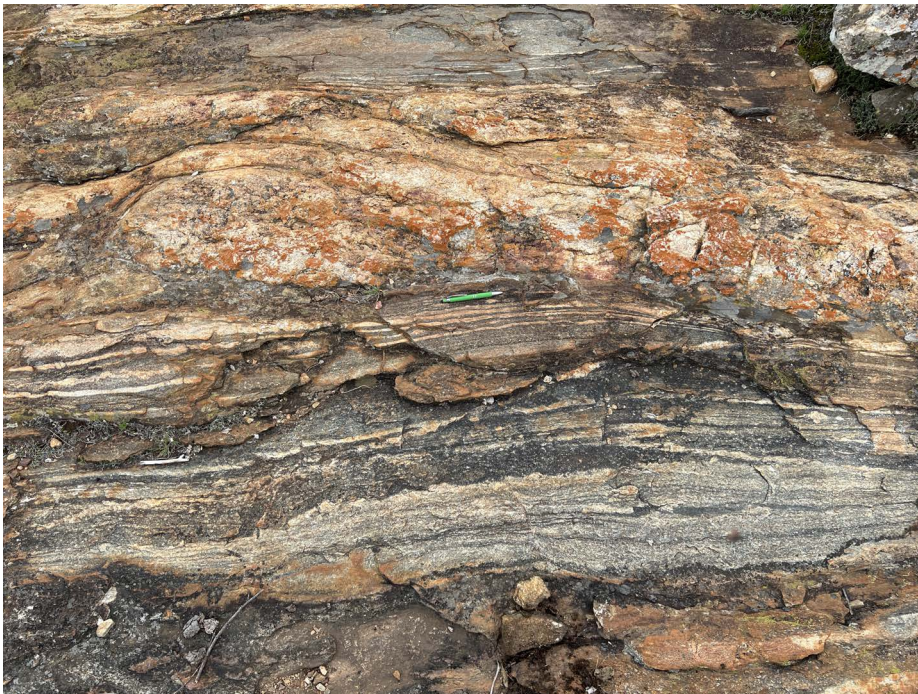


Figure B.40: Locality ZB06, a washed outcrop along the eastern margin of the Fort Rixon – Shangani greenstone belt.



Figure B.41: Locality ZB07, an outcrop 5 km west of the Three Fingers ranch..



Figure B.42: Locality ZL01, a small outcrop within the Machakwe river south of the Zulu mine.



Figure B.43: Locality ZL09, a row of outcrops west of the Zulu mine.



Figure B.44: Locality ZL11, a row of outcrops west of the Zulu mine.



Figure B.45: Locality ZL15, a small patchy outcrop north of the Zulu mine, west of the dam road.



Figure B.46: Locality ZL16, a row of outcrops west of the Zulu mine offices.



Figure B.47: Locality ZL20, a row of outcrops north of the Zulu mine offices.

Appendix C

Geochronological data

C.1 Zircon U-Pb data

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Sample	Spot name	Run	206/238	±2SE	207/235	±2SE	207/206	±2SE	238/P206	±2SE	206/238 Age	±2SE	207/235 Age	±2SE	207/206 Age	±2SE	U/Th	U	Th	Pb
BL05	1	4	0.07177	0.002083	2.351663	0.112161	0.236224	0.008098	13.944911	0.414437	446.8	12.5	1224.1	33.5	3087	54	2.5	3548	1409	1551
	2	4	0.184585	0.007148	5.16113	0.183013	0.202555	0.00412	5.456537	0.220001	1091.3	39	1843.9	30.3	2845.6	33.3	2.63	1585	593	558
	3	4	0.398949	0.016941	14.712965	0.514466	0.269873	0.006959	2.509665	0.106684	2163.4	78	2796.1	33.3	3303.9	40.2	1.22	623	499	794
	4	4	0.195085	0.006423	4.674774	0.132906	0.172814	0.0037	5.138679	0.169799	1148.5	34.6	1762	23.5	2583.3	36	2.17	1089	500	449
	5	4	0.224867	0.007015	5.784036	0.1709	0.186426	0.004096	4.451857	0.134215	1307.2	36.8	1943	25.8	2708.9	35.8	1.86	1089	574	530
	6	4	0.329597	0.01114	10.412337	0.326099	0.229112	0.004961	3.041418	0.099955	1835.4	53.8	2470.6	28.9	3044.1	34.4	1.65	924	548	762
	7	4	0.633414	0.017039	23.812648	0.643508	0.272469	0.005455	1.576269	0.043334	3162.6	67.5	3260.1	26.6	3319.2	31.3	1.42	242	167	311
	8	4	0.623955	0.01727	23.527376	0.709447	0.272334	0.005687	1.600495	0.043558	3125	68.3	3247.6	28.7	3318.1	32.8	1.46	223	150	287
	9	4	0.655101	0.019479	24.739397	0.74198	0.275059	0.006301	1.525147	0.04616	3247.1	76.1	3296.9	29.3	3333.3	36.2	1.43	123	85	148
	10	4	0.238478	0.009673	7.321275	0.335262	0.222679	0.004897	4.222979	0.173697	1377.6	50.4	2145.8	40.9	2998.3	35	2.39	931	395	489
	11	4	0.347208	0.010375	11.167332	0.400983	0.237212	0.005781	2.887283	0.08721	1920.8	49.7	2534.7	33.8	3098.8	39	1.54	328	218	344
	13	4	0.484458	0.016269	17.607646	0.617011	0.272098	0.006821	2.07193	0.071147	2545.8	70.9	2967.1	34.3	3316.3	39.3	1.94	238	125	265
	14	4	0.67806	0.018686	24.442413	0.626442	0.269525	0.005575	1.4794	0.041512	3336.2	72	3285.9	25	3301.9	32.4	1.83	343	192	372
	15	4	0.278158	0.008277	6.867361	0.242264	0.182593	0.003929	3.609192	0.109428	1581.8	41.8	2092.7	30.8	2675	35.8	2.91	1014	391	324
	16	4	0.174652	0.006054	6.719375	0.226496	0.285917	0.006261	5.750149	0.199825	1037.5	33.2	2074.4	29.7	3394.4	33.9	2	2296	1147	960
	17	4	0.285075	0.008652	7.700236	0.218811	0.20173	0.004182	3.524834	0.108623	1616.4	43.4	2195.8	25.3	2838.8	33.7	1.94	871	459	549
	18	4	0.309789	0.009106	8.823579	0.294415	0.212067	0.004523	3.240661	0.098928	1739.2	45	2317.8	31.7	2919.7	34.6	1.88	967	515	505
	19	4	0.393535	0.011102	13.000769	0.371692	0.245638	0.00519	2.547302	0.071723	2138.7	51.3	2678.8	27	3155.3	33.4	3.63	371	105	203
	20	4	0.516792	0.017982	19.231088	0.646225	0.27696	0.005869	1.926267	0.057295	2684.2	77.1	3051.9	32.8	3344.6	32.9	0.98	277	294	595
	21	4	0.318152	0.011721	10.66512	0.485041	0.244	0.006622	3.129816	0.120925	1780	57.6	2491.1	43.7	3143.8	43	1.9	1014	531	656
	22	4	0.138477	0.012139	4.652447	0.368262	0.245033	0.005492	7.758161	0.798521	832.8	69.1	1738.9	69.5	3150.9	35.4	2.05	2445	1197	1055
	23	4	0.371703	0.010651	12.344161	0.404614	0.242894	0.005776	2.671311	0.075851	2037.1	50	2629.8	31.2	3137.3	37.8	1.75	997	571	690
	24	4	0.245929	0.008859	7.279933	0.328321	0.216224	0.005029	4.053459	0.147296	1416.9	45.8	2124.2	39.9	2950.8	37.1	2.27	1343	595	697
	25	4	0.393294	0.01329	15.599196	0.463368	0.289071	0.005875	2.531307	0.082603	2137.3	61.2	2851.8	27.9	3411.6	31.8	1.91	1130	589	1418
	26	4	0.129291	0.004705	4.222125	0.134127	0.241007	0.005807	7.675939	0.30216	783.5	26.8	1677	25.9	3130.7	43.2	1.07	3903	4285	4675
	27	4	0.490678	0.018524	17.288457	0.673008	0.258871	0.005345	2.038063	0.079348	2571.4	80.5	2947.6	38.1	3238.6	32.5	1.73	669	388	624
	28	4	0.354293	0.02397	13.385563	0.79642	0.28156	0.0077	2.884452	0.200231	1949.2	114.2	2698.2	57.7	3368.6	42.1	1.36	155	118	169
	29	4	0.214932	0.006367	6.180805	0.190681	0.213146	0.00478	4.640499	0.14254	1254.8	33.9	2000.4	27.2	2927.5	36.5	1.96	895	460	591
	30	4	0.184825	0.008319	6.918807	0.303334	0.274889	0.006678	5.42925	0.261199	1092.7	45.4	2098.4	38.4	3342.4	48.7	4.69	430	94	218
CH01	1	4	0.312999	0.00939	8.463542	0.247232	0.191344	0.004079	3.183461	0.096626	1755	46.1	2281	26.7	2752.1	35	0.45	727	1575	1876
	2	4	0.388709	0.011892	9.923449	0.343912	0.181633	0.003786	2.559239	0.075436	2116.4	54.9	2426.4	31.5	2666.7	34.4	0.68	13735	19823	13589
	3	4	0.121612	0.004695	2.879279	0.129075	0.167248	0.003346	8.230363	0.3102	739.5	26.9	1372.8	33.5	2529.1	33.6	0.6	13149	22996	10842
	4	4	0.274588	0.011374	7.185476	0.242739	0.187619	0.007729	3.631117	0.152669	1563.4	57.6	2133.8	30	2715.8	65	1.41	1213	839	1159
	5	4	0.592343	0.016953	18.768837	0.557982	0.223688	0.00534	1.677499	0.047044	2998.4	68.4	3029.3	28.8	3005.6	38	2.07	299	140	227
	6	4	0.437983	0.012864	11.341858	0.31901	0.18428	0.004115	2.271949	0.067246	2340.9	57.7	2550.8	26.2	2689.7	36.7	2.04	343	163	255
	7	4	0.061298	0.001808	1.084926	0.033077	0.126276	0.003471	16.24066	0.478765	383.5	11	745.6	15.8	2042	47.5	1.24	16112	12578	6550
	8	4	0.36588	0.010589	8.887612	0.26529	0.174889	0.004128	2.718083	0.078722	2009.8	50	2326.1	27.2	2603.5	38.7	0.3	1438	4628	4409
	9	4	0.213684	0.005856	3.443189	0.104868	0.114855	0.002668	4.659344	0.129547	1248.3	31.1	1513.7	24.1	1875.5	42.2	0.22	1508	6541	4422
	10	4	0.119032	0.004651	4.638297	0.1588	0.280757	0.006784	8.443039	0.340265	724.7	26.8	1754.3	28.4	3364.5	37.7	1.45	1623	1130	1592
	11	4	0.131627	0.004353	3.635378	0.134237	0.201369	0.005118	7.649451	0.250603	797	24.8	1555.8	29.5	2834.9	41.5	1.35	10439	7957	6984
	12	4	0.082373	0.003427	3.039867	0.141545	0.269849	0.006067	12.215942	0.491822	510.2	20.4	1416.6	35.1	3304.1	35.3	2.67	12171	4721	6540
	13	4	0.106559	0.004874	3.357515	0.13537	0.228179	0.004913	9.536881	0.406835	652.3	28.3	1491.7	30.9	3037.6	34.8	0.98	2637	2793	1212
	14	4	0.525767	0.014895	15.893443	0.493871	0.218733	0.005836	1.908804	0.053628	2723.1	62.8	2869.2	30.1	2968.3	43.1	0.52	85	169	176
	15	4	0.568915	0.02087	18.045652	0.627199	0.226341	0.007417	1.764335	0.061061	2902.3	84.7	2991.2	34.1	3023.3	53.3	0.35	131	383	411
	17	4	0.203262	0.007007	5.18389	0.252241	0.183053	0.005613	4.934499	0.164919	1192.5	37.4	1846	41.6	2676.4	51.5	1.83	1790	1113	1324
	18	4	0.408701	0.011006	10.740362	0.286034	0.187911	0.004111	2.445621	0.06624	2208.7	50.4	2500.4	24.7	2722.1	35.9	0.38	240	646	653
	19	4	0.201441	0.007077	6.083441	0.189873	0.214417	0.005819	4.964774	0.175698	1182.8	38	1987.3	27	2937	43.7	1.15	696	612	789
	20	4	0.167879	0.00485	3.597313	0.102692	0.153155	0.003065	5.949304	0.178755	1000.3	26.8	1548.3	23	2380.2	34.2	0.75	12150	16634	10943
	21	4	0.377861	0.015198	9.500062	0.324642	0.184954	0.004424	2.64462	0.106455	2065.4	71.1	2386.7	31.2	2696.2	39.9	0.59	723	1191	1769

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Sample	Spot name	Run	206/238	±2SE	207/235	±2SE	207/206	±2SE	238/P206	±2SE	206/238 Age	±2SE	207/235 Age	±2SE	207/206 Age	±2SE	U/Th	U	Th	Pb	
		22	4	0.177725	0.004825	4.397518	0.119316	0.180952	0.003591	5.610899	0.150386	1054.5	26.4	1711.4	22.3	2660.5	32.7	0.65	15640	23237	17885
		23	4	0.521216	0.014918	15.659446	0.499259	0.220251	0.005379	1.91574	0.055601	2703.6	63.3	2854.4	30.6	2979.6	39.8	0.58	123	208	345
		24	4	0.220163	0.00622	5.25933	0.150624	0.174406	0.004399	4.532788	0.126872	1282.6	32.8	1861.9	24.5	2598.3	41.7	0.39	738	1884	1409
		25	4	0.447817	0.014573	12.696159	0.472	0.210071	0.006539	2.214196	0.054118	2399.5	48.6	2655.6	35.2	2902.5	50.8	0.75	85	112	176
		26	4	0.393186	0.01067	8.972322	0.238164	0.169005	0.003498	2.539389	0.068608	2137.4	49.3	2334.9	24.3	2546.6	34.8	0.26	1009	3861	3194
		27	4	0.625468	0.016272	21.693422	0.584871	0.255501	0.00518	1.596387	0.041884	3131.4	64.6	3169.5	26.3	3217.9	32.2	1.21	427	352	544
		28	4	0.517455	0.013572	12.607933	0.355887	0.178981	0.004085	1.929474	0.051075	2688.2	57.7	2650.1	26.5	2641.4	37.7	0.48	360	748	946
		29	4	0.127934	0.003548	3.562601	0.102168	0.204842	0.004549	7.809125	0.217354	776	20.3	1540.6	22.7	2863.2	35.9	1.69	512	304	245
		30	4	0.38744	0.011316	10.122177	0.297742	0.19211	0.004337	2.576769	0.075499	2110.6	52.6	2445.3	27.6	2758.7	37.2	0.33	1000	3004	3209
DP01		1	7	0.443108	0.009033	11.333301	1.151205	0.18546	0.006798	2.256421	0.044941	2364	40.2	2550.1	93.6	2701.2	60.7	0.02	2734	139109	162701
		2	7	0.435497	0.010604	10.88055	1.117458	0.181598	0.006678	2.300836	0.054742	2329.4	47.4	2511.2	94.3	2666.3	60.4	0.02	1761	95776	108636
		3	7	0.797561	0.026274	34.42554	3.702013	0.31175	0.013188	1.257075	0.044095	3778.1	95.4	3619.3	110.3	3527.1	65.7	0.01	1033	87042	74469
		4	7	1.010966	0.066673	91.863468	10.426001	0.660516	0.026711	1.009881	0.066505	4486	213.5	4593.7	114	4645.1	59.1	0.02	1605	90763	46283
		5	7	0.372607	0.009842	18.99048	1.987618	0.369713	0.015073	2.694953	0.072474	2040.5	46.3	3037.4	102	3786.4	62	0.01	642	84263	49951
		6	7	0.314966	0.017605	6.864722	0.820811	0.157334	0.006568	3.282618	0.203509	1759.4	87.1	2076	111.4	2421.7	72.6	0.01	519	47194	40558
		7	7	0.46508	0.008531	11.900902	1.205265	0.185692	0.00683	2.151923	0.03955	2461.6	37.5	2596.1	95	2705.9	54.4	0.02	1862	98346	113053
		8	7	1.661861	0.099236	158.728514	19.264501	0.68876	0.026725	0.615461	0.036618	6283.2	239.6	5135.7	124.3	4705.7	56.3	0.02	786	39302	45150
		9	7	0.44526	0.009609	12.414591	1.293941	0.202196	0.008277	2.251959	0.049265	2373.4	42.9	2632.9	98.7	2839.6	66.9	0.01	475	36496	40349
		10	7	0.333289	0.007458	7.371269	0.759751	0.160454	0.006122	3.010309	0.068035	1853.8	36.1	2155.5	92.6	2463.2	57.2	2.83	550	195	153
		11	7	0.508454	0.010489	18.147052	1.858306	0.258017	0.009679	1.971321	0.040052	2649.9	44.7	2997.1	98.5	3233.7	59.1	0.02	2339	154424	188960
		12	7	0.505872	0.013629	12.98836	1.364275	0.187316	0.007953	1.984881	0.055866	2638.2	58.7	2677.1	100.2	2716	68.6	4.88	174	36	52
		13	7	0.449081	0.009328	12.127042	1.245265	0.195447	0.007896	2.232893	0.045458	2391	41.4	2613.7	97.3	2786.7	66.6	0.01	767	62336	70749
		14	7	0.527396	0.013143	27.323713	2.980489	0.37593	0.017743	1.902486	0.048093	2730	55.6	3392.6	108.3	3812.5	70.9	0.01	836	117315	88511
		15	7	0.550347	0.018242	13.820864	1.483665	0.183557	0.00961	1.837744	0.061728	2823	75.8	2731.1	101.6	2669.2	91.2	0	22	0	0
		16	7	0.319059	0.012171	15.984419	1.676342	0.364429	0.013608	3.174165	0.132761	1783.3	60	2872.7	103.1	3767.2	56.5	0.02	1788	106630	55785
		17	7	0.443917	0.00794	11.637219	1.178354	0.190828	0.007243	2.259964	0.040666	2367.9	35.5	2575.1	94.7	2747.2	61.8	0.02	2440	134629	151595
		18	7	0.575866	0.029932	17.977074	2.159635	0.228437	0.013569	1.753454	0.088914	2928.7	121.7	2982.8	113.3	3034.5	93.5	1.97	43	21	47
		19	7	0.445525	0.011939	10.85681	1.122488	0.178322	0.006815	2.253243	0.060177	2374.9	53.2	2510.3	97.6	2636.2	63.4	0.01	890	81983	86577
		20	7	0.437492	0.008544	11.296325	1.145574	0.188089	0.006891	2.294922	0.044656	2339	38.3	2547.2	93.9	2724.4	60.5	0.02	2559	132126	146886
		21	7	0.465327	0.013021	12.856701	1.321028	0.201456	0.007495	2.166449	0.060122	2461.2	57.1	2667.2	97.8	2836.5	60.5	4.29	470	109	210
		22	7	0.501924	0.012127	14.264031	1.447248	0.207625	0.008033	1.995068	0.053081	2620.8	51.9	2769.6	106.7	2884.4	63.8	1.58	274	178	266
		23	7	0.529983	0.012529	15.23771	1.585991	0.210099	0.008508	1.897374	0.044065	2740.1	52.6	2827	97.2	2902.3	66.2	1.85	167	89	157
		24	7	0.499241	0.019134	13.867337	1.491186	0.203087	0.00789	2.033858	0.078728	2605.9	82.3	2734	102.9	2848.4	62.7	3.24	296	98	168
		25	7	0.376363	0.009537	10.248299	1.058053	0.199216	0.007618	2.673346	0.064725	2058.3	44.3	2455.2	94.6	2817.5	62.6	8.31	484	56	201
		26	7	0.380736	0.009867	9.990966	1.039852	0.191978	0.007468	2.643544	0.06785	2078.6	46	2430.9	96.2	2756.3	64.2	11.52	631	57	151
		27	7	0.4096	0.008869	11.06279	1.12376	0.197996	0.007711	2.451816	0.052503	2212.5	40.5	2527.6	95.2	2806.9	63.4	0.02	2311	132713	135181
		28	7	0.4191	0.009413	11.724991	1.220304	0.204074	0.008133	2.393291	0.05511	2256.1	42.9	2581.3	98.7	2857.1	64.6	0.02	841	52710	57221
		29	7	0.556549	0.012335	18.273689	1.937545	0.241798	0.011428	1.801132	0.040617	2852.1	51.2	3002.9	101.5	3128.8	74.9	0.01	1387	132619	144773
		30	7	0.709459	0.02014	45.660446	5.18369	0.468889	0.023475	1.42049	0.042495	3453.1	76.7	3886.8	119.7	4135	76.3	0.01	1126	110488	90492
GH02		1	3	0.509965	0.014007	12.986798	0.725326	0.183218	0.003975	1.960495	0.053897	2656	59.8	2677.5	51.7	2679.3	36	1.01	137	220	757
		2	3	0.448371	0.015683	12.811393	0.82373	0.205934	0.006991	2.230748	0.078582	2387.4	69.9	2664	61.1	2870.3	55.7	1.37	246	364	446
		3	3	0.49768	0.015084	13.54202	0.858659	0.195406	0.006006	2.010642	0.062196	2602.8	65.1	2713.4	58	2781.2	49.6	0.97	209	663	274
		4	3	0.498133	0.014334	12.660181	0.714059	0.181908	0.004061	2.005939	0.055559	2605.1	61.2	2653.2	53.4	2667.1	37.3	0.97	263	1996	203
		5	3	0.498407	0.014366	12.887179	0.718054	0.185865	0.004238	2.004046	0.057709	2606.3	61.8	2670.3	52.4	2702.5	37.7	0.92	330	357	188
		7	3	0.410142	0.014704	11.039916	0.627057	0.193674	0.004566	2.44313	0.087154	2213.9	67.1	2524.9	53.4	2770	38.6	0.89	365	412	113
		8	3	0.527254	0.014419	13.365617	0.739419	0.18197	0.004105	1.88972	0.051847	2729.4	60.9	2704.8	52.5	2667.6	37.6	1.01	352	350	118
		9	3	0.524291	0.017793	13.687807	0.854775	0.187683	0.00689	1.904473	0.063919	2715.9	75	2724.8	59.1	2713.6	58.2	1.61	122	76	28
		10	3	0.336122	0.009794	10.410252	0.686524	0.22211	0.008289	2.94781	0.071642	1867.7	47.4	2467.7	60.6	2988.1	60.1	1.16	1199	1032	414
GW02		1	2	0.5094	0.006078	12.953629	0.95216	0.184314	0.004017	1.962168	0.023054	2653.8	25.9	2675.7	69	2691	36	1.1	352	304	407
		2	2	0.449635	0.011083	10.690585	0.850695	0.172059	0.004385	2.233248	0.05811	2392	49.6	2492	77	2575	42.9	2.1	291	130	174
		3	2	0.536182	0.007658	14.000913	1.037032	0.18931	0.004445	1.864859	0.02618	2766.9	32	2748.8	70.7	2734.4	38.8	1.51	204	130	190
		4	2	0.580913	0.017673	15.503582	1.267266	0.193149	0.004835	1.732322	0.050266	2948.9	71.3	2840.9	75.9	2766.7	41.4	1.33	161	117	179
		5	2	0.544076	0.010187	14.092922	1.067836	0.18767	0.004296	1.840059	0.034506	2799.4	42.5	2754	71.9	2724.3	33.2	2.08	390	178	260

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Sample	Spot name	Run	206/238	±2SE	207/235	±2SE	207/206	±2SE	238/P206	±2SE	206/238 Age	±2SE	207/235 Age	±2SE	207/206 Age	±2SE	U/Th	U	Th	Pb
6	2	0.476491	0.006534	12.058556	0.891157	0.18351	0.004422	2.091213	0.025629	2517.4	25.2	2608.2	68.4	2682.8	39.4	1.5	221	151	201	
7	2	0.461172	0.016176	13.698981	1.159157	0.217223	0.00571	2.187878	0.074474	2441.2	70.9	2721.3	79.1	2957.3	41.7	17.98	3275	173	2029	
8	2	0.374856	0.011177	8.788918	0.711229	0.168828	0.004413	2.684059	0.081693	2050.3	52.5	2311.3	74.2	2543	44.3	2.56	243	118	131	
9	2	0.556694	0.009526	14.513715	1.096037	0.18894	0.004618	1.796729	0.031405	2852	39.6	2782	72	2730.7	40.3	1.08	199	179	230	
10	2	0.505448	0.006517	12.665303	0.924921	0.182148	0.004121	1.976216	0.025477	2636.8	27.9	2657	77.4	2671.1	37.6	1.82	338	181	259	
11	2	0.543572	0.006602	13.983645	1.025314	0.186513	0.004235	1.835933	0.022228	2798.1	27.6	2748.1	69.9	2710.2	37.4	1.12	211	189	321	
12	2	0.298583	0.017287	7.502844	0.961002	0.178248	0.0096	3.436561	0.181853	1679	84.7	2135.4	106.6	2614	87.5	4.56	1583	343	1624	
13	2	0.233904	0.003016	5.436861	0.399456	0.168608	0.00421	4.26657	0.056316	1354.8	15.8	1890.2	63.1	2541.3	42.2	2.68	1717	642	1072	
14	2	0.304032	0.005325	6.875053	0.517542	0.163843	0.003745	3.286484	0.057949	1710.9	26.3	2093.9	67.7	2494.1	38.6	5.29	1492	296	699	
15	2	0.331996	0.022277	7.412798	0.881028	0.158874	0.006065	3.138549	0.218977	1839.6	108.1	2127.3	111.1	2433	66.3	4.15	1198	300	447	
16	2	0.515335	0.008317	12.945843	0.961627	0.182115	0.004193	1.937634	0.030811	2678.7	35.3	2674.6	70	2670.6	37.7	1.34	343	257	406	
17	2	0.472179	0.008065	11.488233	0.852825	0.176431	0.004127	2.115276	0.036394	2492.4	35.3	2562.6	69.2	2617.8	39.1	1.29	582	469	688	
18	2	0.328942	0.00839	7.041947	0.554757	0.154604	0.003415	3.045438	0.073117	1832.2	40.4	2113.4	68	2396.2	37.9	3.36	1194	350	466	
19	2	0.522942	0.006893	13.199686	0.973017	0.182979	0.004186	1.902248	0.022588	2711.2	29.3	2693.3	69.6	2682.4	33.5	0.99	487	485	694	
20	2	0.238177	0.00394	4.786482	0.367355	0.144885	0.003499	4.175529	0.076399	1377	20.5	1780.5	63.7	2284.2	41.5	3.09	1833	614	625	
21	2	0.538292	0.008015	13.580436	1.007843	0.18405	0.004051	1.852993	0.028334	2775.7	33.7	2719.8	71.1	2688.5	36.5	1.14	168	147	186	
22	2	0.361957	0.012481	8.753835	0.751624	0.174582	0.004515	2.784101	0.106441	1989	59.7	2303.9	82.6	2599.1	43.6	1.81	1013	552	718	
23	2	0.524417	0.009048	13.23554	0.983917	0.182593	0.003997	1.903298	0.034013	2717	38.5	2695.5	71.2	2678.5	41.1	1.74	281	160	203	
24	2	0.498862	0.009758	12.401705	0.933535	0.180161	0.004004	2.002535	0.041453	2607.8	42.3	2633.6	72	2653.1	36.8	1.49	413	285	356	
25	2	0.446278	0.005823	10.899137	0.800132	0.177064	0.004032	2.233536	0.029403	2378.4	26	2514	68.1	2624	37.9	2.37	498	220	330	
26	2	0.206539	0.004506	4.01246	0.324569	0.139773	0.003935	4.839839	0.105091	1210	24.1	1632.9	64.5	2228.8	53.9	3.99	3179	806	858	
27	2	0.558508	0.006416	14.217782	1.041853	0.184551	0.004228	1.783966	0.020574	2860.2	26.6	2763.9	69.4	2692.6	37.9	1.36	222	163	221	
28	2	0.78782	0.037693	22.700886	2.136297	0.207839	0.005906	1.290626	0.060981	3732.2	135.3	3198.3	91.6	2884.8	46	0.84	227	273	530	
29	2	0.540336	0.006938	13.643712	1.003699	0.183042	0.004179	1.844468	0.023566	2784.5	29	2724.7	69.4	2679.1	37.6	1.07	370	351	440	
30	2	0.532154	0.006194	13.481701	0.992613	0.183222	0.004372	1.868031	0.024205	2750.2	26	2713.3	69.1	2680.3	39	1.33	190	144	185	
31	5	0.524099	0.019361	14.618291	0.956887	0.19149	0.006619	1.921978	0.072127	2716	82.1	2789.5	62.7	2752.4	57.1	1.22	88	63	97	
32	5	0.517415	0.020102	14.397854	0.9455	0.186397	0.006644	1.952356	0.075962	2687.7	85.4	2775.7	62.2	2708.8	59	1.16	86	64	100	
33	5	0.486346	0.021699	13.486195	0.960829	0.184452	0.006023	2.095644	0.095349	2552.2	94.3	2709.7	69.3	2691.9	54.3	1.22	177	125	181	
34	5	0.527451	0.023602	15.594259	1.13929	0.194361	0.006672	1.935792	0.086316	2727.8	99.4	2846.6	69.5	2777.1	56.5	1.06	147	121	207	
35	5	0.484561	0.01858	14.036061	0.944114	0.189182	0.00632	2.100266	0.082585	2546.2	81	2750	63.6	2733.2	54.6	1.19	141	103	150	
36	5	0.46855	0.01775	13.505568	0.899991	0.187574	0.006259	2.171341	0.082074	2476.5	77.8	2713.9	62.4	2719.2	54.9	2.34	156	58	108	
37	5	0.480114	0.017602	13.615845	0.880674	0.185009	0.006129	2.114957	0.07709	2527.2	76.6	2722.6	61.3	2696.6	54.6	0.98	128	113	157	
38	5	0.484985	0.01785	13.656916	0.888164	0.185126	0.006182	2.087108	0.077303	2548.5	77.5	2725.3	61.9	2697.5	55.2	1.59	136	75	106	
39	5	0.223114	0.009246	6.46106	0.444203	0.19445	0.00711	4.477325	0.117754	1298.1	49	2039.5	62.1	2778.2	59.4	6.35	624	86	248	
40	5	0.495382	0.018946	13.182154	0.859702	0.182159	0.005937	2.025879	0.077759	2593.2	81.7	2692	61.6	2671.4	54.1	2.01	151	66	92	
41	5	0.698883	0.025459	15.266267	0.993303	0.183074	0.006294	1.456023	0.054075	3416	96.9	2831.4	62.5	2679.1	57.5	1.41	86	57	86	
42	5	0.654103	0.02331	15.074959	0.963408	0.193661	0.006413	1.554867	0.054512	3243.9	90.5	2819.7	60.8	2771.9	54.4	1.17	194	160	263	
43	5	0.704726	0.028666	15.885934	1.064764	0.186944	0.006676	1.431012	0.043333	3437.4	109.4	2868.9	65.9	2713.5	58.3	1.16	120	97	144	
44	5	0.375328	0.016706	7.414469	0.55969	0.16166	0.00568	2.707706	0.12144	2053.4	78.3	2159	68.3	2471.1	59.8	3.43	347	95	110	
45	5	0.446091	0.017	9.540837	0.643891	0.17286	0.005766	2.265381	0.087254	2377.3	75.9	2389.9	62.4	2584	55.8	2.82	284	95	103	
46	5	0.732334	0.027111	18.272201	1.19664	0.197609	0.006851	1.373567	0.050494	3541.7	100.7	3003.5	63.9	2804.8	56.8	1.54	174	108	209	
47	5	0.545603	0.019949	12.683151	0.839137	0.181732	0.00607	1.841015	0.067398	2806.1	83.1	2654.9	63	2666.9	55.5	2.33	287	117	139	
48	5	0.581412	0.021833	14.110966	0.926589	0.184189	0.007537	1.720907	0.064812	2954.3	89	2756.9	62.3	2688.6	67	1.28	156	115	155	
49	5	0.412853	0.015773	9.651972	0.647526	0.176768	0.006209	2.422767	0.092845	2227.5	72	2401.1	61.5	2620.8	58.1	4.25	543	119	218	
50	5	0.580664	0.020596	14.909843	0.977988	0.18826	0.006489	1.719753	0.061098	2951.2	84	2808.7	62.6	2725.2	57.2	1.51	146	92	119	
51	7	0.258383	0.006123	5.208366	0.538381	0.145145	0.005608	3.866582	0.089037	1481.1	31.3	1852	86.9	2286.7	65.5	31.54	1019	33	321	
52	7	0.657596	0.019501	23.768128	2.545009	0.259964	0.010405	1.525239	0.048451	3254.4	76.7	3252	109.7	3242.9	64.5	6.68	361	54	286	
53	7	0.605701	0.017797	16.1885	1.688235	0.192179	0.007479	1.643751	0.042519	3049.8	71.5	2884.5	99.9	2758.1	63.1	1.53	174	114	188	
54	7	0.740676	0.022302	20.400604	2.157228	0.198846	0.007878	1.353988	0.0408	3569.1	82.5	3105.5	102.7	2813.4	64.3	2.28	122	53	119	
55	7	0.176287	0.005195	2.855825	0.316756	0.1167	0.005318	5.686961	0.168973	1046.2	28.5	1364.2	81.8	1897	79	2.94	758	261	123	
56	7	0.557472	0.014843	13.617089	1.429669	0.176865	0.007496	1.795178	0.046981	2854.2	61.2	2719.4	98.3	2618	70.3	2.54	122	47	75	
57	7	0.527505	0.010659	13.633577	1.396449	0.187328	0.007471	1.891381	0.037965	2730.3	44.9	2723	96.1	2715.2	65.4	1.53	144	95	141	
58	7	0.526072	0.01093	13.319537	1.382207	0.183776	0.008166	1.897085	0.039515	2724.1	46.2	2699.6	98.5	2679.9	72.7	1.19	104	86	113	
59	7	0.607443	0.01973	15.723207	1.683286	0.187627	0.007551	1.653772	0.053789	3055.9	78.9	2854	101.5	2717.7	63.9	1.47	110	75	115	
60	7	0.534814	0.01261	13.68377	1.414517	0.185967	0.007738	1.868635	0.044267	2760.4	53	2725.5	98.3	2701.8	68.3	1.76	92	51	72	

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Sample	Spot name	Run	206/238	±2SE	207/235	±2SE	207/206	±2SE	238/P206	±2SE	206/238 Age	±2SE	207/235 Age	±2SE	207/206 Age	±2SE	U/Th	U	Th	Pb
GW03	1	2	0.519445	0.008831	14.98115	1.127047	0.20889	0.004715	1.934429	0.033206	2696	37.5	2812.3	71.8	2895.6	36.6	0.88	507	628	998
	2	2	0.527157	0.006445	15.566572	1.158989	0.213939	0.005198	1.903259	0.023438	2729.1	27.2	2849.4	70.7	2933.5	39.5	0.84	514	677	1092
	3	2	0.86774	0.036402	44.501783	3.739214	0.371446	0.008443	1.162326	0.052893	4015.4	131.2	3879.5	81.1	3796	34.5	0.95	387	437	2812
	4	2	0.430181	0.008539	28.929336	2.111629	0.489719	0.012932	2.336868	0.045271	2305.6	38.3	3450.8	71.3	4207.9	39.4	0.87	2689	3349	14704
	5	2	0.510516	0.007083	13.589408	1.0067	0.192982	0.004853	1.965195	0.027092	2658.4	30.2	2720.6	69.2	2765.2	41.1	0.65	582	969	1247
	6	2	0.50538	0.005121	13.014769	0.950152	0.187149	0.0042	1.978998	0.023495	2636.7	21.9	2680.4	69.4	2715.9	37	0.61	691	1221	1532
	7	2	0.508566	0.005674	13.019551	0.949266	0.185555	0.00407	1.970797	0.022095	2650.3	24.3	2680.8	68.8	2702	36.2	0.62	740	1294	1660
	8	2	0.513413	0.005928	13.36014	0.978989	0.188579	0.004178	1.952	0.022636	2670.9	25.3	2705	69.7	2728.6	36.4	0.93	560	640	888
	9	2	0.329738	0.013213	7.785524	0.755078	0.169855	0.006876	3.083616	0.129335	1834.2	64.3	2190.1	90.5	2543.6	71.2	2.41	456	246	363
	10	2	0.574503	0.009784	15.667104	1.211917	0.198525	0.005426	1.746393	0.030252	2925.3	40.2	2860.3	65	2810.6	44.2	1.27	58	47	92
GW04	1	2	0.530237	0.005975	13.614289	0.99591	0.186083	0.004127	1.886004	0.021508	2742.2	25.2	2722.9	69.4	2706.6	36.6	1.31	431	334	477
	2	2	0.515233	0.006072	13.196901	0.967039	0.185625	0.004111	1.94057	0.023065	2678.7	25.9	2693.4	69.1	2702.5	36.8	1.35	527	397	530
	3	2	0.53444	0.006513	13.745034	1.004702	0.186406	0.00407	1.87055	0.022978	2759.8	27.4	2731.9	69.5	2709.6	36.2	0.75	634	867	1189
	4	2	0.550116	0.00717	14.190072	1.039146	0.186984	0.004202	1.817139	0.0237	2825.2	29.8	2762.1	69.3	2714.4	37	0.95	459	495	696
	5	2	0.546668	0.007615	14.311695	1.057623	0.189712	0.004275	1.828566	0.025848	2810.8	31.8	2769.7	70.7	2738.2	37.3	1.09	406	383	525
	6	2	0.429071	0.012768	9.829381	0.773188	0.165969	0.003715	2.34554	0.074313	2299.2	58	2415.1	74.6	2516	37.9	1.39	1262	932	1276
	7	2	0.503094	0.011318	13.002851	1.00735	0.187151	0.00452	1.99183	0.045241	2625.5	48.6	2676.7	73.5	2715.2	39.7	1.56	432	361	435
	8	2	0.54983	0.006977	14.384981	1.058222	0.189046	0.00432	1.815956	0.02269	2824.1	28.9	2774.8	69.2	2732.3	37.8	1.83	223	127	203
	9	2	0.499347	0.006896	13.043161	0.960207	0.189299	0.004187	1.999604	0.027793	2610.5	29.7	2682.1	70.2	2734.8	36.6	2.44	489	214	376
	10	2	0.534543	0.005922	13.801351	1.009067	0.186628	0.004028	1.866182	0.020772	2760.3	24.9	2735.8	69	2711.6	35.7	1.09	759	730	1060
	11	2	0.301844	0.022377	6.738675	0.810541	0.159273	0.005171	3.472733	0.257393	1691.5	110.4	2043.7	107.1	2441.4	54.9	0.55	6586	15053	9760
	12	2	0.239417	0.020537	5.093894	0.797191	0.145258	0.007726	4.358326	0.355056	1403.3	117.2	1775.3	126.2	2264.7	98.4	0.26	16984	84955	39299
	13	2	0.430483	0.055428	11.207914	1.899883	0.182822	0.007219	2.685715	0.329867	2263.7	243.4	2445.1	158.5	2667.7	65.3	2.29	2659	3542	3503
	14	2	0.564985	0.008509	14.580639	1.078498	0.187033	0.004193	1.762369	0.027598	2886.5	35.3	2787.4	70.9	2714.9	37	1.14	363	336	491
	15	2	0.453115	0.023106	11.335161	1.087568	0.180199	0.004763	2.250767	0.122552	2401.4	103.3	2534.2	93.1	2651.5	44.4	1.56	580	395	468
	16	2	0.551123	0.007588	14.736837	1.099347	0.193154	0.004413	1.803203	0.025592	2829.3	31.7	2797.1	70.6	2767.7	37.4	1.17	381	351	551
	17	2	0.530076	0.005843	13.671763	0.999319	0.187259	0.004059	1.875402	0.020436	2741.5	24.6	2726.9	68.9	2717.2	35.7	0.77	696	934	1295
	18	2	0.537947	0.00651	13.956122	1.020603	0.188026	0.004069	1.847963	0.022429	2774.5	27.3	2746.4	69.9	2723.9	35.8	0.98	539	571	823
	19	2	0.333313	0.019532	7.756534	0.809691	0.167325	0.004546	3.06318	0.154188	1848.1	92.5	2184	86.5	2527.5	44.9	1.84	793	448	445
	20	2	0.540516	0.005385	14.097129	1.02663	0.188956	0.004213	1.837413	0.018294	2785.5	22.5	2756.1	69.2	2731.8	36.6	0.84	582	730	1054
	21	2	0.537508	0.006206	14.062005	1.034459	0.189566	0.004213	1.844816	0.021344	2772.7	26	2753.3	69.9	2737.1	36.7	1.24	519	475	636
	22	2	0.505935	0.008429	18.790836	1.940215	0.270202	0.021505	1.962395	0.034275	2638.5	36.3	3008.3	97.5	3257.5	124.2	1.25	676	706	1977
	23	2	0.48281	0.005438	12.187405	0.890749	0.183412	0.004001	2.052888	0.022888	2539.3	23.6	2621	77.6	2682.8	36.1	2.07	628	342	422
	24	2	0.630687	0.007787	20.062937	1.471306	0.230556	0.005077	1.571704	0.019836	3151.9	30.9	3093.9	71.2	3054.7	35.2	0.82	651	891	1301
	25	2	0.602792	0.008064	19.057938	1.39774	0.229154	0.004966	1.639851	0.019823	3040.6	32.6	3044.2	71.6	3045.1	34.8	0.76	760	1103	1533
	26	2	0.538942	0.006014	13.861708	1.014888	0.186942	0.004212	1.838135	0.020715	2778.8	25.2	2739.9	69.3	2714	37	0.91	825	1005	1339
	27	2	0.467429	0.007328	11.805012	0.870058	0.183151	0.00437	2.113931	0.029571	2471.7	32.3	2588.4	68.9	2679.6	39.1	2.6	763	328	448
	28	2	0.536875	0.007355	13.987838	1.028059	0.18888	0.004335	1.845761	0.024971	2769.9	30.8	2748.3	69.5	2730.9	37.9	1.31	402	335	478
	29	2	0.236236	0.004946	5.235789	0.401457	0.160539	0.003823	4.204358	0.089023	1366.7	25.8	1856.4	64.7	2459.4	39.6	1	3801	4251	2644
	30	2	0.237038	0.008999	4.729415	0.45322	0.143298	0.0048	4.230234	0.162253	1369.7	46.9	1759.5	80.8	2259.9	58.1	4.03	2245	652	515
GW06	1	6	0.33876	0.007415	9.721441	0.697464	0.21019	0.007276	2.946037	0.062539	1880.4	35.6	2408	65.7	2905.4	56.3	7.41	528	70	339
	2	6	1.233511	0.031439	43.931858	3.266593	0.262254	0.009784	0.813211	0.020404	5173.5	90.3	3859.4	72.2	3263.5	52.7	0.76	38	49	153
	3	6	0.748724	0.02659	23.139639	1.793594	0.226679	0.009792	1.332521	0.04126	3594.9	99.2	3235	70.5	3019.1	68.7	0.69	35	49	90
	4	6	1.091333	0.033119	38.089156	2.933734	0.25731	0.010531	0.915842	0.028655	4753.5	102.9	3720.2	76.4	3227.5	64.2	0.77	33	41	112
	5	6	0.79683	0.033569	26.29378	2.388169	0.238941	0.010844	1.273856	0.053287	3766	119.9	3339.6	89.2	3101.2	73.8	0.82	44	53	116
	6	6	0.275466	0.014169	6.222868	0.504586	0.162461	0.005683	3.66059	0.21311	1566.9	72.3	2004.6	73.3	2479.5	59.3	3.19	1182	376	589
	7	6	0.571843	0.014046	14.377218	1.025294	0.183147	0.005632	1.745593	0.045761	2914.7	58.3	2774.3	69.9	2680.8	50.9	0.35	341064	940552	1262252
	8	6	0.343901	0.013825	8.757067	0.688011	0.183235	0.005609	2.947735	0.119864	1901.8	66.3	2306.4	71.7	2681.5	51	0.34	309572	882040	743303
	9	6	0.122213	0.007637	3.003163	0.276584	0.175735	0.005414	8.440719	0.443976	741.7	43.3	1397.9	65.5	2614.4	43.7	0.32	317217	940228	273157
	10	6	0.133488	0.009486	3.273966	0.338059	0.174254	0.005383	7.885515	0.573632	805.4	53.9	1455.9	80.9	2597.8	51.5	0.32	304134	928912	291266
	11	6	4.349811	0.053062	107.586152	7.435138	0.176859	0.005552	0.229454	0.002763	10807.7	63.6	4759.9	69	2622.2	52.5	6.31	8118	1969	267226
	12	6	3.716031	0.073684	91.851934	6.506011	0.177086	0.005494	0.267986	0.004822	9989.9	101.3	4598.8	71.7	2624.6	51.2	11.75	7660	706	213254
	13	6	2.017763	0.072156	49.020276	3.869948	0.1756	0.005469	0.496582	0.017733	7112.4	154.8	3969.3	78.2	2610.8	51.7	8.1	11478	1463	166074
	14	6	3.783938	0.047575	93.779136	6.541918	0.177583	0.005478	0.263	0.003689	10087	63.4	4620.3	69.5	2629.4	51.3	18.2	6686	382	192383

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Sample	Spot name	Run	206/238	±2SE	207/235	±2SE	207/206	±2SE	238/P206	±2SE	206/238 Age	±2SE	207/235 Age	±2SE	207/206 Age	±2SE	U/Th	U	Th	Pb
	15	6	3.815799	0.129051	94.815724	7.329193	0.178314	0.005476	0.262843	0.008987	10120.3	173.5	4628.9	78	2636.4	51.1	10.07	6660	703	200477
	16	6	3.964096	0.096655	98.424635	7.139622	0.178677	0.005522	0.252806	0.006387	10317.4	127.3	4667.2	74.8	2639.6	51.2	13.33	5628	434	175026
	17	6	3.473075	0.237094	85.114458	8.380377	0.177101	0.005547	0.302735	0.022449	9558.3	352.2	4496.1	103.7	2624.8	51.4	19.89	8906	454	223598
	18	6	3.119095	0.083346	76.775011	5.615624	0.177535	0.005633	0.320694	0.008343	9118.9	129	4418.9	73.2	2628.8	52.1	9.24	9703	1082	237131
	19	6	3.18123	0.094603	77.588554	5.723001	0.176477	0.005489	0.314944	0.009315	9212.4	145.3	4428.9	74.4	2619	51.6	10.77	10165	1001	250593
	20	6	4.30034	0.07963	105.727704	7.47132	0.178151	0.005517	0.232223	0.00433	10746.9	97.1	4741.1	71.2	2634.8	51.4	15.55	5317	349	177430
LP02	1	1	0.289645	0.00561	5.889325	0.420548	0.149728	0.005952	3.434665	0.065384	1639.5	28	1957	61.6	2337.3	67.5	2.01	30079	15817	17182
	2	1	0.424814	0.007666	11.1168	0.748889	0.194552	0.006698	2.339809	0.043045	2282.1	34.8	2532.3	64.2	2779.4	56.1	1.56	2012	1302	2224
	3	1	0.517897	0.009081	12.074991	0.781681	0.172151	0.005402	1.920635	0.03382	2689.8	38.6	2609.8	61.5	2577.3	52.6	5.15	2359	488	1180
	4	1	0.47648	0.007681	12.2587	0.807619	0.190418	0.006014	2.087958	0.033486	2511.5	33.5	2623.2	62.4	2744.1	52.3	4.25	1616	371	763
	5	1	0.493838	0.010235	13.041322	0.849699	0.19662	0.006076	2.018098	0.040949	2586.3	44	2681.8	61.1	2797.1	50.3	1.02	1221	1178	1821
	6	1	0.369263	0.01455	8.848236	0.670507	0.177961	0.005518	2.732014	0.117701	2022.7	69.1	2315.4	73	2632.6	51.6	4.3	2838	760	1165
	7	1	0.464717	0.011549	11.583362	0.752046	0.18536	0.005964	2.149265	0.052693	2458.9	50.7	2570.6	60.6	2699.4	53.1	0.9	1662	1998	1058
	8	1	0.167531	0.007052	2.708841	0.21799	0.119414	0.004365	5.972776	0.250143	998.1	38.9	1328.4	60.4	1944.9	63.9	0.3	46755	156883	53260
	9	1	0.516163	0.007434	12.496353	0.802666	0.179915	0.005525	1.92849	0.028156	2682.6	31.7	2642.1	60.5	2650.9	51.1	3.59	1828	496	648
	10	1	0.201173	0.009715	5.612189	0.437849	0.207051	0.007488	5.010947	0.250826	1180.5	52.2	1913.5	69.8	2879.6	59.7	1.08	8840	8903	10770
	11	1	0.344073	0.006022	7.85983	0.505448	0.170178	0.005231	2.900994	0.051655	1905.9	28.9	2214.8	57.7	2558.2	51.2	3.82	2218	574	1468
	12	1	0.338787	0.014041	6.735459	0.512449	0.147921	0.004601	2.986366	0.12277	1877.7	67.5	2070.9	67.3	2320.6	53.3	1.42	10192	7960	6579
	13	1	0.245909	0.004491	5.369169	0.351203	0.162603	0.005329	4.06162	0.074555	1417.1	23.2	1879.2	56.4	2480.4	56.1	2.68	4909	1855	3003
	14	1	0.362871	0.005824	7.953452	0.514656	0.163091	0.005025	2.750958	0.044109	1995.5	27.5	2225.2	58.4	2486.7	51.7	2.67	2332	1003	814
	15	1	0.225061	0.00504	4.038543	0.263523	0.134119	0.004256	4.445879	0.095495	1308.2	26.4	1644.5	59.3	2150.7	55	5.98	9858	1546	3939
	17	1	0.38253	0.005354	8.854732	0.56775	0.172005	0.006937	0.036251	0.2088	2088	24.9	2322.9	58.6	2576	51.6	3.82	1705	439	743
	18	1	0.315022	0.008861	8.036256	0.550012	0.189456	0.005806	3.186626	0.087788	1764.2	43.3	2322.7	60.5	2736.3	50.4	1.1	3959	3567	4661
	19	1	0.339179	0.010636	8.505107	0.596458	0.186029	0.005668	2.967183	0.098103	1881	51.5	2283	65.4	2706.3	50.3	3.18	3440	1064	3108
	20	1	0.392731	0.005883	8.624727	0.556734	0.162424	0.005046	2.543401	0.038056	2135.3	27.2	2298.7	58.5	2479.6	53.1	1.44	4918	3458	3929
	21	1	0.501028	0.009702	11.612991	0.755915	0.171156	0.005742	1.995479	0.038814	2617.8	41.7	2573.4	60.7	2567.2	55.2	4.04	3097	753	1264
	22	1	0.448797	0.010638	10.076265	0.744219	0.164204	0.005667	2.234428	0.052537	2388.7	47.2	2436.3	67	2495.9	57.3	4	4150	929	1835
	23	1	0.498291	0.010879	14.736624	0.980807	0.216403	0.00672	2.008006	0.044315	2605.8	46.9	2797.5	63.7	2953.1	49.8	10.55	2773	275	2326
	24	1	0.308893	0.008114	8.172609	0.550586	0.193672	0.005975	3.251785	0.089998	1734.3	40.2	2248.4	62.3	2772.3	50.8	2	2005	991	2585
	25	1	0.249274	0.007613	4.907223	0.372028	0.14327	0.00484	4.041087	0.136181	1433.7	39.6	1797.4	68	2263.9	59.4	3.13	7098	3144	1899
	26	1	0.374856	0.006483	9.648645	0.629363	0.187215	0.005833	2.669258	0.046268	2051.8	30.4	2401	59.9	2716.4	51.8	1.16	1288	1092	1242
	27	1	0.311721	0.004768	6.809366	0.439373	0.158758	0.004891	3.207309	0.048658	1749	23.4	2086.6	57	2441.4	52.3	4.23	2042	488	477
	28	1	0.338817	0.008415	6.655797	0.471426	0.142617	0.004633	2.9463	0.066249	1880	40.6	2063.2	62.5	2256.9	56.1	1.36	4829	4197	2913
	29	1	0.17709	0.00653	3.965799	0.2807	0.163087	0.005838	5.70738	0.20605	1050.2	35.7	1624.4	56	2483.7	58.6	2.17	7046	3446	6299
	30	1	0.194598	0.004731	5.03329	0.329511	0.187742	0.005863	5.159258	0.134092	1145.9	25.7	1824.1	56.1	2721	50.5	2.29	2462	1081	2005
LP03	1	1	0.501155	0.007352	12.083997	0.779738	0.173921	0.005312	2.001134	0.029342	2618.5	31.6	2610.5	61.3	2594.6	50.7	0.93	1217	1199	1513
	2	1	0.326643	0.00897	7.705713	0.560495	0.168316	0.005871	3.073462	0.082823	1821.7	43.5	2195.6	64.4	2539.1	58.6	1.32	1478	1048	2201
	3	1	0.435056	0.0066	10.250983	0.661713	0.169682	0.005221	2.304198	0.035211	2328.2	29.7	2459.8	52.6	2553.3	51.6	2.08	1494	674	681
	4	1	0.471653	0.006365	11.459022	0.736093	0.17464	0.005405	2.123828	0.028714	2490.7	27.9	2560.9	60.5	2601.3	51.4	4.34	912	203	293
	5	1	0.272154	0.007404	6.699886	0.466152	0.176334	0.005569	3.699987	0.104404	1550.9	37.6	2069.7	62.2	2616.6	53.8	0.57	519	888	310
	6	1	0.269971	0.004894	5.67683	0.372945	0.150569	0.004667	3.713506	0.067877	1540.4	24.9	1930.5	50.7	2351	53.3	2.05	2311	1124	877
	7	1	0.44448	0.009314	11.249238	0.778477	0.180938	0.005865	2.257082	0.045972	2369.8	41.4	2541	63.6	2659.3	53.7	0.51	1010	1946	2176
	8	1	0.527388	0.009479	12.77144	0.822959	0.173869	0.005562	1.898614	0.033283	2730	39.8	2662.7	60.5	2593.7	53.4	1.38	435	310	427
	9	1	0.232079	0.00546	4.745089	0.317654	0.146327	0.00465	4.325412	0.101591	1344.9	28.5	1773.8	56.8	2301.6	54.4	1.87	1702	901	1144
	10	1	0.14497	0.005175	2.588605	0.206159	0.12815	0.004438	6.922701	0.245481	872.5	29.1	1295.3	58	2070.8	61.5	0.84	4226	5048	2366
	11	1	0.515277	0.007813	12.456973	0.800244	0.173194	0.005271	1.937691	0.029529	2678.9	33.3	2639.3	60.5	2587.8	50.9	4.79	1109	237	370
	12	1	0.494336	0.009393	12.460681	0.817947	0.181876	0.00579	2.020563	0.037961	2589	40.4	2639.1	61.6	2668.7	52.3	1.07	1184	1127	2265
	13	1	0.448359	0.011238	10.964926	0.74933	0.174049	0.006031	2.228212	0.055207	2387.4	49.9	2519.2	64.5	2595.2	57.1	1.29	687	544	817
	14	1	0.711082	0.011164	28.248723	1.839539	0.282994	0.008721	1.403014	0.022066	3462.2	42.1	3427.3	63.8	3378.6	48	1.72	1935	1194	2596
	15	1	0.478229	0.007152	11.79377	0.7596	0.176478	0.005478	2.086368	0.030856	2519.3	31.1	2587.8	60.4	2618.7	51.9	0.8	1033	1314	1896
	17	1	0.528527	0.008574	12.925647	0.837461	0.174885	0.005528	1.885491	0.030866	2735	36.2	2673.9	61.7	2603.7	52.8	0.93	671	710	991
	18	1	0.425661	0.00619	9.870789	0.63665	0.166269	0.005105	2.342819	0.034356	2285.9	28	2422.3	59.5	2519.2	51.9	0.4	2238	5730	5997
	19	1	0.328703	0.0062	7.812043	0.508926	0.17035	0.005302	3.037518	0.058635	1831.7	30.1	2212	52.2	2559.7	51.5	0.79	1036	1321	1133
	20	1	0.368987	0.01506	9.171645	0.713597	0.17792	0.005606	2.739748	0.110532	2021.2	70.7	2347.3	70.5	2631.9	51.8	0.95	881	955	1069

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Sample	Spot name	Run	206/238	±2SE	207/235	±2SE	207/206	±2SE	238/P206	±2SE	206/238 Age	±2SE	207/235 Age	±2SE	207/206 Age	±2SE	U/Th	U	Th	Pb
21	1		0.302693	0.006012	8.266061	0.538202	0.19754	0.006204	3.292872	0.068444	1704.2	29.9	2259.9	59.5	2804.4	51.6	1.77	1087	633	896
22	1		0.475448	0.009537	11.323775	0.760466	0.172711	0.00535	2.09592	0.043023	2506.6	41.8	2548.2	64.2	2582.8	51.8	0.89	876	938	1033
23	1		0.425535	0.008033	10.320212	0.674779	0.175801	0.005383	2.340097	0.043738	2285	36.3	2463	60.4	2612.5	51.1	2.01	1059	517	728
24	1		0.459572	0.011224	11.016575	0.756114	0.173851	0.005372	2.171234	0.052377	2436.3	49.4	2521.9	63.4	2593.8	51.8	1.26	1328	1020	1417
25	1		0.495816	0.010952	11.931321	0.804983	0.174797	0.005411	2.010317	0.045225	2594.6	47.3	2596.9	63.5	2602.8	51.5	1.34	846	630	833
26	1		0.517775	0.007509	12.849549	0.82972	0.180032	0.005689	1.919517	0.02722	2689.5	31.8	2668.2	60.9	2651.5	52.9	1.14	540	461	712
27	1		0.407797	0.007527	9.12089	0.59201	0.163044	0.004994	2.439979	0.045546	2204.3	34.5	2349.5	59.9	2486.3	51.7	1.9	2027	1061	1275
28	1		0.312944	0.011513	7.29126	0.593005	0.169274	0.006063	3.2106	0.12066	1753.1	56.6	2138.3	71.7	2546.2	58.6	0.7	882	1223	1032
29	1		0.226423	0.011303	4.597943	0.415827	0.14748	0.005196	4.477701	0.208246	1313.3	58.9	1736.2	72.2	2312.8	59.9	2.51	2507	1011	1351
30	1		0.380729	0.00862	8.736926	0.573772	0.167269	0.00512	2.615359	0.059205	2079.2	40.2	2310	59.7	2529.4	51.3	1.37	1166	843	1816
LP04	1		0.475409	0.0084	11.57126	0.755884	0.176575	0.005434	2.108838	0.038034	2506.7	36.8	2569.5	60.9	2619.7	50.9	1.05	379	366	569
	2		0.478327	0.006979	11.834762	0.762283	0.179446	0.005526	2.094038	0.03047	2519.7	30.4	2591	60.2	2646.5	51	1.04	448	466	743
	3		0.430149	0.006884	10.147871	0.662472	0.170213	0.005521	2.327946	0.037571	2306.3	31.1	2448	60.6	2558.4	54.1	1.34	532	380	519
	4		0.245528	0.004433	5.887242	0.398618	0.174089	0.006046	4.082282	0.073148	1415.2	22.9	1958	57.6	2594.5	56.8	3.45	1462	422	1419
	5		0.372099	0.007673	8.311272	0.558528	0.161739	0.004942	2.697439	0.056241	2038.6	36.1	2263.7	61.2	2472.8	51.5	2.24	2089	1293	1775
	6		0.388342	0.014467	10.181994	0.668572	0.192339	0.007951	2.607043	0.093484	2112.2	66.7	2454.3	67.5	2754.4	66.6	0.61	1076	8462	1823
	7		0.357955	0.005319	8.422593	0.541348	0.170553	0.005557	2.798221	0.041844	1972.3	25.3	2277.3	58.2	2560.8	53.5	1.06	1570	1577	2793
	8		0.478269	0.006821	12.09832	0.780183	0.182418	0.005564	2.093024	0.030008	2519.6	29.8	2611.8	60.3	2674.1	50.2	1.24	1043	937	1764
	9		0.444438	0.009161	11.351451	0.754619	0.184365	0.005722	2.255173	0.045238	2370	40.7	2551.5	62.6	2691.4	51.4	1.51	1122	834	1824
	10		0.136305	0.011415	2.416823	0.307579	0.126596	0.005281	7.5929	0.59956	822.1	64.3	1232.4	87.3	2045.9	72	0.5	8875	22696	6255
	11		0.277397	0.010188	7.357197	0.512599	0.192505	0.006527	3.638799	0.130244	1576.5	51.2	2152.8	62.2	2760.6	56.2	1.05	10806	156082	290484
	12		0.384991	0.012845	10.524323	0.721311	0.197197	0.006238	2.602969	0.08979	2098.5	60	2480.8	64.9	2802	51.4	1.33	5090	3829	7520
	13		0.349586	0.007737	9.513309	0.651589	0.197907	0.006449	2.855846	0.062793	1932.4	36.9	2388	62.9	2807.8	53.4	0.76	14065	18467	6050
	14		0.443393	0.008648	10.492958	0.676501	0.171416	0.005289	2.253231	0.044287	2365.1	38.6	2478.9	59.9	2570.2	51.4	0.58	3464	5944	14113
	15		0.39181	0.007959	9.865837	0.645911	0.18181	0.005539	2.546635	0.052328	2130.8	36.9	2421.8	60.6	2668.5	50.4	2.27	1602	706	1777
	16		0.43241	0.022873	11.448049	0.990758	0.190915	0.006038	2.370125	0.141903	2309	104.5	2545.1	85.9	2748.4	52.6	0.92	1545	1683	2180
	17		0.227534	0.008474	5.841257	0.451626	0.18521	0.00593	4.434219	0.187939	1320.2	45	1945.1	71.9	2698.1	53.4	1.21	2424	1998	3748
	18		0.24164	0.0039	6.271915	0.406135	0.187094	0.005749	4.119455	0.066853	1395.2	20.3	2104.3	56.7	2715.8	50.4	3.39	1501	443	1849
	19		0.274114	0.011509	6.937286	0.489789	0.182897	0.00607	3.657281	0.152813	1560.4	58.2	2101.6	62.2	2677.6	55	1.79	711	398	693
	20		0.298311	0.009775	7.656954	0.533179	0.18593	0.005729	3.362768	0.113553	1681.4	48.6	2188.5	63.3	2705.2	51.2	1.82	533	293	417
LP10	1		0.537531	0.0131	16.006518	1.083502	0.215603	0.006633	1.852347	0.047143	2771.4	55.3	2874.8	66.1	2947	49.9	1.61	1666	988	1248
	2		0.417478	0.014089	11.560888	0.906547	0.199585	0.006602	2.395062	0.076843	2246.4	63.5	2561.4	70.2	2820.3	52.9	1.84	2055	1034	1252
	3		0.489384	0.008675	14.020071	0.908062	0.206553	0.006524	2.026252	0.036337	2567.7	37.6	2750.7	61.2	2877.4	51.7	1.85	1287	648	810
	4		0.504548	0.007812	12.190235	0.789046	0.174622	0.005579	1.964726	0.030149	2633	33.4	2618.6	60.6	2600.5	53.3	0.77	420	510	648
	5		0.381714	0.005795	9.980466	0.647756	0.189176	0.005828	2.596081	0.039536	2084.1	27	2432.3	60	2733.8	51	1.44	1852	1213	1223
	6		0.524417	0.008633	15.11156	0.980159	0.208899	0.006704	1.888498	0.030918	2717.7	36.5	2822	62	2895.8	52.6	2.3	1570	657	850
	7		0.31713	0.010048	7.79848	0.546807	0.178242	0.00573	3.129878	0.098445	1775.1	49.1	2206.8	63.2	2635.4	53.6	1.32	2314	1722	1122
	8		0.602217	0.01184	19.421444	1.268349	0.232507	0.007191	1.646767	0.032801	3037.7	47.7	3062	64	3068.1	49.6	1.19	879	740	1097
	9		0.391815	0.006588	10.384915	0.670634	0.190999	0.005928	2.527941	0.043078	2130.9	30.6	2469.2	60	2749.3	51.6	2.81	1084	390	406
	10		0.457358	0.012321	13.14079	0.928604	0.206589	0.00644	2.174146	0.058102	2426.1	54.4	2686	65.7	2877.6	50.4	1.74	2054	1215	1630
	11		0.868233	0.038436	34.003075	2.981082	0.28094	0.010177	1.143693	0.051386	4025.1	132.9	3604.8	87.8	3366.4	56.9	1.6	944	634	3026
	12		0.219734	0.003879	4.246803	0.282001	0.138834	0.00436	4.500753	0.078981	1280.3	20.5	1682	54.7	2211	54.7	1.64	2698	1760	1077
	13		0.260576	0.008319	5.664278	0.440331	0.155597	0.005242	3.819672	0.118756	1491.6	42.4	1919.1	66.1	2405.3	56.9	1.79	2137	1320	767
	14		0.381137	0.008456	10.314553	0.728325	0.195614	0.007245	2.592205	0.057817	2081.4	39.5	2462.4	65.4	2788.1	60.8	2.21	1354	662	884
	15		0.239441	0.005391	5.001713	0.33224	0.15045	0.005461	4.136313	0.091683	1383.4	28	1818.4	55.8	2346.3	60.6	2.35	1880	877	737
	16		0.370181	0.005893	9.496883	0.616715	0.184447	0.00572	2.668886	0.04209	2030	27.7	2386.6	59.5	2691.9	51.1	4.94	1757	381	367
	17		0.44033	0.010418	12.25745	0.83115	0.200076	0.006165	2.250899	0.058044	2350.9	47.2	2621.9	66.7	2825.6	51	1.43	1720	1294	1325
	18		0.236847	0.005522	4.535428	0.337419	0.137056	0.00516	4.174933	0.09637	1370	28.7	1734.9	60	2186.8	63.9	1.51	2044	1444	900
	19		0.558794	0.017576	16.992127	1.251946	0.218313	0.007078	1.781238	0.060422	2858.2	73.6	2928.1	73.7	2966.2	52.5	1.25	895	769	998
	20		0.48512	0.013688	13.838117	0.979213	0.205059	0.006469	2.046924	0.059287	2547.4	59.6	2734.6	67.4	2865.3	51.1	3.2	1654	546	652
	21		0.453786	0.013245	13.647996	0.972378	0.216797	0.006935	2.190339	0.068306	2409.9	59.3	2721.2	68.4	2955.2	51.1	5.47	1125	207	374
	22		0.369354	0.006214	8.993154	0.585415	0.175258	0.00569	2.67405	0.044643	2026.2	29.2	2336.9	59.9	2606.9	53.8	1.88	1599	858	842
	23		0.358637	0.005891	8.722531	0.574829	0.17633	0.005703	2.753534	0.045377	1975.6	28	2308.9	60	2617.2	53.8	4.36	2110	494	372
	24		0.492614	0.011909	14.030634	0.954265	0.205667	0.006336	2.012817	0.05207	2580.5	51.9	2749.3	66.2	2870.5	50.2	2.67	942	347	438

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Sample	Spot name	Run	206/238	±2SE	207/235	±2SE	207/206	±2SE	238/P206	±2SE	206/238 Age	±2SE	207/235 Age	±2SE	207/206 Age	±2SE	U/Th	U	Th	Pb
	25	1	0.308556	0.009872	7.643221	0.535175	0.179923	0.005902	3.210449	0.107694	1733	48.9	2188.6	64.9	2650.8	54.1	1.23	2067	1673	1479
	26	1	0.383502	0.011967	10.231156	0.801863	0.193683	0.007285	2.580888	0.080336	2092	55.7	2453.2	73.7	2771.5	61.8	1.73	1282	727	704
	27	1	0.287267	0.004192	6.280679	0.412107	0.158108	0.005096	3.439935	0.050274	1627.7	21	2014.9	56.9	2433.4	54.4	3.73	1622	432	397
	28	1	0.302877	0.01166	8.460246	0.603605	0.201816	0.006579	3.300759	0.12086	1703.4	57.3	2277.8	63.5	2838.8	51.8	1.81	2069	1131	1462
	29	1	0.561376	0.008697	16.526665	1.073697	0.213063	0.006493	1.761366	0.027316	2871.9	35.9	2907	62.8	2928	49.2	1.09	916	831	1157
	30	1	0.299461	0.004044	7.295817	0.467829	0.176448	0.005399	3.300792	0.044381	1688.6	20.1	2148.1	57.4	2618.6	50.8	2.03	1941	954	867
LP11	1	1	0.555427	0.011535	16.269356	1.091803	0.212396	0.006628	1.806198	0.037829	2846.6	47.9	2890.8	65.1	2922.5	50.5	2.51	607	247	602
	2	1	0.507672	0.015119	14.594182	1.056022	0.208529	0.006484	1.983695	0.056176	2644.1	64.1	2784.3	66.6	2892.8	50.6	3.23	908	297	713
	3	1	0.411277	0.01758	11.502155	0.971235	0.202302	0.006773	2.471131	0.105627	2216.4	80.1	2552.1	79.5	2842.1	55	2.37	1838	827	1483
	4	1	0.555413	0.01213	16.873273	1.129835	0.22128	0.006892	1.806335	0.039594	2846.3	50.3	2925.8	64.4	2988.7	50.1	1.83	1101	655	1651
	5	1	0.488516	0.015604	14.46801	1.090368	0.215417	0.006954	2.063686	0.06468	2561.4	67.3	2774.1	70.8	2944.8	52.4	1.77	1112	695	1482
	6	1	0.508376	0.01058	14.455779	0.975711	0.208899	0.006906	1.967901	0.041521	2649.4	45.3	2779.6	65.5	2895.9	54.1	2.96	1257	479	1128
	7	1	0.458794	0.010906	12.522823	0.842309	0.199575	0.006093	2.187754	0.052432	2432.9	48.2	2642.5	63.1	2821.6	49.9	2.43	1203	562	1114
	8	1	0.407545	0.008459	10.991114	0.732205	0.198525	0.006462	2.454441	0.050301	2203.5	38.7	2521.9	61.9	2813	52.8	1.98	1388	797	1291
	9	1	0.468504	0.014738	12.972683	0.97596	0.202422	0.006695	2.149621	0.064395	2474.4	64.1	2671.3	68.5	2843.3	53.1	1.44	1090	873	1356
	10	1	0.57889	0.009437	17.522556	1.130571	0.222133	0.006886	1.728235	0.028386	2943.7	38.6	2963.4	61.8	2994.9	49.6	1.61	1657	1152	1969
	11	1	0.151181	0.005693	2.189151	0.196218	0.105674	0.004542	6.693332	0.276891	906.9	32	1168	65.3	1714.7	80.8	15.05	3300	224	161
	12	1	0.178105	0.00528	2.749075	0.20213	0.113521	0.003728	5.650409	0.190169	1056.1	29.1	1338.1	57.1	1853.8	59.8	18.84	3417	183	160
	13	1	0.103536	0.002437	1.20245	0.091471	0.085905	0.003382	9.657963	0.229461	635.1	14.2	800.7	41.6	1332	74.9	11.99	5673	465	290
	14	1	0.586188	0.012208	17.930392	1.175205	0.226762	0.007609	1.704557	0.034799	2973.5	49.4	2985.6	62.4	3027.9	54.1	2.1	968	455	560
	15	1	0.472122	0.017738	12.982958	1.014062	0.202157	0.006692	2.141168	0.084752	2488.8	78.2	2669.5	75.9	2841.2	54.1	2.84	1349	496	510
	16	1	0.55914	0.009699	16.480206	1.090026	0.217203	0.006737	1.7876	0.031147	2862.4	40.1	2903.7	63.7	2958.7	50.2	2.09	856	410	589
	17	1	0.460973	0.008872	13.416914	0.877868	0.214643	0.00672	2.169429	0.041537	2443.1	39.1	2708.5	61.5	2939.5	50.8	1.98	1243	653	707
	18	1	0.563252	0.00865	16.288029	1.05093	0.213188	0.006639	1.773004	0.026939	2879.7	35.6	2983.4	61.4	2928.5	51.1	1.48	478	334	453
	19	1	0.471363	0.010504	13.044221	0.881464	0.203803	0.006296	2.123708	0.047866	2488.4	46.1	2680.7	64.3	2855.6	50.3	1.7	1642	991	1207
	20	1	0.247659	0.004743	5.680474	0.383446	0.168869	0.005532	4.035975	0.075068	1426.1	24.4	1926.7	57.3	2544.1	54.8	1.9	3704	2013	1688
	21	1	0.524135	0.008334	15.30012	0.995286	0.214203	0.006671	1.902222	0.030463	2716.3	35.3	2833.3	61.8	2936.2	50.3	2.07	999	528	786
	22	1	0.423211	0.00729	11.523742	0.765461	0.199569	0.006361	2.356181	0.041112	2274.6	33.1	2565	62.1	2820.9	51.7	1.59	1283	898	1059
	23	1	0.187528	0.004789	4.549127	0.416741	0.177294	0.009848	5.324396	0.1365	1107.7	26	1730.7	79.1	2613.4	97.2	3.06	4073	1490	2580
	24	1	0.228752	0.004255	5.490699	0.359335	0.175822	0.005618	4.358714	0.080477	1327.7	22.3	1898.3	56.3	2611.9	52.7	2.02	2769	1551	1600
	25	1	0.503034	0.013421	15.115945	1.043366	0.219797	0.006949	1.988597	0.053657	2624.8	57.6	2819.5	66.5	2977.6	51.1	1.1	702	741	903
	26	1	0.573378	0.012877	17.67286	1.176263	0.225368	0.006958	1.740741	0.04009	2920.2	53	2970.4	64.7	3018.2	49.7	1.47	539	421	641
	27	1	0.536154	0.00927	16.037104	1.048996	0.218484	0.006723	1.857373	0.033113	2766.7	39.1	2878	63.1	2968.4	49.9	1.86	999	607	923
	28	1	0.469034	0.018892	12.655867	0.940238	0.197022	0.006137	2.151683	0.093403	2474.6	83.7	2648	72.7	2800.3	50.8	5.35	857	278	407
	29	1	0.453952	0.00995	12.845747	0.865573	0.206283	0.006371	2.19647	0.050572	2411.8	44.4	2666.3	64.8	2875.3	50.1	1.16	1520	1468	1676
	30	1	0.219175	0.008558	6.317234	0.44859	0.210759	0.007247	4.596707	0.191723	1276.1	45.5	2017.2	62.7	2908.1	55.5	34.31	3826	131	2709
MA01	1	4	0.052411	0.002129	3.521528	0.107745	0.483872	0.012094	19.121477	0.741138	329.2	13	1531.1	24	4190	37.5	3.73	28790	7426	17327
	2	4	0.096495	0.004085	4.784704	0.173959	0.356093	0.008221	10.404859	0.439833	593.6	24	1779.8	30.9	3731.1	35.7	9.93	8930	910	5776
	3	4	0.107319	0.004298	2.749701	0.158322	0.184368	0.005312	9.289564	0.367204	657	25	1339.1	42.6	2690	47.6	13.93	13675	981	4519
	4	4	0.162525	0.004926	11.355456	0.41326	0.499547	0.012588	6.129268	0.183033	970.6	27.3	2549.7	34.7	4237	37.8	0.27	7563	30276	14592
	5	4	0.240569	0.006684	12.577219	0.390336	0.374095	0.008507	4.136633	0.114332	1389.5	34.7	2647	29	3806.1	34.2	17.87	3715	210	7876
	6	4	0.583096	0.015657	20.710966	0.584909	0.256406	0.005702	1.705504	0.045831	2961.2	63.8	3124.8	27.2	3223.5	35.4	5.89	4514	817	6318
	7	4	0.489344	0.016475	19.697146	0.620539	0.287783	0.005672	2.040118	0.070661	2566.5	71.6	3075.2	30.7	3404.7	30.8	1.51	3018	2039	2626
	8	4	0.492879	0.014146	19.242039	0.582833	0.28044	0.005572	2.020489	0.059505	2582.8	61.3	3053.3	29.4	3364.5	31.1	1.13	3956	3563	3407
	9	4	0.150883	0.004344	3.785954	0.111027	0.179139	0.004942	6.601884	0.185563	905.9	24.3	1589.4	23.6	2642.6	45.5	2.22	7711	3537	2588
	10	4	0.218355	0.006198	9.678085	0.276468	0.318197	0.006494	4.565842	0.131666	1273	32.8	2404	26.3	3560	31.4	8.25	3672	557	5337
	11	4	0.302429	0.009022	14.768342	0.576695	0.358394	0.009488	3.313032	0.100169	1703.1	44.7	2798.4	36.4	3740.9	39.4	5.39	2470	450	5176
	12	4	0.24846	0.00705	10.918103	0.334457	0.326058	0.006933	4.03866	0.117309	1430.3	36.5	2514.7	28.8	3597.1	32.7	11.9	2464	206	3692
	13	4	0.273875	0.007818	15.035284	0.404155	0.406584	0.008096	3.664237	0.103765	1560.3	39.5	2817.1	25.7	3932.7	29.8	12.85	2472	188	6853
	14	4	0.263365	0.007986	11.662232	0.399925	0.329926	0.006693	3.819401	0.117022	1506.6	40.8	2575.3	31.8	3615.5	30.9	12.67	2734	214	4850
	15	4	0.171216	0.005188	5.833727	0.177178	0.254576	0.005721	5.868208	0.177154	1018.7	28.5	1950.9	26.1	3212.1	35.7	13.15	2821	210	2451
	16	4	0.226131	0.006152	13.183562	0.35994	0.435061	0.008968	4.44256	0.122577	1314.1	32.4	2692.2	25.9	4033.7	30.7	2.76	3112	1093	7272
	17	4	0.264029	0.007267	14.429201	0.414	0.406922	0.008473	3.804506	0.104063	1510.2	37	2777.4	26.9	3933.5	31.2	5.65	3092	651	8255
	18	4	0.15724	0.004706	8.185146	0.256002	0.382741	0.009411	6.380391	0.182811	941.4	26.1	2251.2	28.1	3841.2	37.4	4.74	8395	1740	12249

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Sample	Spot name	Run	206/238	±2SE	207/235	±2SE	207/206	±2SE	238/P206	±2SE	206/238 Age	±2SE	207/235 Age	±2SE	207/206 Age	±2SE	U/Th	U	Th	Pb
	19	4	0.18866	0.005742	7.070546	0.210448	0.276488	0.006017	5.322337	0.161893	1113.9	31.1	2119.2	26.5	3341.5	33.6	14.85	3212	213	3214
	20	4	0.151947	0.006008	5.335103	0.317761	0.255586	0.009196	6.637084	0.257308	911.4	33.5	1865	48.7	3210.8	57.7	3.26	6018	2039	3555
	22	4	0.212953	0.006149	11.153436	0.369803	0.358478	0.00754	4.656393	0.13449	1244.3	32.7	2534	30.8	3742	32	12.53	4196	345	6896
	23	4	0.583802	0.016537	23.336614	0.656119	0.271238	0.005617	1.696398	0.048888	2963.8	67.5	3240.5	27.2	3312.1	32.6	3.62	790	227	335
	24	4	0.270528	0.007536	16.159084	0.69824	0.40728	0.016205	3.661663	0.104464	1543.3	38.3	2880.8	41.1	3926.3	57.8	3.28	4120	1946	9840
	25	4	0.381884	0.015252	14.054737	0.635426	0.250542	0.005162	2.614446	0.1074	2082.8	71.4	2747.2	43.2	3186.8	32.8	16.24	3054	194	1761
	26	4	0.146486	0.004124	3.125507	0.112665	0.144966	0.005011	6.756346	0.188885	881.2	23.2	1438	27.6	2282.4	58	15.73	4308	285	866
	27	4	0.205815	0.005767	10.053791	0.31092	0.334385	0.007109	4.812386	0.137509	1206.3	30.9	2438.2	28.9	3635.8	32.6	17.58	2935	174	4920
	28	4	0.179287	0.005289	7.231029	0.337866	0.276589	0.009737	5.529318	0.159727	1062.9	28.9	2134.7	40.1	3336.1	52.5	14.38	2165	158	2263
	29	4	0.178448	0.005296	4.838201	0.156684	0.18829	0.005028	5.549919	0.156514	1058.4	28.8	1790.8	27	2724.8	43.6	14.51	3808	274	1955
	30	4	0.231729	0.006639	11.559997	0.392676	0.345287	0.008405	4.277426	0.120875	1343.3	34.7	2567.4	31.9	3683.9	37.7	15.53	2532	170	4521
MU03	1	4	0.363259	0.011133	13.38306	0.389366	0.26641	0.005933	2.745148	0.083422	1997.1	52.6	2706.3	27.5	3283.5	34.8	4.31	1079	272	354
	2	4	0.263522	0.008141	10.341623	0.32163	0.285065	0.007561	3.780406	0.116053	1507.6	41.5	2465.1	28.6	3388.8	41	3.48	894	286	462
	3	4	0.386691	0.011367	14.637869	0.44944	0.271842	0.006355	2.576139	0.078064	2107	53.1	2790.9	29	3314.8	36.9	4.25	606	158	242
	4	4	0.173235	0.005619	4.516824	0.132492	0.188379	0.004561	5.764075	0.189889	1029.6	30.9	1733.2	24.6	2725.2	40.1	6.24	2633	472	457
NG01	1	3	0.329285	0.009666	7.95531	0.444468	0.173788	0.002777	3.010711	0.088844	1834.8	46.9	2225.6	50.7	2593.6	26.6	0.03	3776	116777	62915
	2	3	0.349128	0.009145	10.196257	0.55168	0.210545	0.003535	2.84095	0.075019	1930.3	43.7	2452.3	50.2	2908.5	27.5	2.16	679	316	437
	3	3	0.356093	0.010026	10.794644	0.595695	0.219813	0.004241	2.788799	0.080292	1963.4	47.8	2504.9	51.8	2977.6	31.1	2.09	454	219	306
	4	3	0.252386	0.007563	9.202339	0.517605	0.262468	0.004285	3.939318	0.117906	1450.5	38.9	2357.4	50.9	3260.6	25.6	0.54	3028	5755	3698
	5	3	0.225485	0.006459	8.119248	0.44824	0.259464	0.004395	4.411698	0.126049	1310.6	33.9	2243.5	49.9	3242.3	26.7	0.66	1748	2868	1701
	6	3	0.827946	0.078532	35.323565	3.791943	0.306775	0.00505	1.310409	0.134472	3835	280.5	3604.9	112.3	3503.7	25.3	0.01	649	95419	4498
	7	3	0.178858	0.005208	5.40703	0.297838	0.217794	0.003663	5.569374	0.160589	1060.6	28.4	1885.3	46.7	2963.3	27	2.07	1398	673	586
	11	7	0.469616	0.026663	12.727583	1.691673	0.200559	0.015325	2.188676	0.130808	2471.6	117	2646.4	116.8	2785.4	125.4	18.12	10	1	4
	12	7	0.429709	0.029684	10.261383	1.28756	0.180236	0.015332	2.369876	0.139447	2290.8	132.5	2447.4	126.1	2592.9	145.1	9.47	9	1	3
	13	7	0.314344	0.040165	8.758185	1.910531	0.198163	0.024571	3.378134	0.441089	1744.1	198.6	2219.3	204.9	2723.8	222.1	6.89	13	2	5
	14	7	0.18838	0.013568	4.573578	0.557895	0.175898	0.006692	5.541319	0.490137	1109.3	74.5	1729.9	110.8	2612.7	63.2	0.03	1966	60771	28966
	15	7	0.137977	0.005993	3.050895	0.327742	0.159741	0.006109	7.32419	0.313185	832.4	33.8	1416.2	81.9	2450.5	65.6	0.03	1481	57908	24220
NG02	1	3	0.377963	0.02314	15.04535	1.284488	0.281713	0.006574	2.719355	0.185429	2059.5	109.1	2797.7	86.4	3369.1	36.5	6.9	11676	1717	23453
	2	3	0.102712	0.004401	2.199584	0.16291	0.154489	0.003657	9.705788	0.422281	630.2	25.7	1179.9	51.9	2395	40.2	12.94	4965	393	978
	3	3	0.090406	0.003124	3.046203	0.252013	0.240677	0.010776	11.070747	0.37866	557.8	18.5	1408.5	62.6	3108.5	71.7	4.2	9439	2337	3698
	4	3	0.104645	0.004904	4.335673	0.258628	0.300137	0.008684	9.634235	0.465746	641.3	28.6	1698.4	49.7	3466.7	43.8	3.75	6624	1746	4221
	5	3	0.234248	0.007465	6.75159	0.382761	0.206534	0.005521	4.261251	0.134381	1356.6	38.9	2078.8	49.9	2876.1	43.1	7	737	111	477
NM02	1	2	0.293705	0.043726	7.389117	1.181568	0.188389	0.005231	4.528124	0.862504	1624.3	221.8	2069.5	162.1	2724.4	45.7	1.56	736	498	535
	2	2	0.361817	0.019097	9.238843	0.824773	0.185182	0.004485	2.801655	0.142084	1985	89.7	2351.5	80.8	2697.7	39.9	0.91	545	817	419
	3	2	0.207132	0.011421	5.169219	0.471089	0.180808	0.004142	4.918529	0.287766	1210.8	61.1	1836.7	80.5	2658.7	38.1	1.39	1479	1186	952
	4	2	0.413043	0.005507	9.8112	0.722472	0.172174	0.0039	2.396173	0.032106	2228.5	25.1	2416.6	67.5	2577.4	37.6	1.12	816	806	834
	5	2	0.524827	0.009451	13.643369	1.030663	0.188415	0.004742	1.888347	0.034103	2718.7	39.9	2723.4	71.1	2725.8	41.4	0.83	266	353	493
	6	2	0.511211	0.006108	12.734235	0.932746	0.180557	0.004002	1.935214	0.022905	2661.5	26	2659.8	69.4	2656.7	37	0.93	663	792	1101
	7	2	0.538059	0.007401	15.044797	1.152656	0.20268	0.006242	1.839421	0.025216	2774.8	31	2815.6	71.2	2842.7	49	1.25	211	186	424
	8	2	0.395583	0.005115	9.513185	0.695516	0.174379	0.00399	2.501333	0.031992	2148.4	23.6	2388.5	66.7	2598.5	38.3	1.95	680	383	551
	9	2	0.482978	0.007094	12.304329	0.912018	0.183956	0.004171	2.042572	0.034215	2539.7	30.9	2627	68.6	2687.4	37.6	0.88	545	713	928
	10	2	0.518092	0.006401	13.060326	0.960185	0.182804	0.004104	1.909584	0.023447	2690.8	27.1	2683.4	69	2677.1	37.2	1.02	578	632	844
	11	2	0.438258	0.005864	10.817003	0.800076	0.178871	0.004049	2.258535	0.030434	2342.5	26.3	2506.7	69.3	2641	37.4	1.04	652	637	688
	12	2	0.305269	0.004321	6.81577	0.506873	0.16179	0.00372	3.243356	0.046437	1717.1	21.4	2086.8	66.4	2472.8	39.2	1.84	989	549	509
	13	2	0.3567	0.012018	8.292176	0.675544	0.168277	0.003834	2.804109	0.106006	1964.2	57.8	2257.7	78.4	2539	38.3	1.39	817	583	537
	14	2	0.207585	0.003713	3.846322	0.286295	0.134366	0.003199	4.77552	0.084881	1215.7	19.8	1601.7	59.4	2153.7	41.2	0.63	2593	4100	2074
	15	2	0.236894	0.003241	4.584877	0.335975	0.140356	0.003256	4.179841	0.057677	1370.4	16.9	1746.2	60.7	2229.8	40.2	0.61	2991	4893	2821
	16	2	0.505764	0.00621	12.828539	0.938629	0.183982	0.004373	1.952451	0.026652	2643.9	30.1	2666.8	68.6	2691.9	35	1.44	300	204	275
	17	2	0.439678	0.008361	10.857707	0.815517	0.178972	0.003985	2.256408	0.044618	2348.3	37.6	2509.5	70.2	2642.1	36.7	1.24	583	463	525
	18	2	0.481403	0.012586	12.115913	0.929643	0.182471	0.004117	2.06787	0.05734	2531.3	55.3	2610.7	74.4	2674	37.3	1.23	454	364	450
	19	2	0.499332	0.006579	12.804553	0.940431	0.185947	0.004377	1.98342	0.025676	2610.5	28.2	2664.8	69.8	2704.8	39	1.02	325	309	416
	20	2	0.410154	0.007427	9.901121	0.737433	0.175066	0.004192	2.418606	0.044402	2214.9	34	2424.5	69.1	2604.6	40.1	1.44	339	231	271

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Sample	Spot name	Run	206/238	±2SE	207/235	±2SE	207/206	±2SE	238/P206	±2SE	206/238 Age	±2SE	207/235 Age	±2SE	207/206 Age	±2SE	U/Th	U	Th	Pb
NY01	1	3	0.520592	0.014039	12.87205	0.694122	0.178879	0.003153	1.96174	0.053827	2701.3	59.7	2670	50.9	2641	29.5	0.51	296	641	683
	2	3	0.520804	0.013774	12.823504	0.68993	0.17813	0.003026	1.959415	0.050676	2702.3	58.1	2666.5	50.5	2634.2	28.1	0.59	231	421	427
	3	3	0.483458	0.013502	12.853714	0.791515	0.192889	0.006984	2.109304	0.059357	2542.2	58.7	2668.3	57.4	2764.4	59.2	0.62	293	511	596
	4	3	0.527775	0.014736	12.978145	0.712996	0.178014	0.003587	1.93371	0.054615	2731.5	62.3	2677.3	51.8	2632.1	33.4	0.46	120	283	299
	5	3	0.510879	0.013927	12.87341	0.702042	0.182363	0.003172	1.996365	0.055147	2660	59.5	2669.8	51.5	2673.1	28.7	0.66	516	837	868
	7	3	0.392058	0.010838	13.234723	0.796586	0.243979	0.005959	2.600697	0.072228	2132	50.2	2693.2	57.6	3142.2	39.5	3.82	686	188	864
	8	3	0.40891	0.013701	11.295573	0.735506	0.199382	0.004468	2.498766	0.080066	2208.9	62.3	2543.5	58.6	2818.5	36.2	0.9	575	660	1069
	9	3	0.500968	0.014139	12.423953	0.671917	0.179979	0.003152	2.035559	0.059771	2617.4	61.1	2636.6	51.3	2651.2	29.1	0.63	232	372	432
	10	3	0.207003	0.009134	5.799911	0.472683	0.202411	0.006464	4.935841	0.216911	1212.5	48.8	1942.5	71	2843	52.4	9.7	2168	223	2239
	11	3	0.475451	0.013112	11.454097	0.626547	0.174641	0.002832	2.147595	0.061632	2507	57.6	2560.2	51.9	2601.6	26.9	0.74	538	690	998
	12	3	0.425939	0.011553	10.245493	0.55624	0.174689	0.002751	2.397483	0.065335	2287.1	52.2	2456.7	50.4	2602.1	26.3	0.78	505	615	865
	13	3	0.506394	0.01377	13.222481	0.714847	0.189365	0.003188	2.016611	0.054956	2640.9	58.9	2695.4	51	2735.6	27.6	0.59	295	483	797
	14	3	0.455835	0.012823	13.082729	0.712497	0.208411	0.004507	2.242681	0.063409	2420.7	56.8	2685.1	51.6	2890.3	35.7	0.54	366	659	1108
	16	3	0.367415	0.011349	9.513369	0.544749	0.187523	0.00316	2.78403	0.087398	2016.8	53.6	2387.9	52.2	2719.5	27.9	1.8	584	314	525
	17	3	0.466845	0.012324	11.35414	0.614863	0.176469	0.003003	2.18885	0.057747	2469.5	54.1	2552.2	50.7	2618.7	28.4	0.59	349	571	718
	18	3	0.507254	0.013317	12.975148	0.699908	0.185214	0.003143	2.014346	0.052913	2644.7	56.9	2677.5	50.8	2698.8	27.8	0.67	356	515	725
	19	3	0.44055	0.01214	11.451814	0.620975	0.188754	0.003504	2.320777	0.066409	2352.7	54.6	2560.2	50.7	2729.6	30.4	3.21	477	144	287
	20	3	0.471198	0.016249	12.487359	0.758473	0.192218	0.004675	2.172117	0.076216	2488	71.4	2640.1	57.5	2759	39.5	1.38	70	49	85
	22	3	0.325245	0.010738	9.233157	0.569443	0.205962	0.004859	3.133165	0.106769	1814.4	52.4	2357.5	57.7	2870.4	38.2	6.09	786	128	640
	23	3	0.160954	0.004393	3.474162	0.194251	0.15645	0.003441	6.298597	0.169843	962.1	24.4	1521.1	43.8	2416.1	37.3	10.46	1092	102	411
	24	3	0.432645	0.011783	12.038917	0.681883	0.202018	0.004061	2.342631	0.063782	2317.3	53	2605.9	53.4	2840.3	32.8	2.31	503	221	502
	25	3	0.424644	0.011264	11.301492	0.62366	0.192915	0.003975	2.382277	0.062468	2281.5	50.9	2548	52	2765.8	33.8	2.91	236	79	157
	26	3	0.502055	0.013994	12.635458	0.684888	0.182806	0.003082	2.015078	0.05658	2622.2	60.1	2652.5	51.4	2677.3	28	0.77	286	363	467
	27	3	0.462134	0.01539	11.96292	0.660487	0.18863	0.003612	2.195001	0.075263	2447.5	68.1	2600.6	52.1	2728.3	31.3	0.82	280	334	454
	28	3	0.18694	0.006375	5.529441	0.352068	0.215372	0.004549	5.424823	0.191227	1104.4	34.7	1900.6	57.7	2943.7	35.1	2.72	1973	713	1633
	29	3	0.3949	0.011636	14.028854	0.771508	0.258583	0.006277	2.557034	0.07436	2144.9	53.7	2750.8	52	3234.4	37.5	1.33	236	168	473
	30	3	0.40937	0.011231	10.796405	0.584504	0.191677	0.003273	2.465458	0.067166	2211.7	51.3	2505.3	50.2	2755.3	28.2	0.75	508	668	870
NY02	1	4	0.331085	0.009621	7.262073	0.204556	0.162921	0.003377	3.017986	0.089209	1843.2	46.7	2143.4	25.1	2484.7	34.8	0.62	1368	2257	1437
	2	4	0.167291	0.005626	9.806918	0.335769	0.435118	0.010966	5.989945	0.211079	996.8	31.1	2414.5	31.6	4032.2	38	1.26	818	675	2019
	3	4	0.229824	0.006001	8.49968	0.237068	0.273069	0.005886	4.340429	0.112584	1333.5	31.4	2285.2	25.2	3322	34.1	1.29	2808	2227	4488
	4	4	0.245661	0.007348	7.329241	0.252048	0.221359	0.005142	4.065031	0.122833	1415.8	38	2150.9	31.2	2988.8	37.4	1.43	1121	793	1686
	5	4	0.49611	0.018746	15.739515	0.750738	0.232151	0.006547	2.017888	0.077094	2595.5	80.9	2857.3	44.5	3064.1	44.7	0.5	223	456	764
	6	4	0.52411	0.015303	12.487056	0.390595	0.175186	0.004264	1.906167	0.056917	2715.7	64.9	2640.1	29.5	2604.8	40.6	0.36	118	327	316
	7	4	0.200248	0.005954	6.408695	0.203643	0.234567	0.005115	4.990491	0.153481	1176.4	32.1	2031.9	27.9	3081.7	35	1.52	1823	1208	1785
	8	4	0.279429	0.008127	5.681274	0.190356	0.149036	0.003214	3.572225	0.101214	1588.2	40.8	1927.2	28.5	2333.3	36.8	5.43	1113	203	194
	9	4	0.51537	0.014518	12.475308	0.373873	0.176439	0.004231	1.936791	0.053601	2678.9	61.6	2639.6	28	2616.8	40.4	0.31	171	559	590
	10	4	0.508991	0.016531	13.297516	0.446749	0.189456	0.006016	1.963827	0.063496	2651.2	70.5	2699.2	32.7	2732.1	53.3	0.6	78	128	170
	11	4	0.506684	0.014331	12.603406	0.390745	0.176595	0.004361	1.96552	0.054997	2641.8	61.2	2648.9	29.4	2617.9	42	0.47	111	234	413
	12	4	0.320662	0.008575	8.466882	0.23896	0.186962	0.00419	3.101882	0.084016	1792.8	41.9	2281.6	25.9	2713.5	37	1.21	730	610	1161
	13	4	0.27745	0.008979	7.557048	0.304973	0.191524	0.006408	3.588961	0.117667	1578.1	45.4	2177.5	36.7	2750.5	55.9	0.79	119	152	258
	14	4	0.404564	0.011546	9.439032	0.26557	0.165541	0.003587	2.450452	0.062515	2189.5	53.1	2384.4	29.2	2511.2	36.4	0.76	390	516	772
	15	4	0.342095	0.010454	10.261108	0.283039	0.211352	0.004699	2.908965	0.087406	1896.3	50.1	2458.1	25.7	2914.3	36.4	0.5	410	821	1513
	16	4	0.334876	0.009864	11.244168	0.45938	0.236956	0.006373	2.972222	0.088603	1861.5	47.7	2539.2	37.9	3095.8	42.3	0.24	739	4289	1754
	17	4	0.39274	0.012183	11.450612	0.318104	0.207037	0.005124	2.537183	0.080581	2134.7	56.5	2559.8	26	2879.4	40.1	0.32	240	755	947
	18	4	0.235275	0.00819	5.292793	0.154083	0.16033	0.003543	4.247485	0.153651	1361.4	42.9	1866.8	25.1	2457.1	37.8	1.12	2013	1818	1451
	19	4	0.23473	0.00829	8.62364	0.370604	0.261345	0.00693	4.254587	0.152315	1358.7	43.3	2295.3	39.2	3251.9	41.4	1.05	1535	1469	2548
	20	4	0.510639	0.0146	12.566461	0.38478	0.177027	0.004392	1.944645	0.049387	2666.3	55	2646.3	29.1	2622	41	0.3	157	526	666
	23	4	0.311355	0.012605	6.903249	0.239469	0.168339	0.003763	3.24622	0.132178	1746.7	62	2098.2	30.8	2539.9	37.5	3.05	1120	359	594
	24	4	0.244044	0.006929	7.882145	0.223451	0.24311	0.005773	4.135076	0.118634	1407.5	35.9	2216.8	25.6	3137.9	37.3	0.88	1091	1208	2326
	25	4	0.247242	0.006674	6.415594	0.18485	0.195022	0.004856	4.079286	0.112163	1424.1	34.6	2033.5	25.4	2781.9	40.1	0.59	1135	1876	1571
	26	4	0.489211	0.013199	11.64276	0.333417	0.177749	0.003974	2.060506	0.05624	2567	57.2	2575.5	26.8	2630.4	37	0.65	378	565	769
	27	4	0.319312	0.012523	8.203678	0.242086	0.19232	0.005363	3.177018	0.125796	1785	61.2	2252.8	26.6	2758.2	45.3	3.22	1336	411	849
	28	4	0.269945	0.008221	6.948432	0.219557	0.190581	0.005221	3.733758	0.11666	1540.3	41.8	2104	28.5	2744.4	43.9	0.42	353	821	867
	29	4	0.537874	0.014989	12.783562	0.353996	0.176176	0.003836	1.871823	0.053087	2773.9	63	2663.1	26.2	2620	32.3	0.36	124	339	491

Continued on next page

Sample	Spot name	Run	206/238	±2SE	207/235	±2SE	207/206	±2SE	238/P206	±2SE	206/238 Age	±2SE	207/235 Age	±2SE	207/206 Age	±2SE	U/Th	U	Th	Pb	
		30	4	0.441056	0.013355	10.435542	0.375359	0.174904	0.004697	2.280432	0.069775	2354.8	59.8	2472.2	34	2602.1	44.3	0.51	122	234	347
NY03	1	4	0.398987	0.010519	9.475941	0.258259	0.172328	0.003551	2.51144	0.066775	2164.2	48.5	2384.6	25	2578.9	34.6	0.61	1190	2009	1736	
	2	4	0.447133	0.012531	10.928903	0.33031	0.177888	0.003676	2.241974	0.062936	2382	55.8	2515.7	28.7	2631.8	34.7	1.12	659	645	680	
	3	4	0.431901	0.011891	10.973345	0.293562	0.184744	0.004204	2.319475	0.064097	2313.9	53.6	2520.4	25	2693.7	37.5	2.29	502	247	355	
	4	4	0.483349	0.014204	12.353505	0.392679	0.184231	0.004116	2.072777	0.061726	2541.2	61.8	2630.2	30.3	2689.5	36.1	4.36	382	87	164	
	5	4	0.325843	0.011769	7.148095	0.321998	0.158416	0.003899	3.080424	0.113435	1817.4	57.3	2126.6	40.4	2436.5	41.5	1.88	1332	704	595	
	6	4	0.332256	0.008896	7.583575	0.224732	0.166898	0.00359	3.008182	0.082206	1849.2	43.2	2182.5	26.7	2525.6	35.9	1.52	1501	986	758	
	7	4	0.14454	0.003838	2.403626	0.076347	0.121612	0.00348	6.912175	0.181162	870.3	21.6	1243.1	22.9	1976.4	50.7	0.84	3044	3600	1625	
	8	5	0.140579	0.004689	2.901431	0.083704	0.149193	0.003931	7.11719	0.242334	847.8	26.5	1382	21.9	2334.3	45.1	1.91	3210	1688	1206	
	9	5	0.071834	0.002613	1.008906	0.040441	0.101394	0.002106	13.943896	0.511929	447.1	15.7	707.6	20.5	1648.4	38.6	0.77	27492	36280	6979	
	10	5	0.15413	0.004179	3.383295	0.103078	0.159725	0.003757	6.476565	0.172995	924	23.3	1504.3	27.4	2458	45.4	0.77	2783	3725	2709	
	13	5	0.24763	0.007847	6.177713	0.200563	0.180631	0.003892	4.030623	0.131837	1426.1	40.7	2000.6	28.3	2657.5	35.6	2.77	1060	387	681	
	14	5	0.309487	0.009127	7.397873	0.211507	0.17289	0.003444	3.224556	0.095094	1738	44.9	2160.3	25.6	2584.8	33.5	1.25	1470	1202	1360	
	15	5	0.368232	0.010268	9.238641	0.250094	0.182592	0.004219	2.709253	0.076724	2020.9	48.5	2361.6	24.8	2674.9	38.1	3.31	818	254	496	
	16	5	0.16815	0.004701	3.769647	0.124208	0.164299	0.004291	5.934865	0.166054	1001.8	25.9	1585.3	25.8	2497.5	42.7	1.33	2784	2141	2195	
	17	5	0.422229	0.01079	11.574997	0.308799	0.198881	0.004384	2.361636	0.06054	2270.5	48.9	2570.3	25	2815.3	36.3	1.92	457	243	573	
	18	5	0.467792	0.012197	11.503528	0.300155	0.179036	0.003719	2.132368	0.056119	2473.7	53.6	2564.5	24.3	2642.5	34.7	0.93	679	743	908	
	20	5	0.241815	0.007115	11.191873	0.367178	0.336991	0.007343	4.12704	0.119555	1396	36.9	2538.2	30.5	3648	33.3	1.02	1406	1393	3758	
	21	5	0.378132	0.010489	9.915436	0.25325	0.182449	0.003669	2.640686	0.074232	2067.2	49.1	2426.7	23.6	2677.4	37.6	1.67	1159	684	1040	
	22	5	0.253052	0.008487	7.452328	0.288985	0.202183	0.005054	3.956811	0.131878	1453.8	43.6	2165	33.9	2841.5	40.9	1.63	1721	951	1281	
	23	5	0.382797	0.010185	10.699587	0.324139	0.193275	0.004311	2.607854	0.069943	2089.1	47.5	2496.1	28.1	2768.2	36.4	1.04	732	683	764	
	24	5	0.279781	0.008407	6.692759	0.22185	0.163643	0.003797	3.570339	0.106462	1590.1	42.3	2070.7	29.6	2492	39.3	1.1	1996	1739	1626	
	25	5	0.345683	0.010283	9.464224	0.344242	0.186641	0.005022	2.889797	0.086337	1913.7	49.3	2382.7	33.5	2710.4	44.4	1.57	1839	1107	1571	
	26	5	0.382705	0.01051	9.463812	0.250605	0.169513	0.003569	2.610833	0.071885	2088.6	49	2383.6	24.2	2551.2	35.2	1.69	934	508	552	
	27	5	0.369926	0.011584	9.834785	0.291769	0.181728	0.003875	2.707469	0.085784	2028.2	54.6	2418.2	27.9	2667	35.4	0.83	1495	1596	1976	
	28	5	0.427646	0.011907	11.327309	0.301911	0.180942	0.003628	2.337421	0.066278	2294.7	53.9	2549.9	25.2	2660.3	33.1	0.36	1097	2647	3365	
	29	5	0.305388	0.009405	11.172112	0.473755	0.250497	0.007591	3.275552	0.102034	1717.6	46.5	2533.7	41.4	3183.6	48.6	1.36	1686	1111	1910	
	30	5	0.163699	0.005183	4.663799	0.141089	0.194159	0.004975	6.1808928	0.194517	977.2	28.7	1760.1	25.2	2775.3	42.5	3.21	2382	623	1098	
PG01	3.01	3	0.269514	0.007489	6.240093	0.338338	0.165935	0.002747	3.703135	0.104534	1538.1	38.1	2009.7	47.6	2515.9	28	5.33	1409	265	479	
	6.01	6	0.568894	0.005522	14.597697	1.006167	0.187909	0.005988	1.749475	0.018673	2902.9	22.6	2788.9	65.6	2726.1	47.2	2.51	601	236	337	
	3.02	3	0.368146	0.011523	8.772155	0.492775	0.171174	0.002977	2.697736	0.074082	2020.1	54.7	2313.7	52.2	2567.9	28.9	6.06	965	160	301	
	6.02	6	0.546617	0.007026	13.914737	0.968499	0.186172	0.005596	1.826685	0.023669	2810.4	29.3	2742.9	66.5	2706.9	53	2.94	626	212	296	
	3.03	3	0.517015	0.013991	13.225553	0.712906	0.18416	0.003464	1.928405	0.05244	2686.3	59.5	2695.7	50.7	2689.2	31.5	5.69	866	155	295	
	6.03	6	0.551004	0.009007	13.828231	0.97015	0.183353	0.005914	1.814216	0.030451	2828.2	37.6	2736.7	66.9	2681.5	53.4	2.4	592	245	345	
	3.04	3	0.386479	0.010843	9.85868	0.540223	0.18337	0.003113	2.581398	0.073127	2106.1	50.5	2420.9	50.9	2682.3	28	6.07	962	160	387	
	6.04	6	0.5722	0.007267	14.428652	0.996626	0.184618	0.005963	1.744968	0.022159	2916.1	29.8	2777.7	65.9	2692.9	52.9	2.43	580	233	308	
	3.05	3	0.263978	0.007406	7.296074	0.400485	0.198698	0.003636	3.778419	0.105621	1510	37.7	2147.6	49.6	2813.9	29.7	5.78	1386	242	806	
	6.05	6	0.521115	0.008325	12.975312	0.908935	0.181308	0.005797	1.917873	0.030981	2703	35.3	2676.6	66.3	2663.1	52.9	3.13	949	303	420	
	3.06	3	0.271558	0.007382	9.539577	0.515282	0.252503	0.004091	3.670414	0.100053	1548.6	37.4	2391	49.7	3199.5	25.6	0.87	8052	9870	14368	
	6.06	6	0.18278	0.003827	3.980951	0.301179	0.156787	0.005682	5.461335	0.11362	1082	20.8	1629.2	60.8	2419.3	61.7	4.24	2735	632	1013	
	3.07	3	0.28137	0.00581	9.229508	0.512896	0.23619	0.004634	3.555933	0.122819	1597.5	48.3	2360	50.6	3092.4	31.4	1.08	3762	5046	6067	
	6.07	6	0.169872	0.003938	2.696871	0.197817	0.115082	0.003652	5.903017	0.144182	1011	21.8	1325.5	55.6	1879.4	57.4	6.53	2993	454	467	
	3.08	3	0.468751	0.016685	11.289661	0.700354	0.172966	0.003867	2.135793	0.075713	2476.1	73.1	2543.2	57.5	2583.3	37.3	4.63	149	33	54	
	6.08	6	0.298791	0.009491	6.940071	0.559538	0.167616	0.005568	3.371774	0.107303	1683.5	47	2096.4	70.4	2531.4	55.5	5.06	932	192	152	
	3.09	3	0.524551	0.014203	13.504171	0.732747	0.185618	0.003597	1.895081	0.051267	2718.1	60	2717.9	57.8	2701.7	31.7	3.26	156	47	64	
	6.09	6	0.251671	0.008314	5.40297	0.448661	0.154963	0.005445	3.972748	0.130548	1445.6	42.3	1887.1	73.8	2397.4	59.4	3.51	993	287	260	
	3.1	3	0.534682	0.014463	13.693292	0.748722	0.183999	0.003132	1.858642	0.051138	2760.8	60.9	2728	52.1	2688	28.1	14.99	1190	78	176	
	6.1	6	0.224958	0.002976	4.724372	0.354294	0.151346	0.005788	4.439542	0.059087	1307.8	15.7	1767.9	64	2354.5	67.7	4.02	1282	421	462	
	3.11	3	0.780496	0.026	23.636308	1.439572	0.217297	0.005402	1.274067	0.043585	3717.8	94.7	3252	59.2	2958.7	40.4	20.28	905	34	311	
	6.11	6	0.10736	0.00392	3.635192	0.281218	0.246114	0.011021	9.347276	0.354864	657.2	22.8	1555.3	61.7	3154.8	69.5	19.				

Sample	Spot name	Run	206/238	±2SE	207/235	±2SE	207/206	±2SE	238/P206	±2SE	Age/238	±2SE	207/235	±2SE	207/206	±2SE	U/Th	U	Th	Pb	
		3.14	3	0.227512	0.006996	4.680593	0.269642	0.147774	0.002646	4.37281	0.13394	1321.1	36.7	1762.2	48.3	2318.7	31.1	4.58	1162	259	252
		6.14	6	0.211974	0.00702	5.725659	0.441405	0.195097	0.006724	4.761058	0.166594	1238.2	37.5	1930.1	67.2	2782.3	56.4	5.3	326	62	98
		3.15	3	0.205443	0.005996	4.27474	0.241731	0.149524	0.003645	4.835028	0.143454	1204.3	32.1	1687.6	47.1	2337	42	4.67	1074	238	314
		6.15	6	0.526652	0.006547	13.060855	0.902596	0.179793	0.005923	1.895856	0.02304	2726.8	27.5	2686.5	73.9	2654.2	48.2	1.51	424	282	392
		3.16	3	0.314154	0.00954	6.876911	0.387773	0.157648	0.003024	3.159964	0.097755	1760.9	46.9	2095.1	50.3	2429.3	32.8	6.39	1122	181	180
		6.16	6	0.502318	0.00547	12.690575	0.874928	0.182965	0.00586	1.987384	0.021944	2623.3	23.5	2656.4	65.1	2678.2	53	1.48	477	322	450
		3.17	3	0.345718	0.012247	8.656897	0.508767	0.178768	0.003262	2.877545	0.103966	1913.4	58.8	2301.1	54.4	2640.2	30.4	4.37	841	206	404
		6.17	6	0.46603	0.005136	11.68245	0.808435	0.181748	0.005868	2.142198	0.023625	2465.7	22.6	2578.6	64.4	2666.9	53.8	1.47	433	291	402
		3.18	3	0.50272	0.013984	12.337985	0.672421	0.176431	0.002808	1.974316	0.055112	2625	60	2629.9	51.1	2618.6	26.5	6.16	695	117	139
		6.18	6	0.481909	0.006432	12.235378	0.855147	0.184295	0.006043	2.072932	0.028073	2534.9	28	2621.5	66.4	2689.6	54.1	2.02	264	129	201
		6.19	6	0.525947	0.009566	15.200167	1.101213	0.21029	0.008278	1.902139	0.033301	2723	40.1	2825	68.8	2900.9	64.1	2.01	73	75	75
		6.2	6	0.531502	0.008404	13.495155	0.963509	0.185582	0.006641	1.880909	0.029592	2746.8	35.3	2717.8	60.7	2699.1	58.9	1.43	118	81	114
		6.21	6	0.528641	0.008039	13.694032	0.973758	0.188926	0.007235	1.890984	0.029022	2734.8	33.9	2726.8	67.9	2736.1	55.9	1.68	92	57	71
		6.22	6	0.339592	0.012945	11.45851	0.930046	0.243466	0.007987	2.953735	0.120436	1883.6	62.8	2557.9	78.5	3141.5	52	0.69	783	1205	884
		6.23	6	0.52225	0.009207	13.074587	0.953919	0.182554	0.007213	1.915766	0.033851	2707.5	39	2681.9	69.6	2668.9	66.9	1.95	56	30	37
		6.24	6	0.506527	0.007433	12.977412	0.941697	0.187745	0.007738	1.97329	0.029084	2641	31.8	2675.1	68.8	2714.1	67.4	1.37	65	50	58
		6.25	6	0.534121	0.012655	13.720222	1.101287	0.18608	0.007244	1.878007	0.043832	2756.5	53	2727	69.5	2701.2	64.2	1.45	59	44	53
		6.26	6	0.201549	0.004499	4.656262	0.371203	0.167053	0.006226	4.963696	0.107274	1183.3	24.1	1755.7	65.3	2524.6	62.4	5.8	1828	341	604
		6.27	6	0.212739	0.003436	4.41814	0.316724	0.151221	0.004953	4.700835	0.075951	1243.1	18.3	1713.9	60.1	2357.4	55.9	3.95	1487	533	395
		6.31	7	0.46251	0.012863	12.939356	1.3593	0.202123	0.008852	2.153661	0.061495	2449.8	56.9	2673.3	99.3	2839.8	71.3	0.48	126	266	364
		6.32	7	0.494752	0.015973	14.884829	1.599679	0.217007	0.010335	2.011697	0.065727	2590.4	69	2806.1	104.1	2955.5	76.4	0.38	120	322	473
		6.33	7	0.479001	0.013308	12.231697	1.279551	0.184883	0.00716	2.084292	0.058552	2521.1	58	2618.7	98.4	2694.5	62.8	1.94	309	163	238
		6.34	7	0.529661	0.01346	13.793576	1.441122	0.188498	0.007512	1.881734	0.047542	2738.4	56.6	2738.5	89.7	2725.4	67.3	1.86	90	49	72
		6.35	7	0.359033	0.010015	8.372698	0.868933	0.168917	0.006418	2.778617	0.074897	1976.3	47.2	2269.6	92.7	2544.7	64.4	1.58	916	593	793
		6.36	7	0.432972	0.011223	10.865037	1.137451	0.181366	0.007255	2.301528	0.061258	2317.8	50.6	2507.8	100.5	2661.6	67.3	3.76	624	212	378
SC03		1	2	0.30465	0.01134	7.314658	0.608137	0.173904	0.004053	3.31694	0.118991	1712.1	55.8	2144.5	73.3	2593.8	38.7	1.47	1283	879	643
		2	2	0.49262	0.007994	12.720507	0.961305	0.188027	0.004208	2.030645	0.032846	2581.3	34.5	2662.7	61.7	2723.6	36.8	5.98	400	62	98
		3	2	0.290134	0.01035	6.297777	0.538539	0.156873	0.003897	3.481566	0.124259	1640.3	51.7	2010.7	75.4	2419.7	42.4	4.89	1482	450	253
		4	2	0.364508	0.004056	8.549357	0.622719	0.169677	0.00373	2.735695	0.026824	2003.4	19.2	2291	66.1	2553.2	37.1	136.4	1123	8	46
		5	2	0.321838	0.006431	7.078341	0.540565	0.159378	0.003779	3.113534	0.063615	1798.1	31.4	2119.2	68.4	2447.1	40.4	17.73	870	48	62
		6	2	0.289293	0.003154	6.072568	0.443198	0.151931	0.003334	3.454749	0.038053	1637.9	15.8	1986.1	63.7	2369.7	42.1	20.45	1228	57	60
		7	2	0.558806	0.00745	14.981469	1.101565	0.194386	0.004239	1.78972	0.024249	2861.2	30.9	2813.4	70.6	2778.5	35.6	1.87	872	438	485
		8	2	0.16214	0.002517	2.244613	0.168132	0.100334	0.002312	6.172373	0.095584	968.5	14	1194.3	52.5	1628.4	42.9	224.83	2904	13	25
		9	2	0.127427	0.005316	8.338067	0.687529	0.473882	0.010417	7.966358	0.330755	772.5	30.3	2262.7	73.1	4160.9	32.2	1.89	6424	3211	13541
		10	2	0.168765	0.003213	10.398282	0.7814	0.44761	0.009856	5.938385	0.114965	1005.1	17.7	2469.3	70.6	4076.2	32.8	2.57	5607	2064	15879
		11	2	0.169947	0.002911	3.692959	0.302744	0.157365	0.005877	5.895989	0.099223	1011.7	16	1565.6	64.4	2417.7	63.3	19.44	3265	156	2213
		12	2	0.314401	0.005142	7.76505	0.587277	0.17894	0.004177	3.186431	0.051096	1762	25.1	2202.5	67	2641.3	38.7	5.37	739	128	412
		13	2	0.547653	0.007415	14.716262	1.080095	0.194808	0.004296	1.828012	0.025202	2814.9	31	2796.5	70.2	2782	36.3	2.11	450	199	423
		14	2	0.395002	0.006226	9.528818	0.702987	0.174176	0.003773	2.526699	0.036544	2151.8	25.5	2389.6	69.6	2597.1	36.2	22.62	1030	43	88
		15	2	0.500974	0.005708	12.692056	0.92995	0.183691	0.004097	1.997442	0.022945	2617.7	24.5	2656.6	69.2	2685.2	37	15.36	577	58	114
		16	2	0.494815	0.005705	12.541174	0.917789	0.18375	0.003953	2.022425	0.023594	2591.2	24.6	2645.4	69.1	2686	35.6	1.34	1039	760	754
		17	2	0.495236	0.006857	12.432351	0.913829	0.181792	0.004058	2.021815	0.027521	2592.8	29.5	2637	68.4	2668	36.9	14.3	467	37	97
		18	2	0.351781	0.007058	8.058627	0.607653	0.166058	0.003709	2.852712	0.060925	1942.4	33.9	2235.9	69.9	2516.9	37.3	24.35	1273	54	69
		19	2	0.350022	0.0191	8.140205	0.80351	0.167639	0.004483	2.929468	0.143362	1928.8	89.7	2230.6	85.3	2530.9	44.3	11.6	804	66	92
		20	2	0.282145	0.018301	8.793949	0.687245	0.2325	0.013649	3.674065	0.214624	1596.1	90.8	2313.5	72.7	3043	92.9	1.61	1141	883	1590
		21	2	0.373166	0.014145	8.772771	0.757531	0.169891	0.004133	2.714364	0.101029	2041.2	66.2	2306.2	78.3	2554.4	40.7	18.13	1138	74	80
		22	2	0.34534	0.021382	7.856422	0.870637	0.162691	0.005478	3.005935	0.192341	1904.7	102.7	2186.7	103.7	2476.3	58	19.63	1065	135	89
		23	2	0.480975	0.012197	12.085703	0.938931	0.182006	0.003943	2.090974	0.056	2529.6	53.5	2607.7	75.3	2670.2	36.1	8.97	1363	146	160
		24	2	0.505164	0.007119	12.815233	0.941029	0.183949	0.004036	1.981709	0.027582	2635.5	30.4	2665.6	68.9	2687.6	36.5	0.91	723	767	862
		25	2	0.521254	0.007584	13.758377	1.036976	0.191191	0.00443	1.920747	0.028377	2703.9	32.2	2731.5	72.2	2750.8	38.2	2.1	511	237	550
		26	2	0.332752	0.005467	10.352519	1.020233	0.224303	0.013179	3.010461	0.051251	1851.3	26.5	2249.6	88.6	2986	91.6	0.8	1077	1307	3553
		27	2	0.252241	0.003376	4.69528	0.343339	0.135062	0.003303	3.967403	0.053121	1449.9	17.4	1766.1	61.2	2162.4	42.4	120.68	2808	24	382
		28	2	0.432314	0.006771	10.524106	0.780653	0.176484	0.003962	2.307833	0.032877	2315.6	30.7	2481.1	69.3	2618.7	37.7	1.52	560	360	385
		29	2	0.306682	0.003226	7.868711	0.573128	0.186406	0.003983	3.260559	0.034083	1724.3	15.9	2215.9	65.6	2709.8	35.1	31.56	5174	160	1881
		30	2	0.31769	0.004624	7.1494	0.53217	0.163077													

Sample	Spot name	Run	206/238	±2SE	207/235	±2SE	207/206	±2SE	238/P206	±2SE	206/238 Age	±2SE	207/235 Age	±2SE	207/206 Age	±2SE	U/Th	U	Th	Pb
	31	2	0.545758	0.008937	14.617862	1.088298	0.194127	0.004537	1.836561	0.029676	2806.7	37.2	2789.5	71.3	2775.7	38.6	1.36	265	240	334
	32	2	0.379457	0.003942	9.410973	0.686705	0.179736	0.003878	2.637826	0.027213	2073.6	18.4	2378.6	66.8	2649.4	35.8	3.55	1618	501	473
	33	2	0.47964	0.010911	11.913296	0.944167	0.179685	0.004613	2.095652	0.050008	2524.2	47.8	2593.3	75.7	2647.2	42.1	1.79	1308	955	1356
	34	2	0.508878	0.006543	12.864621	0.943657	0.183221	0.004011	1.968479	0.025628	2651.5	28	2669.3	69.6	2681.1	36.1	4.24	771	202	343
	35	2	0.325045	0.026333	7.671221	1.032792	0.167011	0.006297	3.26198	0.24441	1802.6	126.2	2147.3	116	2518.1	62.4	2.51	1191	532	414
	36	2	0.368987	0.035227	9.023304	1.297683	0.171561	0.007373	2.998278	0.327344	2004.2	167.5	2271	146.4	2558.2	75.8	2.41	1209	612	456
	37	2	0.161952	0.008382	4.065813	0.727573	0.174566	0.018168	6.319947	0.293339	966.1	46	1585.1	117.8	2526	144.3	0.73	3138	4801	3909
	38	2	0.282993	0.004912	6.426333	0.47912	0.164534	0.00355	3.546035	0.063611	1606.1	24.8	2034.8	66.7	2501.7	36.2	3.42	1192	461	281
	39	2	0.347877	0.005137	8.149109	0.604484	0.169715	0.003702	2.882618	0.044252	1924.1	24.7	2246.8	68.4	2553.7	36.7	29.76	844	37	81
	40	2	0.32763	0.04046	6.954704	1.252898	0.14637	0.007655	3.488553	0.389639	1799.6	191.2	2007.8	157.7	2279.7	93.9	0.44	5156	17445	8501
	42	5	0.353722	0.029541	8.85667	1.069792	0.181328	0.007878	2.987815	0.253824	1940.3	140.5	2279.7	115.1	2655.8	74.7	4.12	816	213	274
	43	5	0.41476	0.015986	10.26099	0.672959	0.182471	0.005857	2.413431	0.0971	2235.9	73.2	2457.4	62.4	2674.4	53.1	2.98	1211	485	391
	44	5	0.53569	0.019909	13.859379	0.894573	0.190749	0.006112	1.864888	0.070984	2764.9	83.9	2739.6	61.7	2747.7	52.7	60.16	1070	16	35
	45	5	0.159347	0.006286	10.789417	0.709716	0.492466	0.015917	6.266388	0.249776	953.1	35	2504.6	61.6	4218	47.9	69.37	3309	39	4103
	46	5	0.238227	0.008645	5.571814	0.358757	0.170173	0.005466	4.187722	0.152781	1377.4	45	1911.4	55.6	2558.3	53.5	3.85	1964	464	658
	47	5	0.39226	0.014412	9.674902	0.629346	0.179493	0.005768	2.544512	0.095512	2132.9	66.9	2403.4	60.6	2647.2	53.3	4.75	1419	258	365
	48	5	0.267157	0.018266	8.385937	1.138991	0.226117	0.01725	3.760728	0.265857	1525.2	93.2	2259.6	129.2	3013.6	127.5	139.04	1705	10	1354
	49	5	0.469163	0.020261	12.141687	0.832781	0.187752	0.006134	2.138409	0.096321	2477.8	89.4	2612.4	66.3	2721.2	53.7	1.71	407	221	265
	50	5	0.278697	0.01404	4.924615	0.383573	0.127885	0.004288	3.602948	0.19479	1583.8	71.3	1802.7	67.7	2067.6	59	52.92	3450	54	3848
SG01	1	4	0.622889	0.044279	20.094873	1.526099	0.246764	0.005661	1.648531	0.12833	3112.3	178.5	3085	77.6	3162.7	36.6	0.25	2841	11264	8119
	2	4	0.285244	0.013654	6.103593	0.408896	0.163808	0.005098	3.583981	0.161927	1615.2	68	1977.5	56.4	2488.8	52.2	1.45	1801	1321	308
	3	4	0.318755	0.009986	7.010094	0.252638	0.169854	0.004306	3.1722	0.099523	1783	48.8	2110.1	32.3	2552.7	42.7	7.44	1550	230	548
	4	4	0.344155	0.011421	7.743551	0.272085	0.174612	0.003887	2.935601	0.09848	1906.1	54.8	2200.4	32.2	2600.9	37.4	36.21	2130	67	64
	5	4	0.303582	0.008843	7.037752	0.196573	0.178003	0.003735	3.323867	0.096829	1708.8	43.7	2115.8	24.8	2633.1	34.9	3.12	1891	605	465
	6	4	0.30359	0.009495	6.720311	0.248088	0.169604	0.003849	3.325756	0.103477	1708.7	46.9	2073.3	32.3	2551.9	37.9	1.24	3193	2671	1459
	7	4	0.278416	0.007821	5.799518	0.200719	0.159463	0.003725	3.620797	0.104522	1583.1	39.5	1944.1	30.6	2447.4	39.5	1.7	2199	1312	646
	8	4	0.516329	0.014602	13.765166	0.413804	0.202414	0.004046	1.949264	0.055262	2683	62.1	2732.6	28.5	2844.6	32.5	10.78	908	94	83
	9	4	0.219321	0.007235	4.224152	0.151387	0.145555	0.003389	4.556089	0.138013	1278	38.6	1677.1	30.3	2292.3	39.7	1.04	3009	3971	1561
	10	4	0.372021	0.011586	9.359027	0.284207	0.187027	0.00374	2.697174	0.085107	2038.3	54.5	2372.8	27.6	2715.1	33.2	8.41	1769	334	848
	11	4	0.252434	0.008609	5.792718	0.274283	0.161814	0.004253	3.955454	0.136915	1450.3	44.4	1939.4	41.5	2470.7	44.9	2.56	2544	990	1125
	12	4	0.239672	0.015948	5.475199	0.476381	0.161073	0.004577	4.258719	0.294006	1381.9	83.1	1879.4	75.6	2463.1	47.8	4.49	2155	659	534
	13	4	0.281304	0.007753	7.003459	0.195215	0.177728	0.003701	3.539173	0.097874	1597.7	39	2111.2	24.8	2630.3	34.6	3.1	2168	714	571
	14	4	0.455277	0.012097	13.176475	0.345959	0.207783	0.004335	2.187354	0.058192	2418.4	53.6	2691.9	24.8	2886.8	33.7	7.09	1064	150	170
	15	4	0.278848	0.009367	6.664909	0.257746	0.171891	0.003594	3.585676	0.125495	1585	47.4	2065.2	35.1	2574.7	35.1	2.44	2034	947	1469
	16	4	0.332167	0.009077	8.788868	0.235127	0.191747	0.003816	3.003688	0.082732	1848.7	44	2315.9	24.4	2756.1	32.6	3.19	1737	552	558
	17	4	0.256948	0.007487	5.41995	0.186015	0.153728	0.003601	3.886328	0.110228	1474	38.3	1886.8	28.5	2385.9	39.2	2.09	2221	1079	1272
	18	4	0.225464	0.005881	4.286352	0.111814	0.139204	0.00284	4.42833	0.114503	1310.6	30.9	1690.5	21.4	2215.9	35.5	1.71	2801	1643	1073
	19	4	0.465544	0.013179	12.751282	0.414869	0.200397	0.004228	2.1471	0.06099	2463.5	58	2659.8	30.6	2828	34.5	23.6	1354	58	481
	20	4	0.238423	0.007984	4.79654	0.236349	0.148547	0.004442	4.193399	0.140702	1378.3	41.6	1781.9	41.3	2326.4	51.1	2.67	2531	950	1176
	21	4	0.251741	0.01153	5.777068	0.348445	0.162337	0.004178	4.007317	0.198503	1445.6	59.8	1931.6	54.8	2476.4	44.2	10.41	2144	221	232
	22	4	0.424807	0.02854	16.30206	1.187222	0.272455	0.009731	2.364117	0.159761	2279.2	129	2888.7	71.3	3317.2	57.1	4.1	3535	854	8301
	23	4	0.237871	0.008115	5.14962	0.231874	0.153321	0.003939	4.199185	0.144085	1375	42.3	1839.4	38.6	2379.5	44.6	0.79	2292	3078	1045
	24	4	0.141788	0.003735	2.021315	0.054826	0.101196	0.002077	7.014555	0.18598	854.7	21.1	1124.4	16.4	1644.5	38.3	85.44	5196	59	514
	25	4	0.37993	0.011782	10.375007	0.371773	0.19494	0.004094	2.621047	0.082141	2075.5	55.1	2467.3	33.7	2783.1	34.5	8.99	2045	218	308
	26	4	0.572134	0.016888	17.400999	0.5384	0.215509	0.004394	1.74185	0.052835	2915.7	69.6	2955.8	30.2	2946.2	32.8	25.06	1111	43	124
	27	4	0.201611	0.005854	3.94106	0.11629	0.139285	0.002956	4.943182	0.154683	1183.8	31.6	1626.4	17.9	2216.7	36.9	2.81	4782	1793	2984
	28	4	0.407134	0.01123	11.269433	0.310317	0.19686	0.003923	2.447533	0.068553	2201.4	51.5	2545	25.8	2799.2	32.6	9.72	1265	124	331
	29	4	0.328213	0.008668	8.535743	0.225555	0.185172	0.003663	3.035288	0.079628	1829.6	42	2289.3	24	2698.6	32.7	2.47	1857	724	908
	30	4	0.216555	0.006206	4.198966	0.143316	0.138126	0.003109	4.607302	0.1333	1263.4	32.9	1672.1	27.4	2201.5	38.5	2.78	4074	1421	2208
SG02	1	4	0.207614	0.005665	4.460247	0.125113	0.155445	0.003264	4.806721	0.133879	1216	30.3	1723	23.3	2405.1	35.8	1.6	3214	2118	1556
	2	4	0.247111	0.009306	5.449608	0.241934	0.159071	0.003603	4.057962	0.155566	1422.8	48.2	1889.1	38.1	2443.9	38	1.57	2739	1795	1770
	3	4	0.364718	0.011425	9.567816	0.330662	0.190204	0.00382	2.743199	0.088795	2003.7	54.2	2391.7	32	2742.7	33.1	1.37	2090	1548	1394
	5	4	0.125268	0.003916	1.673977	0.057171	0.097564	0.002427	7.975226	0.249217	760.7	22.4	998.1	21.5	1575.6	46	7.66	6539	852	958
	6	4	0.237525	0.013261	6.9977	0.537928	0.214783	0.007782	4.217435	0.243964	1373.3	69.3	2107.4	70.4	2939.6	60.2	2.95	4198	1417	6048

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Sample	Spot name	Run	206/238	±2SE	207/235	±2SE	207/206	±2SE	238/P206	±2SE	206/238 Age	±2SE	207/235 Age	±2SE	207/206 Age	±2SE	U/Th	U	Th	Pb	
		7	4	0.53065	0.014469	16.917387	0.443031	0.232608	0.004707	1.881587	0.052014	2743.7	61.1	2929.7	25.3	3068.8	32.2	0.79	1024	1284	2470
		8	4	0.606795	0.018173	17.85032	0.514476	0.215906	0.004437	1.646063	0.049174	3056.4	72.8	2981.1	27.7	2949.3	33.5	5.56	1781	318	808
		9	4	0.29609	0.0086	7.486539	0.219219	0.184587	0.003579	3.375211	0.102876	1671.5	43	2170.3	26.7	2693.5	32	1.58	2163	1359	1629
		10	4	0.389305	0.01122	10.538034	0.304798	0.196156	0.004109	2.56299	0.073562	2119.4	52	2482.8	26.9	2793.4	34.3	14.36	1519	108	212
		11	4	0.344335	0.009837	8.911715	0.256203	0.188694	0.004022	2.897479	0.084497	1907.1	47.3	2328.1	26.4	2729.1	35	2.87	1420	532	771
		12	4	0.259474	0.006814	5.685401	0.147673	0.159826	0.003154	3.84092	0.099711	1487.1	34.8	1928.9	22.4	2452.7	33.2	10.21	2799	277	504
		13	4	0.407298	0.010695	11.265894	0.28807	0.201909	0.004074	2.447292	0.063313	2202.4	48.9	2545.1	23.9	2840.4	33.1	2.17	1881	900	728
		14	4	0.432047	0.011924	11.667261	0.325567	0.195606	0.004231	2.307281	0.063665	2314.8	53.7	2577.8	26	2788.7	35.5	12.05	1086	101	127
		15	4	0.552297	0.014786	16.285749	0.450401	0.214162	0.004219	1.805283	0.047977	2834.5	61.3	2893.3	26.6	2936.4	31.8	3.23	1270	403	532
		16	4	0.572559	0.017984	16.387994	0.477632	0.207185	0.004663	1.742421	0.054488	2917.9	73.6	2899.4	27.7	2882.6	36.2	1.71	1749	1038	941
		17	4	0.292077	0.009138	7.43899	0.236643	0.187637	0.003886	3.418008	0.109016	1651.7	45.7	2165	28.7	2720.5	34.3	2.25	1804	821	864
		18	4	0.384792	0.013172	10.014177	0.345935	0.190188	0.003755	2.605248	0.092751	2097.3	61.6	2433.6	32.6	2742.7	32.5	3.7	2001	593	783
		19	4	0.307933	0.010967	7.599517	0.298449	0.187785	0.003845	3.258441	0.118207	1729.4	54.1	2181.2	35.4	2652.3	35.3	4.16	2040	496	487
		20	4	0.595504	0.015866	18.901101	0.500117	0.231711	0.004606	1.675811	0.043917	3011.3	63.9	3036.3	25.6	3065.5	27.9	1.45	497	339	547
		21	4	0.26423	0.008747	6.406185	0.226512	0.173906	0.003659	3.781616	0.127089	1510.8	44.7	2030.7	31.1	2594	35	3	1998	745	709
		22	4	0.377856	0.00984	10.239192	0.265579	0.194695	0.003849	2.632068	0.068518	2066.2	46	2456.4	24	2781.2	32.4	2.26	2155	966	965
		23	4	0.621872	0.016605	20.51915	0.536494	0.235762	0.004779	1.599421	0.042916	3116.9	66.1	3115.7	25.2	3090.3	32.4	1.22	414	337	505
		24	4	0.369499	0.011254	9.7533	0.295411	0.187785	0.003788	2.694762	0.082288	2026.5	53	2410.7	28	2721.7	33.2	3.55	2119	604	707
		25	4	0.375657	0.01228	10.930793	0.356937	0.207381	0.00403	2.653076	0.08678	2055.1	57.5	2515.7	30.5	2884.2	31.7	1.91	1202	626	756
		26	4	0.303913	0.0091	7.82842	0.232628	0.183232	0.003905	3.274141	0.098981	1710.3	45	2210.6	26.7	2680.9	35.2	2.94	1471	525	558
		27	4	0.238475	0.006211	5.818998	0.146144	0.174722	0.003449	4.166649	0.107531	1378.7	32.3	1949.1	21.7	2602.3	32.8	0.24	2459	13462	1038
		28	4	0.17092	0.006111	3.577754	0.13395	0.149113	0.00307	5.831069	0.195492	1016.9	33.5	1543.3	29	2334.6	35.4	9.8	2371	243	542
		29	4	0.268386	0.007461	6.599934	0.204041	0.176221	0.003909	3.706456	0.104304	1532.5	38	2058.4	27.7	2615.9	37.4	24.37	2066	85	424
		30	4	0.602797	0.016441	19.237676	0.507615	0.231184	0.004807	1.65175	0.044302	3040.5	65.9	3053.3	25.3	3058.8	33.2	1.35	416	311	522
SR01		1	3	0.405318	0.010641	10.150466	0.54759	0.182616	0.003194	2.475251	0.064603	2193.3	48.8	2448.2	49.8	2675.3	28.9	1.02	374	360	372
		2	3	0.38315	0.010902	9.33467	0.538062	0.17743	0.003728	2.620657	0.075835	2090.7	50.9	2370.1	52.5	2626.9	34.7	1.29	288	217	250
		3	3	0.496433	0.013801	12.223668	0.665343	0.179587	0.003707	2.022831	0.05588	2598	59.4	2621.2	50.8	2647.8	28.5	1.35	223	161	209
		4	3	0.492709	0.014883	12.388721	0.687217	0.183571	0.003413	2.04139	0.061731	2581.4	64.3	2633.3	51.7	2683.6	30.6	1.15	87	74	99
		5	3	0.574471	0.016893	23.132324	1.356253	0.293572	0.006048	1.750317	0.052303	2925.2	69.3	3229.8	57.7	3434.3	31.8	1.46	336	225	338
		6	3	0.417251	0.014936	10.229719	0.611911	0.178889	0.002901	2.419995	0.088862	2246.3	68.2	2452.7	56	2641.5	26.9	1.88	347	181	221
		7	3	0.459179	0.014147	11.259799	0.654984	0.178912	0.003329	2.192182	0.070002	2435	62.8	2542.6	54.9	2641	30.8	2.32	165	70	93
		9	3	0.272758	0.00799	7.517817	0.479353	0.201599	0.006487	3.682832	0.107473	1554.7	40.4	2173.2	56.7	2835.5	51.6	1.53	568	372	413
		10	3	0.319407	0.009672	7.325575	0.424267	0.167295	0.002941	3.14879	0.097421	1786.5	47.4	2150.5	52.1	2529.5	29.4	0.77	550	713	584
		11	3	0.243899	0.009165	5.604224	0.360623	0.167662	0.003157	4.135192	0.152426	1406.3	47.4	1913.8	54.7	2532.9	31.6	1.94	1189	638	521
		12	3	0.508848	0.013553	12.569187	0.68222	0.180744	0.00329	1.973914	0.052397	2651.4	57.9	2647.4	51.5	2658	30.2	1.22	283	242	315
		13	3	0.516512	0.013878	12.81557	0.691919	0.18138	0.00312	1.944698	0.052314	2684	59	2665.8	50.2	2664.2	28.7	1.34	261	203	270
		14	3	0.293614	0.007707	7.523043	0.443786	0.187573	0.005701	3.419571	0.090932	1659.5	38.5	2173.2	53.7	2713.7	52	1.25	624	528	490
		15	3	0.428356	0.011597	10.750212	0.579158	0.183645	0.00331	2.344517	0.062991	2298	52.3	2501.5	50.1	2688.5	26.5	2.3	217	99	141
		16	3	0.362243	0.00981	8.78086	0.472174	0.177243	0.002987	2.770989	0.076049	1992.6	46.5	2315.3	49.1	2626.2	28.1	1.99	774	408	409
		17	3	0.418723	0.012723	9.795147	0.538345	0.171143	0.003016	2.399858	0.0724	2254.1	57.7	2415.2	50.4	2567.6	29.4	1.62	485	312	308
		18	3	0.345327	0.01031	8.334314	0.462859	0.176999	0.003116	2.906735	0.085751	1912	49.3	2267.6	50.5	2623.9	29.6	0.9	664	763	744
		19	3	0.382648	0.0113	9.288755	0.511884	0.177799	0.003254	2.625646	0.078813	2088.1	52.8	2366.1	50.7	2630.7	30.2	1.57	404	264	311
		20	3	0.405898	0.010609	9.334125	0.500368	0.168374	0.002799	2.469837	0.065424	2196	48.7	2371.2	49.4	2540.3	28	2.53	640	256	252
		21	3	0.378998	0.011928	10.749281	0.686585	0.207159	0.005326	2.641049	0.082946	2070.8	55.7	2496.6	59.4	2878.9	42.2	0.82	518	611	660
		22	3	0.639134	0.027388	28.076605	2.113113	0.320633	0.008127	1.572215	0.065882	3181.5	107.1	3413	70.1	3569.8	37.8	1.55	195	120	304
		23	3	0.614972	0.01691	24.885245	1.350007	0.295912	0.005144	1.620323	0.044444	3089.5	67.5	3303.4	52.9	3447.9	27.1	1.08	212	187	304
		24	3	0.509838	0.014271	12.434596	0.675084	0.178657	0.003501	1.954091	0.056248	2655.4	61.2	2637.3	51.3	2638.3	32.3	0.97	75	73	107
		25	3	0.521555	0.014812	12.864631	0.719996	0.180662	0.004172	1.908697	0.053854	2705.1	62.7	2668.5	52.2	2655.4	38.4	1.07	59	53	81
		26	3	0.422324	0.012514	10.036277	0.560794	0.174098	0.002955	2.354573	0.069501	2270.6	56.7	2437.4	51.4	2596.3	28.2	9.61	496	50	75
		27	3	0.46973	0.013107	11.957801	0.643873	0.185787	0.003155	2.10752	0.066665	2481.9	57.5	2600.8	50.3	2703.9	28.1	1.66	443	260	367
		28	3	0.451987	0.01289	10.948978	0.610441	0.177213	0.003142	2.196199	0.06328	2403.5	57.3	2517.7	52.2	2625.4	29.4	1.42	412	283	347
		29	3	0.427835	0.01146	10.182674	0.549057	0.174188	0.003113	2.315763	0.06187	2295.7	51.7	2451.2	49.9	2596.7	29.9	1.18	548	456	513
		30	3	0.289837	0.008791	6.025213	0.342701	0.152015	0.00263	3.421581	0.105876	1640.3	44	1978	50.3	2367.3	29.7	2.24	801	358	327
SR02		1	4	0.441832	0.012936	10.656304	0.337819	0.174284	0.003571	2.257819	0.068386	2358.2	58.1	2491.9	30.1	2597.8	34	1.96	469	245	313

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Sample	Spot name	Run	206/238	±2SE	207/235	±2SE	207/206	±2SE	238/P206	±2SE	206/238 Age	±2SE	207/235 Age	±2SE	207/206 Age	±2SE	U/Th	U	Th	Pb
2	4	0.348557	0.010315	8.448334	0.230748	0.176242	0.004274	2.864378	0.084848	1927.2	49.3	2279.8	24.8	2614.9	40.5	2.04	883	439	694	
3	4	0.424592	0.011549	10.306057	0.291001	0.175982	0.003772	2.34891	0.064506	2281.1	52.3	2462.1	25.9	2614	35.5	1.58	607	390	566	
4	4	0.368413	0.010276	8.50741	0.233533	0.168882	0.003426	2.711286	0.076828	2021.6	48.5	2286.1	25.2	2545.3	34	3.21	773	246	368	
5	4	0.24023	0.008497	5.206252	0.182623	0.159004	0.003289	4.178901	0.14548	1387.2	44.1	1851.5	29.7	2443.6	34.9	0.77	1887	2474	1867	
6	4	0.486103	0.013811	11.529377	0.339318	0.174407	0.003602	2.058028	0.05935	2553.2	60	2565.8	27.6	2598.9	34.6	1.73	515	301	485	
7	4	0.177651	0.008901	4.627544	0.21692	0.192877	0.004103	5.691091	0.278554	1053.2	48.6	1750.7	38.5	2765.5	35.1	3.2	1081	336	700	
8	4	0.358798	0.011464	8.165723	0.271746	0.167612	0.0036	2.794983	0.0872	1975.6	54.2	2247.6	30	2532.1	36.3	1.97	328	174	279	
9	4	0.393938	0.01102	8.866793	0.258193	0.1656	0.003893	2.539524	0.071849	2140.6	51	2323.4	26.6	2511	39.1	1.87	295	160	244	
10	4	0.189723	0.010665	3.898191	0.282163	0.150546	0.003942	5.342971	0.306789	1118.6	57.8	1604.5	57.7	2349.4	43.9	4.03	1427	351	518	
11	4	0.43297	0.012894	12.999537	0.460093	0.210284	0.005879	2.300057	0.068763	2318.8	58	2678.6	33.5	2905.6	45.1	1.44	242	168	387	
12	4	0.343062	0.009423	8.321308	0.292008	0.169642	0.004395	2.904738	0.080104	1901.3	45.3	2265.7	31.5	2552.3	43.4	2.55	432	169	184	
13	4	0.306792	0.009331	7.841801	0.25327	0.181407	0.004596	3.25493	0.09956	1724.6	46	2212	29.2	2663.5	41.7	2.35	301	130	162	
14	4	0.48409	0.013035	12.25357	0.327798	0.17865	0.003777	2.06384	0.056583	2544.7	56.8	2623.5	25.3	2638.7	35.3	0.8	649	811	881	
15	4	0.479886	0.012617	12.183961	0.328256	0.178026	0.003827	2.082795	0.055276	2526.6	55	2618.4	25.2	2633.3	35.8	1.52	464	303	381	
16	4	0.267372	0.00855	6.805626	0.339481	0.178999	0.005109	3.754281	0.116137	1526.9	43.3	2079.9	42.3	2638.8	45.7	2.48	1342	544	885	
17	4	0.19772	0.0054	3.986998	0.106907	0.142446	0.002933	5.067634	0.136247	1162.9	29	1631.1	22	2255.6	35.7	1.71	3787	2196	1633	
18	4	0.16026	0.005643	4.96758	0.19252	0.222206	0.005161	6.257397	0.208463	958.1	31.2	1813	32.1	2995.5	37.5	2.04	3866	1875	2674	
19	4	0.21056	0.005747	3.596409	0.099393	0.120526	0.002482	4.747924	0.145399	1235.3	34.7	1548.2	21.8	1962.4	36.7	0.31	5014	15902	9496	
20	4	0.428808	0.012925	10.392702	0.359421	0.172424	0.005273	2.338899	0.06837	2299.8	58	2468.7	32.3	2576.7	52.8	2.33	116	50	65	
21	4	0.20879	0.006432	6.8162	0.217649	0.228051	0.005286	4.770443	0.144106	1222.3	34.2	2087.3	27.9	3037.1	37.4	3.35	983	292	1002	
22	4	0.290768	0.01054	8.130446	0.295095	0.198141	0.004037	3.435237	0.128258	1644.7	52.8	2243.9	33.1	2809.8	33.4	2.85	959	334	1279	
23	4	0.304593	0.008753	7.842495	0.240085	0.182215	0.003963	3.270103	0.093799	1713.8	43.2	2212.1	27.7	2671.6	36.3	1.57	662	414	535	
24	4	0.387494	0.010587	9.251559	0.251641	0.168493	0.003608	2.57007	0.070171	2110.9	49.2	2362.6	24.9	2541	35.8	1.44	570	395	427	
25	4	0.455789	0.012915	11.671757	0.332351	0.180812	0.003642	2.186331	0.061107	2420.4	57	2577.5	26.4	2659.1	33.3	0.95	558	584	780	
26	4	0.419171	0.01221	10.871411	0.326391	0.181989	0.004655	2.375861	0.069509	2256.5	55.5	2511.7	27.9	2669.2	42.6	2.17	259	118	133	
27	4	0.21433	0.008669	5.636366	0.208091	0.186797	0.004093	4.682953	0.197704	1251	46.2	1919.4	32.4	2712.4	36.6	1.29	2209	1679	1515	
28	4	0.296499	0.016902	11.752442	0.959438	0.278004	0.011413	3.384724	0.197048	1672.6	84.1	2576.5	80.1	3347.3	67.3	3.01	1213	395	2617	
29	4	0.256427	0.006919	5.993261	0.19585	0.165839	0.003704	3.887451	0.104602	1471.4	35.5	1973.2	28.1	2513.9	37.2	3.86	1760	464	456	
30	4	0.332342	0.01178	8.491415	0.396253	0.180606	0.005599	3.009882	0.113595	1849	57.4	2280.8	44.7	2655.9	42.9	2.86	846	288	450	
TR02	1	0.571731	0.016381	20.974232	0.620723	0.264218	0.006075	1.737182	0.048058	2914.6	66.7	3137	28.2	3270.8	35.4	1.84	101	55	73	
	2	0.597618	0.015535	23.071929	0.585549	0.275294	0.00558	1.661667	0.043004	3020	62.6	3229.7	24.8	3335.2	31.9	1.68	127	76	125	
	3	0.245954	0.008988	5.739004	0.264705	0.165854	0.004237	4.065183	0.157106	1416.7	46.7	1931.6	41.2	2512.6	43.4	4.1	857	211	185	
	4	0.659967	0.017487	25.486757	0.681746	0.275359	0.005578	1.502173	0.03471	3266.6	68	3326.5	26.5	3335.6	31.4	1.8	140	78	121	
	5	0.501735	0.014555	17.625449	0.494317	0.251524	0.005561	1.983308	0.055985	2620.6	62.2	2968.8	27.1	3192.7	35.3	1.92	295	155	210	
	6	0.372169	0.018081	12.911824	0.44282	0.248405	0.006349	2.705078	0.130539	2036.6	84.8	2671.4	32.2	3172	40.3	1.92	463	248	381	
	7	0.684633	0.017806	26.481742	0.693295	0.27679	0.005528	1.453754	0.037868	3361.8	68.1	3364.1	25.9	3343.8	31.3	1.1	238	218	361	
	8	0.517955	0.026103	19.650128	0.84025	0.27399	0.005807	1.942975	0.096781	2686.4	110.5	3071	41.3	3327.8	33.6	1.95	315	166	268	
	9	0.682714	0.018122	25.848722	0.679237	0.27197	0.005428	1.46031	0.039001	3354.3	69.5	3340.4	25.8	3316.3	31.4	1.96	210	107	178	
	10	0.246746	0.008238	9.810777	0.3072	0.287727	0.006095	4.051038	0.125641	1421.3	42.3	2416.2	28.2	3404.1	32.7	4.89	776	159	531	
	11	0.224807	0.006385	6.146026	0.163567	0.197737	0.004711	4.447719	0.125244	1307.1	33.6	1996.6	23.2	2805.9	39.4	1.55	1020	665	585	
	12	0.610107	0.016727	21.747389	0.60467	0.259721	0.005584	1.639497	0.044523	3069.8	66.8	3171.7	27	3243.3	33.9	1.72	90	53	92	
	13	0.685168	0.019357	26.302749	0.761449	0.281036	0.006058	1.459705	0.04185	3363.5	74.2	3357.1	28.6	3367.3	33.7	1.85	118	64	111	
	16	0.606398	0.015973	22.081054	0.564692	0.267367	0.005698	1.647354	0.043711	3055.4	64.2	3187.2	24.8	3289.6	33.3	2.18	304	142	263	
	17	0.686053	0.019423	25.921927	0.711597	0.275284	0.006261	1.457651	0.041606	3366.7	74.4	3343	27	3334.5	35.3	0.71	51	73	131	
	18	0.695657	0.018185	26.148396	0.683606	0.2742	0.005371	1.436118	0.037714	3403.8	69.2	3351.7	25.7	3329.2	30.5	2.13	235	111	191	
	19	0.700082	0.018334	26.949451	0.71833	0.280688	0.005627	1.426916	0.037684	3420.6	69.6	3381.1	26.1	3365.6	31.2	1.83	164	90	163	
	20	0.674108	0.018152	25.333116	0.676153	0.273899	0.005533	1.482281	0.039824	3321.2	69.8	3320.6	26	3327.3	31.9	2.81	146	52	91	
	21	0.546675	0.017058	20.182068	0.680584	0.263184	0.00546	1.82915	0.05955	2809.9	71.6	3097.6	33.5	3264.5	32.4	3.09	200	95	158	
	22	0.682531	0.018138	26.771277	0.70594	0.278294	0.005511	1.460587	0.040077	3353.6	69.9	3374.6	26.2	3352.3	31	3.41	163	81	141	
	23	0.683407	0.018671	26.780105	0.726727	0.278536	0.00556	1.458017	0.040314	3357	71.7	3375.1	26.9	3353.8	31.3	2.38	142	122	211	
	24	0.679952	0.019269	26.840367	0.755411	0.27762	0.005481	1.466075	0.042812	3343.4	74.4	3376.9	27.8	3348.6	30.9	4.9	96	54	95	
	25	0.435151	0.02575	15.449951	1.099492	0.25021	0.006481	2.310814	0.139888	2325.9	116	2836.2	68.8	3184.4	40.8	2.11	185	88	77	
	26	0.64271	0.018182	24.940263	0.73785	0.270615	0.006013	1.550954	0.04456	3198.8	71.5	3304.6	28.9	3312.4	30.7	2.39	19	60	104	
	27	0.592397	0.017118	23.60352	0.693218	0.276941	0.005557	1.682977	0.048865	2998.2	69.3	3250.9	29	3347.5	27.6	1.74	98	56	90	
	28	0.639286	0.016783	25.309578	0.66651	0.274509	0.00538	1.557312	0.041512	3185.9	66.2	3319.8	26.1	3331	30.6	2.08	122	58	99	

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Sample	Spot name	Run	206/238	±2SE	207/235	±2SE	207/206	±2SE	238/P206	±2SE	206/238 Age	±2SE	207/235 Age	±2SE	207/206 Age	±2SE	U/Th	U	Th	Pb
	29	4	0.239474	0.010697	11.810756	0.432765	0.339726	0.008194	4.180646	0.180875	1383.2	55.4	2588	34.3	3659.9	36.4	0.99	303	304	329
	30	4	0.544181	0.01805	20.219	0.688015	0.257221	0.006204	1.836803	0.062506	2799.1	75.7	3099.4	32.8	3227.1	37.3	0.54	27	50	84
TR06	1	4	0.292492	0.012779	8.526129	0.529841	0.206306	0.005932	3.461031	0.159443	1651.9	64.1	2275.2	59.3	2871.7	47.6	2.55	606	341	247
	2	4	0.235744	0.006312	5.224355	0.163886	0.157466	0.003682	4.241506	0.11394	1364.4	32.9	1855.5	26.7	2426.4	39.4	2.17	993	454	382
	3	4	0.278024	0.008403	8.704783	0.26846	0.223122	0.00456	3.587575	0.099471	1581	42.6	2306.1	28.5	3002	32.7	1.38	665	484	408
	4	4	0.287794	0.009237	7.775249	0.292187	0.191563	0.004137	3.480754	0.111336	1630	46.2	2203.1	33.8	2754.1	35.5	1.58	627	392	300
	5	4	0.344083	0.010159	10.926615	0.346062	0.226792	0.004919	2.90877	0.084312	1905.8	48.6	2515.5	29.4	3027.9	34.8	1.72	471	274	269
	6	4	0.668341	0.017689	25.602002	0.730317	0.272947	0.005962	1.49543	0.039315	3299.1	68.2	3330.5	28.2	3321.3	34.3	1.19	73	61	108
	7	4	0.650364	0.018126	24.554843	0.681811	0.270797	0.006021	1.537827	0.043571	3228.9	71	3293.5	30.3	3308.7	34.8	1.21	48	40	69
	9	4	0.382386	0.012881	11.858994	0.446311	0.221538	0.0048	2.618363	0.091655	2086.7	60.4	2591.4	35.7	2990.6	34.4	2.27	511	232	302
	10	4	0.251939	0.010566	7.558298	0.45085	0.214531	0.006643	4.005618	0.166217	1447.1	54.3	2168.8	53.2	2934.1	49.6	2.65	676	259	357
	11	4	0.391075	0.011254	10.919138	0.401388	0.202551	0.005879	2.553855	0.072158	2127.5	52	2514.1	34.1	2843	47.2	4.18	1175	282	559
	12	4	0.352917	0.009236	10.329213	0.278982	0.213012	0.004888	2.827002	0.074472	1948.4	44	2464.2	24.9	2926.3	36.8	2.69	1148	429	609
	13	4	0.573893	0.015536	21.574166	0.591411	0.273233	0.006307	1.73885	0.047244	2923.3	63.7	3164	26.6	3322.4	36.3	1.49	59	40	62
	14	4	0.274331	0.009386	8.321618	0.324092	0.222261	0.005228	3.637374	0.123772	1562.6	47.4	2265.8	35.8	2996	38.3	2.26	827	366	344
	15	4	0.306644	0.009496	8.77674	0.328165	0.206925	0.004865	3.259713	0.098094	1723.6	46.7	2312	33.6	2879.1	38.3	2.57	675	263	246
	16	4	0.464631	0.012742	15.636169	0.425003	0.243249	0.004948	2.147159	0.05771	2459.6	55.9	2854.2	25.6	3140.1	32.5	1.61	355	221	154
	17	4	0.23963	0.007088	6.011956	0.224492	0.180747	0.003995	4.168092	0.12094	1384.5	36.8	1974.8	31.9	2657.7	36.5	2.92	924	314	319
	18	4	0.554429	0.019431	32.269721	1.111125	0.424279	0.010286	1.805095	0.06373	2841.8	80.5	3556.6	33.7	3995.4	35.8	6.11	29	5	73
	19	4	0.336936	0.009036	9.907897	0.271275	0.213245	0.004583	2.959901	0.079339	1871.8	43.6	2425.9	25.3	2929.1	35	2.16	561	269	242
	20	4	0.27787	0.007877	8.991872	0.244354	0.23351	0.005117	3.592172	0.100307	1580.5	39.7	2336.9	25.1	3074.9	34.8	5.57	434	169	267
	21	4	0.220141	0.006067	5.120926	0.147894	0.170428	0.003685	4.564372	0.128542	1282.5	32.1	1839.1	24.5	2560.5	36.2	2.33	1532	288	380
	22	4	0.224132	0.00654	4.829717	0.167238	0.160065	0.004374	4.48973	0.129989	1303.5	34.4	1789	28.9	2453.7	45.2	2.95	1057	381	301
	23	4	0.63236	0.020937	22.483328	0.780706	0.263704	0.005803	1.598064	0.052496	3156.7	82.5	3202.1	33.8	3267.1	34.6	1.48	78	51	83
	24	4	0.408071	0.012686	13.067398	0.506521	0.236309	0.005302	2.471375	0.077312	2205.6	58.1	2682.5	36.6	3093.9	35.6	1.78	778	435	525
	25	4	0.209307	0.005585	3.843672	0.10242	0.136602	0.00304	4.815711	0.127604	1225	29.7	1601.6	21.3	2182.5	38.8	3.41	1522	449	322
	26	4	0.382089	0.011811	11.427317	0.363768	0.223888	0.004835	2.641832	0.082278	2085.5	55.1	2557.5	29.6	3007.5	34.5	1.82	351	193	188
	28	4	0.385215	0.011802	12.160664	0.383482	0.233859	0.004906	2.618239	0.08034	2100.1	54.9	2615.7	29.4	3077.4	33.1	1.69	739	434	453
	29	4	0.358899	0.014265	12.713778	0.739191	0.263015	0.007988	2.814362	0.121605	1976	68.3	2652.7	58.2	3261.8	48.8	1.26	1084	857	1133
	30	4	0.338444	0.009746	10.063138	0.316609	0.220993	0.004926	2.953217	0.067295	1879	47.3	2439.7	29.3	2986.5	35.5	1.46	602	411	284
TR07	1	3	0.244367	0.006571	4.368916	0.236119	0.130145	0.002178	4.043451	0.108649	1409.3	34	1706.2	44.7	2098.6	29.3	4.28	1581	370	277
	2	3	0.24623	0.006603	4.891955	0.270384	0.144405	0.002516	4.008859	0.107065	1418.9	34.1	1800.4	46.5	2279.5	29.9	9.78	1754	180	261
	3	3	0.538587	0.018926	19.317992	1.252229	0.261361	0.006254	1.835786	0.061074	2776.5	78.4	3055.1	60.3	3252.9	37.2	5.33	2602	483	2357
	4	3	0.514448	0.023417	17.016417	1.239598	0.240165	0.00659	1.924056	0.085741	2674	99.3	2932.5	67.8	3118.9	44.1	2.17	1287	583	1114
	5	3	0.511803	0.015654	24.210807	1.321821	0.344901	0.006152	1.930522	0.057879	2663.3	66.5	3276.2	53.2	3683.2	27.6	1.81	423	232	575
	7	3	0.412127	0.012621	14.782259	0.841524	0.261893	0.005303	2.391202	0.072933	2224.4	57.6	2800.7	54.9	3256.9	32	7.18	5087	692	5220
	8	3	0.651372	0.018324	22.074877	1.215595	0.246939	0.004469	1.511534	0.041306	3233.3	71.1	3186.8	53	3164.3	28.8	12.9	5482	396	9612
	9	3	0.713168	0.020318	35.462473	1.959108	0.36206	0.006417	1.381341	0.039557	3469.7	76.5	3650.8	54.5	3761.4	22.5	1.95	163	82	364
	10	3	0.322476	0.008355	17.08744	0.916094	0.386194	0.006284	3.051531	0.078859	1801.7	40.7	2939.6	51.3	3855.2	24.6	12.84	3054	252	6574
	11	3	0.548561	0.016174	37.024797	2.059177	0.490921	0.007941	1.794919	0.053136	2818.9	67.4	3693.8	55.3	4213.4	23.8	5.44	1248	223	5863
	13	3	0.344804	0.009064	14.187455	0.768455	0.300161	0.005307	2.856414	0.075715	1909.6	43.5	2761.9	51.1	3469.7	27.2	12.39	1263	100	1491
	14	3	0.683264	0.01869	31.323903	1.696901	0.33431	0.005804	1.442603	0.039724	3356.3	71.6	3528.8	53.7	3635.7	26.6	2.82	130	43	118
	15	3	0.286849	0.008312	8.498111	0.498107	0.215556	0.00441	3.437758	0.099327	1625.6	41.6	2284.2	53.2	2946	32.9	5.88	2145	335	965
	16	3	0.272399	0.011295	7.56672	0.52062	0.202066	0.004149	3.632832	0.147669	1552.3	57.1	2177.7	61.5	2841.4	33.7	4.91	5746	1099	2703
	17	3	0.629705	0.021076	22.839625	1.345501	0.264996	0.005038	1.566668	0.053023	3147.8	83.5	3219.4	57.9	3275.7	29.8	5.02	486	90	203
	18	3	0.472878	0.013462	17.540556	0.997173	0.270467	0.006053	2.086048	0.059595	2495.9	58.9	2964	54.3	3306.9	35	0.98	759	718	1075
	19	3	0.392242	0.011905	15.633507	0.903899	0.290532	0.005582	2.516268	0.074973	2132.9	55	2853.7	55.3	3419.3	29.9	14.42	2019	130	2369
	20	3	0.659908	0.017626	24.743646	1.335452	0.273295	0.00471	1.495528	0.040183	3266.4	68.5	3297.8	52.8	3323.8	27.1	1.26	148	110	189
	21	3	5.535242	1.738246	660.589185	216.406123	0.735363	0.061836	0.871034	0.284312	9385.1	2014.9	5605.3	548.5	4328	154	5.32	2093	452	195547
	22	3	0.328213	0.008926	16.699945	0.918017	0.371126	0.008039	3.009014	0.082474	1829.6	43.4	2917.4	53	3794.5	32.8	11.2	4505	464	9013
	23	3	0.379043	0.010652	17.836512	0.970088	0.342508	0.005599	2.608339	0.073742	2071.4	49.8	2980.4	52.2	3673	24.9	14.55	2453	195	4817
	24	3	0.154249	0.005837	3.407959	0.233607	0.161223	0.00909	6.425888	0.259395	924.5	32.7	1503.5	54.3	2453.3	98.4	2.24	11941	7577	4066
	25	3	0.195578	0.005718	4.890671	0.290124	0.181645	0.003694	5.057988	0.145839	1151.3	30.8	1798.4	49.4	2665.5	33.6	1.45	1402	1102	745
	26	3	0.198538	0.006034	6.535528	0.388714	0.238791	0.005188	4.976768	0.148968	1167.4	32.4	2049.8	51.7	3110.5	34.3	2.59	3239	1391	2732

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Sample	Spot name	Run	206/238	±2SE	207/235	±2SE	207/206	±2SE	238/P206	±2SE	206/238 Age	±2SE	207/235 Age	±2SE	207/206 Age	±2SE	U/Th	U	Th	Pb
27	3	0.39477	0.011046	15.172603	0.834944	0.279277	0.00526	2.503338	0.070322	2144.6	51.1	2825.5	52.8	3357.4	29	2.9	251	94	185	
28	3	0.319757	0.01422	18.98821	1.151538	0.433341	0.013125	3.096884	0.128191	1787.9	68.8	3040.2	58.2	4026.9	46.5	5.73	6859	1289	17914	
29	3	0.179921	0.01177	9.23031	0.822017	0.373011	0.007338	5.545884	0.355058	1065.7	64.2	2355	79.8	3802.6	29.4	6.34	8284	1426	10094	
30	3	1.009751	0.081215	52.212145	6.673454	0.365699	0.017286	1.028723	0.075365	4456.3	253	3974.2	125.3	3757.1	72	2.67	2243	880	11400	
31	5	0.458375	0.029347	14.720028	1.31631	0.21869	0.007518	2.220821	0.140411	2427.3	129.2	2787.7	84	2969.4	55.2	2.57	1476	639	893	
32	5	0.632741	0.027786	45.427771	3.273909	0.487396	0.016321	1.592359	0.070106	3157.9	109.7	3892.8	71.6	4202.2	49.2	17.59	2849	181	17289	
33	5	0.489438	0.027205	17.688546	1.521078	0.245946	0.008986	2.083425	0.111403	2562.1	116.9	2958.8	80.4	3155.3	57.5	0.61	455	2638	990	
34	5	0.389638	0.039756	15.640771	1.711464	0.276928	0.009376	2.734542	0.28212	2106	184	2831.1	107	3344.2	52.9	3.13	224	81	150	
35	5	0.615112	0.024411	25.889506	1.711939	0.288117	0.009317	1.634236	0.06573	3088.7	97.7	3340.7	65	3406.4	50.3	0.73	278	428	445	
36	5	0.301926	0.011035	8.411282	0.540166	0.191686	0.006119	3.319695	0.120619	1700.6	54.6	2276	58.1	2755.7	52.2	6.15	3635	656	711	
37	5	0.757455	0.028968	27.441844	1.776893	0.248816	0.008152	1.322232	0.050478	3634.2	106.2	3399	63.2	3176.3	51.7	1.31	613	510	1719	
38	5	0.598441	0.024336	22.951268	1.534507	0.266111	0.008628	1.67864	0.073397	3021.5	99.4	3222.7	67.4	3282.1	50.9	2.07	169	90	169	
39	5	0.320142	0.022433	9.969942	1.050267	0.218221	0.011346	3.153146	0.219046	1788.5	109.2	2423.5	100.1	2962.6	83.6	9.36	2913	424	1717	
40	5	0.28259	0.016622	8.277	0.822002	0.204035	0.009508	3.581307	0.214979	1602.4	83.6	2247.9	91.4	2851.6	77.4	13.44	3255	261	1509	
TR08	1	3	0.660966	0.017679	25.003773	1.348583	0.273198	0.004766	1.510369	0.040456	3270.5	68.6	3308	53.1	3323.1	27.2	1	91	88	125
	2	3	0.678926	0.017992	25.611704	1.382322	0.272405	0.004648	1.468479	0.038877	3339.9	69.1	3331.4	52.4	3318.7	26.6	0.69	163	232	346
	3	3	0.684598	0.019222	25.818447	1.41524	0.272549	0.005867	1.45551	0.040838	3361.4	73.5	3339.1	53.1	3318.7	33.2	0.98	106	106	164
	4	3	0.689765	0.018652	26.317826	1.423518	0.274955	0.004619	1.443158	0.03818	3381.2	70.8	3358	52.6	3333.4	26.2	1.74	120	68	108
	5	3	0.509768	0.014283	17.636833	0.986077	0.250348	0.0045	1.950354	0.055188	2655.5	61.1	2969.7	54	3186	28.6	5.97	554	93	118
	6	3	0.313026	0.01071	8.14676	0.481908	0.187854	0.003197	3.188972	0.113662	1754.6	52.8	2244.8	54.1	2722.2	28	8.63	1654	193	226
	7	3	0.685241	0.018079	25.834279	1.39326	0.272312	0.004804	1.448943	0.037965	3364.1	69	3339.9	52.6	3318	27.6	0.75	153	205	336
	8	3	0.561698	0.016455	20.267309	1.127155	0.260124	0.004203	1.769251	0.051116	2872.7	67.7	3103	53.3	3246.5	25.4	6.94	568	83	136
	9	3	0.192421	0.005472	4.310976	0.238274	0.161809	0.00282	5.159624	0.146245	1134.3	29.6	1694.7	45.6	2473.2	29.5	2.1	951	456	244
	10	3	0.234415	0.006421	5.174237	0.286938	0.159053	0.0032	4.227451	0.117109	1357.6	33.6	1848.1	47.7	2444.4	34	5.62	882	158	180
	11	3	0.692449	0.018693	26.607742	1.453376	0.278622	0.004979	1.431196	0.038243	3391.4	71	3368.4	53.6	3353.8	27.7	0.79	66	81	132
	12	3	0.21321	0.006147	3.131171	0.177227	0.10631	0.002171	4.64727	0.13194	1245.8	32.6	1440.1	43.5	1735.7	37.5	8.33	2641	309	181
	13	3	0.67687	0.01924	25.240136	1.369498	0.269254	0.005117	1.464726	0.042032	3331.9	74.1	3317.3	53	3300.5	29.8	0.71	259	358	556
	14	3	0.675803	0.017783	25.179085	1.356035	0.269455	0.004424	1.467134	0.039446	3327.9	68.7	3314.9	52.5	3301.8	25.7	0.74	275	362	542
	15	3	0.690601	0.018945	25.989435	1.407981	0.272249	0.004787	1.436947	0.039517	3384.3	72.2	3345.7	52.7	3317.6	27.7	0.98	96	145	145
	16	3	0.678677	0.017983	25.509576	1.390158	0.271194	0.004859	1.461959	0.038339	3338.9	68.9	3327.3	53.4	3311.5	28.1	0.94	93	98	146
	17	3	0.686289	0.019168	25.836759	1.409403	0.272496	0.004917	1.447293	0.04111	3367.7	73.5	3339.7	53.2	3318.9	28.4	0.84	202	237	354
	18	3	0.679668	0.01892	25.618026	1.392914	0.272952	0.005312	1.461173	0.04052	3342.4	72.5	3331.5	53.2	3321.1	30.3	0.95	98	103	154
	19	3	0.682563	0.017873	25.819807	1.392281	0.273791	0.004472	1.454736	0.038079	3353.9	68.5	3339.4	52.7	3326.8	25.7	5.75	444	75	118
	20	3	0.683303	0.018585	25.703849	1.396586	0.272348	0.00453	1.454319	0.039415	3356.5	71.1	3334.8	52.9	3318.5	25.9	1.27	166	129	200
	21	3	0.67568	0.017926	26.465041	1.460386	0.283338	0.005722	1.467646	0.044311	3327.4	69	3362.9	54	3379.2	31.9	1	101	101	195
	22	3	0.69619	0.01924	27.309404	1.490717	0.284615	0.005267	1.428373	0.038813	3405.5	72.8	3393.9	53.9	3386.8	28.6	0.76	92	120	236
	23	3	0.681966	0.018133	25.750583	1.396735	0.273937	0.004639	1.45721	0.039157	3351.5	69.6	3336.6	52.8	3327.5	26.6	4.33	417	96	175
	24	3	0.684151	0.018555	25.914567	1.40266	0.274213	0.004829	1.452646	0.039062	3359.7	70.9	3342.9	52.4	3328.9	27.7	0.93	138	147	255
	25	3	0.685718	0.018503	25.752735	1.398259	0.272624	0.00471	1.448889	0.038745	3365.8	70.6	3336.7	53	3319.9	27.1	0.86	165	192	331
	26	3	0.461789	0.012975	16.052147	0.877273	0.253379	0.004183	2.152032	0.058566	2447	56.9	2879.2	51.7	3204.9	26	1.81	651	810	937
	27	3	0.569039	0.015418	20.678999	1.124674	0.263902	0.004455	1.745098	0.047528	2903.5	63.4	3123.1	53.2	3269	26.5	1.01	540	536	747
	28	3	0.602015	0.015913	21.998806	1.188577	0.265414	0.00434	1.648391	0.043367	3037.7	64	3183.3	52.4	3278.1	25.6	0.74	361	487	750
	29	3	0.695761	0.019537	26.775809	1.460564	0.279725	0.005166	1.426967	0.039774	3403.7	74.1	3377.6	47.5	3359.8	28.9	0.99	178	181	332
	30	3	0.677579	0.018726	26.836254	1.476726	0.287541	0.005515	1.464277	0.040352	3334.4	71.9	3376.6	54.4	3402.5	30.4	1.12	89	81	169
WG02	1	3	0.289276	0.009928	7.248308	0.437885	0.182855	0.005303	3.404056	0.118232	1637.5	49.7	2141.1	54.9	2675.7	48.6	1.27	106	82	104
	2	3	0.410401	0.011539	10.105105	0.562318	0.179725	0.003414	2.394561	0.067685	2216.4	52.8	2443.6	51.1	2648.7	31.5	0.92	74	78	99
	3	3	0.436264	0.011868	10.485639	0.570151	0.175578	0.003077	2.250879	0.06054	2333.6	53.2	2478.1	50.3	2610.1	29.5	1.6	329	198	166
	4	3	0.513643	0.01493	12.748886	0.73359	0.18146	0.004356	1.911326	0.055342	2671.6	63.5	2659.9	54.9	2663.3	40	1.07	42	37	47
	5	3	0.518711	0.01425	12.891151	0.701625	0.181486	0.003207	1.885683	0.045208	2693.3	60.4	2671.2	51.3	2665	29.3	0.98	321	313	428
	6	3	0.330075	0.010992	7.726725	0.446674	0.170868	0.002911	2.980484	0.101137	1837.8	53.4	2197.7	52.5	2564.8	28.5	1.14	485	403	385
	7	3	0.412957	0.012327	9.706259	0.537387	0.171406	0.003184	2.374284	0.072436	2228	56.4	2406.7	50.9	2569.9	31.1	1.89	334	167	187
	8	3	0.365239	0.011022	8.857325	0.495619	0.176934	0.002916	2.685971	0.081486	2006.3	52.1	2322.2	51	2623.2	27.5	2.53	378	140	177
	9	3	0.510546	0.013984	12.581329	0.688539	0.179765	0.003169	1.917889	0.052552	2658.5	59.7	2648.1	51.4	2649.2	29.3	2.32	257	105	146
	10	3	0.492023	0.014541	12.906101	0.716863	0.191621	0.003979	1.990024	0.058525	2578.9	62.8	2672.2	52.7	2754.3	34.6	1.19	87	70	101

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Sample	Spot name	Run	206/238	±2SE	207/235	±2SE	207/206	±2SE	238/P206	±2SE	206/238 Age	±2SE	207/235 Age	±2SE	207/206 Age	±2SE	U/Th	U	Th	Pb
		11	3	0.33478	0.010657	7.922051	0.445654	0.172294	0.002911	2.932422	0.094206	1860.8	51.5	2220.9	51.1	2578.8	28.4	1.82	456	263
		12	3	0.361613	0.010075	8.879038	0.483099	0.177651	0.003142	2.708514	0.075787	1989.6	47.7	2325.3	49.5	2629.8	29.3	1.06	269	300
		13	3	0.408786	0.012174	10.054247	0.547752	0.179077	0.003274	2.400539	0.074522	2208.7	56	2439.2	50.6	2642.6	30.4	1.89	243	175
		14	3	0.223575	0.007835	4.715162	0.266435	0.1538	0.003626	4.391714	0.149604	1300.4	41.2	1769.3	47.1	2386.3	40.1	1.23	809	691
		15	3	0.207663	0.007657	4.738711	0.279797	0.166271	0.003308	4.73136	0.181331	1216.1	41	1773.2	50	2519.1	33	2.23	561	261
		16	3	0.461216	0.0124	11.063636	0.603281	0.173819	0.003053	2.126654	0.05641	2444.6	54.6	2527.9	50.9	2593.3	29.2	2	249	128
		17	3	0.482156	0.013262	11.797167	0.6404	0.178428	0.003263	2.035788	0.05732	2536.2	57.9	2588	50.3	2636.6	30.1	1.86	143	83
		18	3	0.470405	0.013445	11.29326	0.620495	0.174392	0.002962	2.081254	0.053707	2484.7	59.2	2546.9	51.5	2598.9	28.5	1.67	247	150
		19	3	0.29168	0.007812	7.062097	0.383465	0.17546	0.003213	3.364819	0.090778	1649.9	39	2119.1	48.2	2609.3	30.6	2.12	369	175
		20	3	0.324939	0.010553	7.754544	0.478337	0.173585	0.004622	3.024189	0.101075	1813.6	51.5	2201.8	54.3	2590.4	43.3	1.47	511	349
		21	3	0.34966	0.010561	8.397225	0.465616	0.173958	0.002931	2.816216	0.086115	1932.4	50.5	2273.9	50.4	2594.8	28.2	1.62	671	396
		22	3	0.290486	0.009341	6.571107	0.361613	0.163966	0.002982	3.393577	0.111766	1643.3	46.8	2054.7	49	2495.3	30.6	4.11	871	203
		23	3	0.507277	0.013595	12.553222	0.686279	0.179074	0.003222	1.93702	0.051772	2644.7	58.1	2646	51.8	2642.7	29.8	0.69	290	401
		25	3	0.167062	0.004809	4.010668	0.222183	0.173345	0.003033	5.88064	0.173373	995.8	26.6	1636	45.2	2589.1	29.2	2.65	1078	381
		26	3	0.504988	0.014905	12.455521	0.710771	0.178406	0.004083	1.947623	0.058562	2634.4	64	2637.6	52.9	2634.7	37.8	1.01	58	54
		27	3	0.447203	0.012836	11.034743	0.603979	0.178394	0.002977	2.197781	0.063255	2382.2	57.2	2525.4	51.1	2636.8	27.7	1.78	466	249
		28	3	0.427334	0.01143	10.347165	0.558328	0.174981	0.002887	2.297107	0.061628	2293.5	51.6	2466	50	2604.7	27.5	1.9	480	239
		29	3	0.427117	0.014078	10.414045	0.602989	0.176672	0.002928	2.301352	0.079139	2292	64	2471.2	54.2	2620.9	27.6	1.17	470	381
		30	3	0.511708	0.013925	12.726086	0.701758	0.179607	0.003404	1.91792	0.050917	2663.5	59.1	2658.7	51.8	2647.3	31.5	1.21	131	102
WG03		1	3	0.200015	0.008452	4.609222	0.327719	0.167647	0.004021	5.058552	0.214045	1175.1	45.4	1748.4	59.4	2532.4	40.7	1.36	822	612
		2	3	0.193196	0.006722	4.393427	0.272359	0.165534	0.003604	5.232582	0.181133	1138.5	36.3	1709.7	51	2511.3	36.6	1.44	806	564
		3	3	0.242504	0.010424	5.504225	0.356884	0.164829	0.003014	4.197653	0.177238	1398.6	54	1897.6	55.5	2504.3	30.9	1.35	672	504
		4	3	0.50641	0.014055	12.604027	0.699408	0.181006	0.003897	1.996431	0.055936	2640.8	60.2	2649.5	52.1	2659.2	35.4	1.1	58	53
		5	3	0.508047	0.015077	12.556759	0.695266	0.180108	0.004171	1.990237	0.059161	2647.9	64.4	2646.6	51.6	2651.9	38.4	1.26	66	53
		6	3	0.148596	0.007208	3.788978	0.279185	0.184685	0.00345	6.912441	0.331664	892.2	40.4	1582.6	58.8	2693.6	31	5.85	1418	240
		7	3	0.468438	0.018769	11.515104	0.675805	0.178594	0.00396	2.167402	0.088906	2475.3	82.6	2564.7	55.2	2638	37.2	1.97	195	100
		8	3	0.516629	0.01431	12.704099	0.688669	0.179355	0.003173	1.959912	0.054094	2684.4	60.8	2660.2	57.8	2645.4	29.4	1.75	215	125
		9	3	0.186083	0.005513	4.181382	0.239314	0.163591	0.002823	5.440683	0.159138	1100	29.9	1669.7	46.8	2492.1	29.1	2.24	1305	584
		10	3	0.186671	0.006759	4.376886	0.27984	0.17034	0.00389	5.432247	0.192404	1103.1	36.6	1706.3	52	2559.1	38.4	2.39	1082	449
		12	3	0.184879	0.007489	4.271739	0.27994	0.167537	0.003784	5.456716	0.231813	1093.3	40.9	1686.1	54.4	2531.4	38.3	2.68	827	299
		13	3	0.205384	0.006492	4.817341	0.273271	0.17056	0.003175	4.895498	0.15806	1204	34.8	1787.3	48.3	2561.9	31.3	1.64	766	456
		14	3	0.234715	0.006605	5.592721	0.307098	0.173062	0.003075	4.275795	0.121699	1359.1	34.5	1914.6	47.5	2586.2	29.6	2.49	821	322
		15	3	0.383184	0.011492	9.230329	0.533208	0.175448	0.003569	2.615729	0.077579	2091	53.5	2360.6	52.4	2609.1	33.7	2.61	481	181
		17	3	0.35028	0.010188	8.346463	0.463304	0.173157	0.003012	2.858524	0.084176	1935.5	48.7	2268.4	50.8	2587	29.2	1.31	457	347
		18	3	0.277736	0.007342	6.565542	0.352651	0.171801	0.002891	3.59629	0.094509	1579.9	37	2054.5	47.4	2574	28	1.93	586	303
		19	3	0.078077	0.002149	1.567234	0.084634	0.145904	0.002532	12.787809	0.358201	484.6	12.9	957.2	33.5	2297	29.7	3.34	1965	589
		20	3	0.169211	0.00583	3.699507	0.218274	0.159099	0.002681	5.90171	0.196418	1007.6	32.1	1570.4	46.8	2445.1	28.5	2.35	1062	449
		21	3	0.278464	0.008688	6.643476	0.375086	0.173242	0.002985	3.579716	0.11441	1583.1	43.9	2063.8	50.6	2587.8	28.9	1.56	704	458
		22	3	0.311521	0.011851	7.298676	0.436071	0.170506	0.003192	3.198029	0.124545	1747.7	58.4	2147.7	53.8	2561.5	31.3	1.7	642	378
		23	3	0.349332	0.013557	8.303263	0.502782	0.172946	0.003449	2.853859	0.111044	1930.7	64.8	2263.4	54.6	2585	33.7	1.08	375	349
		24	3	0.249639	0.007461	5.605246	0.329844	0.162599	0.003738	3.985142	0.123475	1436.5	38.6	1916.1	51.5	2481.1	38.7	1.71	623	368
		25	3	0.322524	0.009419	7.655648	0.425831	0.172164	0.003064	3.08765	0.091111	1801.7	46	2190.6	50.2	2577.4	29.9	1.59	543	348
		26	3	0.372301	0.010143	8.940743	0.486235	0.174299	0.003004	2.674104	0.072713	2040	47.6	2331.5	49.7	2598	28.7	1.55	596	395
		27	3	0.370398	0.011203	9.007895	0.500559	0.176345	0.0034	2.690812	0.08098	2030.9	52.7	2338.1	50.2	2617.2	32.5	1.6	597	386
		28	3	0.424772	0.011844	10.349678	0.569198	0.176773	0.002999	2.34663	0.067442	2281.7	53.8	2465.7	51.7	2621.5	28.3	1.39	611	457
		29	3	0.272424	0.007938	6.322923	0.34451	0.168551	0.003304	3.663125	0.106812	1552.8	40.2	2021.1	47.9	2541.1	32.6	1.5	969	680
		30	3	0.187301	0.005158	4.254545	0.231193	0.165133	0.003123	5.322384	0.148822	1106.7	28	1684.5	44.8	2507.7	31.7	1.78	1360	805
ZB01		1	2	0.34677	0.004407	7.794743	0.571995	0.163061	0.003648	2.887601	0.036591	1918.9	21.1	2207.1	65.7	2486.3	37.7	2.15	1870	8485
		2	2	0.503446	0.006133	12.623988	0.927169	0.181891	0.004108	1.988881	0.024438	2628.3	26.3	2651.5	69.2	2668.7	37.6	1.48	1103	745
		3	2	0.249902	0.00681	4.995834	0.404418	0.144695	0.003683	4.032196	0.12071	1437.1	35.4	1814.1	70.9	2281.3	43.2	1.41	4478	3178
		4	2	0.286357	0.005123	6.65676	0.490592	0.168743	0.003826	3.502936	0.063967	1622.9	25.7	2066.3	64.8	2543.7	38	3.24	4289	1324
		5	2	0.226103	0.005961	4.469432	0.371825	0.142848	0.004001	4.45445	0.12705	1313.4	31.5	1719.6	72.4	2257.7	49.7	1.33	8132	6133
		6	2	0.531303	0.006411	13.386105	0.98212	0.182756	0.004045	1.884994	0.023096	2746.6	27.1	2706.8	70.1	2676.8	36.8	2.81	1879	669
		7	2	0.438857	0.005745	10.719085	0.783042	0.177217	0.003835	2.282744	0.030414	2345.2	25.8	2498.8	68.4	2625.9	36.1	2.48	3548	1432

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Sample	Spot name	Run	206/238	±2SE	207/235	±2SE	207/206	±2SE	238/P206	±2SE	206/238 Age	±2SE	207/235 Age	±2SE	207/206 Age	±2SE	U/Th	U	Th	Pb
	8	2	0.542928	0.007495	13.698572	1.004594	0.183064	0.004042	1.845588	0.025783	2795.2	31.4	2728.6	69.6	2679.6	36.4	1.82	1283	706	695
	9	2	0.526792	0.007869	13.349557	0.989747	0.183867	0.004431	1.902752	0.028393	2727.3	33.2	2703.7	71.2	2686	39.7	2.64	476	701	672
	10	2	0.522506	0.007482	13.266064	0.976323	0.184224	0.004188	1.918163	0.02806	2709.3	31.8	2698.1	69.6	2689.8	37.5	1.88	540	459	560
	11	2	0.329909	0.005842	7.206819	0.549453	0.158354	0.003769	3.041542	0.054841	1837.5	28.4	2135.4	66.4	2436.2	39.8	1.35	682	383	299
	12	2	0.248494	0.008767	4.736877	0.403087	0.138937	0.003561	4.068925	0.128482	1429.3	44.7	1767.9	66.1	2211.1	43.4	1.69	916	412	246
	13	2	0.26176	0.007506	5.768163	0.525025	0.158884	0.005538	3.850091	0.103439	1497.8	38.1	1931.9	74	2436.1	56.5	2.43	938	293	260
	14	2	0.157164	0.002776	1.984986	0.151641	0.091697	0.002237	6.384377	0.113617	940.9	15.5	1113.6	57.6	1458.4	46.1	1.16	2046	1353	577
	15	2	0.314983	0.007085	6.948495	0.528099	0.160049	0.00367	3.191161	0.07077	1764.4	34.6	2103	67.1	2454.5	38.8	2.02	409	158	128
	16	2	0.540256	0.007446	13.635199	0.998726	0.183173	0.00409	1.849561	0.023322	2784	31.4	2724.3	69.9	2680.5	37.1	3.91	299	61	87
	17	2	0.38439	0.006444	9.40561	0.705339	0.177448	0.003972	2.609659	0.045193	2096.2	30.1	2376.9	69.6	2627.8	37.1	2.66	631	194	213
	18	2	0.418332	0.006626	10.13941	0.748118	0.175897	0.003926	2.396992	0.037587	2252.4	30.1	2446.8	67.8	2613.2	37.1	2.18	549	211	231
	19	2	0.342819	0.00681	8.434261	0.630558	0.178528	0.003904	2.92883	0.057199	1899.6	32.6	2277.6	67.6	2638.1	36	1.73	1153	572	560
	20	2	0.524367	0.006263	13.387453	0.980103	0.18529	0.004261	1.910285	0.022689	2717.4	26.4	2707	69.2	2699.2	37.7	1.43	179	110	143
	21	2	0.27671	0.016323	7.158302	0.687537	0.188366	0.00507	3.691552	0.22355	1592	92.1	2117.1	83.3	2724.6	44	4.09	2080	466	2093
	22	2	0.327552	0.013123	7.677353	0.709384	0.168938	0.004833	3.10113	0.126128	1823.6	63.8	2181.4	84	2542.8	48.1	2	832	368	383
	23	2	0.307335	0.006632	7.261153	0.548591	0.171448	0.00385	3.269215	0.073961	1726.9	32.8	2142.3	68.8	2570.4	37.6	2.57	857	292	374
	24	2	0.542981	0.007549	13.889635	1.023511	0.185625	0.004121	1.845244	0.025457	2795.4	31.5	2741.5	70.1	2702.5	36.4	1.09	510	409	503
	25	2	0.481643	0.010336	11.980723	0.910614	0.180011	0.00404	2.085745	0.04617	2533.1	45.2	2600.7	72.5	2651.6	37.5	2.53	319	112	139
	26	2	0.337365	0.014675	7.721974	0.696136	0.165479	0.003973	3.013104	0.118406	1870.5	69.7	2189.2	76.3	2510.3	40.3	1.92	1067	514	353
	27	2	0.523633	0.007028	13.751676	1.043426	0.190465	0.00501	1.913045	0.025429	2714.1	29.7	2730.7	70.2	2743.1	42.9	2.98	280	90	163
	28	2	0.374288	0.019412	8.653054	0.852065	0.166577	0.004439	2.738543	0.136279	2043.6	90.4	2285.4	88	2520.2	44.8	2.31	932	354	324
	29	2	0.48711	0.014294	12.227712	0.97399	0.181991	0.00437	2.070857	0.065021	2555.5	62.5	2617.3	78.5	2669.1	39.3	2.01	488	211	278
	30	2	0.531052	0.008228	13.364829	0.990033	0.182622	0.004002	1.88739	0.030139	2745.2	34.8	2704.8	70.2	2675.7	36.3	1.2	363	267	346
ZB02	1	2	0.992977	0.037391	27.230564	2.279314	0.198742	0.004449	1.020008	0.036841	4435.5	119.2	3384.3	79.5	2814.6	36.6	0.76	6396	7997	94128
	2	2	0.175466	0.003653	4.067725	0.309405	0.168242	0.004083	5.722261	0.122273	1041.9	20.1	1646.2	62.3	2538	40.3	3.3	1161	313	356
	3	2	0.533549	0.006766	13.459698	0.985815	0.183141	0.004134	1.876795	0.023682	2756	28.4	2712	69	2680.1	37	1.89	336	160	229
	4	2	0.356527	0.005518	9.092966	0.672897	0.185131	0.004171	2.810743	0.043489	1965.3	26.2	2346.6	67.1	2698	37.5	2.84	499	161	222
	5	2	0.1594	0.004666	4.761418	0.350741	0.217894	0.006409	6.324317	0.187272	953	26	2177.6	62.6	2960.7	47.8	0.95	5187	5240	4057
	6	2	0.453825	0.026464	11.337013	1.072984	0.181075	0.004114	2.290167	0.15986	2402	120	2534.6	95.1	2661.2	37.5	1.44	777	465	623
	7	2	0.15754	0.002172	2.750059	0.208279	0.12665	0.003355	6.357801	0.087848	943	12.1	1340.9	55.9	2048.5	46.7	2.68	6492	2225	1558
	8	2	0.418934	0.011202	10.840467	0.842779	0.187159	0.004739	2.387475	0.06377	2253.9	51.7	2506.3	74.3	2721.1	46.1	1.84	530	273	371
	9	2	0.519394	0.007894	13.174939	0.967353	0.184209	0.004173	1.929114	0.029322	2696	33.5	2691.7	69.1	2689.7	37.3	2.51	302	109	149
	10	2	0.437141	0.006966	10.539022	0.782253	0.174987	0.003881	2.292653	0.037317	2337.3	31.3	2482.4	68.9	2604.6	36.6	1.99	701	319	353
	11	2	0.316779	0.007175	8.003661	0.608978	0.183471	0.004516	3.170851	0.071282	1773.2	35.1	2229.4	69.1	2682.2	40	0.84	420	454	250
	12	2	0.511713	0.006755	12.863041	0.942733	0.182529	0.00409	1.956631	0.025887	2663.5	28.8	2669.2	69.1	2674.6	37.3	1.26	425	315	394
	13	2	0.373688	0.005541	9.344563	0.690571	0.181562	0.004159	2.680424	0.039918	2046.4	26	2371.6	67.9	2665.6	37.8	2.08	459	207	184
	14	2	0.280439	0.006585	6.396087	0.514275	0.165187	0.004343	3.583376	0.08707	1592.8	33.2	2027.3	71.7	2506.3	44.3	1.95	671	326	200
	15	2	0.329182	0.016245	8.038668	0.786389	0.175907	0.005073	3.110489	0.159403	1829.9	79.1	2218.4	90.3	2610.3	47.9	2.03	637	335	341
	16	2	0.490272	0.009726	12.567416	0.934886	0.186197	0.004169	2.046462	0.042449	2570.7	42.3	2646.6	71	2707.5	37.1	1.82	430	228	295
	17	2	0.361273	0.00612	8.398775	0.620835	0.168852	0.003856	2.774062	0.048516	1987.7	29.1	2274.2	68	2544.7	38.1	3.15	633	195	168
	18	2	0.371274	0.011493	9.105373	0.758133	0.177527	0.004619	2.718681	0.092888	2033.3	54.6	2341.7	80.1	2626.9	43.3	3.23	620	186	218
	19	2	0.133379	0.004132	2.383961	0.193266	0.129545	0.003099	7.564087	0.246027	806.7	23.6	1234.5	60.1	2089.6	43.1	5.04	3001	564	565
	20	2	0.315356	0.004444	7.499649	0.557347	0.172018	0.003984	3.174562	0.04395	1766.8	21.7	2172	65.3	2575.7	38.7	1.99	679	320	318
	21	2	0.4551	0.014178	11.173613	0.88769	0.178164	0.003888	2.216919	0.075972	2415.1	63.6	2533.2	77.4	2634.7	36.1	1.7	992	447	794
	22	2	0.171335	0.009381	4.297204	0.416087	0.181209	0.004499	5.995573	0.322707	1017.5	51.5	1679.8	79.7	2661.5	41.4	3.91	2296	441	1263
	23	2	0.302461	0.010138	7.154848	0.591275	0.171479	0.00413	3.335581	0.10682	1701.8	49.8	2125.3	71.8	2570	40.5	0.35	1197	2537	528
	24	2	0.301081	0.010744	7.200278	0.617969	0.173021	0.004394	3.359551	0.127926	1694.7	53.5	2128.3	79.3	2584.4	42.5	2.45	906	272	352
	25	2	0.533415	0.007429	13.466377	0.991006	0.183359	0.004182	1.875352	0.026556	2755.3	31.3	2712.3	69.7	2682	37.9	1.54	517	245	417
	26	2	0.262085	0.008731	7.927223	0.675044	0.218829	0.005577	3.851309	0.133143	1499.1	44.7	2215	78.2	2969.5	41	3.19	2273	512	1566
	27	2	0.461871	0.005298	11.754383	0.868309	0.185244	0.004234	2.164205	0.025573	2447.6	23.5	2584.2	70.7	2698.8	37.9	2.88	402	102	158
	28	2	0.454464	0.014087	11.073715	0.894563	0.17623	0.003985	2.21762	0.072221	2412.4	62.8	2524	77.3	2620.3	42.8	1.15	1314	862	1135
	29	2	0.34334	0.009721	8.917689	0.688499	0.18868	0.004192	2.929239	0.082281	1901.2	46.6	2326.9	70.8	2729.4	36.8	2.35	814	260	458
	30	2	0.472447	0.008468	11.990681	0.887006	0.184388	0.004109	2.119089	0.038712	2493.5	37.2	2602.8	69.9	2691.4	36.5	1.32	1209	706	1036
ZB03	1	3	0.27643	0.007541	11.607059	0.646441	0.304507	0.005955	3.634344	0.100045	1573.2	38.1	2572.5	52.1	3491.7	30.4	3.01	1040	343	1615

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Sample	Spot name	Run	206/238	±2SE	207/235	±2SE	207/206	±2SE	238/P206	±2SE	206/238 Age	±2SE	207/235 Age	±2SE	207/206 Age	±2SE	U/Th	U	Th	Pb
2	3		0.459329	0.012874	11.413782	0.625116	0.180564	0.003345	2.18761	0.061403	2436.4	56.9	2557.3	51	2656.9	30.7	2.79	350	125	181
3	3		0.375971	0.011902	8.890218	0.507825	0.171955	0.003251	2.676776	0.086212	2057	55.9	2325.9	52.8	2575.4	31.7	3.15	356	113	144
4	3		0.347516	0.010437	8.207497	0.477931	0.171163	0.003801	2.894031	0.087017	1922.5	49.9	2253.3	52.8	2571.9	37.2	2.26	438	194	210
5	3		0.264616	0.006941	7.496507	0.424653	0.205623	0.004721	3.799431	0.099205	1513.3	35.4	2171.1	51.3	2868	37.7	6.26	1108	184	648
6	3		0.192511	0.005775	3.451892	0.208083	0.129957	0.002761	5.233497	0.157802	1134.8	31.2	1513.9	47.9	2094.4	37.3	1.7	879	521	295
7	3		0.483347	0.014075	12.125895	0.669098	0.182042	0.003165	2.082725	0.062593	2541.2	61.4	2613.5	51.9	2670.3	28.7	2.02	392	196	269
8	3		0.521557	0.014641	13.2809	0.72156	0.184823	0.003194	1.929247	0.055141	2705.3	62.2	2702.3	44.7	2699.1	24.4	2.15	325	153	219
9	3		0.146515	0.004079	1.872408	0.102184	0.092728	0.001525	6.866146	0.193804	881.3	22.9	1071.1	36.2	1481.1	31.1	14.38	3592	253	272
10	3		0.273385	0.007928	5.621986	0.324281	0.149064	0.002718	3.680819	0.106202	1557.7	40.1	1918	49.1	2333.6	30.9	15.25	914	61	64
11	3		0.339876	0.008943	7.956085	0.427472	0.169829	0.002779	2.949559	0.078168	1885.9	43.1	2225.8	48.4	2554.8	27.3	18.63	670	37	59
12	3		0.317372	0.009816	7.129427	0.409578	0.162836	0.002679	3.165115	0.101326	1776.3	48.2	2125.9	52	2484.2	27.9	1.67	671	409	341
13	3		0.315163	0.010167	7.153519	0.411875	0.164498	0.002903	3.186029	0.107001	1765.5	50.1	2129.2	52.4	2501	29.8	1.65	760	470	406
14	3		0.377005	0.011044	9.28923	0.538465	0.178282	0.003605	2.655662	0.078327	2062	51.7	2366	53.7	2635.5	33.5	2.55	593	238	289
15	3		0.342545	0.00939	7.759425	0.433502	0.164104	0.002894	2.921491	0.082134	1898.7	45.2	2202.5	51.3	2496.9	30.1	2.09	739	361	341
17	3		0.244343	0.006619	5.251349	0.29291	0.15552	0.003837	4.08788	0.111575	1409.2	34.3	1860.6	47.7	2405.3	41.8	6.62	991	153	232
18	3		0.352948	0.010384	8.076443	0.449326	0.165753	0.002863	2.832585	0.085901	1948.3	49.6	2238.8	51	2513.9	28.8	1.75	735	429	400
19	3		0.414324	0.012244	9.947723	0.557126	0.173831	0.002848	2.411472	0.074038	2234.2	56.1	2428.9	52.6	2593.8	27.3	1.82	670	373	413
21	3		0.504281	0.017204	12.869569	0.732244	0.185121	0.004207	1.970288	0.069175	2631.5	74	2669.6	53.7	2697.7	37.3	2.35	297	127	168
22	3		0.497354	0.013642	12.592491	0.692201	0.182885	0.003056	1.994016	0.054987	2602.2	58.8	2649.2	52	2678.2	27.7	2.05	503	248	345
23	3		0.184415	0.005126	3.19523	0.175306	0.125336	0.002199	5.378547	0.145766	1090.9	27.8	1455.5	42	2032	30.9	10.29	849	90	68
24	3		0.236246	0.007882	5.999	0.38151	0.183575	0.004257	4.203662	0.140272	1366.8	41.1	1972.8	55.9	2682.8	39	2.16	871	416	530
25	3		0.211981	0.005842	4.899146	0.264918	0.167242	0.003319	4.674298	0.128508	1239.2	31.1	1801.8	45.8	2527.9	33.3	4.12	1096	337	452
26	3		0.489678	0.013505	17.088807	0.978834	0.251994	0.005456	2.021654	0.057232	2569	58.7	2938.7	56.4	3195.4	34.4	2.78	609	227	867
27	3		0.323046	0.010046	9.892276	0.622273	0.221071	0.00542	3.0693	0.098099	1804.1	49.1	2421.1	59.5	2985.2	40	3.3	1775	573	1683
28	3		0.331179	0.009493	7.98339	0.448445	0.174237	0.003668	2.980484	0.077259	1843.7	46.1	2227.9	50.1	2596.1	34.8	1.86	776	430	560
29	3		0.471024	0.012537	11.725184	0.635901	0.179542	0.003216	2.099485	0.056063	2487.9	55	2582.5	50.9	2647.5	29.9	1.75	473	278	376
30	3		0.346446	0.009405	7.670867	0.417449	0.159847	0.002752	2.855639	0.077845	1917.4	45	2192.8	49	2452.7	28.9	1.84	688	379	387
31	7		0.349219	0.010105	7.91745	0.835712	0.166379	0.006424	2.875406	0.078893	1929.5	47.9	2217.6	92.8	2525.7	54.9	1.74	384	223	201
32	7		0.345032	0.009496	7.386237	0.7685	0.156331	0.005939	2.908577	0.081078	1909.7	45.5	2156.5	91.2	2414.1	64.2	1.57	477	322	281
33	7		0.357846	0.006975	7.840954	0.796005	0.159835	0.006037	2.792723	0.053797	1971.6	33.1	2212.2	91.7	2451.9	64.4	1.5	554	374	355
34	7		0.410406	0.008961	9.244543	0.94799	0.163921	0.006191	2.437298	0.054466	2216.1	41.1	2360.9	95.1	2494.5	63.7	1.3	631	492	507
35	7		0.181242	0.003507	3.8405	0.39896	0.153288	0.006591	5.507052	0.106974	1073.7	19.1	1600.5	82.9	2380.1	72.8	11.66	1681	143	426
36	7		0.432712	0.009286	10.975278	1.128884	0.183919	0.007188	2.310236	0.049584	2317.3	41.8	2519.1	95.5	2685.5	65.9	2.42	368	152	199
37	7		0.475248	0.012735	14.71925	1.584815	0.223849	0.010943	2.105848	0.058897	2505.4	56	2792.7	100.8	3000.5	80.5	2.36	178	76	195
38	7		0.603337	0.012089	17.420299	1.771889	0.208653	0.007768	1.655271	0.033696	3042.5	48.7	2957.3	97.5	2893.6	60.3	1.5	522	356	527
39	7		0.592371	0.011002	17.182708	1.74043	0.209341	0.00795	1.684546	0.031324	2998.6	44.5	2944.5	97.5	2898.3	62.3	2.22	337	159	227
40	7		0.598058	0.011422	17.715253	1.803006	0.213262	0.008025	1.668552	0.031869	3021.5	46.1	2973.3	97.8	2928.7	60.7	2.81	361	132	235
41	7		0.296393	0.006358	8.400439	0.855999	0.202774	0.008621	3.367765	0.072876	1673	31.6	2274.2	93.4	2842.9	70	2.22	861	394	712
42	7		0.406587	0.01213	11.711969	1.225792	0.204828	0.007864	2.460325	0.075615	2198.1	55.8	2579.2	101.1	2863.2	62.7	1.03	758	731	1207
43	7		0.567316	0.01112	15.51512	1.574559	0.194853	0.007353	1.757849	0.034511	2896.2	45.7	2846.7	96.9	2781.6	62.3	1.1	377	339	512
44	7		0.481097	0.045089	24.335347	3.767231	0.361069	0.029606	2.114034	0.194966	2524.8	195.5	3264.3	160.9	3743.3	116.5	10.04	3015	295	10140
45	7		0.226655	0.009068	5.463601	0.652726	0.16999	0.007988	4.423305	0.169184	1316.2	47.4	1887.2	101.1	2552.2	77.9	16.73	2915	171	2351
46	7		0.208685	0.004738	4.148043	0.434037	0.141199	0.005757	4.776785	0.111189	1221.7	25.3	1662.7	86.7	2239.7	71.1	10.63	2775	256	1167
47	7		0.212595	0.00514	6.257508	0.701451	0.208733	0.010608	4.694092	0.114645	1242.4	27.3	2006.5	99.6	2887.3	82.4	9.95	3190	315	3123
48	7		0.202103	0.005626	4.013835	0.41759	0.141121	0.005727	4.953585	0.135038	1186.1	30.1	1634.7	82.2	2236.6	69.8	1.96	942	586	399
49	7		0.219513	0.005523	4.243459	0.448554	0.136966	0.005548	4.553639	0.109637	1278.9	29.1	1679	84.3	2184.8	70	1.3	849	647	402
50	7		0.285731	0.006668	6.027966	0.625626	0.14956	0.005853	3.49604	0.084064	1619.7	33.5	1977.3	92.9	2337.7	67.1	1.28	781	601	472
51	7		0.159188	0.004986	3.399943	0.352789	0.151851	0.005666	6.305061	0.204091	951.8	27.8	1502.3	82.2	2365.2	63.9	6.49	9473	1479	4296
52	7		0.172779	0.003597	3.856289	0.405314	0.158441	0.006676	5.774093	0.117913	1027.3	19.7	1601.7	81.4	2433.6	69.8	7.89	7820	991	4020
53	7		0.323052	0.00859	8.151984	0.837085	0.179848	0.007368	3.097214	0.082743	1803.7	41.8	2246.3	93.3	2646.9	68.4	1.96	800	418	614
54	7		0.511782	0.014151	14.595049	1.526615	0.203708	0.009209	1.952696	0.054215	2662.9	60.3	2786.7	99.8	2850.5	74.3	2.02	103	54	77
55	7		0.565358	0.01378	14.530418	1.51001	0.18201	0.007279	1.767949	0.045153	2887.2	57.2	2781.9	98.3	2667.5	69.5	2.99	169	58	71
56	7		0.1869	0.004097	3.016027	0.311362	0.115488	0.004949	5.333868	0.117597	1104.5	22.3	1410.9	77.6	1883.8	76.6	1.49	998	669	337
57	7		0.367366	0.009637	8.367091	0.871652	0.162401	0.006253	2.723126	0.072035	2015.9	45.5	2268.5	95.5	2478.2	64.9	2.73	509	185	184
58	7		0.290136	0.012314	11.170441	1.222724	0.276727	0.010869	3.491039	0.156027	1639.5	61.6	2529	103.7	3341.6	62.1	0.99	2025	2117	3739
59	7		0.165906	0.003221	2.284249	0.233228	0.098398	0.003714												

Sample	Spot name	Run	206/238	$\pm 2SE$	207/235	$\pm 2SE$	207/206	$\pm 2SE$	238/P206	$\pm 2SE$	206/238 Age	$\pm 2SE$	207/235 Age	$\pm 2SE$	207/206 Age	$\pm 2SE$	U/Th	U	Th	Pb
	60	7	0.118461	0.002631	1.446093	0.14923	0.086962	0.003399	8.394087	0.209338	721.6	15.2	907.3	60.9	1355.9	76	10.29	4939	478	223
ZB05	1	7	0.480675	0.010073	11.751159	1.202318	0.178037	0.006791	2.076867	0.044002	2529.5	43.9	2583.3	97	2632.4	64.2	2.08	391	190	251
	2	7	0.458689	0.011747	10.960441	1.141459	0.174028	0.00689	2.181547	0.056952	2432.4	52	2516.7	95.4	2593.4	65.3	1.99	294	150	190
	3	7	0.465364	0.010296	11.129457	1.136711	0.174309	0.006657	2.146552	0.049042	2462.4	45.5	2532.7	95.6	2597.1	63.5	1.95	407	213	296
	4	7	0.488633	0.013925	11.951775	1.239175	0.178261	0.006757	2.051729	0.061296	2562.7	60.7	2597.8	99.8	2634.7	63.1	1.82	393	222	298
	5	7	0.528453	0.011745	13.236433	1.354202	0.182595	0.007114	1.890339	0.042329	2733.9	49.6	2695.2	96	2673.7	63.8	1.86	294	164	252
	6	7	0.522934	0.010955	12.909287	1.311686	0.180029	0.007013	1.909234	0.040263	2710.9	46.4	2672.2	96.4	2650.2	64.4	1.63	348	219	307
	7	7	0.517039	0.011241	12.663045	1.291256	0.178495	0.006814	1.931729	0.042085	2685.7	47.8	2653.7	96.1	2636.7	62.8	1.8	379	218	309
	8	7	0.454681	0.011675	11.557204	1.184553	0.185382	0.007276	2.201117	0.057047	2414.7	51.8	2567.6	95.7	2698.5	64.9	2.07	346	173	266
	9	7	0.51088	0.012005	12.495895	1.271238	0.179043	0.006822	1.95701	0.047637	2659.2	51.5	2641.4	97.2	2641.8	63.7	1.7	389	235	330
	10	7	0.53528	0.011306	13.105567	1.331845	0.178384	0.006858	1.865594	0.039295	2762.9	47.4	2686.4	96.2	2635.4	64.1	1.56	333	225	326
	11	7	0.496957	0.010153	12.219695	1.249811	0.178624	0.0069	2.009765	0.042009	2600.1	43.9	2619.9	96.5	2637.5	63.7	1.82	377	206	280
	12	7	0.490004	0.012462	12.138222	1.24571	0.180573	0.006834	2.043143	0.050981	2569.2	53.7	2613.4	94.8	2656.1	62.6	1.91	366	193	289
	13	7	0.524512	0.011129	13.164512	1.341833	0.182254	0.006995	1.90511	0.040277	2717.5	47	2690.3	96.7	2671	64.2	1.93	299	155	229
	14	7	0.476098	0.010199	12.73268	1.347943	0.193768	0.00831	2.100896	0.048676	2509.2	48	2655.3	97.9	2768.5	68.9	1.67	425	256	458
	15	7	0.478357	0.010181	12.018988	1.234051	0.182275	0.007117	2.090746	0.047658	2519.2	47.2	2604	96.2	2670.7	64.7	2.13	313	174	134
	16	7	0.451604	0.009294	11.824407	1.201321	0.189938	0.00724	2.212756	0.045193	2401.8	41.2	2589.8	95	2739.4	63.4	1.52	401	268	359
	17	7	0.485305	0.00999	12.018519	1.220156	0.179589	0.006827	2.059369	0.041945	2549.7	43.2	2605.1	94.2	2646.9	62.8	1.91	491	260	441
	18	7	0.47522	0.016846	12.820307	1.459241	0.194575	0.009085	2.122199	0.07645	2502.9	73.6	2654.1	104	2772	74.9	1.51	539	355	760
	19	7	0.363972	0.010405	9.282413	0.972715	0.184658	0.007227	2.758052	0.078586	1999.6	49.1	2362.7	94.9	2692.1	65.2	1.62	763	513	619
	20	7	0.313975	0.006357	7.042053	0.716075	0.162498	0.006251	3.184167	0.064479	1759.9	31.2	2115.9	89.6	2479.2	65.1	2.25	632	286	277
	21	7	0.524077	0.009537	12.9605	1.313464	0.178973	0.006749	1.908484	0.034583	2716.1	40.3	2676.2	96	2641.4	63.2	1.63	373	230	314
	22	7	0.526344	0.012175	12.996171	1.326168	0.179192	0.006811	1.902451	0.043879	2725.3	51.4	2678.6	95.5	2643.8	63.2	1.54	369	241	352
	23	7	0.562756	0.012585	13.967394	1.438314	0.179643	0.007156	1.781279	0.039742	2876.9	51.8	2745.3	97	2645.9	67.1	2.41	133	56	82
	24	7	0.523185	0.010027	13.077179	1.336531	0.180192	0.006949	1.913789	0.036883	2712.2	42.5	2683.8	95.8	2652	64.4	1.39	335	249	351
	25	7	0.524304	0.009585	13.053657	1.325682	0.180201	0.006821	1.909387	0.034899	2717.1	40.5	2682.7	95.4	2652.7	63.3	1.73	449	262	370
	26	7	0.399067	0.010965	11.222262	1.148129	0.204001	0.008057	2.521294	0.075044	2163.3	51	2544.4	86	2855.2	64.8	1.1	663	660	826
	27	7	0.21083	0.005876	4.3328	0.447768	0.148824	0.005612	4.771551	0.129735	1232.7	31.2	1697.7	84.7	2330.4	65.1	2.48	1658	684	626
	28	7	0.466244	0.010029	11.433512	1.176482	0.178179	0.007295	2.148627	0.046666	2466.7	44.2	2558.2	96.8	2633.5	66.6	1.68	531	314	371
	29	7	0.432397	0.011156	10.746164	1.104516	0.179348	0.007109	2.31891	0.059098	2316	50.1	2500.6	96.6	2644.9	66.9	1.24	420	338	355
	30	7	0.279236	0.006519	5.941197	0.617799	0.153913	0.005853	3.592899	0.085542	1587.2	32.9	1965.3	91.2	2388	65.2	1.55	2016	1362	892
ZB06	1	7	0.454855	0.014486	12.345324	1.301339	0.199548	0.008231	2.217522	0.0702	2414.3	64	2626.7	97.8	2817.9	66.4	1.83	138	75	96
	2	7	0.553246	0.011879	15.242222	1.551118	0.202548	0.007695	1.812377	0.039101	2837.7	49.3	2829.5	97.8	2844.7	62.9	1.73	236	136	201
	3	7	0.518619	0.012482	14.535627	1.504528	0.206848	0.008992	1.931905	0.046205	2692.7	52.9	2784.1	98.4	2877.5	71.9	1.51	149	99	152
	4	7	0.338125	0.007212	8.624834	0.877377	0.186738	0.007035	2.964448	0.065914	1877.2	34.9	2298.3	92.4	2711.8	61.3	1.53	692	458	459
	5	7	0.544908	0.010346	14.738072	1.49807	0.198923	0.007681	1.837144	0.034579	2803.5	43.1	2797.6	97.2	2814.8	63.4	1.71	239	143	202
	6	7	0.565389	0.011345	15.272772	1.560097	0.198694	0.00781	1.771081	0.03569	2888.2	46.7	2831	96.5	2812.4	65	1.49	165	112	159
	7	7	0.369849	0.00852	9.357463	0.95707	0.185484	0.007134	2.710365	0.061164	2028	39.9	2372.3	92.7	2700	63.4	1.24	399	325	329
	8	7	0.18376	0.00643	4.96211	0.520992	0.197757	0.007973	5.454012	0.17473	1087.2	34.7	1811.9	87.7	2806	66.2	1.47	1304	898	802
	9	7	0.156438	0.005065	2.991897	0.324883	0.140229	0.005623	6.438932	0.201572	936.5	28.1	1400.8	80.4	2226	68.8	3.03	2908	970	641
	10	7	0.477349	0.011982	13.096168	1.352011	0.201151	0.007869	2.097998	0.053421	2514.9	52.4	2685.1	97	2842.1	76.8	1.89	266	142	208
	11	7	0.546271	0.010264	14.717408	1.492831	0.197731	0.007418	1.828349	0.03471	2809.2	42.9	2796.4	96.4	2805.8	61.3	1.54	280	181	247
	12	7	0.514338	0.010438	14.069047	1.426051	0.201292	0.007645	1.942571	0.03923	2674.5	44.4	2753.7	95.5	2834.6	62.5	1.79	379	210	280
	13	7	0.532559	0.011818	14.616286	1.493503	0.201221	0.007886	1.77343	0.041875	2751.2	49.7	2789.2	97.3	2833.1	63.8	1.6	180	112	143
	14	7	0.54757	0.011615	14.808753	1.506989	0.197519	0.007508	1.824714	0.039297	2814.2	48.5	2802	98.2	2808.5	69	1.31	326	249	339
	15	7	0.391161	0.009704	10.84909	1.11456	0.202216	0.007996	2.547899	0.071198	2136.1	50.2	2512.9	86.4	2840.9	64.1	1.87	249	132	179
	16	7	0.474796	0.01489	12.183225	1.29862	0.187341	0.007299	2.116117	0.069381	2502	65.5	2612.9	101.8	2716.1	64	45.78	659	14	28
	17	7	0.44416	0.020205	12.401839	1.349038	0.204298	0.007616	2.289241	0.109694	2363.4	90.7	2627.4	103.9	2859.2	61	1.48	1221	838	642
	18	7	0.563358	0.010603	15.368723	1.559938	0.199547	0.007489	1.765902	0.036897	2880	43.7	2837.6	97.5	2820.7	61.6	1.96	348	177	269
	19	7	0.544005	0.01388	15.118974	1.549534	0.203851	0.008479	1.834415	0.046724	2799.5	57.9	2822	95.8	2854.8	67.4	1.22	214	175	256
	20	7	0.543281																	

Sample	Spot name	Run	206/238	±2SE	207/235	±2SE	207/206	±2SE	238/P206	±2SE	206/238 Age	±2SE	207/235 Age	±2SE	207/206 Age	±2SE	U/Th	U	Th	Pb
		24	0.522576	0.01306	14.542623	1.507317	0.203107	0.008198	1.90877	0.047612	2709.2	55.3	2784.1	98.5	2848.7	66.1	1.24	160	131	177
		25	0.421037	0.014193	11.021376	1.204822	0.189533	0.008437	2.36986	0.080604	2264.4	64.4	2522.6	101.7	2735.4	75.2	1.78	290	166	215
		26	0.543365	0.012738	14.777435	1.516428	0.198613	0.007862	1.834258	0.042432	2796.9	53.1	2800.1	99.7	2812.8	64.8	1.5	324	217	307
		27	0.453242	0.015646	11.95067	1.289493	0.191418	0.007433	2.218625	0.080682	2406.5	69.9	2593.3	105.1	2751.6	63.3	1.5	416	299	322
		28	0.511464	0.012736	14.042957	1.443577	0.199799	0.007498	1.953343	0.049039	2661.4	54.3	2750.6	97.9	2822.8	61.4	2.19	316	149	195
		29	0.488717	0.010586	13.277849	1.359439	0.197785	0.007805	2.033868	0.040565	2564.3	46	2698	96.4	2804.8	64.7	1.76	199	112	145
		30	0.521429	0.011235	14.185768	1.449104	0.198044	0.00774	1.912987	0.041207	2704.4	47.6	2760.8	98.4	2807.2	64.1	1.38	185	134	182
ZB07		1	0.483216	0.012014	11.964418	1.220127	0.18019	0.006855	2.068844	0.051403	2539.9	52.2	2600.4	96.4	2652.4	62.7	2.21	291	132	171
		2	0.476347	0.010726	12.170814	1.246962	0.185876	0.007508	2.096648	0.047925	2510.4	46.9	2616	95.5	2702	65.8	2.87	171	59	73
		3	0.369863	0.009595	9.287256	0.964072	0.182367	0.007003	2.7056	0.07233	2027.7	45.3	2363.9	96.3	2672.1	62.4	4.65	472	101	113
		4	0.529817	0.012023	13.192588	1.368471	0.180929	0.007538	1.885654	0.042752	2739.6	50.6	2690.6	98.2	2656.2	70.8	3.55	107	30	42
		5	0.476873	0.013842	11.980188	1.235532	0.182851	0.00784	2.101804	0.063689	2511.7	60.8	2600.8	97.1	2673.3	70.6	3.44	122	37	35
		6	0.47888	0.012239	12.791832	1.308075	0.193522	0.007286	2.089716	0.052455	2521	53.2	2663	96.7	2770.5	61.5	3.33	420	140	221
		7	0.433754	0.011685	12.440188	1.295455	0.20881	0.008477	2.287815	0.051323	2321.8	52.9	2636.4	98.4	2893.7	65.6	1.31	299	224	304
		8	0.362218	0.018741	9.933954	1.094171	0.199799	0.007998	2.820804	0.141114	1986.9	88	2420.3	98.5	2820.9	65.8	5.55	595	102	320
		9	0.42877	0.009808	11.14647	1.145414	0.188698	0.007493	2.331753	0.052986	2299.3	44.2	2533.6	95.2	2727.4	64.8	0.01	226	32104	33991
		10	0.861815	0.03169	47.914615	5.039751	0.405596	0.016341	1.159865	0.044113	4005.1	110.5	3948.9	105.8	3928.7	60.4	0.01	406	33120	36706
		11	0.455661	0.01439	12.343845	1.286451	0.196053	0.008377	2.198212	0.068029	2419	63.5	2629	99.3	2789.9	70.4	0.01	235	33985	38866
		12	0.488265	0.014353	15.376579	1.58831	0.227277	0.009425	2.045726	0.059765	2562.6	62.1	2838.2	98.7	3031.2	66	0.01	162	29238	31916
		13	0.44028	0.008445	13.396495	1.380779	0.220182	0.008922	2.267027	0.043512	2351.6	37.8	2706.2	98.2	2978.6	66.3	0.01	231	42230	44175
		14	0.449607	0.010303	11.798714	1.212205	0.190305	0.00756	2.223932	0.050461	2392.6	45.7	2586.6	96.8	2741.3	66.2	0	182	48265	50949
		15	0.446431	0.009433	11.222275	1.155961	0.182902	0.007582	2.235248	0.047777	2379.1	42.1	2540.8	96.3	2676.7	68.9	0.01	358	56553	58261
		16	0.467805	0.008429	14.026964	1.442937	0.217275	0.008766	2.132485	0.038009	2473.7	36.9	2749.5	96.9	2956.8	64.9	0.01	655	54451	60048
		17	0.457311	0.010654	11.933722	1.235666	0.188985	0.007148	2.186213	0.051308	2426.7	47.2	2596.6	97	2731.3	62.6	0.01	900	75992	84564
		18	0.414961	0.010985	10.728936	1.093712	0.187913	0.007526	2.413019	0.064659	2236.2	50	2498.8	93.4	2720.3	65.7	0.01	304	44396	46809
		19	0.437911	0.016687	13.689998	1.526451	0.226509	0.009454	2.295691	0.089388	2338.7	75	2721	107.4	3023.9	67.5	0.01	258	46075	48437
		20	0.484621	0.016299	17.568653	1.852639	0.26347	0.011358	2.075272	0.072032	2544.2	71	2961.8	100.3	3261.8	66.8	0.01	276	38433	41236

C.2 Apatite U-Pb data

Sample	Spot name	238/206 $\pm 2\sigma$	207/206 $\pm 2\sigma$	Err. Corr	207/235 $\pm 2\sigma$	206/238 $\pm 2\sigma$	Err. Corr	207/235 Age $\pm 2\sigma$	206/238 Age $\pm 2\sigma$	207/206 Age $\pm 2\sigma$	206/238* Age $\pm 2\sigma$									
GW06	1	1.752	0.115	0.248	0.004	0.471	19.01	1.233	0.556	0.036	0.376	3068.4	57.49	2910.82	153.35	3173.25	26.1	2762.38	137.87	
	2	1.759	0.109	0.247	0.003	0.508	18.755	1.198	0.553	0.035	0.198	3058.1	54.8	2901.74	145.43	3162.48	20.58	2756.27	131	
	3	1.827	0.117	0.233	0.004	0.554	17.123	1.123	0.534	0.036	0.71	2966.61	54.64	2813.71	145.87	3071.98	24.76	2686.76	132.82	
	4	1.779	0.119	0.239	0.005	-0.017	18.023	1.232	0.548	0.036	0.704	3015.64	67.42	2875.54	155.42	3110.41	30.43	2750.44	141.99	
	5	1.851	0.117	0.219	0.003	0.391	15.944	1.026	0.526	0.033	0.483	2895.2	56.67	2783.95	142.74	2973.51	21.25	2696.66	133.81	
	6	1.595	0.101	0.336	0.004	0.202	28.463	1.855	0.611	0.039	0.53	3454.76	60.83	3137.26	157.13	3644.4	17.3	2704.53	116.19	
	8	1.687	0.111	0.286	0.004	0.188	22.83	1.506	0.579	0.038	0.684	3243.21	62.78	3000.08	157.38	3397.2	22.62	2732.81	130.19	
	9	1.819	0.12	0.213	0.004	0.584	15.728	1.022	0.536	0.034	0.07	2884.59	53.95	2824.63	150.47	2926.72	32.68	2776.26	145.28	
	10	1.817	0.116	0.217	0.003	0.281	16.075	1.042	0.538	0.035	0.68	2904.02	58.9	2826.3	146.03	2958.36	19.47	2762.95	139.46	
	11	1.768	0.115	0.253	0.002	0.443	19.28	1.236	0.553	0.035	0.727	3078.95	59.19	2890.22	150.95	3204.49	14.79	2714.71	132.91	
	12	1.785	0.112	0.25	0.004	0.153	18.785	1.227	0.547	0.035	0.493	3056.92	60	2867.91	144.92	3183.61	24.52	2697.04	127.92	
	13	1.687	0.12	0.287	0.007	-0.114	22.939	1.593	0.577	0.037	0.513	3246.64	76.44	3000.41	171.17	3402.46	39.82	2729.26	141.19	
	14	1.784	0.123	0.256	0.004	0.645	19.309	1.27	0.548	0.037	0.727	3079.37	57.63	2868.32	160.2	3220.07	26.97	2675.49	139.08	
	15	1.665	0.116	0.28	0.007	0.145	22.633	1.639	0.587	0.04	0.688	3233.01	68.94	3031.69	168.32	3360.39	41.66	2804.65	143.67	
	16	1.769	0.117	0.248	0.004	0.131	18.933	1.269	0.553	0.036	0.694	3059.68	63.87	2888.34	154.17	3174.2	26.4	2730.62	137.55	
	17	1.758	0.111	0.254	0.003	0.447	19.467	1.248	0.556	0.036	0.548	3087.42	56.84	2903.21	147.45	3209.53	18.47	2729.36	130.07	
	18	1.753	0.112	0.256	0.004	0.428	19.688	1.289	0.558	0.036	0.529	3098.12	56.93	2909.66	150.15	3222.66	27.11	2729.97	131.91	
	19	1.617	0.105	0.322	0.006	0.231	26.872	1.777	0.605	0.039	0.522	3400.37	61.98	3103.38	159.97	3580.34	27.94	2719.95	122.33	
	20	1.603	0.106	0.321	0.006	0.456	26.949	1.758	0.61	0.039	0.392	3404.15	59.09	3124.74	164.14	3572.98	27.27	2755.78	127.11	
	21	1.76	0.116	0.262	0.003	0.331	20.303	1.309	0.556	0.036	0.773	3114.61	60.97	2899.83	153.91	3256.14	18.71	2695.22	132.65	
	22	1.815	0.117	0.23	0.004	0	17.039	1.128	0.539	0.034	0.569	2959.21	64.12	2829.79	147.81	3048.4	27.41	2721.32	136.53	
	23	1.81	0.113	0.234	0.003	0.458	17.416	1.119	0.54	0.034	0.314	2981.19	55.27	2834.91	143.66	3081.41	22.96	2710.55	131.15	
	24	1.775	0.113	0.243	0.004	0.515	18.42	1.187	0.55	0.035	0.241	3035.43	55.16	2880.41	148.08	3139.73	25.29	2740.2	133.8	
	25	1.84	0.116	0.224	0.004	0.368	16.388	1.064	0.531	0.034	0.382	2922.94	56.76	2797.46	143.5	3010.52	25.86	2696.67	133.2	
	26	1.797	0.117	0.231	0.004	0.493	17.268	1.123	0.543	0.035	0.479	2973.55	56.42	2851.72	150.16	3057.02	27.03	2747.08	139.18	
	27	1.819	0.116	0.228	0.003	0.571	16.95	1.085	0.537	0.034	0.522	2952.29	55.34	2824.29	146.34	3040.71	20.47	2717.84	135.35	
	28	1.71	0.114	0.284	0.01	-0.695	22.45	1.948	0.572	0.039	0.906	3222.99	92.57	2968.1	158.61	3385.68	55.88	2698.1	130.66	
	29	1.891	0.121	0.201	0.002	0.154	14.323	0.924	0.516	0.033	0.797	2795.06	60.04	2736.75	142.53	2837.41	14.89	2694.66	138.12	
	30	1.874	0.118	0.207	0.003	-0.327	14.881	0.989	0.521	0.033	0.839	2830.65	64.57	2756.49	140.68	2883.9	20.05	2701.21	135.01	
	31	1.848	0.118	0.205	0.003	0.398	14.878	0.956	0.527	0.034	0.47	2832.42	57.19	2788.33	144.73	2863.95	20.73	2754.61	141.2	
	32	1.757	0.113	0.247	0.003	0.597	18.904	1.216	0.555	0.036	0.423	3062.46	56.27	2904.94	149.88	3167.49	17.17	2757.68	134.85	
	34	1.77	0.112	0.251	0.003	0.045	19.023	1.235	0.55	0.035	0.694	3068.53	61.91	2886.63	147.85	3189.87	18.54	2718.76	130.91	
	35	1.738	0.112	0.276	0.006	-0.09	21.292	1.449	0.56	0.036	0.574	3177.39	67.66	2930.15	151.52	3337.47	33.77	2681.63	126.56	
	36	1.793	0.112	0.236	0.005	0.214	17.635	1.166	0.542	0.034	0.181	2996.94	59.3	2857.1	144.6	3092.14	32.99	2735.19	132.34	
	38	1.783	0.116	0.238	0.004	-0.309	17.887	1.225	0.546	0.035	0.757	3009.44	70.27	2870.25	150.74	3103.78	29.57	2746.87	137.86	
	39	1.827	0.12	0.229	0.003	-0.046	16.901	1.111	0.533	0.034	0.583	2952.08	64.87	2814.12	149.39	3047.5	22.19	2700.31	137.38	
	40	1.761	0.114	0.249	0.003	0.352	18.989	1.228	0.552	0.035	0.433	3065.43	59.22	2899.4	150.8	3176.12	19.74	2744.73	134.91	
	41	1.281	0.083	0.49	0.015	0.076	51.217	3.66	0.758	0.049	0.399	4045.05	69.22	3719.03	182.33	4210.98	45.98	2654.28	91.62	
	42	1.163	0.089	0.543	0.014	0.572	62.161	4.114	0.833	0.056	0.458	4243.68	65.43	4000.37	228.84	4361.08	38.31	2662.56	99.59	
	43	1.167	0.076	0.532	0.014	0.7	60.94	4.01	0.835	0.056	0.289	4219.01	50.29	3988.82	193.31	4330.6	38.3	2730.12	89.04	
	44	1.312	0.087	0.466	0.01	0.046	47.616	3.304	0.741	0.049	0.679	3971.37	68.31	3652.05	184.19	4136.81	32.51	2714.23	100.53	
	45	1.745	0.116	0.239	0.006	0.556	18.27	1.21	0.556	0.036	0.271	3033.85	54.41	2919.93	155.42	3110.21	38.64	2813.32	144.1	
	46	1.812	0.114	0.231	0.003	0.024	17.036	1.1	0.535	0.034	0.743	2967.1	61.35	2833.3	144.69	3059.1	18.2	2720.44	133.23	
	47	1.757	0.117	0.242	0.005	0.511	18.437	1.204	0.552	0.036	0.349	3042.92	56.89	2904.47	155.42	3135.67	30.74	2776.2	141.79	
	48	1.762	0.117	0.249	0.004	0.414	18.89	1.236	0.55	0.036	0.457	3066.64	58.85	2897.16	154.3	3179.61	28.33	2739.5	137.72	
	50	1.754	0.109	0.267	0.004	0.251	20.379	1.354	0.551	0.035	0.343	3137.27	58.22	2908.88	145.29	3286.84	25.72	2687.4	123.7	
	NG01	3	1.537	0.099	0.399	0.008	0.159	34.598	2.326	0.629	0.041	0.534	3661.32	63.45	3230.69	163.31	3906.11	30.39	2534.03	99.67
		4	1.524	0.099	0.385	0.011	-0.337	33.628	2.471	0.634	0.041	0.702	3632.56	78.17	3252.63	165.6	3849.26	43.21	2631.46	107.6
		6	0.56	0.04	0.834	0.018	0.294	199.567	14.448	1.739	0.126	0.819	5410.84	68.59	6602.39	293.35	4980.66	31.31	292.59	0.53

Continued on next page

Sample	Spot name	238/206 $\pm 2\sigma$	207/206 $\pm 2\sigma$	Err. Corr	207/235 $\pm 2\sigma$	206/238 $\pm 2\sigma$	Err. Corr	207/235 $\pm 2\sigma$	206/238 $\pm 2\sigma$	Err. Corr	207/206 $\pm 2\sigma$	206/238* $\pm 2\sigma$
7		0.377	0.025	0.896	0.013	0.252		317.482	21.929	2.574	0.177	0.881
8		0.519	0.034	0.84	0.015	0.063		216.44	15.478	1.87	0.131	0.842
9		0.195	0.017	0.973	0.019	0.375		686.169	61.397	5.14	0.479	0.959
10		0.209	0.018	0.98	0.028	0.083		645.848	61.53	4.789	0.441	0.906
11		0.207	0.02	0.974	0.031	0.486		652.109	62.678	4.802	0.478	0.928
15		1.446	0.094	0.425	0.009	0.363		39.053	2.575	0.667	0.043	0.291
16		1.377	0.093	0.491	0.009	0.021		47.546	3.313	0.702	0.047	0.768
17		0.902	0.069	0.675	0.019	-0.61		101.63	8.977	1.085	0.082	0.937
18		1.306	0.084	0.492	0.009	0.594		50.038	3.248	0.739	0.048	0.376
19		0.91	0.072	0.659	0.022	0.19		96.532	7.859	1.064	0.083	0.781
20		1.074	0.09	0.621	0.021	0.728		77.481	5.553	0.897	0.068	0.756
22		0.829	0.067	0.701	0.014	-0.199		113.049	8.913	1.185	0.094	0.934
24		0.844	0.093	0.704	0.03	-0.264		115.98	14.346	1.188	0.124	0.914
25		1.316	0.095	0.487	0.02	-0.492		49.595	4.578	0.737	0.055	0.855
26		1.597	0.103	0.36	0.007	0.121		29.969	2.006	0.605	0.039	0.604
27		1.898	0.122	0.333	0.012	-0.464		23.38	1.827	0.508	0.033	0.705
29		1.405	0.092	0.446	0.008	0.346		42.17	2.764	0.687	0.044	0.427
30		1.639	0.106	0.375	0.005	0.347		30.414	1.954	0.591	0.038	0.518
31		0.234	0.021	0.964	0.025	0.412		551.632	42.666	4.163	0.324	0.894
32		0.24	0.02	0.795	0.022	-0.029		446.93	40.564	4.078	0.352	0.902
33		0.18	0.017	0.959	0.026	0.1		695.437	54.279	5.508	0.609	0.958
34		0.07	0.012	1.054	0.026	0.203		2434.883	459.086	16.963	3.161	0.979
35		2.015	0.132	0.22	0.002	0.076		14.636	0.963	0.482	0.031	0.862
36		0.09	0.013	1.007	0.043	-0.157		1679.125	257.772	11.985	1.684	0.954
37		0.091	0.012	1.032	0.035	0.205		1652.849	204.698	11.635	1.535	0.973
38		0.167	0.029	0.967	0.052	-0.282		984.315	235.074	6.777	1.292	0.977
39		0.27	0.022	0.954	0.016	0.152		481.805	37.903	3.669	0.287	0.93
40		0.236	0.019	0.976	0.019	0.192		566.727	47.78	4.219	0.356	0.942
41		0.109	0.01	1.036	0.023	-0.127		1319.57	126.87	9.226	0.851	0.948
42		1.704	0.11	0.324	0.006	-0.123		25.267	1.7	0.572	0.038	0.741
43		0.301	0.021	0.942	0.021	0.284		419.844	33.062	3.243	0.256	0.904
44		0.97	0.101	0.653	0.058	-0.789		93.241	16.264	1.021	0.11	0.936
46		1.503	0.098	0.416	0.008	0.364		37.142	2.46	0.649	0.042	0.412
47		1.372	0.088	0.472	0.007	0.154		46.234	3.094	0.711	0.047	0.733
49		0.271	0.04	0.932	0.051	0.323		473.423	58.262	3.703	0.481	0.856
50		0.285	0.022	0.955	0.022	0.159		457.757	38.237	3.481	0.286	0.912
51		1.464	0.095	0.479	0.011	0.133		43.992	2.979	0.666	0.043	0.44
52		1.219	0.085	0.619	0.022	0.426		68.261	4.963	0.802	0.057	0.451
53		0.954	0.078	0.659	0.022	0.391		93.135	7.346	1.028	0.08	0.762
54		1.28	0.074	0.519	0.009	0.001		54.376	3.678	0.76	0.05	0.7
55		1.862	0.119	0.273	0.004	0.371		19.735	1.269	0.524	0.033	0.315
56		1.814	0.123	0.287	0.004	0.362		21.343	1.407	0.539	0.036	0.737
57		1.764	0.113	0.307	0.004	0.266		23.435	1.516	0.553	0.035	0.512
58		0.538	0.036	0.835	0.016	0.148		209.852	14.743	1.826	0.127	0.782
59		0.206	0.021	1	0.043	0.085		674.37	72.811	4.891	0.477	0.889
60		0.355	0.032	0.884	0.023	0.32		339.344	28.295	2.789	0.23	0.891
61		0.242	0.02	0.968	0.025	0.166		553.137	48.658	4.153	0.36	0.907
62		1.456	0.098	0.435	0.007	0.422		40.17	2.619	0.671	0.044	0.588
63		0.256	0.034	0.954	0.021	-0.128		553.834	69.318	4.086	0.461	0.979
64		0.208	0.016	0.986	0.017	0.321		645.19	55.718	4.762	0.419	0.954
65		0.138	0.013	1	0.02	0.447		1000.646	85.995	7.294	0.661	0.957
66		0.101	0.011	1.033	0.023	0.344		1462.546	162.037	10.322	1.16	0.971
67		0.083	0.008	1.008	0.021	0.152		1674.769	167.82	12.068	1.214	0.965
68		0.374	0.026	0.907	0.021	0.248		323.941	22.349	2.619	0.182	0.724
69		0.203	0.018	0.981	0.02	0.423		664.711	54.641	4.936	0.414	0.938
71		0.582	0.045	0.808	0.019	-0.195		187.173	15.78	1.684	0.133	0.94

Continued on next page

Sample	Spot name	238/206 $\pm 2\sigma$	207/206 $\pm 2\sigma$	Err. Corr	207/235 $\pm 2\sigma$	206/238 $\pm 2\sigma$	Err. Corr	207/235 Age $\pm 2\sigma$	206/238 Age $\pm 2\sigma$	207/206 Age $\pm 2\sigma$	206/238* Age $\pm 2\sigma$									
72		1.08	0.1	0.587	0.016	-0.147		71.751	5.121	0.892	0.061	0.861	4396.69	100.58	4226.2	287.53	4475.79	39.44	2607.88	107.1
73		1.89	0.121	0.254	0.004	0.335		18.015	1.17	0.515	0.033	0.525	3017.13	58.71	2737.94	143.39	3208.79	23.33	2513.01	120.51
74		1.881	0.118	0.258	0.003	0.174		18.413	1.192	0.517	0.033	0.608	3038.36	59.68	2748.18	140.59	3236.33	19.1	2510.2	117
76		1.562	0.102	0.411	0.007	0.275		35.322	2.364	0.624	0.041	0.69	3674.11	61.99	3189.24	164.13	3950.16	25.49	2429.28	94.42
77		1.418	0.094	0.446	0.008	0.355		42.188	2.795	0.687	0.045	0.601	3850.7	61.57	3440.28	176.27	4071.79	27.13	2556.79	96.33

C.3 Titanite U-Pb data

Sample	Spot name	238/206	$\pm 2\sigma$	207/206	$\pm 2\sigma$	Err. Corr	207/235	$\pm 2\sigma$	206/238	$\pm 2\sigma$	Err. Corr	207/235	$\pm 2\sigma$	206/238	$\pm 2\sigma$	207/206	$\pm 2\sigma$	207/206*	$\pm 2\sigma$
												Age	$\pm 2\sigma$	Age	$\pm 2\sigma$	Age	$\pm 2\sigma$	Age	$\pm 2\sigma$
ZL181	g1.001	1.905	0.088	0.224	0.01	0.371	0.526	0.026	16.231	0.848	0.536	2888.79	52	2723.65	107.57	3006.3	69.24	2594.01	95.95
	g1.002	1.786	0.062	0.225	0.004	0.217	0.562	0.02	17.401	0.64	0.75	2954.81	35.09	2872.55	81.1	3012.35	27.79	2789.66	33.47
	g1.003	1.847	0.064	0.21	0.004	0.375	0.543	0.018	15.744	0.55	0.578	2859.43	34.01	2795.74	76.76	2905.25	33.75	2737.14	34.73
	g1.004	1.853	0.06	0.21	0.004	0.292	0.541	0.018	15.647	0.561	0.583	2853.28	34.26	2787.15	74.65	2900.73	33.98	2726.93	34.98
	g1.005	1.844	0.064	0.213	0.004	0.415	0.544	0.019	15.976	0.558	0.609	2873.43	33.55	2798.85	77.71	2926.92	32.77	2731.38	34.87
	g1.006	1.776	0.057	0.248	0.005	0.053	0.565	0.019	19.429	0.787	0.781	3059.77	38.81	2884.48	77.47	3165.92	30.83	2719.89	43.94
	g1.007	1.858	0.06	0.221	0.004	0.108	0.54	0.017	16.37	0.603	0.625	2896.12	36.43	2780.42	72.82	2988.84	26.55	2679.03	36.17
	g1.008	1.839	0.065	0.234	0.005	0.337	0.546	0.019	17.6	0.652	0.695	2965.58	35.14	2807.96	81.01	3075.53	32.04	2663.53	45.71
	g1.009	1.816	0.062	0.235	0.005	0.186	0.553	0.019	17.919	0.677	0.601	2982.66	36.62	2834.56	77.67	3084.66	35.33	2698.67	44.6
	g1.010	1.734	0.064	0.257	0.007	0.36	0.58	0.021	20.491	0.834	0.52	3111.27	39.17	2944.88	87.11	3221.59	44.87	2762.19	59.72
	g1.011	1.885	0.065	0.224	0.005	0.248	0.532	0.018	16.414	0.616	0.578	2899.32	36.16	2748.04	75.59	3006.31	36.71	2624.97	46.95
	g1.012	1.8	0.064	0.243	0.005	0.32	0.558	0.02	18.715	0.695	0.72	3024.7	35.57	2857.5	83.01	3139.06	30.7	2696.49	44.67
	g1.013	1.759	0.065	0.254	0.006	0.25	0.571	0.02	20.012	0.785	0.607	3088.66	38.99	2910.28	83.38	3218.85	32.69	2727.5	52.45
	g1.014	1.55	0.067	0.335	0.011	0.329	0.65	0.028	29.986	1.428	0.559	3481.31	46.68	3225.61	108.73	3633.93	53.75	2808.05	90.88
	g1.015	1.615	0.058	0.339	0.009	0.221	0.622	0.023	28.747	1.088	0.593	3442.08	37.21	3116.33	89.89	3652.13	40.85	2651.46	82.97
	g1.016	1.742	0.082	0.296	0.012	0.276	0.579	0.026	23.56	1.289	0.551	3243.98	52.71	2940.88	107.37	3439.24	64.33	2594.12	115.13
	g1.017	1.732	0.065	0.28	0.011	0.243	0.58	0.021	22.328	1.063	0.353	3192.23	46.15	2946.65	84.51	3350.61	59.83	2674.75	99.77
	g1.018	1.856	0.068	0.226	0.005	0.485	0.541	0.02	16.854	0.592	0.589	2924.95	34	2785.86	82.27	3023.1	34.92	2664.39	45.68
	g1.019	1.922	0.083	0.211	0.004	0.333	0.523	0.022	15.244	0.632	0.825	2828.06	39.36	2711.24	92.54	2914.38	30.76	2612.72	37.88
	g1.020	1.886	0.063	0.2	0.004	0.372	0.532	0.017	14.664	0.498	0.202	2792.03	32.27	2748.12	73.44	2822.86	37.37	2707.58	35.46
	g2.001	1.815	0.074	0.211	0.006	0.249	0.553	0.022	16.084	0.724	0.59	2879.87	42.75	2834.7	91.73	2911.97	48.19	2789.67	50.21
	g2.002	1.813	0.064	0.223	0.004	-0.005	0.554	0.02	17.029	0.681	0.776	2933.11	38.64	2839.7	81.74	2998.94	30.34	2748.47	34.46
	g2.003	1.845	0.069	0.218	0.005	0.421	0.545	0.019	16.322	0.605	0.619	2893.25	35.57	2800.53	81.87	2959.68	37.05	2711.53	44.2
	g2.004	1.863	0.061	0.216	0.004	0.107	0.538	0.018	16.04	0.599	0.652	2876.68	35.35	2775.13	73.69	2948.89	32.46	2688.68	35.93
	g2.005	1.805	0.064	0.224	0.006	0.372	0.556	0.019	17.111	0.645	0.537	2938.46	35.93	2848.85	80.18	3001.14	39.26	2758.94	51.31
	g2.006	1.784	0.063	0.238	0.008	-0.264	0.563	0.02	18.435	0.949	0.695	3017.38	54.08	2877.19	83.13	3112.15	56.99	2743.7	69.14
	g2.007	1.827	0.065	0.231	0.005	0.318	0.549	0.019	17.474	0.655	0.537	2958.61	36	2820.02	78.32	3054.88	38.1	2694.62	44.72
	g2.008	1.829	0.056	0.228	0.006	0.21	0.548	0.017	17.177	0.678	0.366	2941.61	37.73	2814.39	70.06	3029.23	45.21	2702.23	53.38
	g2.009	1.806	0.054	0.247	0.006	0.28	0.552	0.019	18.738	0.729	0.552	3025.42	37.33	2831.95	79.83	3157.34	39.41	2671.21	54.55
	g2.010	1.758	0.057	0.261	0.006	0.453	0.57	0.018	20.512	0.714	0.33	3113.84	34.2	2907.13	75.72	3250.12	39.05	2702.15	53.39
	g2.011	1.788	0.063	0.262	0.007	0.072	0.562	0.021	20.296	0.909	0.655	3100.59	43.15	2871.21	84.26	3253.67	43.55	2645.7	64.79
	g2.012	1.769	0.068	0.27	0.014	-0.083	0.568	0.022	21.191	1.445	0.504	3133.55	63.13	2898.71	88.57	3287.01	80.65	2648	129.37
	g2.013	1.696	0.077	0.29	0.011	-0.242	0.589	0.024	23.337	1.276	0.784	3233.08	50.72	3004.24	111.34	3408.1	57.83	2701.33	97.93
	g2.014	1.759	0.065	0.271	0.017	-0.487	0.572	0.022	21.47	1.842	0.789	3141.02	67.17	2911.33	88.69	3315.48	95.15	2661.65	155.6
	g2.015	1.85	0.065	0.236	0.005	0.163	0.542	0.019	17.643	0.682	0.726	2967.58	36.81	2791.43	79.02	3090.17	31.9	2638.26	46.52
	g2.016	1.874	0.061	0.228	0.006	0.219	0.535	0.017	16.807	0.653	0.482	2920.77	37.55	2759.44	73.04	3033.94	41.73	2628.26	56.21
	g2.017	1.627	0.06	0.315	0.009	0.302	0.616	0.023	26.701	1.091	0.5	3369.77	40.02	3093.51	90.54	3539.14	43.39	2731.88	78.43
	g2.018	1.822	0.064	0.25	0.005	0.165	0.551	0.02	19	0.775	0.714	3038.31	38.59	2827.33	83.5	3182.26	34.86	2633.23	46.68
	g2.019	1.846	0.071	0.23	0.005	0.357	0.544	0.021	17.239	0.685	0.662	2945.84	37.7	2796.64	88.4	3050.72	37.54	2666.44	45.61
	g2.020	1.678	0.07	0.286	0.011	-0.506	0.598	0.026	23.649	1.659	0.86	3246.32	64.14	3019.4	104.28	3390.44	56.17	2752.66	94.48
	g2.021	1.83	0.066	0.225	0.004	0.363	0.549	0.02	17.004	0.625	0.732	2932.54	35.24	2817.12	84.25	3014.38	30.47	2711.52	35.36
	g2.022	1.836	0.055	0.239	0.007	0.318	0.545	0.017	17.97	0.686	0.242	2985.34	36.19	2804	69.59	3109.36	44.75	2650.32	64.58
	g2.023	1.67	0.067	0.278	0.014	-0.352	0.6	0.024	22.995	1.679	0.744	3220.48	68.98	3027.07	97.68	3342.92	75.51	2802.25	116.13
	g2.024	1.117	0.042	0.57	0.013	-0.224	0.899	0.035	70.77	3.459	0.85	4332.82	47.14	4130.7	116.79	4429.56	32.62	2611.73	123.21
	g2.025	1.341	0.05	0.46	0.01	0.325	0.747	0.028	47.314	1.821	0.698	3935.54	38.17	3594.41	102.07	4114.79	31.96	2671.13	90.93
	g2.026	1.671	0.071	0.314	0.011	-0.201	0.602	0.025	26.08	1.54	0.736	3340.84	56.85	3033.32	99.78	3532.36	54.17	2649.72	101.53
	g2.027	1.528	0.068	0.359	0.012	0.01	0.658	0.028	32.53	1.835	0.749	3560.25	53.5	3254.72	111.08	3738.61	49.95	2744.29	103.67
g2.028	1.711	0.067	0.306	0.011	-0.318	0.588	0.024	24.904	1.492	0.781	3293.99	56.74	2977.05	95.41	3494.43	52.99	2608.97	104.45	
g2.029	1.809	0.06	0.235	0.005	0.08	0.554	0.019	18.055	0.745	0.653	2988.93	39.82	2839.75	79.2	3080.37	32.34	2710.76	44.22	
g2.030	1.892	0.083	0.226	0.01	-0.053	0.53	0.023	16.53	1.076	0.638	2901.8	57.85	2740.08	96.71	3016.65	66.89	2607.06	95.08	

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Sample	Spot name	238/206 $\pm 2\sigma$	207/206 $\pm 2\sigma$	Err. Corr	207/235 $\pm 2\sigma$	206/238 $\pm 2\sigma$	Err. Corr	207/235 $\pm 2\sigma$	206/238 $\pm 2\sigma$	207/206 $\pm 2\sigma$	207/206* $\pm 2\sigma$								
	g2.031	1.732	0.062	0.256	0.008	-0.267	0.579	0.021	20.479	1.052	0.763	3107.04	49.64	2943.11	86.55	3215.45	46.89	2769.97	67.88
	g2.032	1.866	0.064	0.204	0.005	0.477	0.537	0.019	14.983	0.486	0.498	2812.84	30.95	2769.5	78.58	2854.94	36.38	2726.55	43.74
	g2.033	1.826	0.066	0.213	0.006	-0.636	0.549	0.021	15.866	0.791	0.909	2878.89	55.54	2818.09	86.88	2922.24	43.74	2762.63	51.17
	g2.034	1.869	0.068	0.208	0.004	0.638	0.537	0.02	15.335	0.5	0.545	2834.88	31.45	2766.79	81.8	2885.08	35.04	2707.28	35.46
	g2.035	1.635	0.102	0.342	0.023	-0.523	0.617	0.04	29.358	3.351	0.872	3444.68	113.73	3094	156.51	3657.73	109.39	2599.63	219.83
	g2.036	1.817	0.062	0.26	0.007	0.19	0.551	0.019	19.716	0.82	0.566	3074.07	39.88	2828.72	80.16	3239.08	42.38	2604.65	66.67
	g2.037	1.679	0.063	0.3	0.009	0.116	0.597	0.022	24.706	1.118	0.619	3291.62	43.76	3017.14	89.54	3464.49	44.49	2692.6	80.62
	g2.038	1.765	0.061	0.279	0.006	-0.042	0.564	0.017	21.795	0.854	0.654	3171.51	37.75	2892.9	82.2	3352.9	34.47	2620.38	56.52
	g2.039	1.713	0.063	0.255	0.005	0.546	0.585	0.021	20.553	0.691	0.647	3116.32	32.72	2967.12	85.86	3215.03	30.31	2810.93	41.22
	g2.040	1.79	0.071	0.23	0.006	0.496	0.561	0.021	17.727	0.681	0.562	2972.2	36.79	2868.54	88.88	3044.85	42.48	2763.29	51.15
	g3.001	1.813	0.067	0.228	0.006	0.387	0.553	0.02	17.365	0.669	0.516	2952.14	38.41	2836.37	82.91	3032.96	43.79	2729.85	52.36
	g3.002	1.912	0.064	0.206	0.004	0.264	0.524	0.017	14.838	0.536	0.574	2802.74	33.8	2713.14	72.73	2868.21	33.76	2645.23	37.03
	g3.003	1.841	0.061	0.213	0.004	0.185	0.544	0.018	15.987	0.583	0.653	2873.71	34.1	2798.27	74.34	2927.43	31.22	2736.53	34.75
	g3.004	1.85	0.078	0.22	0.006	0.38	0.538	0.019	16.446	0.665	0.529	2899.67	38.65	2774	80.19	2976.07	45.55	2695.9	53.62
	g3.005	1.767	0.068	0.237	0.004	0.336	0.568	0.021	18.497	0.684	0.817	3013.41	35.99	2896.63	87.92	3094.16	25.5	2778.69	33.73
	g3.006	1.742	0.062	0.253	0.006	0.304	0.575	0.02	20.079	0.777	0.554	3092.12	37.02	2928.27	83.28	3201.29	38.87	2762.92	51.16
	g3.007	1.702	0.062	0.275	0.008	0.1	0.589	0.021	22.31	0.881	0.084	3194.03	38.8	2982.88	84.43	3330.05	47.54	2751.03	68.79
	g3.008	1.586	0.054	0.325	0.009	0.282	0.628	0.024	28.071	1.18	0.582	3417.41	41.88	3137.85	95.08	3587.25	43.04	2773.93	76.15
	g3.009	1.588	0.067	0.361	0.011	0.277	0.633	0.025	31.472	1.373	0.619	3528.93	44.24	3159.29	100.31	3748.17	44.01	2607.8	104.54
	g3.010	1.522	0.065	0.349	0.013	-0.075	0.66	0.026	31.8	1.774	0.652	3535.68	56.7	3263.98	102.52	3694.5	59.08	2806.22	107.54
	g3.011	1.625	0.066	0.347	0.01	0.12	0.619	0.025	29.574	1.411	0.729	3466.83	46.3	3100.91	100.06	3687.85	42.21	2597.17	95.74
	g3.012	1.76	0.081	0.25	0.01	-0.131	0.569	0.028	19.667	1.296	0.742	3070.55	62.97	2903.1	113.59	3182.8	62.29	2741.2	86.58
	g3.013	1.847	0.086	0.251	0.009	0.167	0.543	0.024	18.785	1.022	0.613	3026.97	52.18	2794.24	102.82	3186.51	58.93	2589.07	86.65
	g3.014	1.56	0.086	0.313	0.016	-0.027	0.646	0.035	27.911	2.149	0.707	3406.33	73.85	3206.52	135.87	3528.19	78.56	2888.49	124.91
	g3.015	1.797	0.068	0.225	0.007	0.492	0.558	0.02	17.3	0.661	0.349	2948.79	36.09	2857.21	83.95	3012.88	47.68	2769.57	59.41
	g3.016	0.768	0.04	0.747	0.038	0.328	1.322	0.071	132.081	6.702	0.558	4958.3	52.93	5410.96	195.27	4752.94	51.3	2193.14	480.92
	g3.017	0.981	0.053	0.624	0.027	-0.067	1.021	0.048	89.064	6.121	0.77	4552.26	72.41	4524.13	153.55	4551.07	64.67	2680.36	243.93
	g3.018	0.918	0.044	0.66	0.019	0.221	1.101	0.055	100.008	5.084	0.765	4678.41	52.23	4774.72	165.55	4640.57	42.36	2595.78	182.08
	g3.019	1.856	0.058	0.219	0.006	0.105	0.539	0.017	16.303	0.648	0.53	2891.56	37.7	2778.02	71.32	2971.35	41.19	2689.51	53.86
	g3.020	1.772	0.054	0.264	0.006	0.083	0.562	0.019	20.406	0.824	0.686	3107.13	39.6	2872.3	81.08	3263.5	35.25	2665.75	54.76
	g3.021	0.479	0.045	0.898	0.045	0.338	2.227	0.214	272.04	25.399	0.819	5664.69	92.3	7450.5	419.12	4788.39	98.83		
	g3.022	1.067	0.046	0.565	0.026	0.227	0.937	0.046	72.629	4.219	0.554	4354.07	60.18	4250.32	153.14	4405.98	65.53	2852.22	208.23
	g3.023	1.721	0.059	0.281	0.007	0.226	0.581	0.019	22.664	0.917	0.549	3209.06	39.51	2953.22	77.17	3362.01	41.54	2690.91	62.78
	g3.024	1.812	0.062	0.237	0.007	0.272	0.553	0.019	18.045	0.75	0.418	2988.05	41.25	2834.99	79.09	3092.95	50.11	2698.13	62.46
	g3.025	1.861	0.06	0.222	0.006	0.312	0.537	0.017	16.409	0.612	0.387	2898.43	35.25	2770.96	70.94	2987.98	41.28	2670.53	54.58
ZL182	1.001	1.907	0.07	0.2	0.003	0.226	0.525	0.019	14.421	0.524	0.802	2775.52	35.11	2716.85	79.43	2819.84	25.13	2673.76	27.23
	1.002	1.783	0.054	0.263	0.005	0.331	0.559	0.017	20.261	0.664	0.555	3102.79	31.28	2863.78	71.23	3261.25	27.15	2650.49	46.12
	1.003	1.664	0.056	0.307	0.006	0.232	0.601	0.021	25.437	0.951	0.684	3322.48	36.36	3030.76	83.52	3504.38	31.08	2692.34	53.75
	1.004	1.73	0.059	0.283	0.006	0.249	0.578	0.02	22.541	0.846	0.597	3204.58	36.24	2937.27	79.91	3377.14	34.56	2666.5	54.73
	1.005	1.919	0.067	0.21	0.004	0.249	0.521	0.019	15.052	0.566	0.671	2815.81	37.16	2702.52	78.84	2899.03	35.11	2620.72	37.67
	1.006	1.883	0.063	0.217	0.004	0.531	0.531	0.018	15.847	0.509	0.592	2866.39	30.72	2742.21	75.11	2955.65	28.69	2652.7	36.84
	1.007	1.913	0.06	0.214	0.004	0.31	0.522	0.017	15.369	0.523	0.54	2836.68	32.71	2705.61	70.55	2931.4	31.31	2616.2	37.79
	1.008	1.846	0.06	0.214	0.003	0.442	0.541	0.018	15.956	0.513	0.735	2872.87	30.89	2786.65	74.49	2934.57	23.24	2724.34	26.28
	1.009	1.815	0.068	0.215	0.005	0.334	0.551	0.02	16.285	0.621	0.608	2891.39	36.63	2826.19	82.1	2937.86	37.75	2774.74	42.28
	1.01	1.87	0.067	0.222	0.005	0.279	0.535	0.019	16.357	0.612	0.588	2895.19	36.24	2759.17	78.5	2992.14	36.03	2655.9	45.95
	1.011	1.865	0.055	0.222	0.004	0.261	0.533	0.018	16.273	0.59	0.634	2890.63	34.83	2751.27	75.24	2989.84	32.71	2664	36.55
	1.012	1.848	0.058	0.225	0.004	0.342	0.54	0.017	16.761	0.547	0.569	2919.87	31.25	2781.67	70.36	3016.71	27.35	2681.18	36.12
	1.013	1.831	0.06	0.228	0.004	0.474	0.546	0.018	17.115	0.565	0.619	2939.8	31.52	2805.29	75.54	3033.89	28.83	2698.83	35.67
	1.014	1.842	0.06	0.232	0.004	0.284	0.542	0.017	17.32	0.586	0.631	2951.04	32.87	2789.62	72.51	3063.34	28	2665.81	36.51
	1.015	1.923	0.068	0.201	0.004	0.368	0.52	0.018	14.414	0.513	0.679	2775.25	33.54	2696.68	76.94	2833.97	30.75	2645.3	37.03
	1.016	1.948	0.067	0.199	0.004	0.202	0.513	0.017	14.054	0.512	0.687	2751.08	34.5	2666.96	73.33	2814.02	30.73	2614.09	37.85
	1.017	1.828	0.066	0.216	0.004	0.042	0.546	0.02	16.285	0.676	0.804	2891.62	39.44	2807.74	84.12	2951.07	31.36	2748.05	34.47
	1.018	1.936	0.063	0.2	0.003	0.319	0.515	0.017	14.212	0.475	0.639	2762.33	31.93	2678.69	70.76	2824.45	27.81	2628.8	28.1
	1.019	1.859	0.064	0.227	0.004	0.243	0.537	0.018	16.835	0.598	0.751	2923.41	34.05	2771.15	76.84	3030.88	26.34	2655.92	36.76
	1.02	1.808	0.06	0.231	0.004	0.135	0.552	0.018	17.611	0.644	0.696	2966.29	34.91	2832.64	75.61	3058.72	29.38	2727.43	34.97

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Sample	Spot name	238/206 $\pm 2\sigma$	207/206 $\pm 2\sigma$	Err. Corr	207/235 $\pm 2\sigma$	206/238 $\pm 2\sigma$	Err. Corr	207/235 $\pm 2\sigma$	206/238 $\pm 2\sigma$	Err. Corr	207/235 $\pm 2\sigma$	206/238 $\pm 2\sigma$	207/206 $\pm 2\sigma$	207/206* $\pm 2\sigma$				
1.021	1.851	0.061	0.225	0.004	0.292	0.54	0.019	16.705	0.604	0.693	2915.8	34.34	2780.9	77.3	3011.12	30.43	2676.21	36.24
1.022	1.828	0.062	0.216	0.004	0.431	0.546	0.018	16.277	0.556	0.575	2891.46	32.46	2808.17	76.45	2950.67	31.86	2748.05	34.47
1.023	1.834	0.065	0.217	0.004	0.319	0.545	0.019	16.284	0.577	0.665	2891.53	34.33	2801.59	78.69	2955.71	30.43	2733.97	34.81
1.024	1.848	0.059	0.216	0.003	0.456	0.54	0.017	16.125	0.512	0.572	2883.06	30.62	2782.19	72.41	2947.35	22.79	2713.71	26.48
1.025	1.86	0.062	0.217	0.004	0.33	0.537	0.018	16.006	0.544	0.644	2875.42	33.27	2768.09	74.3	2952.03	29.1	2690.05	35.89
1.026	1.878	0.065	0.204	0.004	0.363	0.532	0.019	14.942	0.53	0.721	2809.53	33.43	2749.15	78.08	2854.21	28.94	2706.61	35.48
1.027	1.832	0.076	0.214	0.007	0.268	0.546	0.022	16.075	0.745	0.595	2877.74	43.54	2806.96	93.09	2928.88	50.03	2748.44	60.3
1.028	1.854	0.065	0.215	0.004	0.267	0.539	0.018	15.942	0.578	0.606	2871.02	35.02	2776.5	76.2	2938.64	33.57	2707.22	35.47
1.029	1.833	0.06	0.214	0.005	0.529	0.545	0.018	16.026	0.531	0.385	2876.81	31.98	2801.37	75.16	2930.48	36.39	2746.7	43.12
1.03	1.86	0.061	0.22	0.004	0.277	0.537	0.018	16.264	0.572	0.561	2890.41	33.8	2768.57	73.53	2976.81	32.63	2679.32	36.16
1.031	1.839	0.056	0.22	0.006	0.478	0.542	0.017	16.412	0.568	0.17	2899.29	33.11	2791.76	69.08	2974.43	40.7	2714.44	52.93
1.032	1.843	0.074	0.205	0.004	0.212	0.543	0.022	15.349	0.64	0.76	2834.93	41.05	2792.28	93.02	2866.55	35.52	2762.23	34.13
1.033	1.846	0.055	0.218	0.005	0.234	0.54	0.016	16.12	0.575	0.458	2881.85	33.87	2782.6	67.13	2959.52	35.19	2709.85	44.25
1.034	1.853	0.061	0.217	0.004	0.503	0.539	0.018	16.135	0.516	0.599	2883.59	30.22	2777.92	73.63	2958.82	27.12	2701.69	35.6
1.035	1.851	0.06	0.217	0.003	0.602	0.539	0.017	16.149	0.49	0.509	2884.82	29.07	2778.85	71.63	2959.99	26.74	2705.04	26.64
1.036	1.903	0.066	0.22	0.004	-0.023	0.525	0.018	15.969	0.643	0.768	2871.61	38.05	2719.67	76.93	2980.82	31.28	2610.89	37.93
1.037	1.892	0.063	0.215	0.004	0.417	0.528	0.018	15.656	0.53	0.564	2854.34	32.13	2731.3	74.02	2943.01	31.19	2645.41	37.03
1.038	1.882	0.064	0.218	0.005	0.569	0.531	0.017	15.898	0.533	0.26	2869.13	32.02	2742.92	73.82	2959.29	40.12	2650.78	46.11
1.039	1.86	0.068	0.223	0.005	0.388	0.538	0.02	16.531	0.609	0.63	2905.52	35.26	2773.15	82.1	2999.98	35.06	2668.59	45.54
1.04	1.809	0.066	0.219	0.004	0.41	0.553	0.02	16.79	0.629	0.623	2920.86	36.46	2834.56	84.53	2981.79	37.74	2770.53	33.93
1.041	1.78	0.059	0.232	0.007	0.253	0.561	0.019	17.946	0.729	0.456	2983.16	39.49	2870.53	77.81	3060.1	46.31	2773.91	59.23
1.042	1.818	0.06	0.222	0.004	0.549	0.55	0.018	16.768	0.545	0.473	2920.31	31.14	2821.84	75.74	2989.43	32.9	2743.38	34.58
1.043	1.737	0.071	0.272	0.014	-0.295	0.577	0.025	21.681	1.567	0.707	3157.71	71.25	2932.92	99.79	3303.91	80.29	2697.24	125
1.044	1.825	0.066	0.221	0.005	0.529	0.548	0.02	16.685	0.579	0.636	2915.06	33.11	2816.35	82.74	2985.33	33.26	2734.85	43.48
1.045	1.831	0.064	0.223	0.005	0.152	0.546	0.019	16.794	0.66	0.663	2920.01	38.08	2807.55	79.78	2999.31	35.63	2717.14	44.02
1.046	1.844	0.059	0.219	0.003	0.228	0.542	0.018	16.237	0.562	0.695	2889.01	33.29	2788.7	74.06	2968	22.93	2709.6	26.55
1.047	1.888	0.062	0.219	0.005	0.481	0.529	0.017	15.941	0.538	0.402	2871.64	32.42	2736.73	73.45	2968.08	36.32	2637.75	46.54
1.048	1.845	0.064	0.215	0.004	0.561	0.542	0.019	16.065	0.54	0.647	2879.06	31.69	2791.13	79.9	2942.39	31.16	2722.42	35.09
1.049	1.844	0.063	0.223	0.004	0.39	0.542	0.018	16.668	0.568	0.566	2914.13	32.96	2790.32	74.96	3001.37	31.58	2695.1	35.77
1.05	1.8	0.068	0.223	0.005	0.13	0.555	0.02	17.104	0.719	0.702	2938.08	40.91	2845.7	85.05	3002.55	38.91	2771.7	42.37
1.051	1.831	0.068	0.221	0.005	0.061	0.547	0.02	16.655	0.707	0.717	2911.23	40.11	2809.26	81.8	2983.52	37.3	2724.48	43.8
1.052	1.89	0.063	0.217	0.004	0.533	0.529	0.017	15.818	0.51	0.499	2864.6	30.96	2734.13	73.63	2958.26	31.26	2641.59	37.13
1.053	1.828	0.065	0.219	0.006	0.416	0.547	0.02	16.467	0.64	0.439	2902.12	37.84	2810.39	83.11	2966.86	46.79	2737.01	52.1
1.054	1.85	0.059	0.215	0.004	0.327	0.54	0.017	15.999	0.533	0.637	2875.19	31.94	2781.47	72.17	2941.83	27.59	2713.95	35.3
1.055	1.852	0.06	0.214	0.004	-0.045	0.54	0.018	15.872	0.574	0.65	2866.83	34.44	2780.27	73.45	2943.3	37.42	2714.19	35.29
1.056	1.898	0.058	0.204	0.004	-0.188	0.524	0.018	14.596	0.575	0.77	2793.94	43.2	2712.73	76.37	2853.68	35.19	2674.21	36.29
1.057	1.866	0.068	0.212	0.005	0.191	0.536	0.019	15.501	0.581	0.61	2843.9	35.44	2766.62	80.83	2914.16	40.46	2698	44.62
1.058	1.929	0.062	0.206	0.004	0.212	0.518	0.017	14.664	0.522	0.669	2791.63	33.28	2688.82	71.82	2867.21	29.54	2619.1	37.72
1.059	1.897	0.062	0.211	0.004	0.647	0.526	0.017	15.263	0.47	0.376	2830.84	29.09	2725.48	72.35	2907.25	32.04	2651.45	36.87
1.06	1.776	0.06	0.234	0.006	0.599	0.562	0.02	18.142	0.629	0.31	2995.96	33.4	2875.24	80.53	3078.29	43.13	2773.6	50.78
1.061	1.869	0.063	0.204	0.004	0.112	0.535	0.018	15.05	0.57	0.707	2815.63	35.95	2760.29	77.01	2856.29	31.2	2721.53	35.11
1.062	1.87	0.059	0.209	0.003	0.184	0.534	0.017	15.411	0.513	0.674	2839.53	31.07	2756.54	69.73	2899.12	24.23	2702.07	26.69
1.063	1.743	0.054	0.28	0.005	0.063	0.57	0.02	21.97	0.856	0.748	3179.14	38.37	2906.45	80.91	3365.64	25.91	2654.98	45.98
1.064	1.885	0.065	0.205	0.005	0.11	0.53	0.018	14.956	0.611	0.563	2809.49	38.81	2740.13	74.71	2859.67	41.82	2691.64	44.82
2.001	1.826	0.085	0.238	0.01	0.115	0.549	0.028	18.003	1.119	0.688	2984.79	59.86	2816.9	115.28	3101.39	67.02	2670.62	90.96
2.002	1.82	0.064	0.219	0.004	0.43	0.548	0.019	16.517	0.557	0.718	2906.11	32.44	2817.77	78.58	2968.36	26.49	2750.99	34.39
2.003	1.868	0.065	0.209	0.004	0.14	0.535	0.018	15.402	0.571	0.684	2837.92	35.96	2761.72	76.5	2893.19	31.15	2705.38	35.51
2.006	1.655	0.062	0.273	0.007	0.502	0.603	0.023	22.684	0.835	0.543	3212.47	35.93	3040.32	91.32	3322.04	39.38	2854.9	55.96
2.007	1.801	0.076	0.262	0.007	-0.271	0.552	0.021	19.848	0.981	0.858	3077.91	45.59	2831.5	87.18	3252.85	43.47	2623.71	65.79
2.008	1.727	0.088	0.281	0.009	0.126	0.584	0.033	22.561	1.332	0.819	3200.97	57.71	2957.24	130.85	3361.96	49.19	2679.89	81.34
2.009	1.74	0.049	0.262	0.005	-0.07	0.57	0.019	20.555	0.823	0.772	3115.28	39.1	2904.94	80.46	3254.22	30.29	2731.15	43.6
2.01	1.702	0.078	0.267	0.012	-0.526	0.589	0.027	21.739	1.739	0.857	3160.24	74.53	2981.29	109.47	3276.83	70.61	2783.7	100.85
2.011	1.724	0.057	0.267	0.004	0.081	0.579	0.019	21.261	0.771	0.759	3148.17	35.94	2945	78.49	3288.88	22.07	2741.25	34.63
2.012	1.828	0.067	0.223	0.005	0.332	0.547	0.019	16.787	0.631	0.596	2920.09	36	2811.14	81.58	2997.17	36.73	2722.3	43.87
2.013	1.831	0.079	0.204	0.006	0.256	0.545	0.024	15.343	0.714	0.756	2835.29	44.82	2802.98	100.17	2858.77	43.74	2787.01	50.31
2.014	1.845	0.087	0.22	0.007	0.242	0.541	0.026	16.424	0.85	0.706	2900	49.8	2786.63	107.75	2980.23	52.8	2704.29	62.19

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Sample	Spot name	238/206 $\pm 2\sigma$	207/206 $\pm 2\sigma$	Err. Corr	207/235 $\pm 2\sigma$	206/238 $\pm 2\sigma$	Err. Corr	207/235 $\pm 2\sigma$	206/238 $\pm 2\sigma$	207/206 $\pm 2\sigma$	207/206* $\pm 2\sigma$							
2.015	1.981	0.085	0.207	0.005	0.226	0.505	0.021	14.404	0.653	0.732	2774.08	43.31	2632.48	91.77	2880.03	41.75	2539.75	49.82
2.016	1.742	0.067	0.286	0.011	-0.402	0.57	0.02	22.748	1.419	0.766	3204.59	59.54	2906.11	80.38	3385.42	62.03	2633.26	102.7
2.017	1.785	0.057	0.239	0.007	-0.011	0.559	0.018	18.446	0.796	0.554	3009	40.42	2861.33	74.74	3109.03	43.68	2738.09	60.74
2.018	1.735	0.066	0.255	0.009	-0.403	0.577	0.021	20.347	1.181	0.792	3098.56	54.29	2932.43	87.38	3208.6	52.24	2768.24	76.46
2.019	1.864	0.067	0.217	0.004	0.231	0.536	0.02	15.986	0.624	0.734	2874.02	37.88	2763.53	84.16	2953.15	33.15	2683.46	36.06
2.02	1.854	0.081	0.245	0.007	-0.183	0.54	0.023	18.248	1.021	0.791	2996.5	52.15	2781.58	95.17	3145.75	45.84	2599.52	66.91
3.001	1.807	0.061	0.233	0.005	0.43	0.553	0.019	17.704	0.62	0.538	2971.82	33.44	2835.15	78.21	3066.39	34.54	2721.71	43.88
3.002	1.884	0.064	0.225	0.007	-0.304	0.53	0.018	16.407	0.794	0.699	2895.16	44.34	2740.33	76.36	3024.18	53.74	2623.04	65.82
3.003	1.869	0.076	0.216	0.004	0.308	0.531	0.019	15.937	0.613	0.731	2870.08	36.83	2761.05	91.97	2949.48	32.31	2678.83	36.18
3.004	1.891	0.065	0.228	0.004	0.362	0.528	0.018	16.555	0.569	0.616	2907.76	32.9	2729.77	74.46	3034	30.05	2601.63	38.18
3.005	1.816	0.069	0.238	0.008	-0.424	0.546	0.018	17.917	0.852	0.732	2979.9	42.65	2807.55	73.31	3097.95	48.71	2687.56	71.91
3.006	1.928	0.067	0.208	0.003	0.461	0.517	0.017	14.823	0.482	0.726	2802.96	31.23	2684.66	73.78	2889.69	24.31	2613.83	28.39
3.007	1.916	0.062	0.206	0.004	-0.075	0.52	0.017	14.804	0.568	0.716	2799.68	37.69	2699.68	72.25	2872.81	32.71	2639.02	37.2
3.008	1.852	0.057	0.223	0.005	0.414	0.538	0.017	16.495	0.562	0.323	2904.21	32.61	2773.87	70.82	2995.82	36.74	2681.76	45.13
3.009	1.847	0.059	0.228	0.004	0.085	0.54	0.017	16.941	0.609	0.63	2929.15	35.2	2780.74	71.52	3032.89	30.99	2672.01	36.35
3.01	1.948	0.06	0.188	0.003	0.515	0.511	0.016	13.268	0.404	0.635	2698.04	29.2	2662.08	68.67	2725.41	23.64	2650.95	27.67
3.011	1.898	0.067	0.201	0.004	0.671	0.526	0.019	14.421	0.43	0.521	2777.11	28.3	2722.95	78.17	2827.24	34.12	2684.64	36.03
3.012	1.895	0.07	0.2	0.004	0.564	0.527	0.019	14.484	0.495	0.705	2780.2	32.05	2728.5	81.53	2819.61	30.42	2692.95	35.82
3.013	1.916	0.066	0.201	0.004	0.527	0.521	0.018	14.482	0.484	0.511	2780.25	32.05	2701.53	76.99	2838.87	34.66	2656.17	36.75
3.014	1.827	0.064	0.229	0.005	0.492	0.546	0.019	17.241	0.6	0.506	2946.44	33.4	2807.55	79.4	3043.61	36.44	2701.97	44.49
3.015	1.932	0.065	0.209	0.004	0.332	0.516	0.018	14.878	0.517	0.687	2805.97	33.01	2680.32	74.46	2898.11	28.52	2604.4	38.1
3.016	1.888	0.062	0.22	0.005	0.38	0.528	0.017	15.989	0.557	0.535	2874.28	32.78	2731.32	72.72	2976.41	33.34	2634.26	46.65
3.017	1.855	0.063	0.217	0.005	0.313	0.538	0.018	16.057	0.581	0.585	2877.96	34.62	2771.71	76.13	2953.79	34.51	2698.36	44.61
3.018	1.879	0.067	0.218	0.003	0.527	0.532	0.019	15.93	0.539	0.784	2870.91	32.11	2745.91	80.41	2961.39	25.3	2655.57	27.58
3.019	1.895	0.07	0.223	0.006	0.272	0.527	0.02	16.172	0.686	0.587	2884.02	40.42	2725.43	83.45	2997.5	45.15	2612.85	56.82
3.02	1.879	0.059	0.224	0.003	0.323	0.53	0.017	16.277	0.553	0.752	2891.44	32.61	2741.3	70.33	3008.13	24.02	2634.44	27.99
3.021	1.913	0.064	0.22	0.004	-0.197	0.521	0.018	15.791	0.653	0.789	2861.35	39.31	2702.13	74.28	2975.68	32.31	2595.6	38.34
3.022	1.895	0.06	0.226	0.005	0.345	0.526	0.016	16.342	0.575	0.377	2894.94	33.59	2721.75	69.19	3017.6	37.19	2602.42	47.7
3.023	1.861	0.063	0.215	0.006	0.448	0.536	0.018	15.857	0.566	0.354	2866.11	33.81	2764.92	75.33	2938.41	40.5	2695.56	53.63
3.024	1.912	0.058	0.197	0.003	-0.391	0.521	0.016	14.451	0.833	0.958	2772.87	43.8	2701.26	67.21	2795.38	27.65	2676.2	27.18
3.025	1.92	0.063	0.206	0.004	0.356	0.519	0.017	14.702	0.505	0.553	2794.35	32.4	2694.4	71.25	2867.66	32.04	2632.85	37.36
4.001	1.966	0.069	0.206	0.006	0.581	0.506	0.018	14.365	0.525	0.231	2773.24	34.64	2639.78	76.2	2871.64	49.54	2564.4	58.77
4.002	1.796	0.062	0.252	0.005	0.465	0.556	0.019	19.243	0.648	0.611	3052.43	32.79	2846.88	79.08	3191.59	30.18	2669.64	45.51
4.003	1.731	0.066	0.274	0.006	0.673	0.578	0.022	21.77	0.738	0.587	3171.81	32.52	2936.92	90.15	3325.87	36.39	2700.33	53.45
4.004	1.801	0.055	0.262	0.007	0.454	0.55	0.019	19.802	0.745	0.454	3079.49	36.19	2822.44	81.49	3252.44	42.18	2623.71	65.79
4.005	1.744	0.06	0.249	0.006	0.56	0.572	0.019	19.627	0.67	0.407	3071.31	33.13	2915.16	79.91	3175.77	38.18	2774.94	50.73
4.006	1.752	0.063	0.249	0.007	0.149	0.57	0.021	19.558	0.887	0.596	3064.96	42.32	2906.25	84.29	3171.44	46.36	2759.9	59.82
4.007	1.759	0.066	0.241	0.007	0.361	0.568	0.021	18.868	0.778	0.586	3031.27	39.63	2898.2	87.76	3122.11	44.62	2778.08	59.06
4.008	1.932	0.067	0.21	0.006	-0.249	0.513	0.02	14.85	0.791	0.817	2799.33	50.05	2665.48	86.21	2898.52	44.04	2601.01	57.29
4.009	1.856	0.085	0.231	0.007	0.367	0.532	0.019	17.111	0.76	0.611	2938.26	41.43	2775.41	107.13	3052.99	48.06	2646.47	64.75
4.01	1.869	0.065	0.22	0.004	0.393	0.534	0.018	16.165	0.549	0.668	2884.9	32.74	2757.74	76.23	2975.76	28.33	2664.62	36.54
4.011	1.9	0.065	0.199	0.003	0.143	0.526	0.019	14.404	0.536	0.805	2774.25	35.06	2722.8	78.31	2813.07	25.67	2688.39	26.95
4.012	1.885	0.065	0.211	0.004	0.215	0.53	0.018	15.373	0.549	0.732	2836.5	33.79	2738.41	75.31	2907.77	27.09	2670.56	36.39
4.013	1.876	0.071	0.23	0.006	0.492	0.532	0.021	16.883	0.625	0.585	2926.73	35.61	2750.15	86.23	3051.36	39.21	2618.04	56.62
4.014	1.803	0.067	0.218	0.005	0.352	0.554	0.02	16.611	0.625	0.634	2910.36	35.81	2841.47	83.63	2959.42	35.38	2785.12	41.98
4.015	1.888	0.064	0.227	0.005	0.51	0.529	0.017	16.542	0.566	0.368	2906.9	32.41	2735.85	73.84	3028.03	38.07	2609.78	47.45
4.016	1.94	0.067	0.202	0.005	0.016	0.514	0.018	14.291	0.615	0.665	2767.65	40.72	2672.03	76.91	2837.93	41.21	2616.01	47.25
4.017	1.89	0.062	0.197	0.004	0.305	0.528	0.017	14.308	0.495	0.568	2768.53	32.89	2733.05	72.46	2794.82	31.93	2711.57	35.36
4.018	1.948	0.079	0.186	0.003	0.108	0.513	0.021	13.171	0.576	0.871	2689.76	40.45	2668.67	89.35	2706.79	28.16	2657.66	27.54
4.019	1.876	0.071	0.201	0.003	0.362	0.534	0.02	14.688	0.509	0.839	2793.4	32.13	2754.73	84.2	2835.36	25.19	2720.54	26.35
4.02	1.796	0.076	0.256	0.008	0.061	0.551	0.018	19.651	0.917	0.612	3070.59	44.77	2828.56	74.56	3218.4	47.62	2654.62	73.59
4.021	1.965	0.069	0.187	0.003	0.462	0.509	0.018	13.124	0.442	0.709	2687.04	31.85	2649.9	75.95	2716.26	27.71	2628.71	28.1
4.022	1.767	0.059	0.256	0.006	-0.046	0.566	0.019	19.757	0.766	0.679	3076.51	37.03	2888.59	78.35	3216.65	39.16	2705.31	53.27
4.023	1.783	0.083	0.227	0.007	0.182	0.56	0.026	17.504	0.917	0.693	2960.75	50.73	2866.83	106.98	3025.73	54	2787.55	58.67
4.024	1.883	0.061	0.214	0.005	0.24	0.53	0.018	15.601	0.599	0.463	2851.1	36.43	2740.32	73.58	2930.13	41.01	2663.24	45.72
4.025	1.703	0.072	0.292	0.012	-0.228	0.59	0.025	23.805	1.479	0.721	3247.38	65.15	2985.5	100.79	3433.41	59.02	2680.04	108.44

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Sample	Spot name	238/206 $\pm 2\sigma$	207/206 $\pm 2\sigma$	Err. Corr	207/235 $\pm 2\sigma$	206/238 $\pm 2\sigma$	Err. Corr	207/235 $\pm 2\sigma$	206/238 $\pm 2\sigma$	Err. Corr	207/235 Age $\pm 2\sigma$	206/238 Age $\pm 2\sigma$	207/206 Age $\pm 2\sigma$	207/206* Age $\pm 2\sigma$					
	5.001	1.849	0.058	0.217	0.005	0.439	0.54	0.017	16.225	0.57	0.308	2888.22	32.8	2783.09	71.52	2962.3	39.18	2708.41	44.29
	5.002	1.839	0.082	0.217	0.013	0.513	0.543	0.024	16.201	0.94	0.098	2886.85	56.32	2793.4	99.14	2952.21	96.54	2725.38	113.81
	5.003	1.824	0.06	0.237	0.007	-0.041	0.548	0.018	17.862	0.797	0.543	2977.46	42.86	2814.45	74.08	3089.16	47.94	2677.66	63.36
	5.004	1.754	0.075	0.264	0.006	0.325	0.57	0.024	20.749	0.908	0.739	3125.42	41.31	2905.76	99.38	3270.33	37.85	2697.71	53.55
	5.005	1.871	0.063	0.215	0.004	0.392	0.534	0.018	15.823	0.53	0.691	2864.54	31.84	2759.02	74.73	2940.37	26.49	2679.12	36.17
	5.006	1.865	0.059	0.214	0.004	0.282	0.536	0.017	15.753	0.539	0.504	2860.15	32.83	2763.81	69.95	2928.7	32.15	2692.52	35.83
	5.007	1.862	0.063	0.214	0.004	0.381	0.537	0.018	15.813	0.541	0.646	2863.83	32.7	2770.25	75.91	2931.08	29.47	2697.48	35.71
	5.008	1.837	0.065	0.218	0.004	0.174	0.545	0.019	16.347	0.613	0.72	2894.66	36.46	2802.1	79.44	2960.67	30.77	2725.15	35.02
	5.009	1.883	0.065	0.213	0.004	0.127	0.532	0.018	15.59	0.583	0.628	2856.84	40.43	2746.4	76.43	2936.18	38.44	2666.75	36.48
	5.01	1.946	0.1	0.201	0.008	-0.68	0.512	0.025	13.968	0.866	0.906	2738.75	53.16	2660.23	102	2851.15	86.17	2610.38	75.89
	5.011	1.821	0.059	0.221	0.005	0.535	0.549	0.018	16.732	0.548	0.469	2918.19	31.5	2820.94	74.97	2986.41	33.32	2741.83	43.27
	5.012	1.888	0.067	0.217	0.004	0.449	0.531	0.019	15.816	0.55	0.657	2863.83	32.92	2742.21	78.19	2951.67	30.9	2644.75	37.05
	5.013	1.851	0.062	0.218	0.004	0.303	0.541	0.018	16.217	0.566	0.716	2887.71	33.91	2785.02	76.54	2960.9	28.28	2701.44	35.61
	5.014	1.911	0.072	0.221	0.005	0.342	0.525	0.02	15.949	0.618	0.678	2870.78	36.42	2718.27	83.47	2981.38	35.1	2595.21	47.94
	5.015	1.86	0.061	0.22	0.004	0.3	0.539	0.018	16.342	0.569	0.574	2895.03	34.04	2776.26	73.95	2979.1	32.34	2679.32	36.16
	5.016	1.928	0.063	0.198	0.003	0.236	0.52	0.017	14.18	0.474	0.74	2760.31	31.71	2696.23	71.12	2807.97	23.38	2647.81	27.73
	5.017	1.917	0.069	0.195	0.003	0.522	0.52	0.016	14.058	0.442	0.74	2752.58	29.7	2697.13	69.02	2785.21	22.72	2675.19	27.2
	5.018	1.919	0.07	0.201	0.004	-0.296	0.523	0.019	14.641	0.708	0.84	2786.52	43.86	2711.44	81.76	2832.4	29.53	2651.5	36.87
	5.019	1.966	0.065	0.196	0.003	0.33	0.51	0.017	13.8	0.469	0.718	2734.43	32.09	2654.78	72.64	2794.52	26.23	2597.46	28.72
	5.02	1.798	0.061	0.228	0.005	0.37	0.557	0.019	17.514	0.637	0.453	2961.11	34.63	2854.67	77.42	3034.68	38.45	2756.43	42.83
	5.021	1.881	0.059	0.212	0.004	0.342	0.533	0.017	15.526	0.515	0.564	2846.53	31.62	2751.21	70.44	2914.91	28.97	2673.49	36.31
	5.022	1.897	0.063	0.209	0.004	-0.215	0.528	0.017	15.258	0.599	0.822	2829.15	37.71	2732.23	73.77	2899.12	28.2	2658.41	36.7
	5.023	1.831	0.052	0.236	0.005	-0.132	0.544	0.018	17.685	0.692	0.744	2969.6	38.19	2799.81	75.15	3087.12	30.58	2669.56	45.51
	5.024	1.927	0.066	0.204	0.005	0.243	0.517	0.016	14.498	0.488	0.451	2787.27	37.19	2697.08	77.4	2853.14	36.64	2628.94	46.82
	5.025	1.861	0.071	0.211	0.003	0.695	0.54	0.021	15.657	0.518	0.807	2854.6	31.51	2782.82	86.74	2907.98	26.86	2709.88	26.55
ZLC32	1.001	1.969	0.098	0.21	0.01	0.699	0.508	0.026	14.658	0.664	0.304	2792.08	43.02	2648.71	109.82	2897.98	79.07	2546.89	99.15
	1.002	1.48	0.066	0.409	0.039	-0.164	0.676	0.029	38.234	4.166	0.478	3712.46	108.35	3328.44	113.51	3925.34	140.9	2607.98	370.59
	1.003	0.757	0.073	0.716	0.06	0.384	1.431	0.163	136.974	14.341	0.66	4960.64	107.96	5626.31	400.45	4598.3	117.75	2906.58	462.49
	1.005	0.662	0.089	0.835	0.085	0.077	1.751	0.284	184.559	26.813	0.754	5273.7	169.75	6311.35	603.43	4649.97	132.02	39.27	4343.14
	1.006	0.701	0.114	0.775	0.063	0.41	1.704	0.277	177.374	27.573	0.797	5163.73	161.1	6143.48	659.57	4626.41	125.81	2084.45	858.77
	1.007	0.786	0.08	0.726	0.046	0.274	1.342	0.149	133.058	15.146	0.902	4955.9	119.2	5399.41	377.57	4685.73	78.35	2458.4	484.97
	1.009	0.808	0.083	0.664	0.034	-0.386	1.305	0.137	117.502	14.244	0.888	4848.43	142.28	5324.29	377.65	4643.23	73.67	3403.33	184.1
	1.01	1.555	0.107	0.408	0.028	-0.043	0.651	0.042	36.703	3.715	0.658	3667.67	93.96	3225.97	164.21	3922.87	102.47	2458.38	295.2
	1.011	0.995	0.095	0.598	0.046	-0.666	1.039	0.1	87.366	13.433	0.94	4500.01	159.17	4562.2	313.49	4482.23	114.52	2879.86	361.31
	1.012	1.849	0.069	0.231	0.007	0.144	0.543	0.021	17.257	0.802	0.65	2944.12	43.2	2791.62	86.96	3051.46	46.34	2657.89	64.24
	1.013	1.912	0.082	0.207	0.01	-0.397	0.525	0.023	15.047	1.124	0.793	2808.85	68.59	2716.09	98.06	2876.38	73.09	2641.79	92.81
	1.014	1.868	0.081	0.22	0.012	-0.216	0.536	0.024	16.254	1.281	0.687	2885.16	74.24	2763.68	100.03	2970.59	91.1	2666.24	109.49
	1.015	1.915	0.087	0.215	0.011	0.359	0.524	0.023	15.204	0.703	0.287	2839.87	56.02	2713.62	97.13	2931.47	81.05	2609.71	104.4
	2.001	1.439	0.107	0.345	0.026	0.325	0.69	0.043	33.439	3.107	0.504	3609.84	104.15	3443.91	222.28	3709.54	144.16	3045.88	181.63
	2.005	1.473	0.1	0.387	0.023	0.359	0.697	0.054	36.86	2.821	0.708	3673.04	77.92	3391.42	196.7	3840.69	88.32	2734.33	200.09
	2.006	1.584	0.076	0.303	0.024	0.507	0.631	0.031	26.254	1.888	-	3353.05	67.59	3152.09	120.7	3475.1	115.21	2878.93	188.64
	2.008	1.873	0.065	0.233	0.007	-0.29	0.535	0.019	17.19	0.848	0.743	2939.16	47.41	2758.99	78.92	3065.24	45.96	2612.18	66.32
	2.009	1.978	0.067	0.196	0.004	0.325	0.506	0.017	13.664	0.486	0.606	2724.64	33.87	2636.81	73.73	2791.13	33.12	2580.05	38.76
	2.01	1.921	0.069	0.206	0.004	-0.082	0.522	0.02	14.96	0.749	0.713	2806.61	45.38	2704.16	83.33	2866.63	33.66	2631.32	37.4
	2.011	1.912	0.075	0.216	0.008	0.089	0.52	0.018	15.56	0.765	0.455	2844.02	46.41	2697.93	75.79	2937.86	58.2	2610.86	75.87
	2.012	1.922	0.058	0.193	0.004	0.222	0.519	0.016	13.843	0.483	0.49	2737.12	32.89	2695.89	66.99	2767.39	33.77	2674.18	36.29
	2.015	1.282	0.079	0.484	0.025	0.083	0.789	0.052	52.565	4.093	0.713	4030.01	78.97	3739.39	184.86	4182.75	77.12	2690.16	224.32
	3.001	1.903	0.123	0.237	0.02	-0.529	0.528	0.039	17.366	2.482	0.867	2937.06	127.51	2730.41	161.54	3082.89	134.85	2552.13	197.58
	3.002	1.653	0.098	0.293	0.02	-0.432	0.607	0.04	24.64	2.825	0.843	3281.73	106.89	3055.27	157.28	3424.37	107.53	2773.42	169.29
	3.003	1.94	0.066	0.198	0.005	0.524	0.516	0.018	14.04	0.49	0.442	2750.48	33.35	2678.52	76.49	2804.52	40.34	2629.48	46.8
	3.004	1.735	0.078	0.251	0.009	-0.41	0.577	0.026	20.53	1.716	0.719	3104.38	77.42	2935.65	104.54	3215.47	84.05	2784.15	75.61
	3.005	1.68	0.08	0.296	0.018	-0.433	0.601	0.03	24.779	2.332	0.781	3270.61	88.31	3028	119.18	3424.49	93.65	2707.24	159.59
	3.006	1.512	0.122	0.371	0.033	-0.732	0.67	0.057	34.784	5.632	0.925	3600.11	155.29	3297.02	216.86	3777.43	133.75	2722.81	289.41
	3.007	0.68	0.062	0.715	0.04	-0.251	1.512	0.123	149.578	14.776	0.835	5066.43	112.24	5903.51	324.5	4709.28	56.91	5830.5	35.87

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Sample	Spot name	238/206 $\pm 2\sigma$	207/206 $\pm 2\sigma$	Err. Corr	207/235 $\pm 2\sigma$	206/238 $\pm 2\sigma$	Err. Corr	207/235 $\pm 2\sigma$	206/238 $\pm 2\sigma$	207/206 $\pm 2\sigma$	207/206* $\pm 2\sigma$							
3.008	1.782	0.086	0.275	0.019	-0.16	0.562	0.027	21.342	1.939	0.569	3143.12	82.03	2873.82	112.51	3320.06	100.66	2606.71	180.7
3.009	0.875	0.095	0.621	0.043	-0.661	1.193	0.139	104.025	16.95	0.953	4673.39	167.49	5015.52	392.38	4541.08	105.59	3459.72	223.6
3.012	0.775	0.07	0.685	0.089	0.587	1.322	0.131	120.981	14.628	0.271	4824.42	113.66	5453.57	399.08	4396.44	177.18	3371.12	493.17
3.014	1.102	0.096	0.54	0.046	0.464	0.952	0.08	69.175	6.18	0.49	4288.29	88.02	4273.3	262.7	4287.97	116.19	2916.89	352.01
3.015	1.417	0.122	0.459	0.033	0.274	0.739	0.06	46.114	4.431	0.594	3878.39	96.54	3539.09	223.09	4074.95	117.04	2490.78	340.19
4.001	1.854	0.064	0.215	0.007	-0.083	0.539	0.018	15.952	0.925	0.7	2864.02	55.36	2778.08	74.49	2946.07	57.44	2707.22	62.06
4.002	1.828	0.07	0.23	0.006	0.051	0.544	0.018	17.134	0.678	0.65	2939.26	37.29	2812.66	87.83	3044.37	43.97	2696.59	53.59
4.003	1.844	0.073	0.217	0.009	0.36	0.542	0.02	16.213	0.78	0.322	2886.06	44.9	2790.69	84.79	2953.27	63.81	2716.86	79.26
4.004	1.962	0.067	0.194	0.005	0.016	0.51	0.018	13.608	0.589	0.444	2718.84	39.17	2654.15	74.85	2767.85	41.49	2609.96	47.45
4.005	2.016	0.069	0.187	0.003	0.205	0.496	0.016	12.773	0.456	0.696	2660.98	33.62	2594.25	70.83	2712.65	29.4	2555.5	29.57
4.006	2.014	0.068	0.192	0.005	0.475	0.496	0.017	13.077	0.48	0.285	2683	33.8	2596.56	71.48	2748.95	45.57	2542.28	49.73
4.007	1.956	0.077	0.215	0.007	-0.156	0.513	0.021	15.257	0.862	0.753	2822.68	51.71	2666.68	86.96	2937.39	53.66	2548.95	69.31
4.008	1.68	0.075	0.275	0.012	-0.21	0.592	0.023	22.437	1.258	0.741	3193.26	55.11	2994.64	93.7	3339.28	76.01	2794.52	100.08
4.009	1.893	0.067	0.206	0.007	-0.142	0.529	0.019	15.055	0.813	0.671	2811.31	48.79	2734.77	80.99	2866.85	54.59	2675.23	63.47
4.01	1.93	0.078	0.207	0.008	-0.121	0.52	0.021	14.903	0.929	0.678	2798.89	54.51	2697.53	89.86	2873.87	61.83	2614.2	75.69
4.011	1.929	0.062	0.192	0.007	-0.699	0.518	0.017	13.755	0.778	0.895	2725.01	48.42	2689.62	70.96	2750.21	55.56	2666.66	63.85
4.012	1.894	0.061	0.209	0.006	0.431	0.528	0.017	15.154	0.552	0.311	2822.62	34.52	2730.48	73.14	2889.24	43.15	2663.17	54.86
4.013	1.907	0.081	0.22	0.009	-0.482	0.522	0.02	16.077	1.15	0.834	2866.03	65.07	2705.09	84.15	2967.92	66.54	2604.75	85.71
4.014	1.921	0.079	0.215	0.011	0.231	0.523	0.022	15.272	0.808	0.41	2836.44	55.38	2708.84	91.36	2929.11	76.62	2600.59	105.06
4.015	1.983	0.062	0.193	0.003	0.253	0.504	0.016	13.395	0.454	0.553	2706.27	32.11	2628.93	67.62	2764.58	30.61	2582.68	29.01
4.016	1.993	0.081	0.204	0.007	0.332	0.503	0.02	14.14	0.629	0.526	2755.84	41.95	2623.29	85.22	2855.43	52.15	2532.71	70.09
4.017	1.875	0.068	0.202	0.006	0.44	0.535	0.02	14.833	0.57	0.504	2801.72	37.09	2758.53	84.02	2834.24	44.88	2718.65	52.77
4.018	1.849	0.052	0.209	0.006	0.002	0.537	0.017	15.521	0.662	0.573	2843.75	39.67	2783.22	63.2	2895.17	42.9	2737.35	52.09
4.019	1.863	0.062	0.221	0.007	-0.104	0.537	0.018	16.402	0.805	0.634	2894.5	46	2769.52	76.36	2982.39	51.53	2670.83	63.66
4.02	1.756	0.078	0.24	0.009	0.063	0.567	0.022	18.946	1.067	0.658	3029.85	53.64	2890.82	89.3	3108.58	59.72	2787.63	75.43
5.001	1.964	0.069	0.18	0.003	0.28	0.509	0.017	12.536	0.416	0.655	2643.95	31.06	2652.53	73.01	2648.48	30.15	2653.4	27.62
5.002	1.957	0.064	0.181	0.003	0.118	0.511	0.016	12.767	0.451	0.697	2660.62	33.42	2658.37	70.32	2662.62	28.23	2660.67	27.48
5.003	1.934	0.065	0.18	0.003	0.435	0.517	0.017	12.822	0.423	0.68	2665.18	30.98	2686.04	72.7	2650.07	27.22	2700	26.73
5.004	1.876	0.063	0.184	0.004	-0.174	0.533	0.018	13.492	0.546	0.856	2711.56	36.81	2754.19	76.13	2686.79	38.08	2782.6	33.64
5.005	1.91	0.097	0.19	0.004	0.281	0.524	0.025	13.739	0.671	0.843	2729.84	46.41	2715.68	107.76	2741.75	37.52	2703.52	35.56
5.006	1.906	0.07	0.182	0.004	0.507	0.524	0.018	13.13	0.454	0.552	2688.05	32.77	2716.34	77.6	2667.07	36.12	2738.54	34.7
5.007	1.904	0.064	0.177	0.003	0.307	0.526	0.018	12.806	0.446	0.76	2663.62	32.98	2721.21	76.3	2621.12	26.55	2761.84	25.6
5.008	1.931	0.065	0.178	0.003	0.466	0.518	0.018	12.719	0.412	0.743	2658.04	30.93	2688.89	74.96	2635.27	24.22	2712.22	26.51
5.01	1.932	0.061	0.182	0.003	0.036	0.517	0.016	12.954	0.451	0.742	2674.47	32.64	2686.4	68.99	2665.59	25.46	2696.07	26.81
5.011	1.949	0.067	0.177	0.003	0.436	0.514	0.018	12.546	0.43	0.777	2644.47	31.91	2671.31	77.58	2625.39	25.59	2686.98	26.98
5.012	1.94	0.067	0.182	0.004	0.345	0.516	0.018	12.943	0.47	0.602	2673.25	33.92	2680.31	74.88	2668.6	35.46	2683.47	36.06
5.013	1.886	0.071	0.186	0.003	0.186	0.532	0.02	13.62	0.545	0.794	2720.42	37.75	2747.23	84.83	2702.36	30.85	2757.37	25.68
5.014	1.987	0.063	0.181	0.002	0.408	0.503	0.016	12.547	0.389	0.702	2649.53	25.21	2625.38	67.14	2660.73	21.92	2616.06	18.9
5.015	1.879	0.072	0.185	0.004	0.204	0.533	0.021	13.609	0.57	0.781	2720.33	39	2752.71	88.24	2697.62	34.09	2773.33	33.86
5.016	1.912	0.064	0.183	0.004	0.074	0.523	0.018	13.237	0.51	0.716	2693.94	35.69	2712.13	74.6	2680.84	32.32	2724.84	35.03
5.017	1.916	0.067	0.186	0.005	-0.027	0.523	0.018	13.401	0.59	0.699	2703.9	40.94	2709.54	77.31	2700.15	41.84	2707.69	44.32
5.018	1.978	0.081	0.177	0.005	0.309	0.506	0.02	12.337	0.541	0.594	2628.48	40.15	2638	86.37	2621.68	47.44	2642.42	46.38
5.019	1.924	0.104	0.182	0.005	0.451	0.521	0.026	13.038	0.615	0.76	2680.3	45.91	2703.18	113.37	2665.05	47.88	2708.89	44.28

C.4 Dioctahedral (White) mica Rb–Sr data

Sample	Spot name	Rb ppm	Sr	$^{87}\text{Rb}/^{87}\text{Sr}$	$\pm 2\text{SE}$ (%)	$^{87}\text{Rb}/^{86}\text{Sr}$	$\pm 2\text{SE}$ (%)	$^{87}\text{Sr}/^{86}\text{Sr}$	$\pm 2\text{SE}$ (%)	ρ
Z6007	Fs-1	15.8	35	2.08	6.2	0	6.6	0.7302	0.72	0.33
	Fs-2	24	30	2.11	4.5	0	4.3	1.232	1.5	0.036
	Fs-3	1.04	10	0.278	22	0	22	1.942	1.2	0.02
	Fs-5	1.15	4.1	0.79	23	0	23	1.746	2	0.18
	Fs-6	18.5	4.7	6.5	14	0	15	1.982	2.6	0.43
	Fs-7	0.688	7.5	0.33	14	0	13	1.659	1.3	0.028
	Fs-8	0.556	14	0.1759	4	0	4.1	1.712	0.82	0.13
	Fs-9	25.7	4	8.2	13	0	13	2.535	1.5	0.071
	Mc-1	7510	0.065	28.8	3.7	288200	9.8	12400	9.1	0.98
	Mc-2	7350	0.24	30.26	2.5	81500	10	3410	10	0.95
	Mc-3	9520	0.068	30.45	0.91	364300	7.2	15000	7.4	0.99
	Mc-4	8740	0.19	31.3	0.93	119800	5.7	4780	5.8	0.99
	Mc-5	8190	0.1	29.39	3.2	208200	6.5	8920	7.5	0.87
	Mc-6	9980	0.087	30.97	1.2	286800	4.6	11570	4.5	0.97
	Mc-7	5070	0.32	32.24	0.82	40280	4.6	1562	4.8	0.98
	Mc-8	7870	0.079	29.85	3.3	276600	29	11500	28	1
	Mc-9	9210	0.18	30.46	0.79	139700	9.9	5740	10	1
	Fs-10	75.3	18	7.23	6.8	0	10	1.795	1.7	0.38
	Fs-11	53.2	13	7.15	9.2	0	9.5	1.775	1.5	0.36
	Fs-12	0.788	8.6	0.431	4	0	4.1	1.214	0.97	0.26
	Fs-13	4.12	15	0.79	18	0	19	1.267	1.2	0.42
	Fs-15	0.252	1.3	1.677	4.5	0	3.7	1.367	2.3	-0.11
	Mc-10	10900	0.22	30.52	0.98	127700	11	5230	11	1
	Mc-11	5920	0.12	29.1	5.3	129300	13	5490	9.4	0.95
	Mc-12	8310	0.11	30.68	2.1	200100	8.6	7910	6.4	0.97
	Mc-13	8300	0.18	31.91	1.1	114000	3.8	4460	3.8	0.97
	Mc-14	5290	0.57	33.55	0.97	23600	5.2	879	4.7	0.99
	Mc-15	6370	0.13	30.7	4.5	128300	10	5240	9.6	0.9
	Mc-16	9920	0.13	30.19	1.4	192500	8.2	8000	7.2	0.98
	Mc-17	8510	0.24	30.34	2.2	90700	8.7	3730	7.7	0.97
	Mc-18	5210	0.16	29	10	81340	8.9	3550	11	0.54
	Mc-19	8120	0.23	31.14	1.9	89400	5	3640	5.5	0.86
	Mc-20	12500	0.44	31.9	0.7	71300	6.2	2800	6.2	0.99
	Mc-21	10400	0.23	30.9	0.98	112500	4.6	4560	4.7	0.98
	Mc-22	11800	0.066	31.42	0.75	444760	5.9	17650	5.5	0.99
	Mc-23	8770	13	29.33	2.4	1700	6.9	72.7	5.9	0.96
	Mc-24	10700	0.075	30.64	1.3	361900	7.7	14800	7.8	0.99
	Mc-25	10900	0.087	31.42	1.1	353800	12	14200	12	0.99
	Mc-26	11800	0.11	31.28	0.88	263552.846	8.8	10520	8.4	1
	Mc-27	12200	0.084	31.88	1	358337.67	5.7	14070	5.8	0.99
	Mc-28	6380	0.27	32.02	0.75	59198.53	3.7	2314	3.7	0.98
	Mc-29	9230	0.12	31.57	2.3	196311.13	5.8	7630	5	0.84
	Mc-30	8130	0.2	31.4	0.84	103045.9	4.7	4130	4.5	0.96
	Mc-31	6190	0.24	32.39	0.77	63520.617	3.8	2459	3.4	0.97
	Mc-32	10000	0.16	31.72	0.66	157791.661	4.9	6220	4.9	0.99
	Mc-33	12100	0.29	32.38	2.9	106582	6.5	4040	3.6	0.77
	Mc-34	9990	0.18	32.13	0.95	137299.5	4.6	5340	4.3	0.98
	Mc-35	8700	0.25	31.45	0.83	89159.3	4.1	3540	4.1	0.97
	Mc-36	10000	0.45	32.06	1	55884.973	3.5	2179	3.4	0.95
	Mc-37	5240	0.27	32.85	0.91	48363.06	4.4	1838	4	0.96
	Mc-38	8350	0.39	32.44	0.58	52314.547	3.4	2013	3.5	0.98
	Mc-39	5440	0.24	31.95	0.98	55500	4.1	2171	3.6	0.97
	Mc-40	8370	0.099	30.97	1.5	210100	8.6	8480	8.5	0.99
	Mc-41	7660	0.21	32.49	0.96	90200	6.7	3480	7.1	0.99
	Mc-42	5380	1.1	28.2	6.4	12400	7.4	556	10	0.74
	Mc-43	11100	0.35	31.1	1.2	80800	5.8	3260	6.3	0.98
	Mc-44	9980	0.23	30.55	1	109000	4.6	4460	4.7	0.97
	Mc-45	5010	0.26	31.91	1	48200	4.7	1861	4.4	0.98
	Mc-47	8330	0.5	32.7	4.3	42400	8.2	1630	8.1	0.87
	Mc-48	11700	0.49	31.13	1.2	59900	7.1	2410	7.6	0.99
	Mc-49	7880	0.37	31.66	1.1	53500	3.9	2112	3.7	0.96
	Mc-50	11100	0.66	33.85	0.85	41800	3	1543	3.2	0.72
	Mc-51	6410	0.21	32.65	1.1	77900	4.3	2980	4.4	0.97
	Mc-52	8560	0.32	30.25	1.2	65800	5.7	2650	6.6	0.69
	Mc-53	11800	0.047	30.54	1.4	634600	7.7	26000	7.5	0.98
	Mc-54	12200	0.28	31.46	1.1	109130	4.3	4320	4.5	0.95
	Mc-55	11000	1.4	31.83	0.94	19375	2.8	761	2.9	0.95
	Mc-56	11600	0.091	31.16	0.94	315128	4.6	12650	5	0.99
Mc-57	7440	0.2	31.12	0.98	91400	3.8	3700	4.5	0.95	
Mc-58	8850	0.22	31.54	0.91	98200	4.5	3890	4.2	0.98	
Mc-59	4410	0.25	32.02	0.74	43600	3.1	1703	3.2	0.98	
Mc-61	1750	0.72	25.2	14	6400	12	332	21	0.84	
Mc-62	5340	0.41	31	3.4	32600	6.7	1311	5.9	0.88	
Mc-65	10600	0.11	31.5	0.92	247800	6.4	9810	5.9	0.99	
Mc-66	8230	0.27	31.8	0.73	76000	3.6	2990	3.6	0.98	
Mc-68	7610	0.19	30	3.9	101600	7.4	4250	7.8	0.84	
Z6029	Fs-1	5930	0.35	33.54	0.8	42400	3.7	1580	3.8	0.98
	Fs-2	5620	0.34	33.37	1.1	40700	3.4	1529	3.7	0.95
	Fs-3	4810	0.36	30.8	7.2	33200	7.2	1352	6.9	0.52
	Mc-1	7020	0.14	31.7	1.1	129700	5.6	5120	6.1	0.99

Continued on next page

Sample	Spot name	Rb ppm	Sr	$^{87}\text{Rb}/^{87}\text{Sr}$	$\pm 2\text{SE}$	$^{87}\text{Rb}/^{86}\text{Sr}$	$\pm 2\text{SE}$	$^{87}\text{Sr}/^{86}\text{Sr}$	$\pm 2\text{SE}$	ρ
	Mc-3	4810	0.24	31.42	1.8	51000	4.2	2030	5.3	0.95
	Mc-4	5340	0.2	32.46	1.3	66100	3.6	2541	3.2	0.93
	Mc-5	5820	0.21	33.42	0.74	68200	4.2	2530	4	0.93
	Mc-6	5350	0.23	32.58	1.3	58100	3.2	2231	3.2	0.92
	Mc-7	6070	0.19	32.31	0.88	77600	4.3	3010	4	0.98
	Mc-9	5930	0.17	32.93	1.3	89600	4.2	3410	4.9	0.97
	Mc-10	4950	0.16	32.83	1.2	77900	5.5	2960	5.2	0.98
	Mc-11	7250	0.38	33.4	0.96	47500	2.4	1765	2.2	0.92
	Mc-12	4880	0.28	33.74	0.96	43400	4.4	1622	4.7	0.93
	Mc-13	6950	0.23	33.54	0.93	73900	3.2	2778	3.4	0.96
	Mc-14	8160	0.15	33.88	0.92	137700	5.2	5080	4.9	0.98
	Mc-15	5450	0.28	33.75	0.81	48700	2.6	1804	2.6	0.98
	Mc-16	8190	0.11	32.26	0.97	180000	5	6970	4.9	0.98
	Mc-17	4750	0.62	31.79	1.4	18900	2.9	743	3	0.89
	Mc-18	6580	6.2	32.87	1.2	2600	1.3	99.3	1.4	0.56
	Mc-19	4650	5.5	32.6	1	2100	1.1	79.75	0.94	0.54
	Mc-20	5500	0.31	32.01	1	43700	3.1	1688	2.6	0.96
	Mc-21	10100	0.11	32.37	0.95	222400	4.1	8490	3.7	0.97
	Mc-22	5710	0.18	33.13	1	79700	4.3	3010	4.3	0.97
	Mc-23	8970	0.16	33.05	0.91	137100	5.1	5110	4.1	0.98
	Mc-24	8570	0.16	32.33	1.8	137800	7.2	5320	6.8	0.97
	Mc-25	8620	0.28	32.7	0.64	78100	7.5	2990	7.8	1
	Mc-26	5880	0.25	33.01	0.82	58600	4.1	2219	4	0.98
	Mc-27	4350	0.36	30.4	4.9	29400	4.2	1215	5.5	0.41
	Mc-28	6990	0.52	34.7	3.3	33800	9.6	1188	5.6	0.86
	Mc-29	6140	0.44	31.68	0.93	35200	4.1	1388	4.4	0.98
	Mc-30	5240	0.16	33.23	1	80500	5.6	3030	5.5	0.98
	Mc-31	7040	0.18	32.75	1.1	99800	4.3	3810	4.3	0.97
	Mc-32	5620	0.15	33.09	0.94	94500	3.8	3570	3.8	0.97
	Mc-33	6000	0.35	33.55	0.78	43500	4.2	1614	3.9	0.97
	Mc-34	5480	0.19	33.14	1	74500	4.7	2810	4.8	0.98
	Mc-35	5180	0.19	32.63	1.1	69000	4.4	2640	4.4	0.97
	Mc-36	8840	0.12	32.04	1.3	181800	5.4	7000	4.8	0.97
	Mc-37	5480	0.21	33.74	0.74	65800	4.7	2420	4.3	0.95
	Mc-38	5290	0.19	33.68	0.86	69600	4.5	2610	4.8	0.91
	Mc-39	4950	0.19	32.5	1.1	66200	3.1	2545	3	0.93
	Mc-40	5910	0.23	32.86	1.2	63100	3.5	2400	3.3	0.96
	Mc-41	5070	0.18	32.93	0.76	70200	4.3	2660	4.2	0.99
	Mc-42	5370	0.11	32.8	1.1	127400	8.9	4860	9.2	0.99
	Mc-43	7970	0.19	32.76	0.88	106900	6.1	4080	5.9	0.99
	Mc-44	5890	0.2	33.38	0.8	72800	2.8	2731	3.1	0.97
	Mc-45	4940	0.21	32.69	1.2	58800	4.4	2246	3.8	0.97
	Mc-46	5690	0.16	32.79	1.1	92200	4.7	3520	4.6	0.97
	Mc-47	5530	0.24	33.05	1.1	57400	4.1	2198	4.3	0.97
	Mc-48	5230	0.27	32.8	1.1	48300	3.6	1842	3.3	0.95
	Mc-49	9130	0.14	32.58	1.1	167700	6	6420	5.2	0.99
	Mc-50	4720	0.16	32.38	1.1	73600	4.6	2840	4.1	0.97
	Mc-51	5250	0.22	32.21	1.1	60100	5.4	2330	5.1	0.98
	Mc-52	4870	0.24	32.25	1.3	51400	3.9	1989	3.1	0.95
	Mc-53	5190	0.24	33.31	1	54900	4.3	2061	4.3	0.97
	Mc-54	5310	0.2	32.35	1.1	66400	4.6	2560	4.1	0.98
	Mc-55	4860	0.12	31.84	0.98	100200	5.1	3930	4.8	0.98
	Mc-56	6400	0.55	33.07	0.93	28900	4.5	1097	4.1	0.97
	Mc-57	5540	0.21	31.72	1.4	65400	4.2	2580	3.9	0.95
	Mc-58	5110	0.18	32.65	0.71	70900	3.7	2694	3.4	0.9
	Mc-59	5770	0.17	32.33	1	81500	4.5	3150	4.3	0.98
	Mc-60	6410	0.19	32.1	1.1	85800	4.4	3340	4.3	0.96
	Mc-61	6220	0.21	32.97	0.82	74700	5.1	2830	5	0.99
	Mc-62	5470	0.19	32.45	0.72	71400	3.6	2732	3.2	0.96
	Mc-63	6960	0.46	31.91	0.87	38200	3.8	1496	3.4	0.98
	Mc-64	5710	0.22	32.81	0.81	65800	5.1	2480	4.2	0.9
	Mc-65	5930	0.19	32.55	1	77900	4.2	2990	4.2	0.97
	Mc-66	7480	0.18	33.3	0.78	103200	3.9	3890	4	0.87
	Mc-67	5690	0.2	33.04	1	70100	4.1	2650	3.9	0.97
	Mc-68	9910	0.08	31.32	1.1	294800	4.9	11760	5	0.98
	Mc-69	4910	0.17	32.71	0.94	71400	5.2	2730	5.3	0.99
	Mc-70	5000	0.13	32.3	1	93600	5.4	3620	5.3	0.98
	Mc-71	5520	0.26	32.95	0.92	54400	3.9	2061	3.6	0.97
ZU008	Fs-1	14.1	37	1.72	6.4	1.168813848	6.9	0.726	1.4	0.45
	Fs-2	0.992	4.6	0.3587	1.4	1.090329689	2	3.258	1.9	0.75
	Fs-3	9.57	34	1.334	1.1	0.879647093	0.7	0.7103	0.75	0.3
	Fs-4	6210	14	40.8	3	1090.749327	2.6	33.37	1.7	0.084
	Fs-5	27.2	7.8	5.23	14	10.46710974	14	2.146	2.9	0.19
	Fs-6	5460	7.6	35	2.9	1785.788175	2.9	63.6	4.1	0.69
	Fs-7	9.84	34	1.362	1.3	0.901275809	0.91	0.7098	0.79	-0.12
	Fs-8	12.3	38	1.456	3.4	0.972432526	3.2	0.7163	0.97	-0.019
	Fs-9	12.2	21	1.97	12	1.766197769	12	0.966	1.7	-0.091
	Mc-1	6290	0.26	32.17	0.68	60900	6	2370	5.9	0.99
	Mc-2	5800	0.19	31.79	0.86	76800	4	3050	4.3	0.98
	Mc-3	8060	0.18	32.22	0.84	113600	4.7	4410	4.6	0.99
	Mc-4	6450	0.2	32.22	0.82	80000	4.2	3100	4.1	0.98
	Mc-5	7820	0.21	32.56	0.86	92100	4.4	3540	4.3	0.98
	Mc-6	7610	0.18	32.35	0.88	106800	4.7	4120	4.4	0.98
	Mc-7	8390	0.24	31.98	0.74	88900	4.2	3480	4.3	0.99
	Mc-8	9410	0.15	32.74	0.87	158900	4.1	6070	4.2	0.98

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Sample	Spot name	Rb ppm	Sr	$^{87}\text{Rb}/^{87}\text{Sr}$	$\pm 2\text{SE}$	$^{87}\text{Rb}/^{86}\text{Sr}$	$\pm 2\text{SE}$	$^{87}\text{Sr}/^{86}\text{Sr}$	$\pm 2\text{SE}$	ρ
	Mc-9	8180	0.093	31.67	0.79	222290	5.4	8770	5.2	0.99
	Fs-10	28.2	33	3.24	16	2.474576304	20	0.758	2.4	0.37
	Fs-11	11.2	36	1.424	1.5	0.945789146	1.5	0.7134	0.8	0.34
	Fs-12	10.7	32	1.459	5.7	1.000188231	7.1	0.7192	1.1	0.14
	Mc-10	7030	0.21	32.82	0.66	84338.891	4.9	3210	4.9	0.99
	Mc-11	8530	0.12	31.85	0.75	175094.626	4.8	6810	4.6	0.91
	Mc-12	6680	0.21	32.77	0.77	79376.8445	4.1	3030	3.9	0.99
	Mc-13	6060	0.25	33	0.94	59425	3.8	2254	4	0.97
	Mc-14	6900	0.17	32.44	0.94	103673.9	5.5	3990	5.3	0.99
	Mc-15	6990	0.18	32.43	0.97	95206	3.7	3670	3.5	0.98
	Mc-16	6060	0.23	30.94	1.3	67534.7163	8.6	2750	8.5	0.98
	Mc-17	7290	0.1	31.46	0.91	174334.482	5.7	6930	5.8	0.99
	Mc-18	6120	0.17	32.77	1.3	89013.45	3.7	3370	3.5	0.91
	Mc-19	8210	0.13	31.81	0.79	155406.72	6.8	6100	6.5	0.99
	Mc-20	5260	0.29	32.75	0.67	45162.701	3.1	1725	3.1	0.98
	Mc-21	5150	0.29	32.72	0.76	44177.146	4	1687	3.6	0.99
	Mc-22	6010	0.15	34.09	2	97978.19227	6.7	3600	7.1	0.96
	Mc-23	4790	0.28	32.04	0.95	43955.74819	5	1716	5	0.98
	Mc-24	3610	0.17	31.61	0.94	52047.25055	3.6	2060	3.5	0.98
	Mc-27	5130	0.19	31.95	0.89	68016.29746	3.7	2681	3.7	0.91
	Mc-28	6240	0.09	31.74	1	174287.0152	5.1	6860	4.9	0.99
	Mc-29	7410	0.11	31.6	1	175896.8561	3.8	6970	3.9	0.96
	Mc-30	5460	0.12	31.6	0.88	114962.6728	5.7	4540	5.4	0.99
	Mc-31	8420	0.08	31.12	1.2	258298.6845	4.7	10370	4.4	0.97
	Mc-32	3880	0.22	37.12	2.6	45348.38162	6.4	1528	6	0.92
	Mc-33	6830	0.15	32.07	0.8	113681.7929	4.2	4450	4.1	0.85
	Mc-34	7150	0.22	31.62	1.3	81566.62112	4.9	3230	5.2	0.97
	Mc-35	8210	0.22	32.82	0.52	96332.65221	4.4	3660	4.3	0.98
	Mc-36	8320	0.11	31.37	1.1	195812.8116	7.1	7810	7.1	0.99
	Mc-37	6830	0.24	31.92	1.1	72515.75379	3.9	2836	3.2	0.97
	Mc-38	6460	0.21	32.46	1	80252.5535	5.4	3110	5.9	0.97
	Mc-39	7760	0.13	31.5	1.1	151146.5473	5.3	6010	5.7	0.98
	Mc-40	8600	0.058	30.91	1.1	371261.6918	9.1	15000	8.6	0.99
	Mc-41	5070	0.51	33.07	0.74	24868.28128	5	940	5.1	0.99
	Mc-42	8260	0.11	32.13	1.2	188385.338	5.4	7320	5.4	0.98
	Mc-43	6680	0.14	31.65	0.93	120786.3163	4.1	4770	4	0.97
	Mc-44	7440	0.12	31.7	1.1	159627.83	5.9	6310	6.3	0.99
	Mc-45	6380	0.19	31.46	0.89	85090.98504	4.2	3370	3.8	0.98
	Mc-46	7130	0.11	31.39	0.8	169921.4256	5.8	6770	6.1	0.99
	Mc-47	8390	0.12	31.67	0.97	186198.7409	7.2	7310	6.9	0.99
	Mc-48	6560	0.12	31.37	0.9	137369.5624	4.6	5510	4.5	0.95
	Mc-49	7690	0.19	31.59	1.1	100074.8584	4.2	3930	4.4	0.9
	Mc-50	7710	0.12	31.66	0.89	154571.7999	4.4	6100	4.6	0.98
	Mc-51	6890	0.13	31.2	0.85	126104.4531	7.2	5050	7.3	1
	Mc-52	7430	0.096	31.59	0.66	198874.0853	4.7	7860	4.9	0.99
	Mc-53	5420	0.15	32.72	0.84	89543.49045	5.2	3410	4.9	0.99
	Mc-54	8430	0.31	31.68	1.1	67119.13124	3.7	2650	3.9	0.98
	Mc-55	7780	0.08	31.25	0.9	240476.7026	5	9670	4.6	0.95
	Mc-56	8300	0.084	31.89	0.57	246017.1776	5.8	9520	5.5	0.97
	Mc-57	5620	0.12	32.7	0.69	115674.7034	4	4410	4	0.99
	Mc-58	8610	0.085	31.22	0.89	257416.0338	5.7	10290	5.7	0.99
	Mc-59	7910	0.2	31.59	0.82	98075.5794	4.4	3870	4.1	0.98
	Mc-60	6170	0.18	31.37	1.1	84334.92262	3.4	3350	3.3	0.96
	Mc-61	6130	0.22	31.56	0.97	70091.81242	3.1	2772	3	0.97
	Mc-62	6250	0.18	31.57	1.1	89094.76061	4.3	3520	4.2	0.96
	Mc-63	3460	0.33	32.01	0.87	26244.32412	2.4	1033	3.5	0.97
Z6035	1	2620	0.36	19218	3.4			753	3.5	0.97
	2	5440	0.26	54180	2.8			2010	2.6	0.96
	3	2760	0.42	16965	4			735	4	0.99
	4	2930	1.2	6500	2.9			281.8	2.8	0.99
	5	2520	0.73	8950	2.5			337.2	2.6	0.89
	6	2310	0.65	9204	2.2			354.4	2	0.94
	7	2430	0.67	9330	2.4			354.7	1.8	0.91
	8	2710	0.4	17566	2.7			642	2.5	0.96
	9	2660	0.34	20520	3.5			746	2.9	0.86
	10	2650	0.78	8830	2.8			372.9	2.5	0.83
	11	3160	0.38	21400	3.6			920	3.2	0.99
	12	2970	0.52	14660	2.5			627	2.6	0.96
	13	2570	2.4	2780	3.8			126	3.5	0.98
	14	1010	0.55	4750	3.3			267.6	3	0.99
	15	2960	0.98	7860	2.3			287	2	0.94
	16	2790	1	7050	2			261.7	1.3	0.6
	17	3710	0.98	9740	1.8			355.7	2.1	0.89
	18	2840	0.92	7900	1.9			283.7	1.8	0.92
	19	2580	0.64	10420	2.2			384.8	2	0.96
	20	2120	1	5260	1.3			192.6	1.6	0.87
	21	2160	0.96	5830	1.7			213.2	1.5	0.84
	22	2640	0.47	14570	3.2			541	2.9	0.95
	23	2290	1.8	3325	1.6			122	1.2	0.79
	24	2180	1.2	4860	2.2			183.5	1.6	0.87
	25	2740	0.51	13970	2.8			521	3	0.93
	26	2490	1.6	4010	1.6			144.7	1.5	0.91
	27	2320	0.48	12340	2.2			451.3	2	0.94
	28	2570	1.8	3630	1.3			129.2	1.2	0.91
	29	2770	0.74	9605	1.9			350.5	1.9	0.91

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Sample	Spot name	Rb ppm	Sr	$^{87}\text{Rb}/^{87}\text{Sr}$	$\pm 2\text{SE}$	$^{87}\text{Rb}/^{86}\text{Sr}$	$\pm 2\text{SE}$	$^{87}\text{Sr}/^{86}\text{Sr}$	$\pm 2\text{SE}$	ρ
	30	2750	1	7100	3			256.2	2.8	0.96
	31	2600	0.46	14680	2.6			555	2.3	0.94
	32	2660	0.65	10600	2.9			397.4	2.5	0.95
	33	2470	0.5	12710	2.8			473	2.6	0.96
	34	2990	0.63	12270	2.6			459	2.4	0.95
	35	2300	0.96	6260	1.9			227.7	1.7	0.85
	36	2480	0.78	8230	2			299.2	1.8	0.92
	37	2450	0.93	6795	2.5			248.6	2.2	0.97
	38	2520	0.64	10075	2.9			374.2	2.6	0.96
	39	3700	0.56	17160	2.4			621	2.4	0.96
	40	2760	0.75	9540	2.1			359.2	1.9	0.93
	41	2180	0.59	9470	3.1			360.5	2.4	0.91
	42	2610	0.96	7100	2.6			257.4	1.7	0.94
	43	2410	1.4	4610	1.8			171.6	1.2	0.91
	44	2400	1.2	5335	1.6			200.9	1.8	0.77
	45	2490	1	6190	1.9			224.1	1.9	0.91
	46	2300	0.56	10590	2.3			385.1	2.1	0.95
	47	2930	0.47	15970	1.5			578.2	1.6	0.91
	48	2380	0.83	7470	2			272.8	2.1	0.94
	49	2660	1.6	4330	1.8			157.5	1.3	0.89
	50	3280	0.26	33020	3.2			1198	3.3	0.91
	51	2680	0.88	7817.584	1.9			289.8	1.9	0.79
	52	2500	0.95	6797.821	2.1			257.1	1.8	0.9
	53	3030	0.72	10866.4	2.1			392.6	2.1	0.79
	54	2160	0.82	6870.32	2.7			250	2.3	0.95
	55	2650	0.6	11562.709	2.2			424.1	1.9	0.87
	56	2640	0.72	9418.897	2			349.6	1.7	0.89
	57	2170	1.5	3711.682	1.5			135.6	1.5	0.89
	58	2100	2.2	2477.64	1.2			94	1.1	0.85
	59	2470	0.55	11774.385	2.4			429	2.4	0.97
	60	2860	0.6	12382.047	2.9			447	2.8	0.98
	61	2090	0.84	6438.58	2.2			242.9	2	0.94
	62	2960	1.1	7113.249	1.9			253.5	2	0.93
	63	2670	0.86	8063.66	2.4			294.5	2.4	0.82
	64	2220	1.7	3350	1.8			126.2	1.6	0.93
	65	2590	0.43	15759.53	3.3			584	3.3	0.99
	66	2730	0.61	11609	2.5			420	2.5	0.97
	67	2640	0.56	12016	2.7			440	2.7	0.91
	68	2950	1.1	7140.6	2.8			258.5	1.8	0.92
	69	2510	0.98	6638.2	1.8			252.4	1.8	0.93
	70	3020	0.56	13895	2.7			510	2.6	0.86

C.5 Trioctahedral (Brown) mica Rb–Sr data

Sample	Spot name	Rb ppm	Sr	⁸⁷ Rb/ ⁸⁷ Sr	±2SE (%)	⁸⁷ Rb/ ⁸⁶ Sr	±2SE (%)	⁸⁷ Sr/ ⁸⁶ Sr	±2SE (%)	rho	Fe ppm	Mg	Ti	Mg#	Ti (a.p.f.u.)	Ti in Biotite Temperature
ZU001	1	660	7.1	27.83	1.2	3840	6.3	158.9	6	0.98	154659	30549	16102	0.16	0.32	635
	3	634	5.8	32.24	1.7	6390	5.1	223	7.9	0.99	166352	32162	15599	0.16	0.31	629
	5	549	4.5	4.21	3.3	3.145	3.1	0.866	3.4	0.47	155012	20488	2758	0.12	0.05	
	6	621	6.6	32.51	1.6	1290	9.8	45.7	9.7	0.99	183793	34641	17125	0.16	0.34	645
	7	568	4.3	38.8	3.1	13100	15	390	15	0.98	178281	35661	15282	0.17	0.3	626
	8	614	14	29.84	2.7	179.8	4.8	6.93	3.9	0.84	164923	26290	12755	0.14	0.25	591
	9	571	5.3	32.4	5.5	7800	15	274	12	0.92	188305	38456	19198	0.17	0.38	664
	10	668	6.3	31.3	2.2	5170	10	190	10	0.98	156258	33099	18884	0.17	0.37	662
	11	598	5.4	31.95	2.5	8070	7.1	291	6.7	0.94	174149	34381	15949	0.16	0.31	633
	12	633	6	30.44	1.4	10360	6.1	392	6.3	0.98	168959	34663	16321	0.17	0.32	637
	13	508	4.9	32.78	1.6	2130	6.8	74.6	6.6	0.97	176578	35158	16178	0.17	0.32	636
	14	626	6.2	29.16	0.89	7710	8.3	302	8.2	0.99	175167	30595	16661	0.15	0.33	641
	15	638	5.8	33.03	1.8	3650	5.7	126.6	5.6	0.95	160320	30793	14241	0.16	0.28	613
	17	557	5	32.76	1.6	5350	7.8	188	8.3	0.98	176820	33546	15079	0.16	0.3	623
	18	589	5.7	30.58	1.5	4460	5.9	165.5	5	0.92	163137	31844	15174	0.16	0.3	624
	19	576	5.3	32.05	1.9	6350	5	231	5.6	0.94	175732	33870	17587	0.16	0.34	650
	20	578	5.2	32.82	1.4	6250	8	218	7.5	0.99	161435	32307	18801	0.17	0.37	661
	21	459	5.3	31.34	1.5	951	4	34.8	4.2	0.93	207658	33858	15999	0.14	0.31	633
	22	489	42	19.07	1.3	33.12	1.8	1.991	1.9	0.71	129982	25451	12757	0.16	0.25	591
	25	633	6	32.02	1.3	3510	4.6	125.5	4.6	0.96	156458	31194	14734	0.17	0.29	619
	27	589	5.2	35.53	2	2590	5.5	83.6	5.6	0.93	186743	36751	16762	0.16	0.33	642
	28	600	5.5	31.08	1.2	11260	6.6	415	6.5	0.99	182268	35868	15249	0.16	0.3	625
	29	628	6.1	30.09	1.5	6940	12	261	12	0.97	178610	32747	14259	0.15	0.28	613
	30	623	6.2	28.76	1.3	7390	5.1	295	5.4	0.97	170821	29590	13700	0.15	0.27	605
	31	614	6.1	28.73	1.9	8990	8.8	359	8.5	0.98	181471	32339	14764	0.15	0.29	619
	32	695	6.8	31.33	1.6	2650	8.9	97.3	9.2	0.99	190102	33936	16707	0.15	0.33	641
	33	660	9.6	28.25	1.2	498	3.2	20.06	2.8	0.93	177122	32253	15397	0.15	0.3	627
	34	668	13	27.86	1.5	254	4	10.45	3.6	0.93	182439	33222	16001	0.15	0.31	634
	35	574	12	28.57	2.8	215	5.6	8.63	5	0.85	155918	26526	12373	0.15	0.24	585
	36	619	6.3	29.54	1.6	3090	4.5	120.6	5.2	0.96	181832	31629	15488	0.15	0.3	628
	37	611	5.9	29.28	1.6	9350	6.8	367	6.8	0.97	166406	28918	15411	0.15	0.3	627
	38	462	34	20.92	1.9	38.4	2.9	2.108	2.1	0.75	128043	22789	10743	0.15	0.21	554
	39	615	6.8	36.86	2.4	656	3.3	20.55	4.4	0.85	181396	32075	16086	0.15	0.32	635
	40	607	21	24.66	2.1	97.5	2.5	4.549	2.1	0.58	160336	26607	13402	0.14	0.26	601
	41	649	6.1	30.5	1.8	6540	7.7	246	7.2	0.98	165961	28185	15635	0.15	0.31	629
	42	724	6.9	29.98	1.2	6580	6.7	251	7	0.97	205278	34196	18266	0.14	0.36	656
	44	690	7	29.36	1.3	4180	19	164	19	1	166191	28657	14082	0.15	0.28	610
	46	685	6.7	29.14	1.3	8150	11	323	11	0.99	181899	32451	15135	0.15	0.3	624
	47	609	6.2	30.14	1.2	2280	6.1	87.4	6.1	0.98	183633	31725	19477	0.15	0.38	666
	48	619	5.5	31.83	1.7	10200	11	370	12	0.99	190323	35666	16832	0.16	0.33	642
	49	510	32	22.83	1.3	45.92	1.7	2.325	1.9	0.72	150214	24724	12400	0.14	0.24	585
	50	651	6.6	29.05	1.5	3890	5.6	155	5.2	0.98	179936	31448	14845	0.15	0.29	620
	51	588	25	23.87	2.1	75.3	3	3.646	2	0.76	163521	26577	14743	0.14	0.29	619
	52	705	7.4	29.23	1.4	2170	7	85.8	6.9	0.98	186199	31571	14575	0.14	0.29	617
	53	691	6.5	30.36	1.5	6240	5.4	237	4.7	0.97	187342	31265	16498	0.14	0.32	639
	54	637	6	30.1	1.4	7150	6	274	5.9	0.98	173068	29655	13876	0.15	0.27	607
	55	654	6.5	28.75	1.4	6580	5.5	265	5.5	0.98	188979	31286	16678	0.14	0.33	641
	56	598	5.6	30.8	1.9	4980	7.8	187	7.4	0.97	178346	32910	14611	0.16	0.29	617
	58	524	18	29.38	1.7	91.5	1.8	3.627	2.2	0.64	148886	25648	12174	0.15	0.24	581
	59	682	11	30.36	2.2	327.1	2.3	12.48	2.2	0.5	176301	31630	17048	0.15	0.33	645

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Sample	Spot name	Rb ppm	Sr	$^{87}\text{Rb}/^{87}\text{Sr}$	$\pm 2\text{SE}$ (%)	$^{87}\text{Rb}/^{86}\text{Sr}$	$\pm 2\text{SE}$ (%)	$^{87}\text{Sr}/^{86}\text{Sr}$	$\pm 2\text{SE}$ (%)	rho	Fe ppm	Mg	Ti	Mg#	Ti (a.p.f.u.)	Ti in Biotite Temperature
	60	621	16	25.41	2	145.6	1.9	6.65	2.1	0.49	167644	27673	14136	0.14	0.28	611
	61	632	6.3	28.59	1.7	6480	8.1	263	8.2	0.98	180963	28792	17487	0.14	0.34	649
	62	668	6.6	28.72	1.3	9110	7.8	367	7.8	0.99	197119	33065	16226	0.14	0.32	636
	63	688	6.5	30.07	1.9	6840	9.7	264	9.4	0.98	182885	30555	16485	0.14	0.32	639
	64	666	6.6	28.7	1.5	9010	4.8	363	4.8	0.97	181179	31463	16120	0.15	0.32	635
	65	673	6.1	31.16	1.4	11120	5.8	411	5.9	0.94	182266	29921	14908	0.14	0.29	621
	66	657	6.6	28.06	1.1	10270	6.7	424	7	0.99	175838	29626	15938	0.14	0.31	633
	67	660	6.8	31.59	1.3	1516	3.7	55.6	3.3	0.94	176282	29925	17552	0.15	0.34	649
	68	740	6.5	32.94	1.8	5820	6.2	205	6.1	0.95	182629	32610	16679	0.15	0.33	641
	69	610	5.3	32.81	1.9	6480	7	228	7.2	0.95	184514	30875	13890	0.14	0.27	608
	70	526	4	37.37	1.6	7960	7.6	247	7.8	0.98	194803	29753	12826	0.13	0.25	592
	71	621	6.5	28.63	1.7	3000	5.5	121.4	5.7	0.95	184364	29809	16768	0.14	0.33	641
	73	679	6.4	29.74	1.3	11760	6.4	458	6.1	0.99	180122	31404	15585	0.15	0.31	629
	74	549	4.9	33.09	1.5	4050	5.3	141.8	5.2	0.96	199684	36134	15132	0.15	0.3	624
	75	701	11	28.61	2.9	360	5.2	14.55	3.7	0.84	182374	32890	17319	0.15	0.34	647
	78	670	5.6	33.67	1	15500	6.8	532	6.6	0.99	187209	34135	16055	0.15	0.31	634
	79	712	9.9	27.65	1.9	628	2.9	26.23	3.2	0.8	175222	30918	15703	0.15	0.31	630
	80	666	8	30.51	1.7	801	2.3	30.3	2.4	0.71	176410	30946	14631	0.15	0.29	617
	81	670	6.1	31.53	1.7	10340	7.9	378	7.7	0.97	187178	34425	16593	0.16	0.33	640
	82	624	14	27.23	1.5	180	7	7.78	5.1	0.73	175123	32150	15194	0.16	0.3	624
	83	701	6.7	31.08	1.8	4850	7.3	180	6.7	0.97	187278	34566	15701	0.16	0.31	630
	84	684	6.5	30.54	1.5	6430	5.8	243	6.2	0.98	173061	32265	14673	0.16	0.29	618
	85	696	6.7	30.43	1.3	6980	7.8	264	7.8	0.99	193327	36368	16466	0.16	0.32	639
	86	748	7.2	29.81	1.7	9760	6.7	377	6.5	0.97	178559	34745	16403	0.16	0.32	638
	87	710	6.6	31.31	2.6	6670	12	245	13	0.98	175942	31717	13320	0.15	0.26	600
	88	728	6.6	32.6	2.4	6510	6.3	229	5.7	0.92	189095	34433	17449	0.15	0.34	648
	89	735	7.9	28.34	2.2	2990	4.8	123.7	5.4	0.96	190030	37308	18611	0.16	0.36	659
	90	723	7.6	31.71	2.6	1519	5.5	55	5.3	0.9	181884	34031	15895	0.16	0.31	633
	91	686	6.5	31.97	1.9	5480	6.5	195	5.8	0.96	182836	34632	15326	0.16	0.3	626
	92	6.38	5.9	3.87	2.7	2.809	1.8	0.828	2.6	0.31	163408	23891	3837	0.13	0.08	
	93	6.67	5.8	4.22	3.2	2.989	2	0.814	2.8	0.14	167452	23531	3697	0.12	0.07	
	94	4.82	5.2	3.44	3.4	2.434	2.6	0.813	3	0.23	157846	22425	3060	0.12	0.06	
	95	6.72	5.7	4.28	3.1	3.07	2.9	0.82	2.9	0.4	156948	23454	3999	0.13	0.08	
	96	6.4	5.8	3.85	3.3	2.875	2.3	0.857	3	0.19	154013	21543	3654	0.12	0.07	
	97	5.67	5.2	3.88	3.2	2.842	2.6	0.838	2.6	0.15	154907	21869	3573	0.12	0.07	
	98	5.7	5.4	3.94	2.8	2.776	2.2	0.805	2.8	0.35	152694	21652	3435	0.12	0.07	
	99	5.91	5.1	4.25	3.5	3.062	1.7	0.83	3.7	0.32	165907	21897	3367	0.12	0.07	
	100	6.43	5.6	4.05	2.9	3.014	2	0.846	3	0.086	170075	23361	3801	0.12	0.07	
	101	7.61	6	4.55	2.6	3.334	2.2	0.836	2.9	0.45	157835	21870	3745	0.12	0.07	
	102	6.74	5.4	4.54	3	3.259	2.1	0.821	3.4	0.44	159107	21456	3376	0.12	0.07	
	103	7.02	5.6	4.36	2.8	3.286	2.1	0.855	3.1	0.29	158519	20604	3221	0.12	0.06	
	104	6.91	5.7	4.43	3.4	3.173	2.3	0.821	2.9	0.22	157179	21422	3388	0.12	0.07	
	105	7.17	5.7	4.4	3.2	3.308	2.3	0.857	3	0.23	163400	23332	3778	0.12	0.07	
ZU005	1	1030	9.4	31.3	1.4	21100	9.9	814	11	0.99	129614	54746	10678	0.3	0.21	562
	2	1020	9.6	30.24	1.4	16200	6.9	627	6.8	0.98	125331	56797	11306	0.31	0.22	576
	3	1490	11	38.63	1.9	18600	11	565	10	0.98	124404	56597	11035	0.31	0.22	571
	4	1160	11	31.04	1.8	14300	8.4	542	8.7	0.97	141758	65955	11461	0.32	0.22	579
	5	1350	12	32.48	1.5	25500	8	921	7.9	0.99	133599	62403	10368	0.32	0.2	558
	6	1320	12	30.61	1.1	20200	7.8	770	7.6	0.99	129513	59788	10215	0.32	0.2	554
	7	1090	9.7	32.1	1.8	18600	8.7	683	8.1	0.99	128829	59442	11276	0.32	0.22	576
	8	867	8.2	30.73	1.4	10260	6.7	392	6.3	0.98	123648	56287	11218	0.31	0.22	574
	9	915	7.9	34.5	1.4	4970	6.3	172	6.6	0.98	134078	64236	9286	0.32	0.18	532
	10	1300	11	32.74	1.4	15500	7.1	556	6.9	0.99	128359	58800	10460	0.31	0.2	559
	11	1300	11	34.95	1.8	9590	5.2	325	5.1	0.94	127715	55140	8541	0.3	0.17	508

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Sample	Spot name	Rb ppm	Sr	$^{87}\text{Rb}/^{87}\text{Sr}$	$\pm 2\text{SE}$ (%)	$^{87}\text{Rb}/^{86}\text{Sr}$	$\pm 2\text{SE}$ (%)	$^{87}\text{Sr}/^{86}\text{Sr}$	$\pm 2\text{SE}$ (%)	rho	Fe ppm	Mg	Ti	Mg#	Ti (a.p.f.u.)	Ti in Biotite Temperature
12	1380	12	33.46	1.8	18800	7.8	663	6.9	0.97	121810	53253	9318	0.3	0.18	531	
13	1190	12	29.38	1.2	13210	7.2	531	6.8	0.98	134740	60967	11682	0.31	0.23	583	
14	1370	11	34.77	1.5	20600	6.8	702	7.1	0.99	129297	59101	10434	0.31	0.2	558	
15	1230	12	30.36	1.4	20100	8.6	779	7.9	0.99	129860	58429	11019	0.31	0.22	570	
16	1280	11	32.53	1.3	23000	7.3	830	6.6	0.96	134633	62421	9964	0.32	0.2	548	
17	1220	11	31.89	1.2	15690	6.3	581	5.9	0.98	125373	56975	9800	0.31	0.19	544	
18	1370	12	33.64	1.1	19900	6.5	699	6.2	0.98	131226	62745	10925	0.32	0.21	570	
19	1010	9.1	31.63	1.4	24200	9.2	907	9.1	0.99	116217	56246	9318	0.33	0.18	533	
20	1300	12	30.42	1.4	24600	11	960	11	0.99	129001	57225	9896	0.31	0.19	546	
21	1360	10	37.34	1.4	52300	14	1650	13	1	150964	50135	8941	0.25	0.18	514	
22	1080	9.4	32.95	1.5	15600	10	545	8.9	0.99	157482	51839	11505	0.25	0.23	574	
23	1300	12	31.81	1.3	14500	7.4	541	7.4	0.99	146024	47896	10307	0.25	0.2	549	
24	1380	11	36.94	1.2	14040	7	451	7	0.99	150145	47809	10279	0.24	0.2	548	
25	1140	9.7	33.75	1.4	14900	9.5	522	9.2	0.99	141821	48461	9495	0.25	0.19	530	
26	1110	10	31.55	1.1	20300	11	756	11	1	128401	61737	9545	0.32	0.19	539	
27	1300	11	33.84	1.1	22800	9.5	800	9	0.99	147651	52823	8783	0.26	0.17	511	
28	1430	11	35.78	1.3	11690	7.2	389	7	0.98	142636	51137	9639	0.26	0.19	535	
29	772	7.4	30.37	1.3	6720	8.6	263	8.3	0.99	136966	63569	11642	0.32	0.23	583	
30	914	8.5	31.05	1.6	11700	9	447	8.7	0.98	126248	61882	10424	0.33	0.2	560	
31	1350	11	33.19	2	30200	12	1080	11	0.98	135305	64425	9194	0.32	0.18	530	
32	890	8.1	31.39	1.5	12700	7.9	481	7.7	0.99	129337	62519	11136	0.33	0.22	574	
33	1380	12	33.55	2	27800	14	990	14	0.99	133713	62077	8602	0.32	0.17	512	
34	1330	11	35.88	1.6	13820	6.2	456	6.1	0.95	140962	63331	10981	0.31	0.22	569	
35	1290	9.9	37.65	1.4	11410	7.9	360	7.2	0.98	128690	58839	8006	0.31	0.16	491	
36	1040	8.8	34.38	1.3	8710	5.3	302	5.2	0.97	142867	56769	9246	0.28	0.18	526	
37	1280	11	33.31	1.6	15400	6.6	552	6.6	0.98	143628	55179	10235	0.28	0.2	550	
38	1270	11	31.46	1.2	19900	11	752	11	1	146899	58888	11868	0.29	0.23	583	
39	1390	11	36	1.6	24000	10	795	9.9	0.98	138892	57655	10336	0.29	0.2	554	
40	1150	10	32.29	1.3	22400	12	817	11	0.98	138751	57264	10876	0.29	0.21	565	
41	1190	11	30.57	1.6	21300	8.2	830	8.3	0.99	145004	55066	11606	0.28	0.23	578	
42	786	7.2	32.13	1.4	9170	6.6	339	6	0.97	137451	52814	11190	0.28	0.22	570	
43	1010	8.7	34.63	1	7120	4.7	245	4.5	0.97	136091	50224	9153	0.27	0.18	522	
44	941	8.8	31.47	1.1	9600	7.5	361	6.9	0.96	145669	56531	12124	0.28	0.24	587	
45	1030	9.6	31.2	1.2	17000	7.7	659	8.4	0.98	140600	50916	10164	0.27	0.2	548	
46	882	8.3	31.03	1.3	11110	6.6	426	6.5	0.99	132856	53319	10313	0.29	0.2	553	
47	1250	11	34	1.4	28700	9	998	8.8	0.99	134314	49530	10179	0.27	0.2	548	
48	1100	10	32.23	1.6	11300	9.3	420	9.6	0.98	136590	52733	10520	0.28	0.21	557	
49	1590	13	35.87	1.4	35100	10	1160	9.6	0.99	144762	59403	11886	0.29	0.23	584	
50	1450	12	34.42	1.3	32700	10	1130	10	0.99	130374	53205	10861	0.29	0.21	565	
51	1080	9.9	32.1	1.5	14900	8.1	553	8.1	0.99	131772	53528	11374	0.29	0.22	575	
52	1190	11	33.29	1.1	9300	9.6	329	9.1	0.98	138972	60968	8760	0.3	0.17	515	
53	1190	11	31.18	1.1	15100	7.4	576	7.1	0.99	136773	56034	12096	0.29	0.24	588	
54	1330	13	31.2	1.9	9590	8	356	6.4	0.96	128864	53976	11698	0.3	0.23	581	
55	1440	13	32.82	1.5	21300	12	775	12	1	130212	54672	11432	0.3	0.22	577	
56	1080	9.6	33.09	1.5	10800	11	387	11	0.99	129817	55810	11824	0.3	0.23	584	
57	1350	13	30.89	1.4	18200	9.8	702	9.7	0.99	130213	56803	12140	0.3	0.24	590	
58	1390	12	33.59	1.4	15400	10	546	10	0.99	136432	58854	11463	0.3	0.22	578	
59	1280	11	32.48	1.4	21800	9.6	798	9.3	0.99	133450	58133	10792	0.3	0.21	565	
60	1150	10	31.85	1.4	18300	9.4	662	7.4	0.99	139323	61147	12162	0.31	0.24	590	
61	902	8	33.06	1.6	11900	9.4	434	9.8	0.94	129002	53600	10626	0.29	0.21	560	
62	943	8.6	31.96	1.3	9380	7.3	349	7.2	0.98	132828	57682	10653	0.3	0.21	562	
63	1340	11	35.7	1.3	16400	7.7	544	7.5	0.99	129176	59514	11287	0.32	0.22	576	
64	998	9.5	31.14	1.3	5890	5.2	225	4.9	0.96	131574	57076	11633	0.3	0.23	581	
65	1360	11	34.72	1.3	24000	11	826	12	0.99	133944	61103	11515	0.31	0.23	580	
66	1090	9.8	32.43	1.4	9480	5.9	354	6.7	0.98	134651	63583	12163	0.32	0.24	592	

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Sample	Spot name	Rb ppm	Sr	$^{87}\text{Rb}/^{87}\text{Sr}$	$\pm 2\text{SE}$ (%)	$^{87}\text{Rb}/^{86}\text{Sr}$	$\pm 2\text{SE}$ (%)	$^{87}\text{Sr}/^{86}\text{Sr}$	$\pm 2\text{SE}$ (%)	rho	Fe ppm	Mg	Ti	Mg#	Ti (a.p.f.u.)	Ti in Biotite Temperature
67		1290	12	31.72	1.3	16600	11	619	10	0.99	127682	54701	11260	0.3	0.22	574
68		1150	9.8	33.65	1.3	12800	8	451	7.7	0.98	128755	58604	10994	0.31	0.22	570
69		816	7.8	29.95	1.5	8610	7.5	340	7.4	0.98	133561	61520	12220	0.32	0.24	592
70		1520	11	38.33	1.5	26400	10	819	10	0.99	129839	60047	10270	0.32	0.2	555
71		1130	10	31.24	1.6	14600	7.7	552	7.1	0.98	125727	55594	12534	0.31	0.25	596
72		1120	9.8	32.34	1.2	12040	6.2	442	5.7	0.99	131683	56659	11742	0.3	0.23	583
73		1040	9.9	30.86	1.5	3910	4.3	150.1	4	0.94	121976	54938	25761	0.31	0.5	714
74		1260	11	32.12	1.4	14770	6.5	545	6.3	0.98	125695	52351	12143	0.29	0.24	589
75		1150	10	31.03	1.3	23900	5.6	909	5.4	0.98	134998	56229	11792	0.29	0.23	583
76		1160	10	32.52	1.2	19000	7.6	691	7.4	0.99	136815	58816	11651	0.3	0.23	581
77		792	6.6	33.94	1.4	10050	5.9	358	6.4	0.98	146813	51524	9766	0.26	0.19	538
78		1100	9.5	32.75	1.5	14700	9.2	533	8.8	0.99	153897	55456	10311	0.26	0.2	551
79		871	170	12.49	1.4	12.84	2.5	1.217	1.6	0.7	137239	64538	8921	0.32	0.17	521
80		1210	9.9	34.46	1.3	15900	11	545	11	0.99	128107	54774	11341	0.3	0.22	575
81		1090	10	30.44	1.3	16400	8.4	625	7.5	0.99	130062	54675	11708	0.3	0.23	582
82		906	7.9	32.51	1.3	10430	5.6	379	5.1	0.97	152025	54691	16210	0.26	0.32	640
83		942	8.3	32.03	1.5	11740	7.1	436	7.3	0.98	145193	53291	11413	0.27	0.22	574
84		818	6.8	34.6	1.6	8040	5.8	275	5.7	0.97	158457	62339	11597	0.28	0.23	578
85		1180	11	30.65	1.6	18000	9.1	718	10	0.99	139473	51154	12069	0.27	0.24	585
86		1090	9.3	34	1.7	7520	6.5	261	6.2	0.96	137909	53870	13141	0.28	0.26	603
87		792	7.5	30.19	1.9	12800	8.7	500	8.8	0.98	131550	59914	13062	0.31	0.26	605
88		1430	12	34.22	1.9	16100	11	554	11	0.99	137094	61057	11265	0.31	0.22	575
89		1140	9.8	32.84	1.5	25000	8.8	898	8.5	0.99	135194	64698	12398	0.32	0.24	596
90		945	8.7	31.07	1.4	16600	7.8	630	7.5	0.99	136065	58235	12842	0.3	0.25	601
91		0.147	240	0.00238	28	0.00142	28	0.7032	1	0.13	304	10	42	0.032	0	
93		0.0384	220	0.00073	81	0.00044	81	0.7093	0.87	-0.066	230	3	1	0.013	0	
94		5.64	240	0.098	11	0.0589	11	0.7118	1.1	0.13	360	27	6	0.07	0	
96		0.401	210	0.00782	11	0.00467	11	0.7034	1.3	-0.06	207	3	79	0.012	0	
97		0.202	230	0.00358	18	0.00216	18	0.7125	1.1	0.034	150	3	36	0.018	0	
98		0.0867	240	0.00148	28	0.00089	28	0.7095	0.97	-0.11	254	3	38	0.011	0	
99		0.0542	220	0.001	42	0.0006	42	0.7048	0.97	0.13	223	3	46	0.013	0	
100		0.0105	220	0.00018	260	0.00011	260	0.7066	0.83		162	3	28	0.016	0	
102		0.158	240	0.00269	24	0.00161	24	0.7071	0.98	-0.21	210	4	9	0.02	0	
103		0.86	210	0.0162	12	0.0098	12	0.7119	1.1	-0.09	178	4		0.022	0	
104		0.0447	240	0.00075	41	0.00046	41	0.7187	1.1	0.11	90	2		0.016	0	
105		0.117	210	0.0022	20	0.00138	21	0.7059	0.96	-0.087	213	3	9	0.015	0	

Appendix D

Geochemical data

The following tables list the geochemical data from samples that have collected during this study. Type corresponds to whether the sample was collected from the basement (bas) or a greenstone belt (gst).

Sample	BL05	BL17	CH01	DP01	GH02	GH03	GW01	GW02	GW03	GW04
Type	bas	bas	bas	bas	bas	bas	bas	bas	bas	bas
<i>wt. %</i>										
SiO ₂	75.94	76	75.54	75.44	72.9	72.66	50.32	63.18	74.79	69.34
TiO ₂	0.14	0.08	0.07	0.12	0.26	0.19	1.36	0.44	0.02	0.42
Al ₂ O ₃	12.07	13.03	13.6	14.72	14.13	13.84	13.05	15.41	13.74	16.67
Fe ₂ O ₃	1.27	0.67	0.67	0.89	2.32	1.73	15.96	5.61	0.59	2.18
MnO	0.02	0.03	0.03	0.03	0.04	0.04	0.21	0.09	0.02	0.02
MgO	0.02	n.d.	0.06	0.21	0.31	0.3	5.1	3.38	n.d.	0.95
CaO	0.58	0.74	0.55	1.02	1.81	1.78	8.16	4.46	0.98	3.58
Na ₂ O	3.69	4.66	4.34	4.21	4.46	4.78	2.54	4	4.84	4.97
K ₂ O	5.46	4.54	4.24	1.89	2.85	3.95	1.49	1.88	4.44	1.08
P ₂ O ₅	0.05	n.d.	0.06	0.05	0.1	0.07	0.1	0.32	0.03	0.09
LOI	0.53	0.11	0.67	1.16	0.6	0.44	0.84	0.92	0.36	0.5
Sum	99.77	99.84	99.84	99.73	99.77	99.77	99.13	99.68	99.82	99.81
<i>ppm</i>										
Li	12	18	37	700	28	33	18	33	1	17
B										
Sc	2	1	2	1	3	2	14	7	1	3
V					13	11	348	62	1	43
Cr					35	62	57	173	32	41
Co					5	5	54	21	3	10
Ni					4	6	59	75	3	10
Cu	3	1	1	3	4	8	308	20	1	9
Zn					58	38	122	85	16	40
Ga										
Ge	2	1	2	2	1	1	3	2	1	1
As	2	1	1	3	1	1	1	1	0	0
Rb	135	150	186	831	47	105	41	26	117	15
Sr	44	86	88	346	286	178	98	259	75	373
Y	28	2	8	11	12	8	29	24	9	8
Zr	184	62	50	76	215	101	99	105	39	155
Nb	12	6	10	5	6	6	318	9	3	4
Mo	1	1	0	0	0	0	9	0	0	0
Sn	8	7	8	7	1	1	3	2	1	1

Continued on next page

Sample	BL05	BL17	CH01	DP01	GH02	GH03	GW01	GW02	GW03	GW04
Cs	2	2	3	451	n.d	n.d	0	n.d	n.d	n.d
Ba	448	292	440	448	737	655	229	433	348	273
La	70	12	11	87	36	23	7	18	8	4
Ce	136	25	22	90	66	42	20	45	16	9
Pr	15	3	2	18	7	5	3	6	2	1
Nd	51	8	8	58	26	17	15	24	5	5
Sm	8	1	2	9	4	3	4	6	1	1
Eu	1	0	0	2	1	1	1	1	0	1
Gd	6	1	2	5	3	2	5	6	1	1
Tb	1	0	0	1	0	0	1	1	0	0
Dy	5	0	1	3	2	2	6	5	1	1
Ho	1	0	0	0	0	0	1	1	0	0
Er	3	0	1	1	1	1	4	3	1	1
Tm	0	0	0	0	0	0	1	0	0	0
Yb	3	0	1	1	1	1	4	2	1	1
Lu	0	0	0	0	0	0	1	0	0	0
Hf	6	2	2	3	5	3	3	3	2	3
Ta	1	1	1	1	0	1	2	1	1	1
W					0	0	1	0	0	0
Pb	25	46	38	36	16	21	6	10	28	7
Th	22	25	12	7	5	8	1	4	5	2
U	3	9	6	2	1	2	2	2	2	1

Sample	GW06	LP02	LP03	LP04	LP10	LP11	MA01	MO01	MO03	MO04
Type	bas	bas	bas	bas	bas	bas	bas	gst	gst	gst
<i>wt.</i>										
<i>%</i>										
SiO ₂	75.28	73.15	69.51	73.86	65.75	66.89	74.68	87.38	54.97	41.4
TiO ₂	0.12	0.26	0.4	0.21	0.53	0.49	0.03	0	0.27	0.42
Al ₂ O ₃	13.99	13.5	14.46	13.87	14.28	14.55	14.62	0.16	4.97	17.85
Fe ₂ O ₃	1.07	2.91	3.06	1.07	5.73	4.87	0.45	9.39	7.04	16.51
MnO	0.01	0.03	0.04	0.01	0.1	0.13	0.01	0.31	1	2.44
MgO	0.23	0.38	1.23	0.42	3.8	1.98	0.2	1.94	10.23	17.51
CaO	1.15	1.25	2.22	1.49	3.03	3.76	0.18	0.78	18.48	3.71
Na ₂ O	3.3	2.98	3.57	3.3	2.77	2.72	7.71	0	0.07	0.03
K ₂ O	4.21	4.92	4.5	5.04	2.33	1.68	1.36	0.01	2.83	0.05
P ₂ O ₅	0.03	0.04	0.18	0.09	0.03	0.08	0.02	0.01	0.01	0.01
LOI	0.45	0.35	0.51	0.35	1.4	1.4	0.59	2.82	0.82	11.35
Sum	99.83	99.76	99.66	99.72	99.74	98.55	99.86	99.98	99.87	99.93
<i>ppm</i>										
Li	2	11	14	8	19	21	3	12	141	228
B								8	36	1
Sc	0	1	5	1	6	9	1	0	4	63
V	2	43	38	13	92		3	2	66	153
Cr	9	48	39	29	59		14	54	744	4539
Co	1	7	10	5	17		3	2	28	913
Ni	1	9	17	10	21		4	32	201	3657
Cu	1	2	10	9	0	104	8	1	0	126
Zn	5	38	59	23	140		11	9	48	123
Ga								1	48	60
Ge	0	2	2	2	2	2	0	1	1	2
As	0	1	2	1	1	1	0	14	67	159
Rb	23	182	104	116	82	109	47	1	212	2
Sr	38	185	370	312	108	120	89	2	27	5
Y	1	19	21	25	10	16	2	1	8	10
Zr	25	103	228	125	172	185	39	1	116	9
Nb	1	16	14	13	7	8	2	n.d	2	2
Mo	0	0	1	0	1	3	0	0	0	0
Sn	0	3	3	2	3	12	0	n.d	5	1
Cs	n.d	n.d	n.d	n.d	n.d	1	n.d	1	43	5
Ba	175	720	1107	1307	291	248	528	5	259	325
La	3	27	79	46	28	24	6	0	6	3
Ce	6	56	142	82	48	43	10	n.d	11	44
Pr	1	6	14	11	5	5	1	0	1	1
Nd	2	22	45	37	17	16	3	n.d	4	7
Sm	0	5	6	7	3	3	1	0	1	2
Eu	0	1	1	1	1	1	0	0	0	1
Gd	0	4	4	5	2	3	0	0	1	2
Tb	0	1	1	1	0	0	0	0	0	0
Dy	0	4	3	4	2	3	0	0	1	3
Ho	0	1	1	1	0	1	0	0	0	1
Er	0	2	2	3	1	2	0	0	1	2
Tm	0	0	0	0	0	0	0	n.d	0	0
Yb	0	2	2	3	1	2	0	n.d	1	2
Lu	0	0	0	0	0	0	0	0	0	0
Hf	1	3	6	4	4	5	1	n.d	3	0
Ta	0	1	2	2	1	1	0	0	0	1
W	0	0	0	0	1		0	0	1	28
Pb	6	42	37	43	28	49	3	0	8	0
Th	2	58	39	57	7	7	2	n.d	5	0
U	2	5	5	7	0	1	1	0	2	2

Sample	MO05	MU01	MU03	NG01	NG02	NG03	NM02	NM03	NY01	NY02
Type	gst	bas	bas	bas	bas	bas	bas	bas	bas	bas
<i>wt.</i>										
<i>%</i>										
SiO ₂	18.64	76	73.12	75.63	77.1	75.32	70.35	48.76	74.81	73.79
TiO ₂	0.26	0.01	0.2	0.11	0.07	0.09	0.39	2.52	0.12	0.22
Al ₂ O ₃	5.11	13.56	14.44	13.38	12.82	13.6	15.29	11.58	13.52	14.03
Fe ₂ O ₃	18.57	0.36	1.85	1.24	0.81	0.93	2.71	21.18	1.1	1.71
MnO	0.84	0.01	0.04	0.05	0.04	0.04	0.03	0.26	0.02	0.03
MgO	10.37	n.d.	0.3	0.15	0.15	0.24	0.8	3.85	0.24	0.35
CaO	45.81	0.81	2.15	0.74	0.6	0.8	2.59	8.47	1.18	1.45
Na ₂ O	0	3.54	5.85	3.68	3.42	3.7	4.85	1.36	3.25	3.58
K ₂ O	0.07	5.13	1.01	4.69	4.39	4.71	1.75	0.81	5.11	4.44
P ₂ O ₅	0.02	n.d.	0.06	0.04	0.03	0.03	0.1	0.13	0.04	0.06
LOI	29.72	0.44	0.81	0.6	0.55	1.1	0.9	0.68	0.52	0.7
Sum	99.69	99.84	99.83	99.71	99.43	99.46	99.76	99.6	99.39	99.66
<i>ppm</i>										
Li	14	2	2	106	234	80	22	10	18	23
B	1			6	6				5	4
Sc	3	0	2	2	2	2	3	30	2	2
V	103	1	15	6	5	5	28	746	6	12
Cr	3463	12	12	39	31	18	20	36	41	26
Co	137	3	6	1	1	3	9	68	2	2
Ni	1749	3	5	2	2	11	10	34	2	2
Cu	57	5	1	2	19	6	4	78	1	5
Zn	62	4	45	40	39	43	54	149	20	44
Ga	41			54	49				84	221
Ge	1	2	1	4	3	2	1	2	3	3
As	185	0	1	1	1	1	1	2	1	1
Rb	1	141	56	477	662	547	23	21	180	107
Sr	16	67	245	47	48	54	401	95	89	144
Y	10	40	5	21	16	22	6	40	4	2
Zr	6	46	149	103	84	100	159	105	139	197
Nb	0	2	3	24	21	17	4	5	6	4
Mo	0	0	0	1	1	1	0	0	1	1
Sn	0	1	1	5	7	5	1	1	1	1
Cs	1	n.d.	n.d.	30	53	55	n.d.	n.d.	1	2
Ba	221	645	351	215	184	228	699	115	397	1169
La	3	7	15	27	20	25	21	5	23	23
Ce	2	13	28	53	39	52	44	15	49	43
Pr	1	1	3	6	5	6	5	2	5	4
Nd	3	4	10	21	16	20	15	13	18	14
Sm	1	2	2	4	3	4	3	4	3	2
Eu	0	0	0	0	0	0	1	1	0	0
Gd	1	3	1	4	3	4	2	5	2	1
Tb	0	1	0	1	1	1	0	1	0	0
Dy	1	5	1	4	3	4	1	6	1	0
Ho	0	1	0	1	1	1	0	1	0	0
Er	1	4	0	2	2	2	1	4	1	0
Tm	0	1	0	0	0	0	0	1	0	0
Yb	1	5	0	2	2	3	1	4	1	0
Lu	0	1	0	0	0	0	0	1	0	0
Hf	0	3	4	4	4	4	4	2	4	6
Ta	0	0	0	5	8	3	1	1	0	1
W	4	0	0	1	1	1	0	0	1	0
Pb	1	25	15	42	41	40	11	4	46	38
Th	0	9	4	24	25	25	5	1	31	23
U	0	3	1	13	22	16	1	0	7	4

Sample	NY03	PG01	SC03	SC04	SG01	SG02	SR01	SR02	TR02	TR04
Type	bas	bas	bas	bas	bas	bas	bas	bas	bas	bas
<i>wt.</i>										
<i>%</i>										
SiO ₂	75.52	73.3	66.75	75.4	73.42	71.97	69.24	68.25	64.26	51.22
TiO ₂	0.15	0.15	0.43	0.04	0.86	0.28	0.38	0.52	0.42	0.72
Al ₂ O ₃	13.28	14.92	15.42	13.46	9.31	14.73	15.7	15.37	15	9.84
Fe ₂ O ₃	0.95	0.94	4.23	0.73	5.67	2.6	3.1	3.91	4.72	13.36
MnO	0.03	0.02	0.05	0.02	0.03	0.02	0.04	0.05	0.09	0.24
MgO	0.17	0.37	2.06	n.d.	2.09	0.69	1.29	1.59	3.42	9.37
CaO	0.77	1.47	4.2	0.7	2	3.28	3.68	3.57	5.78	11.47
Na ₂ O	3.4	4.46	4.61	4.69	2.7	4.6	4.56	3.81	4.33	1.44
K ₂ O	5.08	3.94	1.08	4.57	2.41	1.05	1.58	2.21	0.76	0.56
P ₂ O ₅	0.07	0.07	0.1	n.d.	0.13	0.11	0.15	0.19	0.06	0.08
LOI	0.58	0.41	0.79	0.28	0.82	0.44	0.72	0.87	0.95	1.29
Sum	99.42	99.64	99.72	99.89	99.43	99.78	99.72	99.47	99.78	99.59
<i>ppm</i>										
Li	46	40	12	n.d	33	14	120	346	9	5
B	5	9					7	6		
Sc	2	1	8	2	7	3	3	5	16	7
V	6	11		1			39	59	122	194
Cr	23	20		9			49	70	248	812
Co	1	3		3			8	10	28	66
Ni	1	5		4			15	17	104	258
Cu	5	2	33	1	2	3	1	2	55	45
Zn	26	32		13			52	62	54	109
Ga	60	332					104	115		
Ge	3	3	1	2	3	1	2	4	1	3
As	1	1	0	0	4	1	6	3	0	1
Rb	278	106	37	139	90	41	23	50	37	21
Sr	45	682	436	15	136	209	423	352	234	220
Y	6	3	7	11	13	14	7	8	9	17
Zr	141	106	114	72	1131	220	161	170	61	142
Nb	14	4	4	5	9	7	7	6	2	3
Mo	1	0	0	0	0	0	1	1	0	0
Sn	2	4	7	1	7	7	2	1	1	1
Cs	2	7	1	n.d	4	1	6	8	n.d	n.d
Ba	263	1752	460	52	925	494	488	554	150	126
La	19	14	20	3	248	32	15	24	7	10
Ce	53	29	35	8	440	55	29	49	14	27
Pr	5	3	3	1	44	6	3	5	2	3
Nd	15	13	13	5	141	21	11	20	6	14
Sm	2	2	2	2	17	4	2	3	1	4
Eu	0	1	1	0	2	2	1	1	1	1
Gd	2	1	2	2	10	3	2	3	2	4
Tb	0	0	0	0	1	1	0	0	0	1
Dy	1	1	2	2	4	3	2	2	2	4
Ho	0	0	0	0	1	1	0	0	0	1
Er	1	0	1	1	2	2	1	1	1	2
Tm	0	0	0	0	0	0	0	0	0	0
Yb	1	0	1	1	2	1	1	1	1	2
Lu	0	0	0	0	0	0	0	0	0	0
Hf	5	3	3	3	30	5	4	5	2	3
Ta	1	0	0	1	1	1	1	1	0	0
W	0	0		0			1	0	0	0
Pb	52	67	10	30	13	9	18	25	7	4
Th	29	7	3	7	47	8	5	11	1	1
U	5	2	0	2	3	1	2	3	0	0

Sample	TR06	TR07	TR08	WG01	WG02	WG03	ZB01	ZB02	ZB03	ZB04
Type	bas	bas	bas	bas	bas	bas	bas	bas	bas	bas
<i>wt.</i>										
%										
SiO ₂	66.09	76.79	65.14	53.36	69.44	65.66	72.39	73.6	70	72.06
TiO ₂	0.53	0.01	0.62	0.16	0.36	0.42	0.29	0.21	0.34	0.3
Al ₂ O ₃	15.5	14.51	15.25	2.9	15.68	16.9	15.01	14.76	14.98	15.5
Fe ₂ O ₃	4.21	0.99	4.87	12.86	2.75	2.62	2.07	1.55	3.68	2.1
MnO	0.06	0.04	0.08	0.48	0.04	0.05	0.03	0.02	0.06	0.03
MgO	1.32	0.13	2.23	17.24	1.33	2.19	0.68	0.48	1.67	0.73
CaO	4.23	0.84	6.9	12.3	3.53	2.31	3.03	2.09	3.29	2.75
Na ₂ O	4.68	1.68	3.27	0.37	4.38	6.11	4.43	3.89	3.97	4.75
K ₂ O	1.02	3.23	0.43	0.32	2.31	3.12	1.63	3.34	1.8	1.34
P ₂ O ₅	1.12	0.02	0.07	0.01	0.14	0.15	0.09	0.08	0.1	0.11
LOI	0.97	1.59	0.94	0.93	0.73	2.24	0.53	0.69	1.54	0.96
Sum	99.72	99.84	99.8	100	99.96	99.53	99.65	100.02	99.89	99.67
<i>ppm</i>										
Li	12	1	2	12	23	10	46	113	36	37
B				5	6	6	6	5	7	6
Sc	5	4	6	12	3	3	2	2	9	2
V	56	3	94	65	42	32	23	19	50	30
Cr	42	15	122	1705	49	37	36	46	115	40
Co	13	4	21	43	8	6	5	4	10	5
Ni	14	11	62	1206	16	14	6	4	28	11
Cu	16	3	23	1	3	1	1	1	13	8
Zn	69	4	53	260	43	46	48	34	61	27
Ga				11	142	226	77	119	58	213
Ge	1	2	1	5	2	3	2	2	3	2
As	0	0	0	0	1	1	1	0	1	0
Rb	6	60	10	16	24	21	27	60	16	106
Sr	319	37	145	13	404	186	213	177	235	172
Y	6	26	13	6	6	9	3	1	14	10
Zr	136	30	74	13	145	146	144	121	130	124
Nb	2	0	4	2	6	7	4	3	8	10
Mo	0	1	0	0	1	0	1	4	1	0
Sn	1	0	1	2	1	2	1	0	2	1
Cs	n.d	n.d	n.d	1	1	0	2	10	2	8
Ba	670	215	113	15	714	1179	344	572	232	1116
La	16	14	14	2	11	20	11	4	6	7
Ce	30	21	26	9	26	43	24	9	16	21
Pr	3	2	3	1	3	5	3	1	2	2
Nd	12	5	11	6	13	17	9	3	9	5
Sm	2	1	2	2	3	3	2	0	3	1
Eu	1	0	1	0	1	1	0	0	1	0
Gd	2	1	3	1	2	2	1	0	3	1
Tb	0	0	0	0	0	0	0	0	1	0
Dy	1	3	2	1	2	2	1	0	3	2
Ho	0	1	1	0	0	0	0	0	1	1
Er	1	3	2	1	1	1	1	0	2	2
Tm	0	1	0	0	0	0	0	n.d	0	0
Yb	1	4	1	1	1	1	0	0	2	2
Lu	0	1	0	0	0	0	0	0	0	0
Hf	3	1	2	0	4	4	4	3	4	4
Ta	0	0	0	0	1	1	0	0	1	2
W	0	0	0	0	0	0	1	n.d	0	0
Pb	6	34	5	1	20	12	10	15	12	21
Th	1	2	2	0	7	11	4	3	3	7
U	0	1	0	0	2	2	0	2	1	1

Sample	ZB05	ZB06	ZB07	ZL01	ZL011	ZL015	ZL09	ZL16	ZL20	ZLC06
Type	bas	bas	bas	gst	gst	gst	gst	gst	gst	gst
<i>wt.</i>										
<i>%</i>										
SiO ₂	73.15	73.11	74.57	65.04	50.89	79.28	51.5	48.23	48.64	65.46
TiO ₂	0.21	0.25	0.11	0.61	0.6	0.43	0.53	0.32	0.32	0.13
Al ₂ O ₃	14.49	14.18	15.11	14.54	13.54	13.24	13.12	8.52	8.37	19.56
Fe ₂ O ₃	1.38	1.99	1	6.76	10.65	0.42	10.59	11.88	11.76	0.78
MnO	0.02	0.03	0.01	0.08	0.17	0.02	0.18	0.2	0.21	0.03
MgO	0.5	0.71	0.24	4.13	8.61	0.04	9.05	21.65	20.86	0.29
CaO	1.68	2.13	1.31	2.82	11.65	0.03	10.96	8.59	9.28	1.52
Na ₂ O	4.4	3.57	4.64	4.02	2.21	0.78	1.7	0.26	0.24	4.92
K ₂ O	3.59	3.31	2.86	1.33	0.06	3.67	0.06	0.02	0.03	7.3
P ₂ O ₅	0.09	0.08	0.05	0.15	0.03	0.04	0.01	0.02	0.02	0.03
LOI	0.61	0.8	0.98	1.08	0.98	1.8	1.81	4.63	4.31	1.96
Sum	99.51	99.36	99.9	99.48	99.4	99.75	99.5	99.69	99.73	100.02
<i>ppm</i>										
Li	866	38	48	123	5	21	22	30	53	104
B	7	6	9	6				0	0	5
Sc	2	1	1	7	45	9	44	28	30	2
V	21	8	20	109	229	77	243	146	144	12
Cr	55	44	55	51	452	89	463	3079	2638	5
Co	5	2	4	22	55	5	48	87	78	1
Ni	5	2	8	54	151	9	138	921	777	2
Cu	4	2	6	111	93	1	108	17	38	1
Zn	62	35	44	66	78	4	74	71	74	7
Ga	39	133	160	130				7	7	64
Ge	2	1	3	3	0	2	1	1	1	4
As	0	1	2	1	1	1	1	0	0	1
Rb	27	64	110	44	1	346	3	1	1	653
Sr	196	302	438	296	100	62	106	28	19	52
Y	4	2	4	17	12	8	14	7	7	18
Zr	206	84	120	152	16	122	14	9	11	133
Nb	8	2	5	6	1	4	1	1	1	25
Mo	3	1	1	1	0	0	0	0	0	1
Sn	1	1	1	1	0	3	0	0	0	6
Cs	26	2	6	78	n.d	13	n.d	1	0	33
Ba	123	659	792	673	27	170	18	3	6	272
La	7	4	24	16	1	29	2	1	1	18
Ce	15	9	38	29	4	24	4	2	2	40
Pr	2	1	4	4	1	5	1	0	0	5
Nd	6	3	15	14	3	18	4	1	1	16
Sm	1	1	2	3	1	3	1	1	1	4
Eu	0	0	1	1	0	1	1	0	0	0
Gd	1	1	2	4	1	2	2	1	1	4
Tb	0	0	0	1	0	0	0	0	0	1
Dy	1	0	1	3	2	2	2	1	1	4
Ho	0	0	0	1	0	0	0	0	0	1
Er	1	0	0	2	1	1	1	1	1	3
Tm	0	0	0	0	0	0	0	0	0	0
Yb	1	0	0	2	1	1	1	1	1	3
Lu	0	0	0	0	0	0	0	0	0	0
Hf	5	3	4	4	0	3	0	0	0	5
Ta	1	0	1	1	0	0	0	0	0	6
W	0	0	1	1	0	6	0	0	0	1
Pb	16	17	47	19	1	7	1	0	0	39
Th	4	1	15	4	0	6	0	n.d	n.d	26
U	1	1	5	1	0	2	0	n.d	0	17

Sample	ZLC07	ZLC15	ZLC17/15	ZLC20	ZLC32	ZLD01	ZLD02	ZLD03	ZLD05	ZLD06
Type	gst	gst	gst	gst	gst	bas	bas	gst	gst	gst
<i>wt.</i>										
<i>%</i>										
SiO ₂	66.98	48.76	64.22	64.65	50.08	75.45	76.58	34	43.9	65
TiO ₂	0.01	0.95	0.76	0.67	0.71	0.05	0.03	0.07	0.29	0.55
Al ₂ O ₃	18.81	21.55	14.66	14.84	12.27	14	13.97	0.6	5.74	15.01
Fe ₂ O ₃	0.44	10.03	4.75	6.11	15.57	0.74	0.32	7.58	11.8	6.87
MnO	0.02	0.18	0.11	0.11	0.23	0.02	0.02	0.15	0.19	0.2
MgO	0.08	3.75	2.29	2.92	7.68	0.03	n.d.	39.4	25.78	0.96
CaO	1.26	9.78	5.37	5.57	11.44	0.63	1.05	2.34	5.8	4.69
Na ₂ O	7.85	2.49	1.8	3.54	1.19	3.18	4.55	0.33	0.29	3.3
K ₂ O	4.6	2.08	5.52	1.29	0.44	5.27	3.02	0.01	0.01	2.16
P ₂ O ₅	0.09	0.24	0.06	0.16	0.19	0.02	0.04	n.d.	0.03	0.13
LOI	0.97	1.75	0.67	1.04	0.86	0.4	0.22	14.61	5.34	0.65
Sum	100.14	99.81	99.54	99.86	99.8	99.79	99.79	99.07	99.18	99.51
<i>ppm</i>										
Li	11	287	42	131	156	248	69	n.d	28	352
B	6	6	7	5	1					
Sc	0	6	9	3	19	3	2	5	14	12
V	3	137	79	103	194			15	100	85
Cr	2	98	22	92	282			1021	2276	55
Co	1	31	18	17	52			114	68	23
Ni	1	61	19	57	207			2686	1027	32
Cu	2	51	22	3	50	5	10	4	19	36
Zn	4	164	93	296	106			17	82	67
Ga	22	74	257	88	30					
Ge	4	5	2	5	4	2	2	0	3	1
As	1	1	0	1	1	0	0	5	0	0
Rb	692	100	38	55	29	291	158	0	1	81
Sr	27	294	187	199	197	36	54	68	21	111
Y	8	16	7	17	21	17	18	1	12	11
Zr	57	216	58	195	110	57	66	7	25	143
Nb	14	5	13	7	4	12	9	1	2	5
Mo	12	1	6	1	0	0	5	0	0	2
Sn	1	1	2	1	0	9	7	1	2	1
Cs	39	33	4	10	5	31	35	n.d	n.d	90
Ba	59	299	1401	418	104	223	155	n.d	n.d	326
La	9	25	3	25	22	14	11	0	4	12
Ce	20	52	7	52	45	29	22	0	10	24
Pr	2	6	1	6	5	3	3	0	1	3
Nd	8	24	3	24	22	12	9	n.d	6	10
Sm	2	5	1	5	4	2	2	0	2	2
Eu	0	1	0	1	1	0	0	0	1	1
Gd	2	4	1	4	4	3	3	0	2	2
Tb	0	1	0	1	1	0	0	0	0	0
Dy	2	4	1	4	4	3	3	0	2	2
Ho	0	1	0	1	1	1	1	0	0	0
Er	1	2	1	2	3	2	2	0	1	1
Tm	0	0	0	0	0	0	0	0	0	0
Yb	1	2	2	2	2	2	2	0	1	1
Lu	0	0	0	0	0	0	0	0	0	0
Hf	4	5	2	5	3	2	3	0	1	4
Ta	37	1	1	1	0	2	3	0	0	1
W	0	1	3	3	9			8	0	1
Pb	24	27	49	69	5	37	31	1	1	13
Th	13	5	1	6	4	15	11	0	1	5
U	19	2	1	2	1	5	11	0	0	1

Sample	ZLD07	ZLD08	ZU06
Type	gst	gst	gst
<i>wt.</i>			
<i>%</i>			
SiO ₂	63.51	39.4	74.09
TiO ₂	0.65	0.38	0.31
Al ₂ O ₃	15.66	14.64	14.28
Fe ₂ O ₃	5.81	4.63	1.77
MnO	0.08	0.12	0.02
MgO	1.97	3.87	0.81
CaO	5.21	15.97	1.18
Na ₂ O	3.84	3.83	6.31
K ₂ O	1.47	3.08	0.77
P ₂ O ₅	0.17	0.07	0.08
LOI	0.94	13.33	0.49
Sum	99.33	99.32	99.62
<i>ppm</i>			
Li	347	216	304
B			6
Sc	9	11	3
V	112	125	42
Cr	17	95	13
Co	17	16	3
Ni	14	60	76
Cu	36	2	9
Zn	97	40	30
Ga			30
Ge	1	0	3
As	0	1	1
Rb	85	299	28
Sr	239	146	74
Y	12	12	12
Zr	158	80	164
Nb	6	3	6
Mo	0	0	1
Sn	1	1	1
Cs	58	15	9
Ba	288	524	101
La	13	15	16
Ce	29	29	36
Pr	3	3	4
Nd	13	13	16
Sm	3	2	3
Eu	1	1	1
Gd	3	2	3
Tb	0	0	0
Dy	3	2	2
Ho	1	0	1
Er	2	1	2
Tm	0	0	0
Yb	2	1	2
Lu	0	0	0
Hf	4	2	4
Ta	1	1	1
W	1	3	1
Pb	18	2	20
Th	4	3	10
U	2	1	3

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